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January 23, 2009

Acknowledgements

The Niagara Peninsula Conservation Authority (NPCA) wishes to acknowledge the support of those who helped to prepare this Water Availability Study for the Source Water Protection Tier 1 Water Budget.

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1. INTRODUCTION

1.1 Background and Objectives

The Niagara Peninsula Conservation Authority (NPCA) and AquaResource Inc. have completed this Water Availability Study (WAS) of the Central Welland River (CWR), Big Forks Creek (BFC) and Beaverdams Shriners Creeks (BDSC) as part of the Niagara Peninsula Source Protection Area (NPSP Area) Source Water Protection (SWP) Tier 1 Water Budget.

The NPSP Area Tier 1 Water Budget is funded by the province of Ontario. The study methodology was developed by NPCA in consultation with the Ministry of Natural Resources (MNR), based upon the March 2007 Draft Guidance Module – Water Budget and Water Quantity Risk Assessment.

The purpose of the WAS was to determine the water available for surface water flow, groundwater recharge and evapotranspiration on a monthly basis for the time period 1991 to 2005. This time period was chosen to best suit available datasets (e.g. Statistics Canada) and meet the minimum World Meteorological Organization climate normal criterion of fifteen (15) years.

This report documents the WAS for three (3) Watershed Planning Areas (WSPAs) CWR, BFC and BDSC.

The CWR WSPA is 304 km² and is located entirely within the Regional Municipality of Niagara (Figure 1.1). It is located in the following lower tier municipalities, Township of Wainfleet, Township of West Lincoln, Town of Pelham, City of Welland, City of Port Colborne and a small portion of the Town of Thorold.

The BFC WSPA is 167 km² and is located within Haldimand County and the Township of Wainfleet within the Regional Municipality of Niagara (Figure 1.1).

The BDSC WSPA is 76 km² and is entirely located within the Regional Municipality of Niagara (Figure 1.1). It is located in the following lower tier municipalities, Town of Pelham, City of Thorold, City of Niagara Falls, Town of Niagara-on-the-Lake and the City of St. Catharines.

1.2 Study Team and Approach

AquaResource Inc. was awarded the contract to complete the Water Availability Study and has previously worked with NPCA on the Conceptual Water Budget Report (Franz Environmental Inc. et al., 2007). In late 2007 they also completed a technical memo on NPCA Water Survey of Canada stations regarding Baseflow Separation and Streamflow Recession. AquaResource Inc. is involved in the Source Water Protection Water Budget process at a number of levels, including the development of the Water Budget Guidance Module, ongoing technical support for the Ministry of Natural Resources and completing Tier 1, 2 and 3 Water Budget projects for conservation authorities and municipalities.

Peer review of the WAS project was provided primarily by Robert Muir of Dillon Consulting Limited. Mr. Muir previously assisted NPCA in development of the Tier 1 Water Budget work program. He is a Water Resources Engineer with almost two decades of experience and has provided peer review for the Lake Simcoe Region Conservation Authority Water Budget as well as surface water vulnerability studies for a number of conservation authorities.

NPCA staff from three (3) departments were involved throughout the study. These included Jeff Lee and Geoff Verkade from the Geographic Information Systems (GIS) group, Guangli Zhang from Engineering and Jayme Campbell and Brian Wright from Source Water Protection.

The project approach was designed to take advantage of NPCA's GIS expertise and datasets (e.g. soils, land use and digital elevation model) and NPCA's Engineering Department's experience with HEC-HMS. HEC-HMS is the U.S. Army Corps of Engineers (USACE) Hydrologic Engineering Center Hydrologic Modelling System. This is the current software package utilized by the NPCA Engineering Department for its inhouse floodplain mapping. HEC-GeoHMS was used by NPCA GIS specialists throughout the project to develop the hydrologic modelling inputs for HEC-HMS.

This report describes the work completed as part of the WAS of Central Welland River, Big Forks Creek and Beaverdams Shriners Creek Watershed Plan Areas.

1.3 Project Tasks

The principal objectives of the Water Availability Study are to derive monthly estimates of (i) evapotranspiration, (ii) groundwater recharge and (iii) water available for surface water flow for the time period 1991 to 2005.

The project tasks are:

- Initial parameterization of the numerical models to simulate watershed conditions;
- Calibration of the models to observed surface water flow data (where available) with an emphasis on volumes as opposed to peak flow rates; and
- Continuous model HEC-HMS hourly simulation runs for 1991-2005.

Reporting was completed on the model development, calibration uncertainty and outputs and recommendations for future work.

1.4 Relevant Reference Documents

A variety of previous studies provide details regarding the hydrologic conceptual model; these include, most notably:

- Water Budget Conceptual Understanding for the Niagara Peninsula Source Protection Area (Franz et al, 2007); and
- Baseflow Separation and Streamflow Recession (AquaResource Inc., 2007).

These studies are referenced throughout this report. Additional information was also gathered from the Welland River Watershed Strategy (NPCA, 1999) to assist with the Watershed Characteristics section.

1.5 Document Organization

The sections within the report are organized as follows:

- Chapter 2 Watershed Characteristics;
- Chapter 3 Watershed Modelling; and
- Chapter 4 References.

2. WATERSHED CHARACTERISTICS

2.1 General Description of the Watersheds

The Central Welland River WSPA drains 304 km² and is entirely located within the Regional Municipality of Niagara (Figure 1.1). It includes portions of West Lincoln, Township of Wainfleet, Town of Pelham, City of Welland, City of Port Colborne and a small portion of the City of Thorold. The fifteen (15) subwatersheds include: Black Ash Creek (BAC), Beaver Creek (BCW), Bierderman Drain (BND), Coyle Creek (CC), Draper Creek (DRC), Indian Creek (IC), Lyons Creek Drain (LCD), Little Forks Creek (LFC), Parker Creek (PC), Sucker Creek (SU), Toe Path Drain (TPD), Unnamed Creek (UNC), Welland Canal South (WCS), Welland River (WR) and Welland River between Canal (WRC). The CWR receives inflow from the upstream WSPA Upper Welland River. The portion of the CWR shown on Figure 3.2a discharges to the downstream Lower Welland River WSPA through catchment WRC_W100 (Figure 3.2b). The portion of the CWR shown on Figure 3.2b discharges to the Welland Canal.

The Big Forks Creek WSPA drains 167 km² and includes a portion of Haldimand County and most of the Township of Wainfleet in the Regional Municipality of Niagara (Figure 1.1). The seven (7) subwatersheds include Big Forks Creek (BFC), Beezor Drain (BED), East Kelly Drain (EKD), Ellsworth Drain (ELD), Mill Race Creek (MRC), North Forks Drain (NFD) and Wolf Creek Drain West (WDC) (Figures 3.2a and 3.2b). Big Forks Creek outlets to the Central Welland River WSPA.

The Beaverdams Shriners Creek WSPA drains 76 km² and includes portions of the City of Niagara Falls, City of Thorold and small portions of the Town of Pelham, the Town of Niagara-on-the-Lake and City of St.Catharines, all located within the Regional Municipality of Niagara. The four (4) subwatersheds are Beaver Dam Creek (BRDC), Shriners Creek (SSC), Ten Mile Creek (TMC) and Welland Canal North (WCN) (Figure 3.2b). Generally BDSC WSPA catchments outlet to the Welland Canal. However catchments BRDC_W200 and BRDC_W100 output to Lake Gibson through a siphon and flushing culvert operated and maintained by the St. Lawrence Seaway.

2.2 Climate Setting

The climate of Southern Ontario is characterized as having warm summers, mild winters, a long growing season, and usually reliable rainfall. The climate within southern Ontario differs somewhat from one location to another, and from one year to the next. Spatial variations are caused by the topography and varying exposure to the prevailing winds in relation to the Great Lakes (Schroeter et al, 1998).

According to Brown et al. (1980), CWR, BFC and BDSC are located in the Niagara Fruit Belt climatic region. The stations used for this WAS as shown on Figure 2.1 are Environment Canada Canboro (6131165), Ridgeville (6137161), Welland (6139445), Port Colborne (6136606), Power Glen (6137306) and St.Catharines Airport (6137287). Figures 2.2 and 2.3 show the 1991-2005 mean monthly precipitation and mean monthly temperatures (Schroeter and Associates, 2007) for these stations. Mean monthly

precipitation ranged from a low of 44 mm at the Power Glen Environment Canada station 6137306 in February to a high of 102 mm at Port Colborne Environment Canada station 6136606 in September. The mean annual range in temperature was 26.8 degrees Celsius (°C) as shown on Figure 2.3. Mean temperatures are warmer moving downstream from CWR/BFC to BDSC, likely related to the effect of Lake Ontario and decreasing elevation.

Spatial variations in mean annual snowfall, air temperature and mean annual precipitation across CWR/BFC/BDSC in relation to the entire NPCA jurisdiction are illustrated in Figures 2.4, 2.5 and 2.6 and tabulated in Table 2.1. Annual precipitation and snow in CWR/BFC/BDSC appear to range from 990-870 mm per year and 155 to 115 mm, respectively, on average across the WSPAs. Mean annual temperatures range from 8.4 to 9.3 °C.

Figures 2.7, 2.8 and 2.9 show the annual precipitation, annual snow water equivalent and mean annual temperature for the 1991-2005 period respectively for the six (6) study stations. The total annual precipitation ranged from a 1998 low of 593 mm (at Canboro) to a high of 1296 mm in 1996 (at St. Catharines Airport), over double the amount. On mean the annual precipitation was 923 mm (1991-2005). The amount of snow water equivalent ranged from a low of 62 mm in 1998 (Port Colborne) to a high in 2005 of 235 mm (Welland). Overall 135 mm (15%) of precipitation is delivered as snowfall. The mean annual temperature was coolest in 1996 at 8.0°C and highest in 1998 at 10.9°C.

Brown et al. (1980) previously estimated the regional mean annual actual evapotranspiration between 533-559 mm and mean annual water surplus as about 279 mm.

2.2.1 Net Solar Radiation

Six (6) solar radiation and two (2) sunshine station locations were located in and near NPCA ranging from Buffalo, New York to the Hamilton Royal Botanical Gardens (RBG), shown in Figure 2.1. Annual values of net radiation ranged from 26.29 KW/m² at Niagara Falls, New York in 2004 to 33.89 KW/m² at Hamilton RBG in 1991 (Figure 2.10). Overall all stations had their lowest annual net radiation results in 2004, for the period 1991-2005. A review of the results however indicates sunshine station results tend to be slightly higher than those measuring incoming radiation directly. The greatest monthly variation between station measurements occurs during the summer period (Figure 2.11) and shows an increase in net solar radiation going to the northwest from Buffalo to Hamilton RBG in the July means.

2.3 Topography

The topography of CWR WSPA varies from 258 metres above sea level (m ASL) to 171 m ASL. Higher elevations areas are on the northern margin of the WSPA with the highest in Pelham at the Fonthill Kame-Delta physiographic region (catchment CC_W120 Figure 3.2a). The channel slope is on average 3 % (Table 3.1a) and varies little between different channels with the exception of Little Forks Creek (Figure 2.12a).

The topography of BFC WSPA is fairly flat and varies from 180 to 171 m ASL. The channel slope is on average 1.9 % (Table 3.1b) and varies little between the different channels (Figure 2.12b).

The topography of BDSC WSPA indicates two higher elevation areas creating contrast to the clay plain average elevation of 180 mASL. These are at the western WSPA side at Fonthill (240 mASL) and eastern side at Niagara Falls (202 mASL). The channel slope is on average 3.9 % (Table 3.1b).

2.4 Physiography

The CWR WSPA is located predominantly within the Haldimand Clay Plain physiographic region (Figures 2.13a and 2.13b) with the northeastern portion part of the Fonthill Kame-Delta Complex.

The BFC WSPA is largely part of the Dunnville Sand Plain with the southeastern portion part of the Wainfleet Bog (Figure 2.13a).

The BDSC WSPA crosses a number of physiographic features. In the south the Fonthill Kame-Delta Complex, moving north the Haldimand Clay Plain, the Niagara Escarpment and the Iroquois Sand Plain (Figure 2.13b).

2.5 Soils

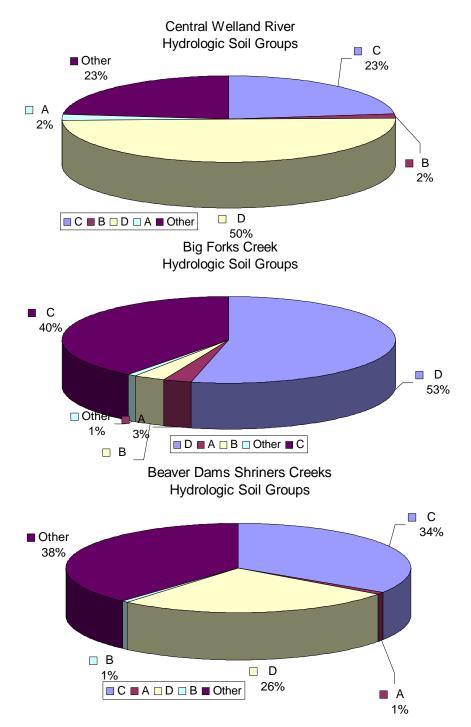
The mapped soils (Figures 2.14a and 2.14b) are classified into four hydrologic soil groups (A, B, C and D) or other. The grouping is according to the soil's minimum infiltration rate, obtained for bare soil after prolonged wetting. According to the United States Department of Agriculture the Hydrologic Soil Groups (HSG) are as follows:

- Group A: sand, loamy sand or sandy loam;
- Group B: silt loam or loam;
- Group C: sandy clay loam; and
- Group D: clay loam, silty clay loam, sandy clay, silty clay or clay.

The category of "other" consists of soils that were not mapped or coincident with an area of high runoff, i.e. urban areas, water bodies, bedrock at surface. These polygons were not assigned HSG values.

The overall dominant soil groups are C and D; CWR 23% and 50%, BFC 40% and 53%, and BDSC 34% and 26%, respectively. The amount of "other" is considerable for CWR (23%) and BDSC (38%). The percentage of "other" is however close to the amount of SOLRIS built-up, transportation and open water areas (CWR 18%, BDSC 35%). The hydrologic soil group data inputs for the model are summarized on Tables 2.2a and 2.2b.

Water Availability Study for the Central Welland River, Big Forks Creek, and Beaverdams Shriners Creeks Watershed Plan Areas - Niagara Peninsula Source Protection Area



2.6 Surficial geology

The surficial geology of the CWR, BFC and BDSC WSPAs (Figures 2.15a and 2.15b) is largely fine-textured glaciolacustrine deposits, matching the overlying clayey soils (Figures 2.14a and 2.14b). However a number of coarse-textured deposits are also mapped, which are somewhat reflected in the soils mapping, i.e. the Dunnville Sand Plain and the Fonthill Kame-Delta Complex. Also a large portion of the surficial geology

adjacent the Welland Canal consists of man-made deposits consisting of overburden excavated for the construction of the canal.

2.7 Land Cover

Land cover information was provided by the Ministry of Natural Resources as part of its Southern Ontario Land Resource Information System (SOLRIS).

Twenty-two (22) to twenty-three categores were provided as shown on Tables 2.3a, 2.3b and 2.3c for CWR, BFC and BDSC WSPAs as shown on Figures 2.16a and 2.16b. Built-up impervious areas account for 14 and 8 % of BDSC and CWR, respectively.

The largest land use categories making up 86% of CWR were (i) rural land use 25%, (ii) monoculture 24%, (iii) swamp 9%, (iv) built-up impervious 8%, (v) mixed agriculture 8%, (vi) deciduous forest 6% and (vi) transportation 5%.

The largest land use categories making up 86% of BFC were (i) monoculture 37%, (ii) rural land use 31%, (iii) swamp 11% and (iv) mixed agriculture 8%.

The largest land use categories making up 85% of BDSC were (i) monoculture 18%, (ii) rural land use 14%, (iii) built-up impervious 14%, (iv) built-up pervious 8%, (v) transportation 7%, (vi) open water 6% (vii) idle land 6% (viii) deciduous forest 6% and (ix) swamp 5%.

2.8 Streamflow

There were no streamflow gauges located within the WSPAs to provide measurements of surface water flow for this particular study. There is a gauge located at the Old Welland Canal Siphon but which only measures water level elevations. Two gauges were previously operated by the Water Survey of Canada (Welland River at Wellandport and Big Forks Creek near Wainfleet) but were not analyzed as their period of record was less than one year (AquaResource, 2007).

Three notable man-made modifications to the Welland River are the (i) Old Welland Canal, (ii) New Welland Canal and (iii) Ontario Power Generation operations at the Niagara River. Two inverted siphons were built to convey the flow of Welland River water beneath the Old and New Welland Ship Canals. These structures flow full under pressure and create backwater pools during floods in a manner similar to dams (NPCA, 1999).

Originally, the Welland River drained directly into the Niagara River at Niagara Falls. However, its flow is now diverted entirely into the Queenston-Chippawa Power Canal. Since 1953, the lower portion of the Welland River now flows in reverse, drawing Niagara River water to the Power Canal. This regulated diversion of water in the lower Welland River creates a pattern of regular diurnal fluctuations in water levels that extend approximately 60 km upstream of the diversion, upstream of the CWR WSPA western boundary (Philips Engineering Ltd., 2004).

Also under normal conditions, the Old Welland Canal provides additional flow to the Welland River at, and immediately downstream of the old siphon. Flow enters the Welland River from the Old Welland Canal through two pathways. First are a series of ports in the roof of the old siphon which allow 14.2 m³/s of higher quality Lake Erie water to dilute the Welland River water. Second is a bypass flow at the Welland Water Treatment Plant which allows 4.5 m³/s of flow (Stantec, 2008). These additional flows enter at catchment WRC_W100 (Figure 3.2b).

3. WATERSHED MODELLING

The following sections describe the construction and calibration of the HEC-HMS models, and present the water balance estimates, as well as the Water Quantity Stress Assessment components.

3.1 Model Description

As outlined in the NPCA WAS Terms of Reference (TOR), HEC-HMS was chosen to model the hydrology of the fourteen (14) Watershed Protection Areas (WSPAs) within the NPCA official boundary. HEC-HMS is a numerical simulation model, supported by the U.S. Army Corps of Engineers, and is designed to simulate the precipitation-runoff processes of a watershed. The program is an integrated work environment, including a database management system, data entry utilities, a computation engine, results reporting tools, and a graphical user interface. A companion product, HEC-GeoHMS, is a software package for use with ArcView, and was employed to develop a significant portion of the required HEC-HMS inputs.

The Tier 1 Water Quantity Stress Assessment is performed on each WSPA; however, the WSPA boundaries are not solely based on surface water drainage patterns. This introduces an inconsistency in the spatial boundaries between WSPAs and HEC-HMS models (which are solely based on surface water drainage boundaries). To address this inconsistency, this report will document the construction of, and results from, two different models for three different WSPAs. The three WSPAs included in this report are: 1) Central Welland River (CWR); 2) Big Forks Creek (BFC); and 3) Beaverdams Shriners Creek (BDSC). The majority of the area for the CWR and BFC WSPAs has been represented within the CWR/BFC model. The second model (known as the BDSC model), includes all of the area for the BDSC WSPA, and ten catchments contained within the CWR/BFC model and the BDSC model. Water balance and Stress Assessment results will be presented in Section 3.5 specific to the three WSPAs, not the two models.

HEC-HMS can be run at a variety of time steps, from 1 minute to 1 day. For the CWR/BFC and BDSC models, and other models created for this study, HEC-HMS was run on the hourly time step.

For complete documentation of the HEC-HMS program, as well as individual hydrologic processes included in HEC-HMS, please refer to the HEC-HMS User Manual and/or Technical Reference Manual (USACE, 2006, 2000).

3.1.1 HEC-HMS Hydrologic Processes

HEC-HMS includes a variety of algorithms for representing the dominant hydrologic processes. This allows the modelling approach to be tailored both to the available data and the overall goals of the study.

The modeler can specify the appropriate algorithm for the following processes:

- Evapotranspiration;
- Snowmelt;
- Loss (infiltration method);
- Baseflow Routing;
- Catchment Hydrograph Transform; and
- Channel Routing.

The algorithms used in the NPCA WAS (specified by the NPCA Water Availability Study Terms of Reference) for each of the six major hydrologic processes are described in the following sections. A conceptualization of the hydrologic processes simulated by HEC-HMS is included in Figure 3.1.

3.1.1.1 Evapotranspiration

The Priestly-Taylor evapotranspiration routine was specified for use in this project. The Priestly-Taylor method relies upon solar radiation and temperature to generate estimates of potential evapotranspiration (PET).

The Priestley Taylor equation is as follows:

$$PET = \alpha \frac{s(\mathcal{I}_{u})}{s(\mathcal{I}_{u}) + \gamma} (K_{u} + L_{u}) \cdot \frac{1}{\rho_{u} \lambda_{v}}$$

Where;

$$\begin{split} K_n &= \text{Short wave radiation} \\ L_n &= \text{Long wave radiation} \\ s(T_a) &= \text{Slope of the saturation-vapour pressure vs. temperature curve} \\ \alpha &= \text{Dryness coefficient} \\ \rho_w &= \text{Mass density of water} \\ \gamma &= \text{Psychrometric constant (ratio of the heat capacity of the air to the latent heat of vaporization)} \\ \lambda_v &= \text{Latent head of vaporization} \end{split}$$

Once the Priestley Taylor PET estimate is generated, HEC-HMS applies crop coefficients to reflect cropping practices or vegetative cover. The crop coefficients are applied as multipliers to scale the Priestley-Taylor PET estimate for that time step.

Evapotranspiration rates are generated by applying the estimated potential evapotranspiration rates to the soil-water reservoir represented within HEC-HMS. Actual evapotranspiration is limited by the amount of water within the soil-water reservoir. When the soil-water reservoir is saturated, actual evapotranspiration is equal to potential evapotranspiration. When the soil-water reservoir is empty (water content is zero), evapotranspiration can no longer be supported bringing the actual evapotranspiration to zero. It remains at zero, until a precipitation event replenishes the soil-water reservoir.

3.1.1.2 Snowmelt

The ability to simulate snow processes is critical to represent the hydrology of cold climate watersheds. The spring snowmelt period (March/April in Southern Ontario) is

the season with the highest typical streamflow, and is also responsible for the majority of streamflow volume. This is also the period of time where saturated soil conditions are common producing groundwater recharge.

HEC-HMS considers snow processes by tracking changes to the snowpack. A snowpack is formed when precipitation occurs and the air temperature is below 0°C. HEC-HMS tracks the accumulation and melt of the snowpack through use of the Temperature Index Method. This method utilizes precipitation and temperature to simulate snow accumulation and melt processes. Water content of the snowpack can be increased by snow or rain falling on the snowpack.

Snowmelt is generated when temperatures rise to the point where there is sufficient energy to transform frozen water into liquid water. The amount of melt experienced by the snowpack is dependent for each degree above the freezing point. Snowmelt is held within the snowpack until the snowpack's point of saturation is reached. When the snowpack becomes saturated (specified by the water capacity of the snowpack), liquid water is then provided to the soil surface as water available for infiltration or runoff.

Sublimation is the direct loss of water from the snowpack to the atmosphere. It is not represented within HEC-HMS. Over the winter season, sublimation can result in a significant loss of water content from the snowpack. Schroeter and Associates have estimated this loss to be 0.33 mm/day (Schroeter and Associates, 2004). This is considered a limitation of the HEC-HMS model, and may lead to an over-estimation of water content held within the snowpack.

For a detailed discussion on the snowmelt processes included in HEC-HMS, please refer to Appendix A for an unpublished presentation provided by the USACE.

3.1.1.3 Loss Method (Infiltration)

The infiltration method, or as HEC-HMS terms it, the "loss method", is responsible for partitioning liquid precipitation into direct overland runoff, evapotranspiration, or percolation. The Deficit and Constant Loss method is utilized for this project, and is carried out on a catchment by catchment basis.

Liquid precipitation that falls as rainfall or snowmelt is input into a storage reservoir. This storage reservoir represents all storage elements within each catchment. This includes, but is not limited to, soil water storage, depression storage, and interception storage. The depth of water held within this element is specified by the user.

Water held within the storage reservoir can be removed by evaporation or by percolation. Evaporation, at the rate estimated by the Priestly-Taylor equation, can remove water held within the storage reservoir. If the storage reservoir is empty, actual evapotranspiration is zero for that time step. Water can also leave the reservoir via percolation, which is determined by the Constant Rate. Percolation can only occur when the storage reservoir is completely saturated, and ceases when the storage reservoir drops below the point of saturation. At this point, evapotranspiration is the sole process that is able to reduce the amount of water held in the storage reservoir. Direct overland runoff is only generated when the storage reservoir is full, and liquid precipitation falls at a rate faster than the Constant Rate.

A limitation of this method is the unlimited acceptance of precipitation into the storage element. Provided there is sufficient storage, the reservoir can accept all precipitation, and produce no runoff or recharge, regardless of the intensity of the event. This can result in an under-prediction of flow, particularly when the reservoir is near empty. The impact of this limitation would be most significant when comparing simulated and observed hydrographs for a particular event. Due to the modelling focus being on regional water budgeting, and not flood flow estimation, it is anticipated that this limitation will be not be a major factor for the purposes of this exercise.

3.1.1.4 Baseflow Method

Once the loss method generates estimates of percolation, this water is passed onto the Baseflow Method for a representation of the subsurface processes (see Figure 3.1). The Baseflow Method selected for this study is the Linear Reservoir Method.

Routing flows through a linear storage element is calculated by the following equations: (Schroeter and Watt, 1980)

$$Q_{t} = C \times Q_{t-1} + (1 - C) \times I_{t-1}$$
$$C = e^{\binom{-dt}{KR}}$$

Where:

 $Q_{t-1}, Q_t = Outflow$ dt = time step KR = recession constant (hr) I = Inflow

The Linear Reservoir method uses two linear reservoirs to model the recession of baseflow after a precipitation event. The first linear reservoir is meant to represent a rapidly responding system, often termed "interflow". Interflow is commonly understood to be subsurface stormflow moving through a shallow unsaturated soil horizon, towards a watercourse (Bedient and Huber, 2002).

The second linear reservoir is meant to represent a slower responding groundwater system, in comparison to the first reservoir. This is the system most commonly associated with baseflow and groundwater recharge.

Previous interpretations of the hydrologic/hydrogeologic system within the NPCA, carried out as part of the Conceptual Water Budget, have indicated that there is very little evidence of a regional groundwater flow system with strong interactions with the surface water system (Franz et al., 2007). The Conceptual Water Budget also stated there was minimal recharge to a deeper regional groundwater system, and that any groundwater discharge that did occur was "fed by localized groundwater recharge, which does not enter the regional aquifer system". This localized groundwater discharge was termed, perhaps mistakenly, "interflow". The term interflow, as it was used in the Conceptual Water Budget, meant to indicate discharge that was not sourced from a larger regional system, but rather from localized, near surface, aquifers. It was not meant to describe the shallow stormflow as described by Bedient and Huber (2002).

For the purposes of this study, flow from the first linear reservoir (interflow) will be considered to be part of the storm response, which travels laterally through the unsaturated soil horizon, before discharging into a watercourse. Flow that enters the second reservoir, and is discharged as baseflow, will represent the amount of water that percolates and reaches the saturated soil layer as groundwater recharge.

The percolation computed from the Deficit and Constant Loss method, is split evenly between both reservoirs. The proportion of water supplied to each reservoir, is specified by the program itself, and can not be modified.

Discharges from both of the linear reservoirs are added with any direct runoff, which create the catchment outflow hydrograph. As this method conserves mass within the catchment, there is no ability to route a portion of baseflow to a downstream catchment, or to remove water from the entirety of the watershed representing "deep recharge". There are two sources of error associated with this limitation. The first source of error is that outflows of groundwater to downstream catchments cannot be represented. This may result in headwater catchments having too much groundwater discharge, with downstream catchments having too little. The error associated with this limitation is inversely proportional to the watershed area. This is due to net groundwater inflows/outflows becoming negligible as the area of interest increases.

The second source of error is that the loss of water to regional groundwater flow systems (removal of water from the watershed) is not able to be represented. By neglecting this loss, other water balance parameters could be over-estimated (ET, runoff, baseflow). Due to the conceptualization reported in the Conceptual Water Budget (Franz et al., 2007) of minimal interaction between the regional groundwater system and the surface water system, it is anticipated this will not be a significant source of error.

3.1.1.5 Catchment Hydrograph Transform

Whereas the Baseflow Method is responsible for the routing of percolated water, the Transform Method is responsible for the routing of overland runoff. For this study, the SCS Unit Hydrograph Method was specified for the transform method.

The SCS Unit Hydrograph Method was originally developed from observed data collected in small, agricultural watersheds. The observed data has been generalized as dimensionless hydrographs, and a best-approximate hydrograph was developed for general application. The SCS method scales the generalized hydrograph by a user specified time lag to produce the unit hydrograph. The time lag is approximated by taking 60% of the time of concentration.

3.1.1.6 Channel Routing

As catchment outflow hydrographs are generated and added to the main channel, the resulting hydrographs must be routed downstream. While HEC-HMS has a number of methods available for routing, the Muskingum-Cunge method has been specified for this study. The Muskingum-Cunge routing method is based on the combination of the conservation of mass and the diffusion representation of the conservation of momentum. It represents the attenuation of flood waves and can be used in river reaches with a small slope.

The attenuation of hydrographs is calculated by specifying the characteristics of the channel. These characteristics include length, slope, Manning's n, and channel geometry (cross section).

3.2 Model Set-up

HEC-HMS requires a number of datasets to represent the hydrology of a watershed. A large portion of the model set-up was completed by NPCA, utilizing both GIS and HEC-GeoHMS processing. The following sections summarize the methodology for the initial parameterization of the HEC-HMS model. For the detailed description on the implementation of HEC-GeoHMS, please refer to Appendix B.

3.2.1 Meteorological information

To properly represent streamflow and significant hydrologic processes, climate and climate variability must be represented within the hydrologic model.

Climate data from six meteorological stations were considered when constructing the CWR/BFC and BDSC HEC-HMS models (refer to Figure 2.1). All of the stations: Port Colborne, Canboro, Welland, Ridgeville, St Catharines Power Glen, and St Catharines Airport, are owned and operated by Environment Canada. As such, they are operated to a national standard, and undergo significant quality assurance/quality control procedures to ensure accurate data collection. The Canboro station was operated by Environment Canada until 1971, after which data was filled in by Schroeter and Associates (2007).

Catchments in the southern portions of the CWR/BFC and BDSC models were assigned climate data from the Port Colborne station, while catchments in the western portion of CWR/BFC model were assigned climate data from the Canboro station. The climate for the central part of CWR/BFC and BDSC was represented by the Ridgeville, Welland and St Catharines Power Glen stations. Below the Escarpment, a single catchment for BDSC was assigned climate data from the St Catharines Airport station.

To ensure each climate station had a complete period of record, each dataset was cleaned up and filled-in by Schroeter and Associates (2007). The in-fill procedure was carried out on both the daily datasets (max/min temperatures, rainfall/snowfall totals), and the hourly rainfall datasets.

With the hourly modelling time interval, hourly data inputs were required. To produce hourly precipitation, daily snowfall depths were evenly distributed throughout the day and added to the hourly rainfall dataset. While it is unlikely that the reported daily snowfall depths are evenly distributed throughout a day, the fact that snowfall does not generate an immediate streamflow response, means the impact of this assumption is negligible.

Synthetic hourly temperatures were generated using the maximum and minimum daily temperatures and a generalized synoptic curve (Schroeter and Associates, 2004).

As discussed in the Upper Welland River WSPA model, the Canboro temperature dataset was found to be significantly warmer than the Hamilton Airport temperature (up to 1.5°C in February), for only the winter months. The differential for warm season months was

minimal. The Canboro dataset was filled in using data from two other stations, Dunnville and Brantford MOE. The latter is located at the Brantford sewage treatment plant (STP), and likely experiences higher ambient temperatures due local heat inputs from the STP. This effect would be largest during the winter months, when the temperature differential between the sewage and local air masses would be the largest. Based on these suspicions, the temperature dataset for the Canboro station was adjusted by lowering temperatures by 0.75°C for the months of December through March.

Hourly net solar radiation stations were created for the CWR/BFC and the BDSC models using a number of datasets. The CWR/BFC station utilized data generally from Environment Canada sunshine station St.Catharines (1990-1994) and Weather Innovations Incorporated station Jordan (1996, 1998-2000, 2003-2005), with gaps filled by Weather Innovation Incorporated stations Vineland (1995), Grimsby (2001-2002) and the Northeast Regional Climate Center station Niagara Falls, New York (1997). The BDSC station utilized data from Environment Canada sunshine station St.Catharines (1990-1994), Weather Innovations Incorporated Virgil (1995-1996) and Northeast Regional Climate Centre Niagara Falls, New York (1997-2005). The incoming solar radiation at the St.Catharines Airport station was calculated using the methodology of Selirio et al. (1971). The overall hourly net radiation was calculated using the methodology of Allen et al. (2005).

3.2.2 Streamflow Information

As described in Section 2.8, there are no active stream gauges in the CWR/BFC and BDSC WSPAs. Outflow from the Upper Welland River model was input as a source to the CWR/BFC model, at the confluence of Welland River and Beaver Creek. Outflow from the CWR/BFC model was input as a source into the BDSC model.

3.2.3 Catchment boundaries and characteristics

General catchment parameters and specifically parameters for the transform and loss methods are shown on Table 3.1. Catchments were delineated by NPCA GIS specialists and AquaResource Inc., using the NPCA 2 m DEM. The catchments ranged in size from 5 to 18 km². Smaller catchments were explored but were not considered possible without the reduction of the model time steps to less than one hour. This constraint is a modelling limitation related to the size of catchment and the model time step within the Transform algorithm (see Section 3.2.8). The model schematics and catchment boundaries are included in Figures 3.2a and 3.2b.

Due to the complicated nature of these WSPAs, the BDSC model includes three (3) outlets. An outlet was created for the Welland River, which flows under the Welland Canal and continues to the Lower Welland River WSPA. This outlet comprises of outflow from the entire CWR/BFC model, combined with flow from catchment CWR_WRC_W100. A second outlet represents flow into Lake Gibson from Beaver Dam Creek (catchments BDSC_BRDC_W100 and W200). The remaining catchments drain to Lake Ontario via the Welland Canal.

3.2.4 Initial Parameterization – Loss Method

The Loss Method relies on three parameters to determine the amount of water that infiltrates, or is available to become overland runoff. These parameters are the constant rate, the catchment storage capacity, and the percentage of impervious cover.

The Deficit and Constant Loss Method assumes that the soil has a constant infiltration rate approximated by the saturated soil hydraulic conductivity. Using the soil and water holding capacity information in Appendix C, average maximum infiltration rates were assigned to each polygon in the soil layer based on their soil type. The catchment average constant rate was determined by area weighting each of the soil polygons in the specific catchment.

HEC-HMS assumes that the soil has a fixed water holding capacity, based on the active rooting depth of vegetation and soil type. The soil water holding capacity layer was built by intersecting the SOLRIS land cover and the OMAF soils layers and by assigning soil water holding capacity values from Appendix C to each unique combination of land cover class and soil type. Like the methodology employed for the constant rate, the area weighted average for each polygon within the catchment, was used to calculate the catchment average.

HEC-HMS considers an impervious surface as an area in a watershed for which all contributing precipitation runs off, with no infiltration, no evaporation, and no other volume losses. This surface was built by assuming SOLRIS built-up impervious and transportation polygons were 100% impervious, with built-up pervious polygons being 50% impervious. All other polygons were assumed to have zero impervious cover.

3.2.5 Initial Parameterization – Evapotranspiration

In the Deficit and Constant Loss Method, water is removed from the soil to simulate evapotranspiration. Potential evapotranspiration is calculated through use of the Priestly-Taylor method. This method uses a crop coefficient, K_c, indicating the ratio of crop potential and short grass reference evapotranspiration. Daily land use layers were created and assigned crop coefficients from Appendix D.

Combined with the solar radiation and temperature datasets, outlined in Section 3.2.1, and the crop coefficients, HEC-HMS calculates the potential evapotranspiration for each time step. This potential evapotranspiration value is then applied to the catchment storage reservoir to generate actual evapotranspiration.

3.2.6 Initial Parameterization – Snowmelt

The following parameters are required to represent snowmelt. These generalized parameters are referenced from the HEC-HMS User Manual.

- Temperature at which precipitation falls as snow;
- Temperature at which the snowpack begins to melt;
- Water capacity of the snowpack;

- Amount of melt that occurs due to heat transfer from the underlying ground;
- Rate at which snow melts when rain occurs;
- Rate at which snow melts during rainfall-free periods.

3.2.7 Initial Parameterization – Baseflow

Once water percolates through the soil column, HEC-HMS routes this water back to the stream as interflow or baseflow. The Linear Reservoir Method, specified for use by the WAS TOR, approximates the discharge by use of a linear reservoir. Groundwater recession constants, estimated via streamflow analysis, represent the reservoir response time and are used as the reservoir constant (also called the time constant) for the linear reservoir in each layer. There are two linear reservoirs that can be represented within HEC-HMS.

The first linear reservoir was parameterized with the intent to represent interflow. A groundwater coefficient of 18 hours was assigned to this reservoir for all catchments, as this was used in the final calibration parameters for Twenty Mile Creek and Upper Welland River.

The 2nd linear reservoir, meant to represent groundwater discharge to the watercourse, was parameterized based on streamflow recession analysis completed by AquaResource Inc. (2007). The streamflow recession analysis estimated the reservoir constant for nine streamflow gauges located within the NPCA. The median reservoir constants from the 1991-2005 period from two of these stations, Twenty Mile Creek at Smithville and Twelve Mile Creek near Power Glen, were assigned to the 2nd linear reservoir within the CWR/BFC and BDSC models. Four catchments located in Big Forks Creek_(BFC-W300, EKD-W100, ELD-W100, and NFD-W100) have underlying geology consisting primarily of sand. The streamflow recession analysis of Twelve Mile Creek near Power Glen reflects similar geologic conditions due to the presence of the Fonthill Kame. Thus, the reservoir constant for this station (708 hours) was applied to these four catchments in Big Forks Creek. The remaining catchments in the CWR/BFC model and all catchments in the BDSC model were parameterized with the recession constant from Twenty Mile Creek at Smithville (278 hours), similar to the Upper Welland River model catchments.

3.2.8 Initial Parameterization – Transform

The lag time associated with the SCS transform method is a function of the Soil Conservation Service (SCS) Curve Number (Figures 3.3a and 3.3b), the hydraulic length, and the catchment slope. This time lag is used to produce the unit hydrograph that allows precipitation excess (precipitation-infiltration) to be transformed into an overland runoff hydrograph. For adequate definition of the unit hydrograph ordinates, a modelling time step that is less than 29% of the time lag must be used. This constraint effectively places a minimum size requirement on the catchments represented within the model.

Curve Number (CN) values are used in the calculation of CN lag time for the SCS Unit Transform Method. The factors influencing CN values are land cover type, soil type and

Antecedent Soil Moisture Condition (AMC). AMC is an estimate of soil water content prior to the beginning of the simulation period, and has 3 levels:

- AMC I reflects soils that are dry but with water content not below the wilting point.
- AMC II reflects soils having average soil water content, and
- AMC III reflects soils that have experienced rainfall in the five days previous to the simulation period.

CN values in the study area were assumed to reflect average soil water content (AMC II). The CN layer was built by intersecting the SOLRIS land cover and OMAF soil layer and by assigning CN values from Appendix C to each unique combination of land use class and soil type. Built-up impervious, built-up pervious and transportation SOLRIS polygons were considered under the impervious surface data field and not assigned CN values.

3.2.9 Initial Parameterization – Routing

To simulate the effects of channel geometry on hydrograph shape, the traditional Muskingum-Cunge Routing Method was used assuming trapezoidal channel geometry. The following inputs are required:

- Channel Bottom Width. The channel width for each of the routing reaches was estimated by digitizing cross sections. This channel width estimation assumed that the water surface width on digital air photos approximated the width of the channel bed.
- Channel Side Slope. The channel side slope was approximated by digitizing two points at the end of each digitized channel width cross sections using a 2m resolution DEM as a guide. Slope values were extracted at the location where the points intersected a slope grid.
- Channel Manning's Roughness Coefficient. Appropriate Manning's roughness coefficients were assigned (Appendix C) to channel routing reaches based on a visual stream bed condition assessment of 10-20cm resolution digital air photos.

3.3 Model Calibration/Verification

3.3.1 Overview of Procedures

Due to the lack of observed streamflow data for the CWR, BFC and BDSC WSPAs, the calibration and verification process undertaken for the Upper Welland River and Twenty Mile Creek models could not be completed. Instead, the parameters within the CWR/BFC and BDSC models were modified using the parameter adjustments carried out for the Upper Welland River and Twenty Mile Creek models. This methodology assumes that adjustments to model parameters, which result in an acceptable calibration, are transferable between WSPAs. This assumption is validated by the fact that the adjustments required for both Upper Welland and Twenty Mile Creek were extremely

similar, which suggests that these adjustments were regional in nature. Given the geologic homogeneity of the NPCA, this is to be expected.

To reduce the reliance on the user specified initial conditions, and to allow the model to "self-initialize", the modelling period was extended to 1990. No data from the 1990 extension was included in the final results.

3.3.2 Model Parameters

As described in the Twenty Mile and Upper Welland modelling reports, parameters related to the Constant Rate and Maximum Storage values in the Deficit and Constant Loss method were adjusted. These values affected the amount of overland runoff, baseflow and interflow, and the amount of evapotranspiration. The modelled values of the Maximum Storage and the Constant Rate for each catchment are shown in Table 3.2. These values provided the base case for the sensitivity analysis which is documented in Section 3.4.

Model	Catchment	Constant Rate (mm/h)	Max Storage (mm)
CWR/BFC	BFC_BED_W100	0.212	166
	BFC_BFC_W100	0.126	163
	BFC_BFC_W110	0.101	150
	BFC_BFC_W200	0.261	161
	BFC_BFC_W300	0.345	150
	BFC_EKD_W100	0.374	150
	BFC_ELD_W100	0.386	162
	BFC_MRC_W100	0.084	150
	BFC_MRC_W110	0.085	170
	BFC_MRC_W111	0.103	164
	BFC_MRC_W120	0.112	150
	BFC_MRC_W200	0.140	150
	BFC_MRC_W210	0.118	150
	BFC_MRC_W300	0.295	150
	BFC_MRC_W310	0.255	150
	BFC_NFD_W100	0.344	150
	BFC_WDC_W100	0.189	164
	CWR_BAC_W100	0.156	162
	CWR_BCW_W100	0.179	161
	CWR_BCW_W200	0.161	150
	CWR_CC_W100	0.318	160
	CWR_CC_W110	0.293	161
	CWR_CC_W120	0.497	150
	CWR_CC_W130	0.436	150
	CWR_CC_W200	0.201	150
	CWR_DRC_W100	0.126	150
	CWR_LFC_W100	0.167	163
	CWR_PC_W100	0.151	161

 Table 3.2 - Calibrated Constant Rate and Maximum Storage Terms

Model	Catchment	Constant Rate (mm/h)	Max Storage (mm)
	CWR_SU_W100	0.168	162
	CWR_UNC_W100	0.174	150
	CWR_UNC_W110	0.183	162
	CWR_UNC_W200	0.148	162
	CWR_WR_W100	0.043	150
	CWR_WR_W110	0.057	150
	CWR_WR_W200	0.144	150
	CWR_WR_W300	0.173	150
	CWR_WR_W310	0.153	150
	CWR_WR_W320	0.128	150
	CWR_WR_W400	0.202	150
	CWR_BND_W100	0.127	163
	CWR_BND_W200	0.147	200
	CWR_IC_W100	0.149	200
	CWR_LCD_W100	0.043	129
	CWR_TPD_W100	0.195	135
	CWR_WCS_W100	0.047	125
	CWR_WCS_W200	0.037	125
	CWR_WCS_W300	0.061	125
	CWR_WCS_W310	0.031	125
BDSC	CWR_WRC_W100	0.057	125
DDSC	BDSC_BRDC_W100	0.157	134
	BDSC_BRDC_W200	0.184	170
	BDSC_SSC_W100	0.213	177
	BDSC_SSC_W200	0.089	125
	BDSC_TMC_W100	0.171	139
	BDSC_WCN_W100	0.079	125
	BDSC_WCN_W200	0.062	125
	BDSC_WCN_W300	0.115	125
	BDSC_WCN_W310	0.160	200
	BDSC_WCN_W320	0.224	171

As per previous models, crop coefficients were modified to adjust the intensity of evapotranspiration applied to the storage element within the Deficit and Constant Loss Method. These modifications are a means to adjust evapotranspiration to account for issues with temperature data, solar radiation, the potential evapotranspiration method itself, or the lack of a sublimation process. They are not due to the crop coefficients being non-representative of their respective catchments. All month's crop coefficients were increased, with the exception of October, which was slightly lowered from the original estimate. As the crop coefficients are direct multipliers to the potential evapotranspiration method, increasing the crop coefficients resulted in an increase in evapotranspiration, with a corresponding decrease in streamflow (runoff, interflow and baseflow). Table 3.3 displays the final adjustments applied to the original crop coefficients.

Month	Crop Coefficient Adjustment
January	12.42
February	7.45
March	2.48
April	1.38
May	1.38
June	1.21
July	1.38
August	1.38
September	1.38
October	0.92
November	4.97
December	6.21

Table 3.3 – Monthly Crop Coefficient Adjustments

The groundwater coefficients with the Linear Reservoir Baseflow Method were not adjusted, but are presented in Table 3.4 below. While these are simply routing parameters, and are not used in partitioning precipitation, they are important to properly represent how infiltrated water is returning to the watercourse.

 Table 3.4 - Groundwater Coefficients in Linear Reservoir Baseflow Model

Catchment ID	GW 1 Coefficient GW 2 Coefficient				
	(hr)	(hr)			
BFC_BFC_W300					
BFC_EKD_W100	18	708			
BFC_ELD_W100	10	708			
BFC_NFD_W100					
All Others	18	278			

3.4 Model Sensitivity

A sensitivity analysis was carried out on both models to determine the model sensitivity to variations in hydrologic parameters.

Previous HEC-HMS studies have shown that simulated streamflow is most sensitive to two parameters. They are: (1) the maximum infiltration rate (equivalent to the Constant Loss in the Deficit and Constant Loss Method), and (2) the water content available for evapotranspiration (Deficit term in the Deficit and Constant Loss Method) (Fleming and Neary, 2004).

Based on this finding from Fleming and Neary, four scenarios were tested to judge the sensitivity of model outputs to variations in the Constant Rate and Maximum Storage terms, included in the Deficit and Constant Loss Method. It is recognized that many other parameters and inputs can have an impact on simulated streamflow (snowmelt parameters, temperature, crop coefficients, precipitation, baseflow recession constants, etc.). However, due to the constraints in the scope of this project, only a limited sensitivity analysis was possible.

Both the Constant Rate and the Maximum Storage were varied by $\pm 25\%$ independently, resulting in the four scenarios. Changes in total outflow, mean evapotranspiration, runoff and recharge were calculated and tabulated in the following tables:

- Tables 3.5 and 3.6 list the percent change in total outflow for each scenario, over the base case for the CWR/BFC model and BDSC model, respectively.
- Tables 3.7 and 3.8 display the percent change in total outflow, evapotranspiration, runoff and recharge for each scenario, over the base case for the CWR/BFC model and BDSC model, respectively.

Month	Constant Rate +25%	Constant Rate -25%	Max Storage +25%	Max Storage -25%
Jan	0.3%	-0.3%	-5.2%	2.4%
Feb	-0.1%	0.1%	-3.2%	1.4%
Mar	-0.1%	0.1%	-0.4%	0.5%
Apr	0.4%	-0.4%	-0.5%	0.1%
May	0.8%	-0.9%	-0.1%	0.0%
Jun	0.7%	-0.8%	0.0%	0.0%
Jul	0.7%	-0.8%	0.0%	0.0%
Aug	0.2%	-0.2%	0.0%	0.0%
Sep	0.0%	-0.1%	-2.7%	19.0%
Oct	-0.2%	0.2%	-16.4%	47.5%
Nov	-0.3%	0.3%	-19.1%	20.2%
Dec	0.0%	0.1%	-9.6%	8.0%

 Table 3.5 - Sensitivity Analysis Results – Change in Outflow in the Central Welland River/Big Forks

 Creek Model

 Table 3.6 - Sensitivity Analysis Results – Change in Outflow in the Beaverdams Shriners Creeks

 Model

Month Constant Rate +25%		Constant Rate -25%	Max Storage +25%	Max Storage -25%
Jan	0.5%	-0.5%	-8.1%	3.9%
Feb	0.4%	-0.4%	-6.2%	4.5%
Mar	0.1%	-0.1%	-1.9%	0.7%
Apr	1.0%	-1.0%	-0.4%	0.0%
May	0.7%	-0.8%	0.0%	0.0%
Jun	0.4%	-0.4%	0.0%	0.0%
Jul	0.2%	-0.2%	0.0%	0.0%
Aug	0.1%	-0.1%	0.0%	0.7%
Sep	0.0%	0.0%	-5.4%	13.5%
Oct	0.1%	-0.1%	-13.6%	27.8%
Nov	0.0%	0.0%	-19.9%	26.8%
Dec	0.3%	-0.3%	-16.4%	12.9%

Table 3.7- Sensitivity Analysis Results - Change in Water Balance Estimates in the Central Welland
River/Big Forks Creek Model

Scenario	ЕТ	Baseflow	Interflow	Runoff	
1: Constant Rate +25%	0.0%	15.8%	15.8%	-5.1%	
2: Constant Rate -25%	0.0%	-18.2%	-18.2%	5.9%	
3: Max Storage +25%	5.0%	-12.4%	-12.4%	-4.8%	
4: Max Storage -25%	-6.3%	9.8%	9.8%	9.7%	

Scenario	ET	Baseflow	Interflow	Runoff
1: Constant Rate +25%	0.0%	17.9%	17.9%	-2.6%
2: Constant Rate -25%	0.0%	-19.7%	-19.7%	2.8%
3: Max Storage +25%	5.1%	-11.4%	-11.4%	-5.7%
4: Max Storage -25%	-5.5%	10.0%	10.0%	6.4%

 Table 3.8- Sensitivity Analysis Results – Change in Water Balance Estimates in the Beaverdams

 Shriners Creeks Model

As is shown by Tables 3.5 to 3.8, variations in the Constant Rate did not significantly affect overall streamflow volume, but did cause significant changes in water balance estimates. Given that the Constant Rate controls the drainage of the storage reservoir (when fully saturated) to the groundwater reservoirs, increasing the Constant Rate results in an increase in both baseflow and interflow, with a corresponding decrease in runoff. Alternatively, decreasing the Constant Rate, has the effect of increasing runoff, and decreasing baseflow and interflow. Due to the Constant Rate not impacting the amount of water that can be held in storage, evapotranspiration is not affected. Figures 3.4 and 3.5 illustrate the percent change in the mean monthly outflow of the model with a 25% increase and decrease in the Constant Rate for the CWR/BFC and BDSC models, respectively. The dotted line at $\pm 10\%$ represents the uncertainty associated with streamflow estimates (Winter, 1981). As shown in the figures, the percent change for both variations in the Constant Rate, is well within these boundaries, which suggests that estimated streamflow is insensitive to changes in the Constant Rate.

Since percolation and runoff only occur when the storage reservoir is full (i.e. when the soil is saturated), increasing the Maximum Storage results in decreases in baseflow, interflow, and runoff. Actual evapotranspiration increases due to a higher volume of water being held in the storage element. A decrease in the Maximum Storage has the reverse effect: increasing baseflow, interflow and runoff and decreasing evapotranspiration, as less water is required to reach the storage reservoir's point of saturation (refer to Tables 3.7 and 3.8). As illustrated in Figures 3.6 and 3.7 and Tables 3.5 and 3.6, the model outflow is highly sensitive to variations in Maximum Storage in the fall and early winter months, but insensitive to these variations in the spring and summer. This is due to the storage element either being completely empty (summer) or completely full (spring) during these seasons, regardless of the size of the storage element. Very large variations in Maximum Storage would be required to change streamflow during these seasons. Flows during the fall season do exhibit sensitivity to variations in the Maximum Storage term. This is due to the storage reservoir being filled during these months. A smaller storage reservoir would cause the storage reservoir to be filled quicker, resulting in more volume directed to percolation and direct overland runoff. Increases in the storage reservoir will yield the opposite effect: lowered flows, runoff, and percolation.

It is important to note that variations in the Constant Rate and Maximum Storage term, do not impact flows during the summer months. This suggests that uncertainty with these terms will not likely add significant levels of uncertainty to the Tier 1 Surface Water Stress Assessment.

The results of the sensitivity analysis suggest that the model solutions for CWR/BFC and BDSC are non-unique, particularly with respect to the Constant Rate. In a non-unique solution, it is possible to calibrate the model to streamflow volumes and obtain a good fit with a number of differing sets of parameters. Frequently with non-unique solutions it is likely that compensating errors are present; whereby the model is simulating the correct streamflow, but incorrectly replicating the underlying physical processes.

In the case of CWR/BFC and BDSC, the Constant Rate can vary by as much as 25%, with a negligible change in streamflow volume. While streamflow is not sensitive to the Constant Rate variation, there is a significant impact on the water balance parameters estimated by the model (+20% baseflow). Water balance estimates (runoff, baseflow) therefore have a greater degree of uncertainty than the streamflow estimates.

To reduce the level of uncertainty, it is recommended that a more detailed Loss Method, such as the Soil Moisture Accounting Method, be tested on a WSPA to validate the water balance estimates made via the Deficit and Constant Loss Method. Such a test would preferably be carried out within a WSPA with observed streamflow data. The modular approach of HEC-HMS would easily facilitate replacing the Deficit and Constant Loss method, with the Soil Moisture Accounting method. Should the more detailed Soil Moisture Accounting Method generate water balance estimates similar to the Deficit and Constant Loss, a higher level of certainty could be attached to estimates generated for other WSPAs (including CWR/BFC/BDSC). Additionally, the Soil Moisture Accounting Loss method allows the modeller to account for the proportion of percolated water that is lost from the surface water system as "deep recharge", a key limitation of the Deficit and Constant Loss method identified in Section 3.1.1.4.

3.5 Results and Discussion

3.5.1 Water Balance Results

HEC-HMS outputs a number of water balance parameters at the catchment level. These include, but are not limited to: total flow, runoff, percolated water, evapotranspiration, snow water equivalent and hydrographs at catchment or reach junctions. These values are output to a HEC DSS file at an hourly time step.

Output from HEC-HMS is summarized in Table 3.9 for each WSPA, presenting the mean annual water balance on a catchment basis and an overall WSPA basis for the 1991-2005 time period. The water balance terms are defined below:

- Precipitation Climate data used to represent the precipitation over each of the catchments is summarized by HEC-HMS and is presented here.
- ET Estimated actual evapotranspiration.
- Interflow Outflow from 1st linear reservoir (half of percolated water); percolated water which moves laterally through the unsaturated soil horizon.
- Baseflow Outflow from 2nd linear reservoir (half of percolated water); slow responding groundwater system. Consists of water which reaches the saturated soil zone.

- Overland Runoff Depth of water that does not infiltrate, and reaches the surface water system via overland runoff.
- Total Outflow Total annual outflow from the catchment; is the sum of Baseflow, Interflow and Runoff.

WSPA	Catchment	Precipitation	ET	Interflow	Baseflow	Runoff	Outflow
	ID	(mm)	(mm)	(mm)	(mm)	(mm)	(mm)
	CWR BAC W100	894	594	36	36	227	299
	CWR BCW W100	894	593	40	40	220	300
	CWR BCW W200	894	579	39	39	237	314
	CWR CC W100	917	573	60	60	222	343
	CWR CC W110	917	532	56	56	272	384
	CWR CC W120	917	514	81	81	239	401
	CWR CC W130	917	547	74	74	220	369
	CWR_CC_W200	917	531	46	46	292	384
	CWR_DRC_W100	968	353	35	35	544	615
	CWR_LFC_W100	917	595	38	38	244	320
	CWR_PC_W100	894	602	35	35	221	291
0	CWR_SU_W100	917	595	38	38	244	321
Ē	CWR_UNC_W100	894	559	41	41	251	334
TR	CWR_UNC_W110	894	606	41	41	206	287
AL	CWR_UNC_W200	894	595	35	35	229	299
, W	CWR_WR_W100	968	204	13	13	737	764
EL	CWR_WR_W110	968	524	20	20	404	444
CENTRAL WELLAND RIVER	CWR_WR_W200	968	569	43	43	313	399
ND	CWR_WR_W300	917	570	41	41	264	345
RI	CWR_WR_W310	917	575	38	38	264	340
VE	CWR_WR_W320	917	583	33	33	267	332
R	CWR_WR_W400	917	586	45	45	240	330
	CWR_BND_W100	968	597	36	36	298	370
	CWR_BND_W200	968	676	34	34	223	291
	CWR_IC_W100	968	663	35	35	234	304
	CWR_LCD_W100	968	569	17	17	365	399
	CWR_TPD_W100	968	469	49	49	401	499
	CWR_WCS_W100	968	333	14	14	608	635
	CWR_WCS_W200	968	444	13	13	499	524
	CWR_WCS_W300	968	410	17	17	523	557
	CWR_WCS_W310	968	498	12	12	446	470
	CWR_WRC_W100	968	517	17	17	416	451
	Overall CWR WSPA	934	542	36	36	319	391
В	BFC_BED_W100	917	599	45	45	226	316
CFIG	BFC_BFC_W100	917	588	31	31	266	328
G FORH CREEK	BFC_BFC_W110	968	582	34	34	317	386
BIG FORKS CREEK	BFC_BFC_W200	894	566	52	52	222	327
S	BFC_BFC_W300	894	561	64	64	203	332

 Table 3.9 - Summary of Water Balance Model Results

	BFC_EKD_W100	894	565	67	67	193	328
	BFC_ELD_W100	894	568	66	66	193	325
	BFC_MRC_W100	968	571	29	29	337	396
	BFC_MRC_W110	968	623	27	27	291	344
	BFC_MRC_W111	968	595	31	31	310	373
	BFC_MRC_W120	968	575	37	37	319	393
	BFC_MRC_W200	968	562	44	44	318	405
	BFC_MRC_W210	968	578	38	38	313	390
	BFC_MRC_W300	968	579	74	74	241	388
	BFC_MRC_W310	968	578	66	66	257	389
	BFC_NFD_W100	894	570	63	63	196	323
	BFC_WDC_W100	894	588	42	42	221	305
		004		40	40	0.50	254
	Overall BFC WSPA	934	577	48	48	259	356
	Overall BFC WSPA BDSC_BRDC_W100	934 866	5 77 518	48 29	48 29	2 59 288	356 347
BE SH	BDSC_BRDC_W100	866	518	29	29	288	347
BEAV SHRI	BDSC_BRDC_W100 BDSC_BRDC_W200	866 866	518 609	29 31	29 31	288 196	347 258
BEAVER SHRINE	BDSC_BRDC_W100 BDSC_BRDC_W200 BDSC_SSC_W100	866 866 866	518 609 620	29 31 33	29 31 33	288 196 180	347 258 247
BEAVERDA SHRINERS	BDSC_BRDC_W100 BDSC_BRDC_W200 BDSC_SSC_W100 BDSC_SSC_W200	866 866 866 866	518 609 620 191	29 31 33 23	29 31 33 23	288 196 180 627	347 258 247 673
	BDSC_BRDC_W100BDSC_BRDC_W200BDSC_SSC_W100BDSC_SSC_W200BDSC_TMC_W100	866 866 866 866 866	518 609 620 191 522	29 31 33 23 38	29 31 33 23 38	288 196 180 627 265	347 258 247 673 342
AMS	BDSC_BRDC_W100BDSC_BRDC_W200BDSC_SSC_W100BDSC_SSC_W200BDSC_TMC_W100BDSC_WCN_W100	866 866 866 866 866 876	518 609 620 191 522 588	29 31 33 23 38 13	29 31 33 23 38 13	288 196 180 627 265 263	347 258 247 673 342 289
BEAVERDAMS AND SHRINERS CREEKS	BDSC_BRDC_W100BDSC_BRDC_W200BDSC_SSC_W100BDSC_SSC_W200BDSC_TMC_W100BDSC_WCN_W100BDSC_WCN_W200	866 866 866 866 866 876 866	518 609 620 191 522 588 562	29 31 33 23 38 13 13	29 31 33 23 38 13 13 13	288 196 180 627 265 263 275	347 258 247 673 342 289 302
AMS	BDSC_BRDC_W100BDSC_BRDC_W200BDSC_SSC_W100BDSC_SSC_W200BDSC_TMC_W100BDSC_WCN_W100BDSC_WCN_W200BDSC_WCN_W300	866 866 866 866 866 866 866 866 866 866 866	518 609 620 191 522 588 562 580	29 31 33 23 38 13 13 24	29 31 33 23 38 13 13 24	288 196 180 627 265 263 275 235	347 258 247 673 342 289 302 284

The estimated values for evapotranspiration, baseflow and interflow are very similar for most of the catchments. This is to be expected due to the homogeneity of geologic conditions found within the WSPAs. The standard deviation and the corresponding percent of mean annual precipitation for the range of baseflow and runoff estimates are shown in Table 3.10 for each WSPA. The deviations in baseflow estimates represent approximately 2% of annual precipitation for each respected WSPA. The standard deviations for the range of direct overland runoff estimates are higher in CWR and BDSC (14% of annual precipitation) than BFC (5% of precipitation), and are almost entirely related to catchments with high impervious percentages.

The largest differences in baseflow for CWR are shown in two Coyle Creek catchments (CC_W120 and CC_W130), which are attributed to a higher constant rate due to the presence of the Fonthill Kame. In BFC WSPA, four catchments previously discussed as having underlying sand geology (BFC-W300, EKD-W100, ELD-W100, and NFD-W100) also have higher baseflow estimates. In BDSC WSPA, the largest differences are two catchments (WCN_W100 and WCN_W200) which have low constant rates due to the Haldimand Clay Plain present below the Escarpment and thus have lower baseflow estimates.

The largest deviations in runoff are shown in catchments which contain significant portions of impervious areas: CWR_WR_W100 (42%), CWR_DRC_W100 (29%), CWR_WCS_W100 (33%), and BDSC_SCC_W200 (44%). Evapotranspiration for these

catchments is lower than the mean due to a smaller proportion of the catchment being pervious and able to support evapotranspiration. Typical with impervious areas, runoff is higher within urbanized catchments.

	BASEFLOW		RUNOFF	
WSPA	Standard Deviation (mm)	Percent of Precipitation	Standard Deviation (mm)	Percent of Precipitation
CWR	16	1.7%	132	14.1%
BFC	15	1.6%	49	5.3%
BDSC	22	2.5%	123	13.9%

 Table 3.10 – Standard Deviation and Percent Precipitation for Baseflow and Runoff Estimates for CWR, BFC and BDSC WSPAs

3.5.2 Stress Assessment

As discussed in Section 1.0, the primary objective of this modelling is to determine water supply and reserve flows for use in the Tier 1 Water Quantity Stress Assessment. The Stress Assessment will be completed both for groundwater and surface water systems, and identifies those subwatersheds where there may be a potential for water taking related stress.

Following the methodology in Guidance Module 7, a subwatershed's stress is estimated by comparing the amount of water consumed with the amount of available water. This comparison is made by calculating "Percent Water Demand" as follows, with the terms defined in Table 3.11:

% Water Demand = $\frac{Q_{DEMAND}}{Q_{SUPPLY} \cdot Q_{RESERVE}} \times 100\%$

Term	Definition	Calculation	
<i>Q_{DEMAND}</i>	Consumptive Demand	Mean annual or monthly consumptive demand is calculated as the estimated rate of locally consumptive takings. Water demands are grouped into surface and groundwater takings.	
		Estimates of consumptive demand will be made from PTTW analysis, agricultural water use coefficients and private well usage. This portion of the Stress Assessment is outside the scope of the Water Availability Study, and will be completed by NPCA staff.	
Q_{SUPPLY}	Water Supply	For surface water, the supply is calculated as the monthly median outflow for the area to be assessed.	
		Groundwater supply is calculated as the estimated annual recharge rate plus the estimated groundwater inflow into a subwatershed.	
Q _{RESERVE}	Water Reserve	Water Reserve is a specified amount of water that is not considered as part of the available water supply.	
		For surface water supplies, water reserve is estimated using the 90 th percentile monthly outflow, at a minimum. The 90 th percentile flow is defined as the flow that is equaled or exceeded 90% of the time.	
		Groundwater reserve is calculated as 10% of the total estimated groundwater	

 Table 3.11 – Percent Water Demand Components

discharge within a subwatershed.

It is noted that baseflow is considered in both the surface water supply (baseflow within the outflow hydrograph) and groundwater supply (recharge, which sustains baseflow) terms of the Water Quantity Stress Assessment. While this may seem to "double count" baseflow, one should keep in mind the original purpose of the Stress Assessment, which is only to identify areas that have a high proportion of consumptive water taking, in comparison to the water flowing through the system. Identified areas, particularly at the Tier 1 scale, may not necessarily be experiencing hydrologic or ecologic stress, but rather are identified as requiring additional study to better understand the cumulative impacts of water use. The Stress Assessment methodology should not be utilized as a design/allocation tool, in an attempt to determine the total amount of water available to be withdrawn within a subwatershed, as double counting of the baseflow term would then be a consideration.

For surface water systems, the Percent Water Demand equation is based on a mean monthly basis. The maximum percent water demand for all months is then used to estimate the Potential for Surface Water Stress as shown on Table 3.12.

Surface Water Potential Stress Level Assignment	Maximum Monthly % Water Demand
Significant	> 50%
Moderate	20% - 50%
Low	<20 %

Table 3.12– Potential for Surface Water Stress Thresholds

For groundwater systems, the stress assessment calculation is based on mean annual demand conditions, as well as for monthly maximum demand conditions. The stress level for groundwater systems is calculated according to the thresholds shown on Table 3.13.

Table 3.13 – Potential for Groundwater Stress Thresholds

Groundwater Potential Stress Level Assignment	Mean Annual	Monthly Maximum
Significant	> 25%	> 50%
Moderate	> 10%	> 25%
Low	0-10%	0 - 25%

3.5.2.1 Surface Water Supply Components

The monthly median and 90th percentile flows, as estimated by HEC-HMS for the outlets of CWR, BFC and BDSC WSPAs are included in Tables 3.14, 3.15 and 3.16, respectively. These flow estimates include the direct overland runoff calculated from the upstream drainage area, and the interflow and baseflow components.

Table 3.14 – Surface Water Percent Water Demand Components for CWR

CENTRAL WELLAND RIVER WSPA			
Month	Water Supply (Median Flow) (m ³ /s)	Water Reserve (90 th % Flow) (m ³ /s)	
Jan	4.26	0.90	
Feb	5.52	1.86	

Mar	11.16	3.35
Apr	8.41	2.36
May	2.48	0.75
Jun	0.93	0.21
Jul	0.52	0.14
Aug	0.43	0.08
Sep	0.26	0.05
Oct	0.59	0.06
Nov	2.00	0.17
Dec	4.65	0.43

Table 3.15 – Surface Water Percent Water Demand Componen	ts for BFC
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	BIG FORKS CREEK WSPA				
Month	Water Supply (Median Flow)	Water Reserve (90 th % Flow)			
	$(\mathbf{m}^{3}/\mathbf{s})$	$(\mathbf{m}^{3}/\mathbf{s})$			
Jan	0.72	0.25			
Feb	0.94	0.34			
Mar	1.68	0.52			
Apr	1.01	0.33			
May	0.35	0.12			
Jun	0.12	0.04			
Jul	0.06	0.01			
Aug	0.04	0.00			
Sep	0.02	0.00			
Oct	0.05	0.00			
Nov	0.31	0.01			
Dec	0.83	0.11			

BEAV	ERDAMS SHRINE	RS CREEKS WSPA
Month	Water Supply (Median Flow)	Water Reserve (90 th % Flow)
	(m^3/s)	(m ³ /s)
Jan	0.25	0.08
Feb	0.35	0.11
Mar	0.73	0.15
Apr	0.46	0.10
May	0.19	0.03
Jun	0.07	0
Jul	0.04	0
Aug	0.03	0
Sep	0.03	0
Oct	0.06	0
Nov	0.24	0.01
Dec	0.27	0.05

Table 3.16 – Surface Water Percent Water Demand Components for BDSC

There is a greater amount of uncertainty with respect to the 90th percentile flows than with the median flows. The 90th percentile flow, being observed at the extreme low end of flows, may be affected by processes not considered by HEC-HMS. These processes may include, but are not limited to: water takings, evaporation from the stream channel, online ponds, and regional groundwater discharge. Due to the magnitude of these processes not being well known, the net effect of these processes (additional or less flow) is not able to be determined, but does introduce a level of uncertainty into the 90th percentile flows.

3.5.2.2 Groundwater Supply Components

The determination of the groundwater supply term is slightly more complex, due to HEC-HMS producing estimates of both interflow and baseflow. As described in Section 3.1.1.4, interflow is the portion of stormflow that moves through a shallow, unsaturated soil horizon towards a watercourse. Based on this description, the portion of percolated water that is directed into the interflow array will not be considered as part of the available groundwater supply.

The portion of percolated water that is directed to the baseflow array within HEC-HMS, and is meant to represent a slower, deeper groundwater system (only relative to the interflow component), will be used to infer groundwater recharge. It is recognized that within the Haldimand Clay Plain, there is very little evidence of a regional aquifer that has strong interconnections with the surface water system (Franz et al., 2007). Groundwater recharge estimates, inferred from HEC-HMS baseflow estimates, should not be considered recharge to deep, confined aquifers below the Haldimand Clay Plain, but rather recharge to shallow and localized aquifers near the surface. At the scale of a Tier 1 Water Quantity Stress Assessment, no distinction is made for recharge that supplies a specific aquifer unit; rather the stress assessment is carried out on the groundwater system as a whole. This may result in percent water demand being under-estimated for a confined water source whose primary source of water is lateral groundwater inflow.

It is recognized that there is uncertainty associated with HEC-HMS's arbitrary proportioning of percolated water to half baseflow and half interflow. Actually this division would shift from year to year, and season to season, with possibly some periods experiencing all percolated water returning to the watercourse as either interflow or baseflow. Determining the exact proportion of percolated water that reaches the uppermost water table (groundwater recharge) is not an obtainable goal for the scope of this project. By considering half of percolated water that is directed to the baseflow array as available for groundwater taking, the Stress Assessment will be conservative in nature.

Included in Table 3.17 is the estimated 1991-2005 annual mean groundwater recharge rate for CWR, BFC and BDSC WSPAs. Also included is the groundwater reserve value, which is equal to 10% of estimated groundwater discharge (baseflow).

WSPA	Water Supply (Groundwater Recharge) (mm)	Water Reserve (10% Discharge) (mm)
CWR	36	3.6
BFC	48	4.8
BDSC	29	2.9

Table 3.17 – Groundwater Percent Water Demand Components

To complete the groundwater Stress Assessment, groundwater inflow to CWR, BFC and BDSC must be quantified. It is anticipated that NPCA staff will complete this portion of the Stress Assessment as part of a separate project.

3.6 Uncertainty

Any model of a natural system is a simplification of reality, and as such, is inherently uncertain. Although the calibration and verification processes are performed in an attempt to reduce uncertainty, the model results and water budgets reflect the uncertainty in the input parameters.

The certainty of the water balance estimates is inexorably tied to the ability of the climate stations used in the model to accurately represent the average climatic conditions over the WSPAs. The current density of climate stations with long term datasets is likely not sufficient to fully reflect spatial climate variability, particularly during the summer months where extremely localized precipitation events are common (thunderstorms).

Further climate-related uncertainty is introduced into the process by the measurement error in climate observations. Uncertainty with the precipitation measurement has been estimated by Cumming Coburn Limited (2000) to be approximately $\pm 10\%$, with uncertainty during winter months reaching $\pm 20\%$. Precipitation measurement in winter months has a higher degree of uncertainty due to the difficulty of measuring snowfall.

Snow accumulation, ablation, redistribution and melt are significant hydrologic processes in Canadian watersheds. The rates of these processes are determined by the inter-relation of many factors, including: land cover, albedo, solar radiation, wind speed/direction, cloud cover, temperature fluctuations, rainfall amount/temperature, and new snow density. Lack of available data and a complete understanding on the interrelations and

implications of these factors introduces a level of uncertainty into hydrologic modelling. The simplified snow processes within HEC-HMS reflects this level of uncertainty.

With no streamflow data to calibrate to, and verify the performance of the model against, there is a greater degree of uncertainty associated with results from these two models, than the Twenty Mile or Upper Welland models. Given the geologic homogeneity, and the resulting likelihood that adjustments for Twenty Mile or Upper Welland are transferable to the CWR/BFC and BDSC model areas, this uncertainty is minimized.

Care should be taken when interpreting results from extreme events, such as the 1998-1999 drought. During extreme events certain processes that may be insignificant under "average" conditions, and therefore not considered in regional scale modelling, may become dominant and affect the hydrologic response of a watershed. This can lead to model estimates deviating from observed conditions for such periods.

Due to outflows from the Upper Welland being specified in the CWR/BFC model as a source, uncertainties associated with the representation of Binbrook Reservoir within the Upper Welland River model are present in Central Welland River outflows. Due to the decrease in reservoir effects as total drainage area increases, possible errors in the representation of Binbrook Reservoir are minimized, and are likely negligible, at the outfall of the Central Welland WSPA.

All modelling algorithms are simplified and are unable to accurately reflect the host of processes that can affect the hydrologic response of a catchment to a precipitation event. These limitations are not solely specific to HEC-HMS, nor the algorithms contained within HEC-HMS, but are rather a limitation of hydrologic science. With an insufficient ability to conceptualize and replicate all hydrologic processes, hydrologic model algorithms group or average, many processes into one. This averaging has the ability to introduce error into water balance estimates, and often the solution reached by an averaged approach results in a non-unique solution (as is shown to be the case in CWR/BFC and BDSC). This limitation should be kept in mind for utilizing water balance estimates generated as part of this study.

The climate of southwestern Ontario significantly varies from season to season. As a result of these changing seasons, hydrologic parameters (e.g. infiltration, depression storage, overland runoff routing) also vary. In the case of the freezing and thawing of soils, this can have a significant impact on the ability of soil to infiltrate water. HEC-HMS does not have the ability to vary parameters with season, and as such, is a source of uncertainty. Due to the dominance of the Haldimand Clay Plain in the area, and its limited ability to infiltrate water, even under warm conditions, it is expected that the uncertainty associated with this limitation is less significant than for highly pervious watersheds.

While any modelling exercise contains inherent uncertainties, it should be noted that the constructed HEC-HMS models produce estimates of streamflow and water balance values that exceed the level of accuracy expected for a Tier 1 Water Quantity Stress Assessment.

4. Conclusions and Recommendations

A HEC-HMS continuous hydrologic model has been constructed for Central Welland River (CWR) and Big Forks Creek (BFC) and a second model for Beaverdams Shriners Creeks (BDSC). The models have been adjusted based on a successful calibration and verification exercise carried out for both Twenty Mile Creek and the Upper Welland River. Due to the lack of observed streamflow data, the performance of the CWR/BFC and BDSC models in predicting streamflow is not able to be determined. However, based on the performance of Twenty Mile and Upper Welland, it is very likely that the CWR/BFC and BDSC models are predicting reasonable streamflow and water balance estimates.

The generated water balance and streamflow estimates reflect the most complete understanding of the hydrologic system that is available, and represent the best available estimates. Significant uncertainties do remain, however there is insufficient information to quantify the net impact of these uncertainties on the water balance and streamflow estimates. These estimates will form the foundation of a future Tier 1 Stress Assessment.

To advance the Tier 1 Water Quantity Stress Assessment as well as the basic understanding of the significant hydrologic processes, the following recommendations are made:

- 1. That groundwater inflow volumes to CWR, BFC and BDSC be approximated by use of regional groundwater mapping products. Groundwater inflows are required to fully quantify the water supply term of the Groundwater Stress Assessment.
- 2. That monthly consumptive surface and groundwater demand (non-Great Lakes sources only) be approximated from Permits To Take Water, Census of Agriculture, and Census of Population. These consumptive demands are required to complete the Water Quantity Stress Assessment.
- 3. Utilizing the estimated consumptive demands, the groundwater inflow volumes, and values presented in Section 3.5.2, that the Tier 1 Water Quantity Stress Assessment be carried out. This will identify WSPAs that have a potential for hydrologic stress related to water takings.
- 4. That the water balance estimates generated from the Deficit and Constant Loss Method for CWR, BFC and BDSC, be validated against estimates generated from a more detailed loss Method (Soil Moisture Accounting Method). Should the more detailed Soil Moisture Accounting Method generate water balance estimates similar to the Deficit and Constant Loss, a higher level of certainty could be attached to estimates generated for other WSPAs. The need for further model refinement could be re-evaluated following the subsequent stress assessment.

Despite the uncertainties inherent with any modelling exercise, the CWR/BFC and BDSC HEC-HMS models are excellent tools for estimating the water supply components of a Tier 1 Water Quantity Stress Assessment. In addition to exceeding the expectations of a Tier 1 Stress Assessment, it will greatly assist NPCA staff in characterizing and

understanding the fundamental hydrologic processes occurring within CWR/BFC and BDSC.

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TABLES

TABLE 2.1 MEAN ANNUAL CLIMATE STATION VALUES WATER AVAILABILITY STUDY

MSC ID	NAME	AVE	RAGE ANNUAL (1991	-2005)
		PRECIPITATION	SNOW WATER	TEMPERATURE (°C)
		(mm)	EQUIVALENT (mm)	
6132148	DUNNVILLE PUMPING STN	948	91.1	8.6
6132470	FORT ERIE	1053	193	8.7
6133055	GRIMSBY MOUNTAIN	923	184	9.3
6133120	HAGERSVILLE	947	102	8.4
6153194	HAMILTON A	898	160	8.0
6153298	HAMILTON PSYCH HOSPITAL	850	108	8.8
6153290	HAMILTON MUNICIPAL LAB	793	101	9.5
6135657	NIAGARA FALLS NPCSH	948	160	9.4
6155097	MIDDLEPORT TS	896	110	8.5
6135FF4	NIAGARA ON THE LAKE	838	110	9.3
6136606	PORT COLBORNE	971	147	9.3
6136626	PORT DALHOUSIE	895	140	9.5
6137161	RIDGEVILLE	918	117	9.1
6137287	ST CATHARINES A	897	145	9.3
6139141	VINELAND	875	137	9.5
6139143	VINELAND RITTENHOUSE	850	115	9.4
6137306	ST CATHARINES POWER GLEN	890	135	9.2
6139445	WELLAND	969	148	9.0
6139148	VINELAND STATION RCS	840	146	9.1
6131165	CANBORO	894	120	8.4

Table Notes:

MSC - Meteorological Survey of Canada



TABLE 2.2a HYDROLOGIC SOIL GROUPS BY CATCHMENT CENTRAL WELLAND RIVER WATERSHED PLANNING AREA WATER AVAILABILITY STUDY

Subcatchment	Area	· · · · · · · · · · · · · · · · · · ·								
CWR ID	(km²)	Α	В	C	D	Other ¹				
CWR_BAC_W100	12.1			15.72	83.32	0.96				
CWR_BCW_W100	13.2			25.79	73.06	1.16				
CWR_BCW_W200	11.5		15.83		81.57	2.60				
CWR_BND_W100	8.0		12.20	0.29	61.08	26.43				
CWR_BND_W200	10.2		12.03	3.14	23.57	61.27				
CWR_CC_W100	8.2	14.50		46.86	29.43	9.21				
CWR_CC_W110	7.1	4.07	6.60	55.57	26.34	7.42				
CWR_CC_W120	6.5	33.21	3.74	37.52	17.79	7.73				
CWR_CC_W130	8.0	24.22	3.41	40.75	28.97	2.64				
CWR_CC_W200	11.0	3.11	0.69	23.57	66.22	6.42				
CWR_DRC_W100	8.9	0.53	0.33	38.83	16.96	43.35				
CWR_IC_W100	13.9		9.12	12.88	72.56	5.44				
CWR_LCD_W100	13.5				61.59	38.40				
CWR_LFC_W100	6.5			36.83	62.91	0.26				
CWR_PC_W100	8.3			22.80	75.07	2.13				
CWR_SU_W100	11.2			44.71	52.40	2.89				
CWR_TPD_W100	5.0	2.75		52.27	15.10	29.87				
CWR_UNC_W100	7.7			40.28	52.91	6.81				
CWR_UNC_W110	7.7			47.51	48.36	4.13				
CWR_UNC_W200	12.3			13.29	83.60	3.11				
CWR_WCS_W100	12.4			13.02	4.94	82.04				
CWR_WCS_W200	17.9		0.95	5.10	24.86	69.09				
CWR_WCS_W300	9.6		7.60	2.09	8.06	82.25				
CWR_WCS_W310	6.0		0.85		36.42	62.72				
CWR_WR_W100	7.2			7.72	12.31	79.97				
CWR_WR_W110	11.9			0.05	79.21	20.74				
CWR_WR_W200	15.4			24.12	63.56	12.32				
CWR_WR_W300	9.5			42.28	50.09	7.63				
CWR_WR_W310	6.9			25.55	68.72	5.73				
CWR_WR_W320	6.0			13.85	82.26	3.89				
CWR_WR_W400	6.0			66.05	28.67	5.28				
CWR_WRC_W100	5.2			13.27	22.73	64.00				
Average %		11.8	5.2	26.3	47.3	23.7				
% of CWR	100.0%	2.0%	1.8%	22.9%	49.8%	23.4%				
Area (km ²)	304.4	6.08	5.51	69.82	151.71	71.31				

Table Notes:

BAC - Black Ash Creek, BCW - Beaver Creek, BND - Bierderman Drain, CC - Coyle Creek DRC - Draper Creek, IC - Indian Creek, LCD - Lyons Creek Drain, LFC - Little Forks Creek PC - Parker Creek, SU - Sucker Creek, TPD - Toe Path Drain, UNC - Unnamed Creek WCS - Welland Canal South, WR - Welland River, WRC - Welland River between Canal 1 - Where soils unmapped or an area of high runoff, i.e. urban areas, water bodies, bedrock at surface





TABLE 2.2b HYDROLOGIC SOIL GROUPS BY CATCHMENT BIG FORKS CREEK and BEAVER DAMS SHRINERS CREEK WATERSHED PLANNING AREAS WATER AVAILABILITY STUDY

Subcatchment	Area		Hydro	logic Soil Grou	ups (%)	
BFC ID	(km ²)	Α	В	С	D	Other ¹
BFC_BED_W100	12.1			65.7	34.3	
BFC_BFC_W100	5.7			19.8	77.1	3.1
BFC_BFC_W110	9.1			13.5	86.0	0.5
BFC_BFC_W200	12.9	1.4		83.0	15.6	
BFC_BFC_W300	14.1	13.1	4.1	53.3	29.1	0.5
BFC_EKD_W100	15.0	5.3	10.8	77.1	6.7	
BFC_ELD_W100	5.5	13.0	5.0	76.6	5.4	
BFC_MRC_W100	8.7			5.7	93.5	0.8
BFC_MRC_W110	6.0			3.8	96.2	
BFC_MRC_W111	8.4		4.3	3.2	85.6	6.9
BFC_MRC_W120	16.3	0.3		9.6	89.9	0.2
BFC_MRC_W200	13.5			22.9	75.9	1.2
BFC_MRC_W210	10.6	0.0	0.1	2.9	93.8	3.2
BFC_MRC_W300	6.8	1.1	17.7	45.3	34.8	1.0
BFC_MRC_W310	6.4		11.0	47.2	41.4	0.4
BFC_NFD_W100	8.9	13.5	0.6	72.6	13.2	
BFC_WDC_W100	6.5			50.4	49.6	
Average %		6.0	6.7	38.4	54.6	1.8
% of BFC		2.9%	2.9%	39.7%	53.5%	0.9%
Area (km ²)	166.5	4.86	4.80	66.10	89.14	1.56

Big Forks Creek (BFC) Table Notes:

BED - Beezor Drain, BFC - Big Forks Creek, EKD - East Kelly Drain, ELD - Ellsworth Drain

MRC - Mill Race Creek, NFD - North Forks Drain, WDC - Wolf Creek Drain West

1 - Where soils unmapped or an area of high runoff, i.e. urban areas, water bodies, bedrock at surface

Subcatchment	Area		Hydro	logic Soil Groι	ıps (%)	-
BDSC ID	(km²)	Α	В	С	D	Other ¹
BDSC_BRDC_W100	9.6			50.5	14.7	34.8
BDSC_BRDC_W200	5.9			41.7	42.7	15.6
BDSC_SSC_W100	5.5			58.4	28.4	13.2
BDSC_SSC_W200	9.7		0.0	26.9	11.7	61.5
BDSC_TMC_W100	6.5			48.3	25.9	25.8
BDSC_WCN_W100	7.6	0.1	4.0	10.3	14.6	71.1
BDSC_WCN_W200	7.6	1.2		11.4	9.0	78.5
BDSC_WCN_W300	8.4			32.5	26.0	41.5
BDSC_WCN_W310	6.5			31.2	65.1	3.8
BDSC_WCN_W320	8.5	7.5	1.1	37.1	37.8	16.4

Average %		2.9	1.7	34.8	27.6	36.2
% of BFC		1.0%	0.5%	34.0%	26.0%	38.4%
Area (km ²)	75.8	0.73	0.40	25.82	19.74	29.14

Beaver Dams Shriners Creek (BDSC)Table Notes:

BRDC - Beaver Dams Creek, SSC - Shriners Creek, TMC - Ten Mile Creek, WCN - Welland Canal North

1 - Where soils unmapped or an area of high runoff, i.e. urban areas, water bodies, bedrock at surface





TABLE 2.3a LAND COVER BY CATCHMENT CENTRAL WELLAND RIVER WATERSHED PLANNING AREA WATER AVAILABILITY STUDY

Subcatchment CWR	Annual	Mixed	Mixed	Monoculture	Orchards	Perennial	Plantations	Vineyards	Coniferous	Deciduous	Forest	Hedge	Mixed	Built Up	Idle	Rural	Marsh	Swamp	Bog	Open	Built Up	Extraction	Transportation
ID	Crop	Agriculture	Crop			Crop			Forest	Forest		Rows	Forest	Pervious	Land	Land Use				Water	Impervious		
CWR_BAC_W100		17.8		12.1		0.0				10.1	0.9	0.2			0.0	46.4	0.9	8.5		0.0			2.9
CWR_BCW_W100		15.3		21.1			0.0			10.2	1.0	0.5		0.2	0.0	36.8	3.4	9.1			0.0		2.3
CWR_BCW_W200		16.0		20.3			0.1			7.0	1.2	0.3		0.1	0.1	46.3	0.7	4.9		0.0	0.0		2.9
CWR_BND_W100		5.8		35.9						2.4	0.8	0.4		1.1	0.2	20.4	3.4	12.1	11.2	0.0	2.8		3.3
CWR_BND_W200		2.2		9.3						4.7	0.4	0.1		0.6	0.7	8.2	0.8	7.4	64.3			0.4	0.9
CWR_CC_W100	0.5	12.3	4.8	22.2	3.6		2.0	0.4	0.4	11.1	1.9	0.6	1.1	1.4	4.8	11.8	1.1	14.2		0.3	1.8		3.6
CWR_CC_W110		4.2	2.6	21.9	0.4		3.2		0.5	11.4	2.1	1.0	2.7	2.6	6.1	16.3	0.4	10.7			7.7		6.1
CWR_CC_W120	4.4	3.7	7.5	18.8	7.8	0.3	2.1	0.3	0.1	8.9	1.6	0.7	1.2	2.3	13.0	8.6	0.1	4.3		0.0	9.9	0.0	4.5
CWR_CC_W130	0.4	5.1	11.5	18.5	2.7	1.6	3.6	0.3	0.5	8.4	1.7	2.0	2.4	2.6	9.5	7.1	0.8	12.7			3.9		4.7
CWR_CC_W200	3.2	6.1	8.0	20.7	0.6	1.3	0.9		0.4	5.7	1.5	0.8	0.4	1.9	4.7	20.2	0.6	10.5		0.3	8.4		3.8
CWR_DRC_W100		5.6		9.7		0.0	0.5		0.7	5.5	2.5	0.2	0.2	6.4	0.3	9.2	1.5	11.4		0.0	34.3		11.9
CWR_IC_W100		10.7		22.1		0.1	0.4			6.4	1.9	0.4		3.8	0.1	35.2		14.7		0.4	0.1		3.7
CWR_LCD_W100		2.7		52.5						3.7	1.4	0.1		2.0		20.3	0.1	13.0		0.6	0.1		3.6
CWR_LFC_W100		6.5		26.8			0.6			7.8	2.0	0.8		0.1	0.0	37.3	0.7	14.1		0.1	0.1		3.2
CWR_PC_W100		20.7		5.8						7.2	1.0	0.5		0.0		54.2	0.6	7.8		0.0			2.1
CWR_SU_W100		7.7		26.5		0.0	0.1			8.4	1.1	0.1		0.3		40.4	1.1	10.9		0.1			3.3
CWR_TPD_W100	0.1	2.5	2.7	24.4					0.2	3.1	1.4	0.7		10.0	1.4	11.9	0.3	13.6		0.3	19.3		8.0
CWR_UNC_W100		6.6		34.0			1.4			6.8	1.3	0.2			0.0	32.2	2.1	9.6		0.1	3.0		2.6
CWR_UNC_W110		8.3		19.7						10.0	0.6	0.0				49.5	2.3	7.6		0.1	0.3		1.5
CWR_UNC_W200		9.8		18.8			0.4			9.8	1.3	0.2		0.1		43.7	0.3	12.5			0.5		2.6
CWR_WCS_W100		2.1		9.5			1.0			1.8	0.7	0.0		13.5	0.2	4.2	0.2	5.1		10.1	37.9		13.7
CWR_WCS_W200	0.0	2.9		32.0		0.0	0.2			3.0	0.5	0.2		8.3	0.1	8.0	0.3	4.7		9.2	20.2		10.2
CWR_WCS_W300		5.0		13.9		0.0			0.2	5.9	0.3	0.1		9.4	0.1	7.9		2.2		13.2	29.3	0.5	11.9
CWR_WCS_W310		6.4		29.6		0.0	0.1			3.9	1.7			2.5	0.5	15.8	0.2	11.0		7.3	15.5		5.4
CWR_WR_W100		0.7		6.7			0.1		0.0	2.1	0.4	0.3		6.9	0.1	6.1	1.5	5.4		3.5	47.2		18.9
CWR_WR_W110		8.7		31.2			0.2			4.7	2.2	1.3		6.6	0.1	15.4	1.2	15.4		0.0	6.5		6.4
CWR_WR_W200		5.3		42.5		0.0	0.5		0.1	3.4	0.7	0.8	0.1	3.6		23.5	2.8	9.3		2.6	0.7		4.1
CWR_WR_W300		17.3		23.4			2.9		0.0	2.8	1.1	1.6		0.3	0.1	33.2	3.5	5.9		3.0			4.7
CWR_WR_W310		10.6		36.8			0.6			7.0	0.7	0.8		0.4	0.1	29.9	0.3	9.6			0.1		3.0
CWR_WR_W320		8.5		28.9			0.2			6.1	0.9	0.2	0.2	0.7		39.3	0.2	11.4		0.3			3.1
CWR_WR_W400		3.5		42.8			0.5			2.9	0.4	1.5		0.0	0.0	36.4	3.4	2.0		3.7			3.0
CWR_WRC_W100		1.5		32.8		0.0				2.5	0.6	0.2		19.1	1.5	6.8	2.0	6.1		5.5	13.9		7.5
Average %	1.4	7.6	6.2	24.1	3.0	0.3	0.9	0.3	0.3	6.1	1.2	0.5	1.0	3.7	1.7	24.5	1.2	9.3	37.7	2.3	10.5	0.3	5.3
% of CWR	0.2%	7.9%	1.0%	23.7%	0.4%	0.1%	0.6%	0.0%	0.1%	6.1%	1.2%	0.5%	0.2%	3.3%	1.1%	25.2%	1.1%	9.3%	2.5%	2.1%	8.0%	0.0%	5.3%
Area/Land Cover (km ²)	0.7	24.1	3.0	72.3	1.1	0.3	1.9	0.1	0.3	18.7	3.6	1.6	0.6	10.0	3.3	76.6	3.5	28.4	7.5	6.3	24.2	0.1	16.1

Table Notes:

CWR - Central Welland River, BAC - Black Ash Creek, BCW - Beaver Creek, BND - Bierderman Drain, CC - Coyle Creek, DRC - Draper Creek, IC - Indian Creek, LCD - Lyons Creek Drain, LFC - Little Forks Creek, 'PC - Parker Creek, SU - Sucker Creek, TPD - Toe Path Drain,



TABLE 2.3b LAND COVER BY CATCHMENT BIG FORKS CREEK and BEAVERDAMS SHRINERS CREEK WATERSHED PLANNING AREA WATER AVAILABILITY STUDY

Subcatchment BFC ID	Annual	Mixed	Mixed	Monoculture	Orchards	Perennial	Plantations	Vineyards	Coniferous	Deciduous	Forest	Hedge	Mixed	Built Up	Idle	Rural	Marsh	Swamp	Open	Built Up	Bog	Transportation
	Crop	Agriculture				Crop		•	Forest	Forest		Rows	Forest	Pervious	Land	Land Use			Water	Impervious)	
BFC_BED_W100		7.7		30.6			0.5			8.0	1.9	0.6		0.1	0.0	32.8	0.1	15.0				2.6
BFC_BFC_W100		9.7		23.9			1.1			9.9	2.1	0.3			0.1	25.9	3.1	20.4	0.0	0.0		3.6
BFC_BFC_W110		10.7		31.9						5.1	1.0	0.7				37.3	1.3	9.6				2.3
BFC_BFC_W200		12.6		33.1			0.2			3.5	1.2	1.0		0.5	0.0	34.1	1.4	7.3	0.0	1.4		3.8
BFC_BFC_W300	1.3	5.4		40.5		0.0	1.0		0.1	5.1	2.3	0.3		0.4	0.2	27.7	0.2	11.5	0.0	0.4		3.5
BFC_EKD_W100		8.3		40.0		0.0	0.9			3.7	1.0	0.2		0.2		35.7	0.0	6.5	0.0	0.2		3.2
BFC_ELD_W100	4.9	5.1		40.9		0.0	0.0			8.1	1.9			0.2		19.7	0.3	14.4		0.7		3.7
BFC_MRC_W100		10.2		39.7			0.1		0.1	3.8	0.9	0.1		0.3		32.9	0.3	6.3	0.0	0.0	1.4	3.8
BFC_MRC_W110		2.9		22.6						0.0	0.0					21.2	0.8	1.9			50.0	0.5
BFC_MRC_W111		12.4		16.9		0.0				0.8	0.3	0.2				19.3	1.6	22.4	0.8	0.5	21.9	2.9
BFC_MRC_W120		5.3		42.2			0.5			3.2	1.0			0.1	0.0	33.8	0.0	10.0	0.0	0.3		3.4
BFC_MRC_W200		6.2		43.2						5.5	1.3	0.3		0.5	0.0	24.8	0.5	12.4	0.0	1.1		4.2
BFC_MRC_W210	3.0	6.1		38.0			1.0			2.6	1.0	0.3	0.1	0.0	0.0	29.2	0.4	15.1	0.0	0.3		2.8
BFC_MRC_W300		6.1		42.3			0.5			3.8	1.6	0.1		0.1	0.0	31.7	0.0	10.8	0.0			3.1
BFC_MRC_W310		5.8		39.3			0.1			1.1	1.1	0.5		0.1		37.2	1.1	9.8				3.8
BFC_NFD_W100		6.4		45.7			0.1			3.0	0.9	0.6		1.8		31.8		5.8		1.7		2.3
BFC_WDC_W100		6.0		33.6			0.7			5.5	2.4			0.0	0.0	33.6	0.0	15.6	0.0			2.7
																			-			
Average %	3.1	7.5	na	35.5	na	0.0	0.5	na	0.1	4.3	1.3	0.4	0.1	0.3	0.0	29.9	0.7	11.5	0.1	0.6	24.5	3.1
% of BFC	0.5%	7.5%		36.5%		0.0%	0.4%		0.0%	4.2%	1.3%	0.3%	0.0%	0.3%	0.0%	30.5%	0.6%	11.1%	0.1%	0.4%	3.0%	3.2%
Area/Land Cover (km ²)	0.8	12.6		60.8		0.0	0.7		0.0	7.1	2.1	0.5	0.0	0.5	0.0	50.8	0.9	18.5	0.1	0.7	5.0	5.3
Table Notes:					1						1							1				

Table Notes:

BFC - Big Forks Creek, BED - Beezor Drain, BFC - Big Forks Creek, EKD - East Kelly Drain, ELD - Ellsworth Drain, MRC - Mill Race Creek, NFD - North Forks Drain, WDC - Wolf Creek Drain West na - not applicable

Subcatchment BDSC ID	Annual	Mixed	Mixed	Monoculture	Orchards	Perennial	Plantations	Vineyards	Coniferous	Deciduous	Forest	Hedge	Mixed	Built Up	Idle	Rural	Marsh	Swamp	Open	Built Up	Extraction	Transportation
	Crop	Agriculture	Crop			Crop			Forest	Forest		Rows	Forest	Pervious	Land	Land Use			Water	Impervious		
BDSC_BRDC_W100		5.3		16.6		0.0			0.0	3.8	0.6	0.4		14.4	0.2	30.6	1.0	1.8	2.5	14.1		8.6
BDSC_BRDC_W200		10.0		33.5			0.1			3.8	0.2	1.1		12.2	0.1	25.7	0.5	2.9	0.0	5.9		4.0
BDSC_SSC_W100	0.2	3.7	0.4	41.1		1.9	0.2			5.5	0.7	0.4		7.4	4.4	18.7	2.3	2.9	2.2	3.6		4.3
BDSC_SSC_W200		6.3		6.6						3.1	0.2	0.1		5.1	0.1	9.2	0.1	0.6		49.8		18.8
BDSC_TMC_W100	12.5	1.8		25.7		0.0	0.1		0.1	6.1	0.6	0.8		4.7	14.4	3.5	0.2	2.0	3.3	8.1	10.1	6.0
BDSC_WCN_W100	0.7		11.6		3.2	0.0	0.1	1.6	2.4	6.7	1.5	0.2	1.2	8.9	12.4		0.2	4.6	27.1	9.6		8.1
BDSC_WCN_W200	1.6		4.7	0.3		0.4			0.1	13.3	1.6	0.0		8.7	17.5		0.4	4.8	20.0	13.0	8.6	5.0
BDSC_WCN_W300		3.9		29.2			0.3		0.3	6.4	0.6	0.4		4.4	6.4	22.9	1.2	4.1	8.5	8.7		2.7
BDSC_WCN_W310	0.1	7.7	3.9	25.5	1.4		0.5	1.0		9.2	1.2	0.6		2.2	0.4	26.5	1.5	14.4	0.0	0.5		3.5
BDSC_WCN_W320	9.1	3.2	12.1	14.2	3.1	0.0	1.1	0.1	0.0	2.0	1.3	0.4	0.4	6.8	9.7	5.5	0.7	12.9		10.9		6.6
Average %	4.0	5.2	6.6	21.4	2.6	0.4	0.3	0.9	0.5	6.0	0.8	0.4	0.8	7.5	6.6	17.8	0.8	5.1	7.9	12.4	9.4	6.8
% of BDSC	2.3%	4.1%	3.4%	17.8%	0.8%	0.2%	0.2%	0.3%	0.3%	5.8%	0.8%	0.4%	0.2%	7.6%	6.4%	14.1%	0.8%	5.0%	6.4%	14.1%	1.7%	7.3%
Area/Land Cover (km ²)	1.8	3.1	2.6	13.5	0.6	0.1	0.2	0.2	0.2	4.4	0.6	0.3	0.1	5.7	4.9	10.7	0.6	3.8	4.9	10.7	1.3	5.5

Table Notes:

BDSC - Beaverdams Shriners Creek, BRDC - Beaverdams Creek, SSC - Shriners Creek, TMC - Ten Mile Creek, WCN - Welland Canal North



TABLE 3.1a **CATCHMENT PARAMETERS CENTRAL WELLAND RIVER WATERSHED PLANNING AREA** WATER AVAILABILITY STUDY

Catchment CWR ID	Area (km²)	Slope (%)	Impervious Area (%)	Curve Number	Basin Time	Maximum storage	Infiltation Rate
	()	, ,		(CN)	Lag	(mm)	(mm/hour)
				· · /	(hours)		````
CWR_BCW_W100	13.2	2.7	2.3	83.4	4.8	220.2	1.6
CWR_BCW_W200	11.5	2.4	2.9	84.5	3.0	207.8	1.5
CWR_BAC_W100	12.1	2.4	2.9	83.6	3.9	222.0	1.4
CWR_CC_W110	7.1	3.6	9.9	78.6	2.6	219.1	2.7
CWR_CC_W120	6.5	4.6	9.5	72.1	3.3	177.2	4.5
CWR_CC_W130	8.0	4.0	6.6	75.0	2.3	210.9	4.0
CWR_CC_W200	11.0	2.3	8.0	82.2	2.9	198.4	1.8
CWR_CC_W100	8.2	4.0	4.5	77.4	3.8	213.6	2.9
CWR_DRC_W100	8.9	2.8	28.9	81.7	2.7	110.9	1.1
CWR_LFC_W100	13.5	2.4	3.3	82.3	4.4	231.7	1.5
CWR_PC_W100	8.3	2.7	2.2	83.9	3.0	218.9	1.4
CWR_SU_W100	11.2	2.3	3.3	83.1	3.5	222.1	1.5
CWR_UNC_W200	12.3	2.6	2.8	83.1	3.2	225.1	1.3
CWR_UNC_W110	7.7	2.6	1.7	83.1	2.6	223.1	1.7
CWR_UNC_W100	7.7	2.9	4.2	83.5	2.6	205.3	1.6
CWR_WR_W320	6.0	2.1	3.1	84.1	2.1	209.1	1.2
CWR_WR_W400	6.0	3.1	3.0	84.9	2.4	194.4	1.8
CWR_WR_W310	6.9	2.2	3.0	84.0	2.9	205.8	1.4
CWR_WR_W300	9.5	2.7	4.7	84.4	2.8	189.3	1.6
CWR_WR_W100	7.2	2.7	42.4	86.1	1.7	44.0	0.4
CWR_WR_W110	11.9	2.5	9.7	84.5	2.9	161.3	0.5
CWR_WR_W200	15.4	2.3	4.5	85.0	3.3	176.9	1.3
CWR_BND_W100	8.0	2.0	4.8	82.3	2.3	163.4	1.2
CWR_BND_W200	10.2	1.6	0.9	74.3	3.8	294.4	1.3
CWR_IC_W100	13.9	1.9	3.8	83.1	3.2	200.3	1.4
CWR_LCD_W100	6.5	3.5	3.6	85.6	1.6	129.1	0.4
CWR_TPD_W100	5.0	3.3	17.4	82.8	2.2	134.9	1.8
CWR_WCS_W100	12.4	3.7	32.6	89.4	2.6	32.4	0.4
CWR_WCS_W200	17.9	4.9	20.4	88.3	2.6	56.3	0.3
CWR_WCS_W300	9.6	3.3	26.4	89.3	2.3	28.8	0.6
CWR_WCS_W310	6.0	3.1	13.2	86.1	1.6	66.0	0.3
CWR_WRC_W100	5.2	6.0	14.4	88.2	1.2	65.8	0.5
Minimum	5.0	1.6	0.9	72.1	1.2	28.8	0.3
Maximum	17.9	6.0	42.4	89.4	4.8	294.4	4.5
	0.5			00.4		170.0	

Table Notes:

Average

BAC - Black Ash Creek, BCW - Beaver Creek, BND - Bierderman Drain, CC - Coyle Creek DRC - Draper Creek, IC - Indian Creek, LCD - Lyons Creek Drain, LFC - Little Forks Creek PC - Parker Creek, SU - Sucker Creek, TPD - Toe Path Drain, UNC - Unnamed Creek WCS - Welland Canal South, WR - Welland River, WRC - Welland River between Canal

9.4



3.0

9.5



83.1

2.8

170.6

1.5

TABLE 3.1b **CATCHMENT PARAMETERS BIG FORKS CREEK AND BEAVERDAMS SHRINERS CREEK WATERSHED PLANNING AREAS** WATER AVAILABILITY STUDY

Catchment ID	Area (km²)	Slope (%)	Impervious Area (%)	Curve Number	Basin Time	Maximum storage	Infiltation Rate
	()	. ,		(CN)	Lag	(mm)	(mm/hour)
					(hours)		
BFC_BED_W100	12.1	1.7	2.6	81.4	4.4	246.8	1.9
BFC_BFC_W100	5.7	1.9	3.7	81.5	2.3	227.3	1.1
BFC_BFC_W300	14.1	2.6	3.7	78.9	2.9	209.1	3.1
BFC_BFC_W110	9.1	1.4	2.2	84.5	3.0	202.6	0.9
BFC_BFC_W200	12.9	2.0	4.5	82.3	4.1	221.1	2.4
BFC_EKD_W100	15.0	2.0	3.4	79.8	4.7	212.9	3.4
BFC_ELD_W100	5.5	2.2	4.1	78.4	3.1	225.9	3.5
BFC_MRC_W100	8.7	2.0	3.9	85.3	3.1	188.3	0.8
BFC_MRC_W200	13.5	1.9	4.8	83.3	3.2	209.8	1.3
BFC_MRC_W120	16.3	1.5	3.5	84.4	4.2	198.1	1.0
BFC_MRC_W110	6.0	1.0	0.5	79.1	4.3	264.5	0.8
BFC_MRC_W111	8.4	1.6	3.1	79.7	3.6	234.8	0.9
BFC_MRC_W210	10.6	1.8	2.9	83.5	3.2	204.6	1.1
BFC_MRC_W300	6.8	2.0	3.1	80.6	3.4	211.0	2.7
BFC_MRC_W310	6.4	1.9	4.0	82.3	3.3	205.5	2.3
BFC_NFD_W100	8.9	2.0	3.1	80.8	3.5	199.6	3.1
BFC_WDC_W100	6.5	2.3	2.7	82.0	2.3	236.9	1.7
Minimum	5.5	1.0	0.5	78.4	2.3	188.3	0.8
Maximum	16.3	2.6	4.8	85.3	4.7	264.5	3.5
Average	9.8	1.9	3.3	81.6	3.4	217.6	1.9

Big Forks Creek (BFC) Table Notes:

BED - Beezor Drain, BFC - Big Forks Creek, EKD - East Kelly Drain, ELD - Ellsworth Drain MRC - Mill Race Creek, NFD - North Forks Drain, WDC - Wolf Creek Drain West

BDSC_BRDC_W100	9.6	3.4	15.7	86.1	2.1	133.7	1.4
BDSC_BRDC_W200	5.9	2.1	7.0	85.7	2.2	170.5	1.7
BDSC_SSC_W100	5.5	3.1	6.2	84.5	1.7	177.2	1.9
BDSC_SSC_W200	9.7	3.1	43.8	85.3	2.3	52.1	0.8
BDSC_TMC_W100	6.5	4.9	10.1	85.3	1.8	138.9	1.6
BDSC_WCN_W100	7.7	4.9	13.0	87.7	2.2	65.7	0.7
BDSC_WCN_W200	7.6	8.4	11.4	87.2	1.1	46.1	0.6
BDSC_WCN_W300	8.4	4.3	7.0	85.5	2.0	120.4	1.0
BDSC_WCN_W310	6.5	2.6	3.6	82.7	1.8	213.9	1.5
BDSC_WCN_W320	8.5	2.7	12.0	80.9	2.3	171.0	2.0

Minimum	5.5	2.1	3.6	80.9	1.1	46.1	0.6
Maximum	9.7	8.4	43.8	87.7	2.3	213.9	2.0
Average	7.6	3.9	13.0	85.1	2.0	128.9	1.3

Beaverdams Shriners Creek (BDSC)Table Notes:

BRDC - Beaverdams Creek, SSC - Shriners Creek, TMC - Ten Mile Creek, WCN - Welland Canal Nort

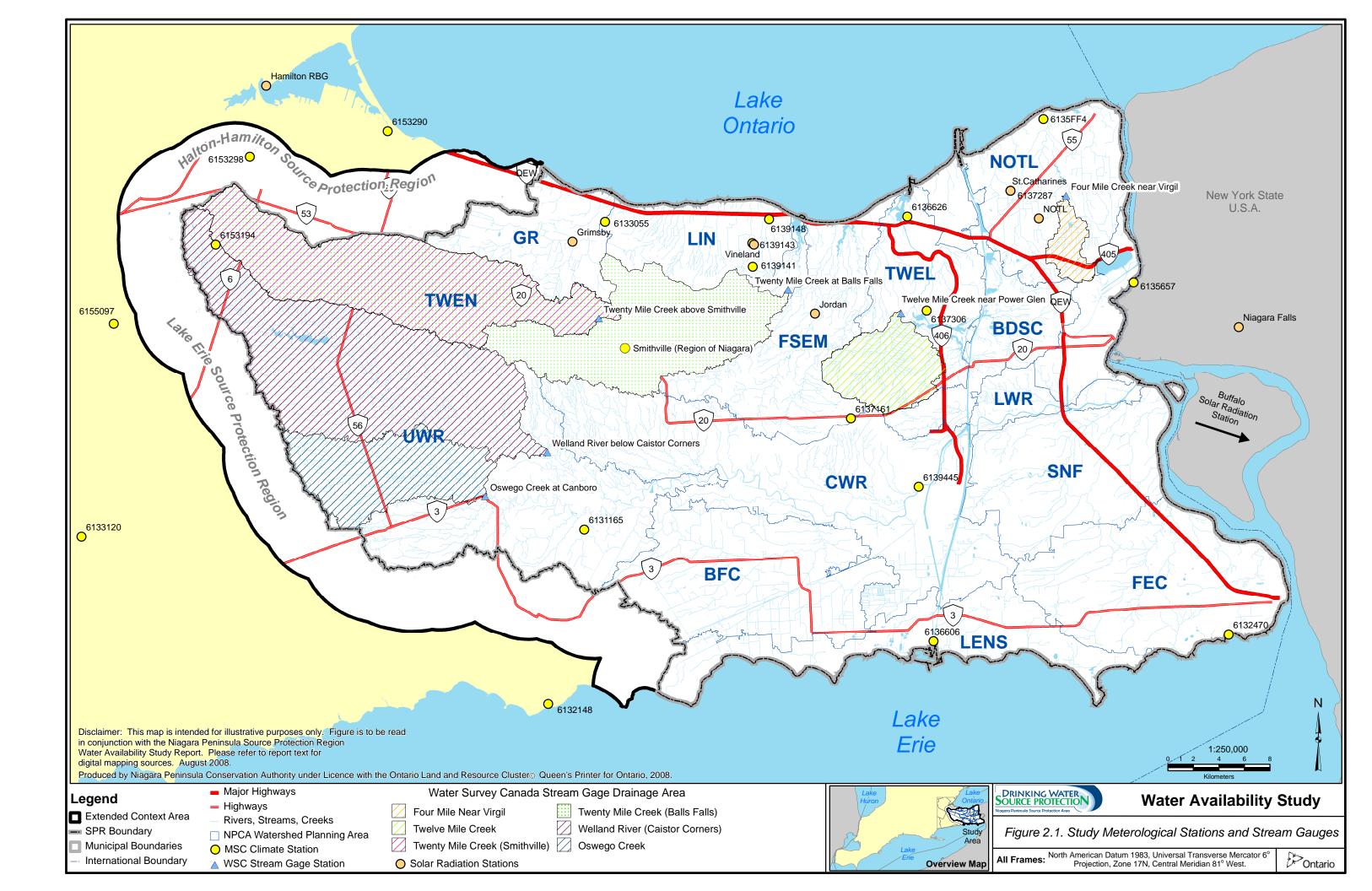


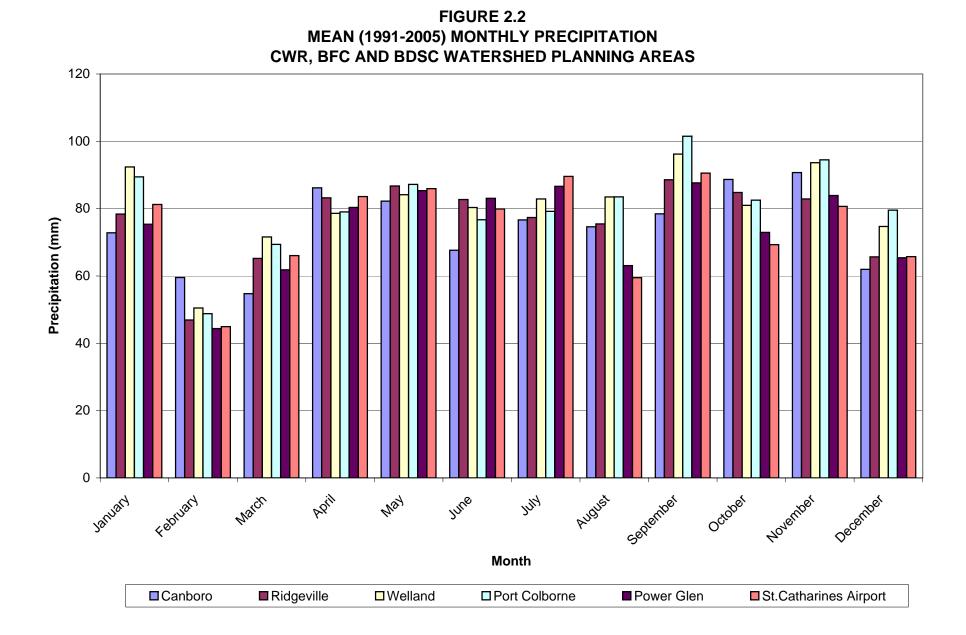


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FIGURES







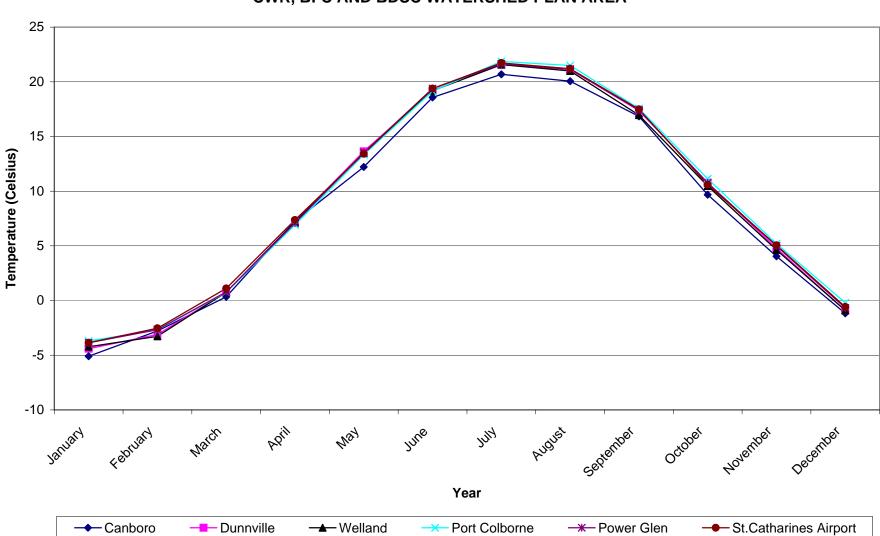
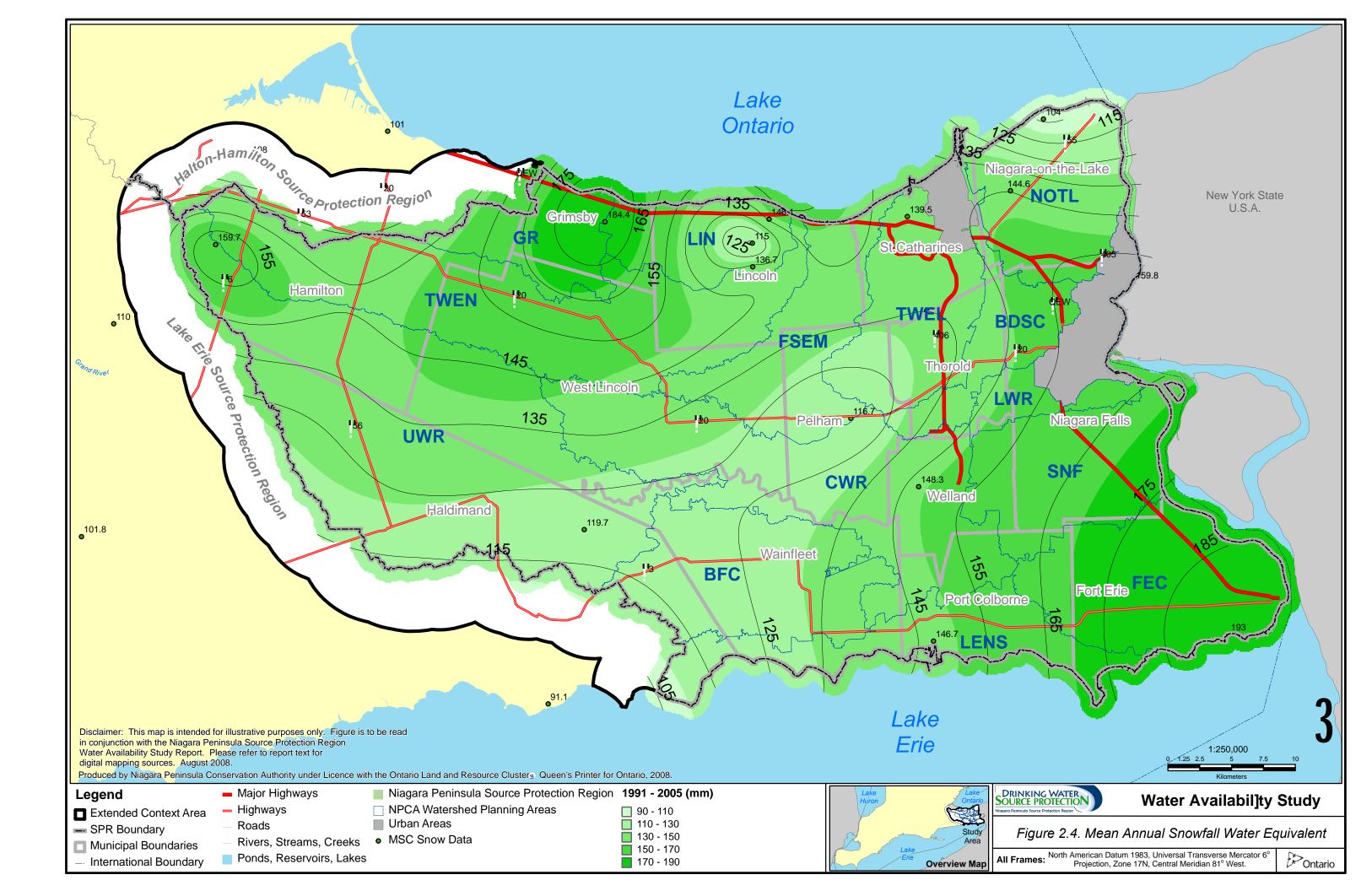
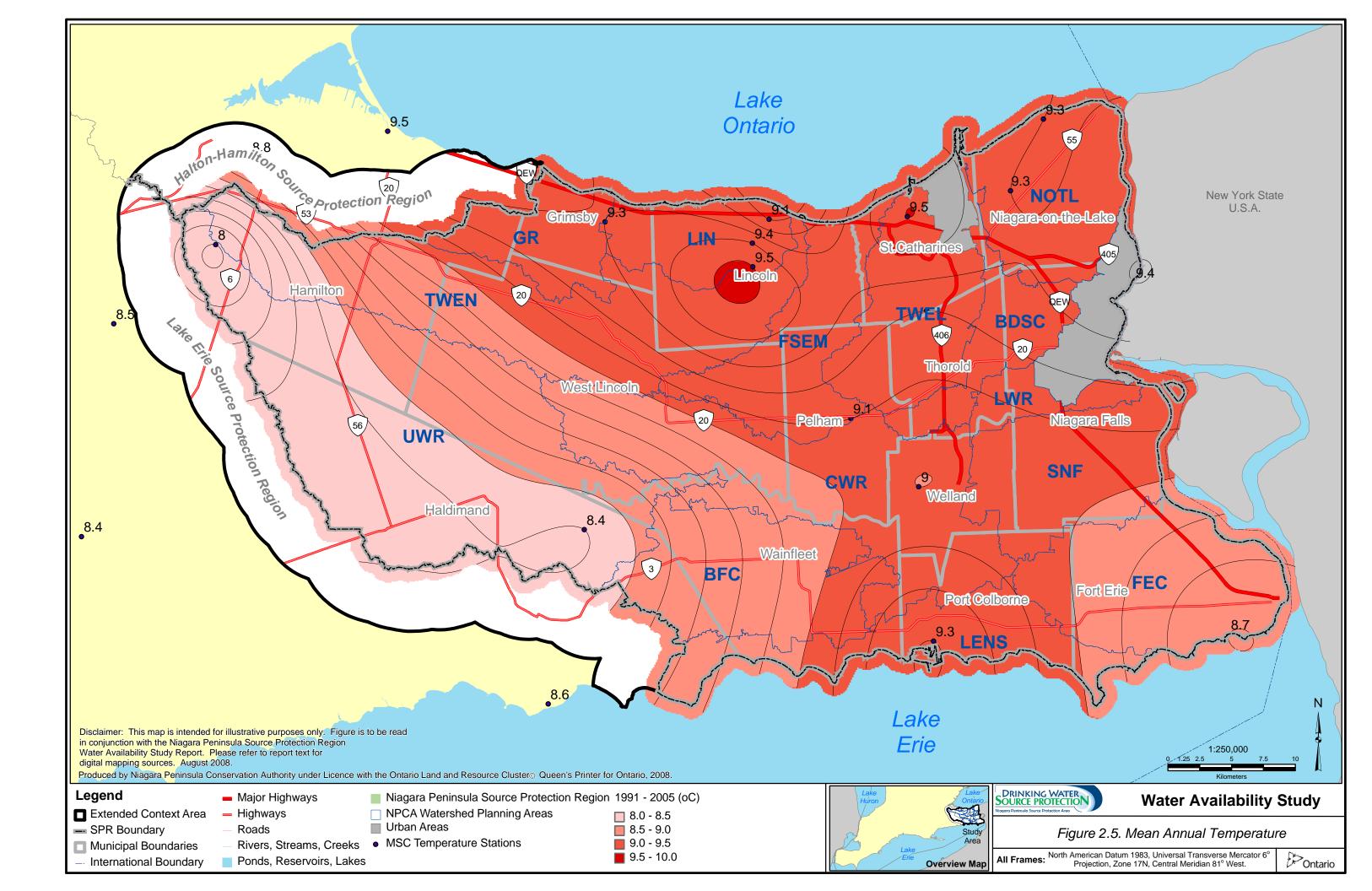
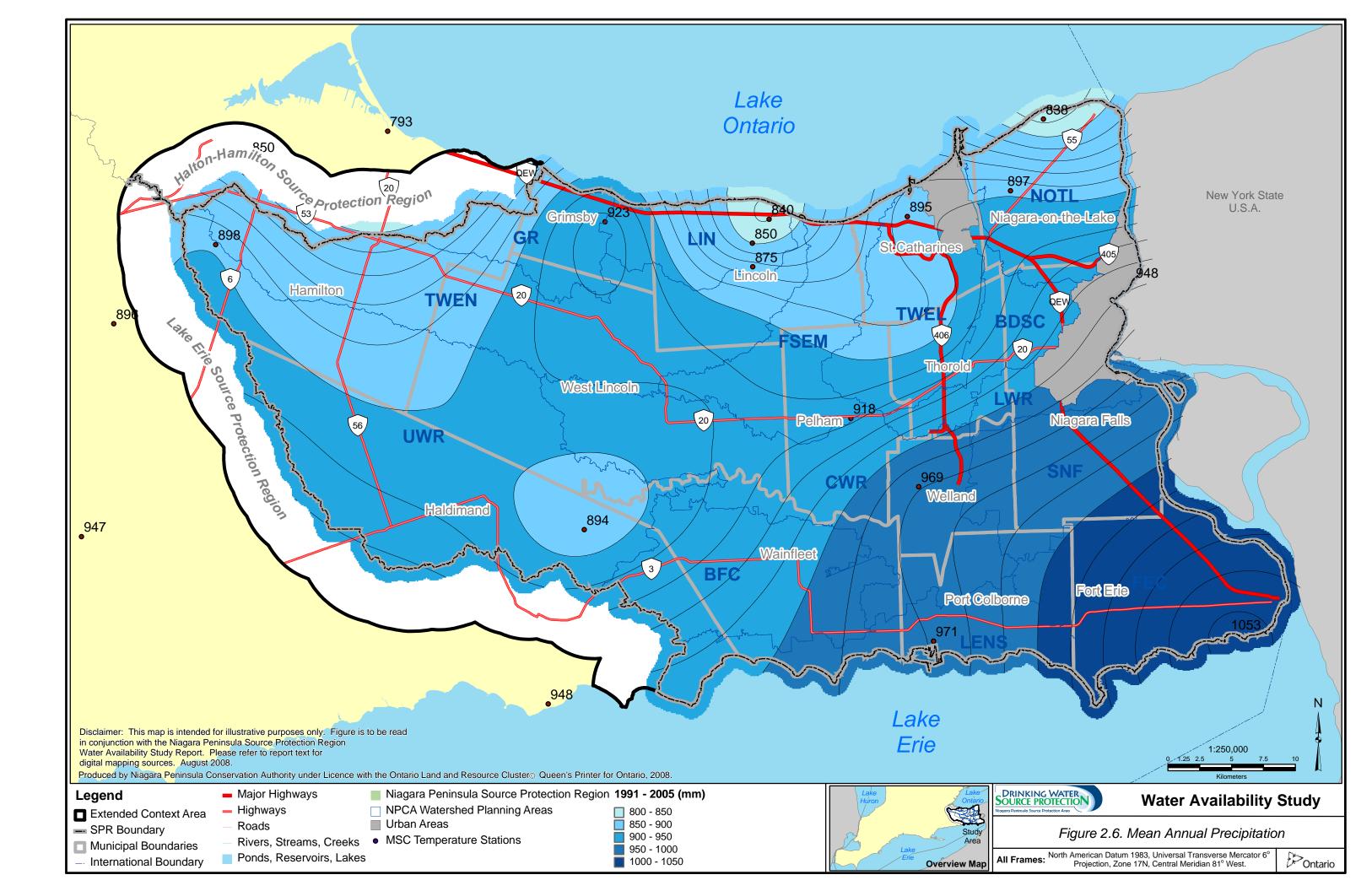
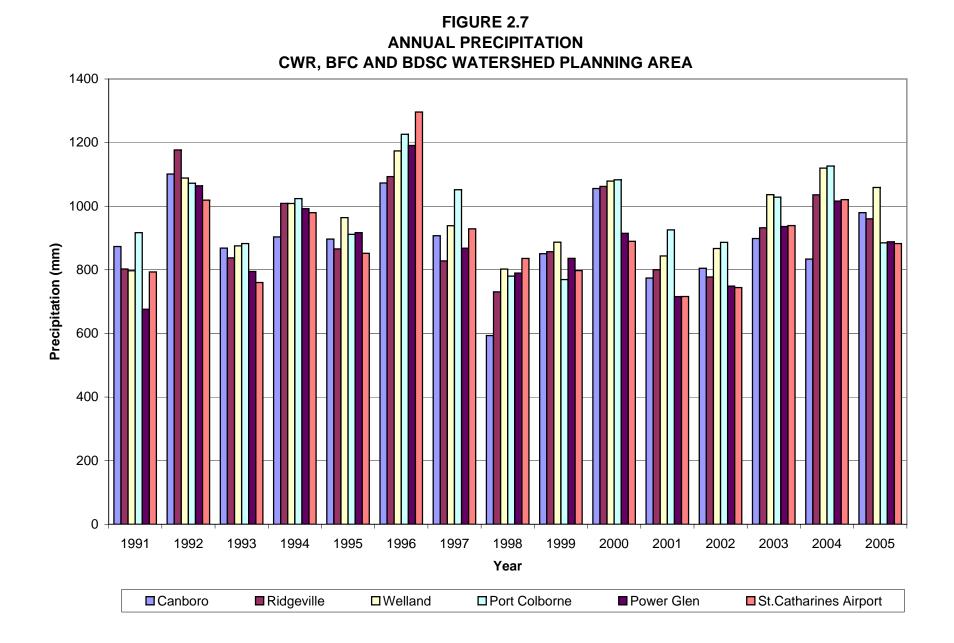


FIGURE 2.3 MEAN (1991-2005) MONTHLY TEMPERATURE CWR, BFC AND BDSC WATERSHED PLAN AREA









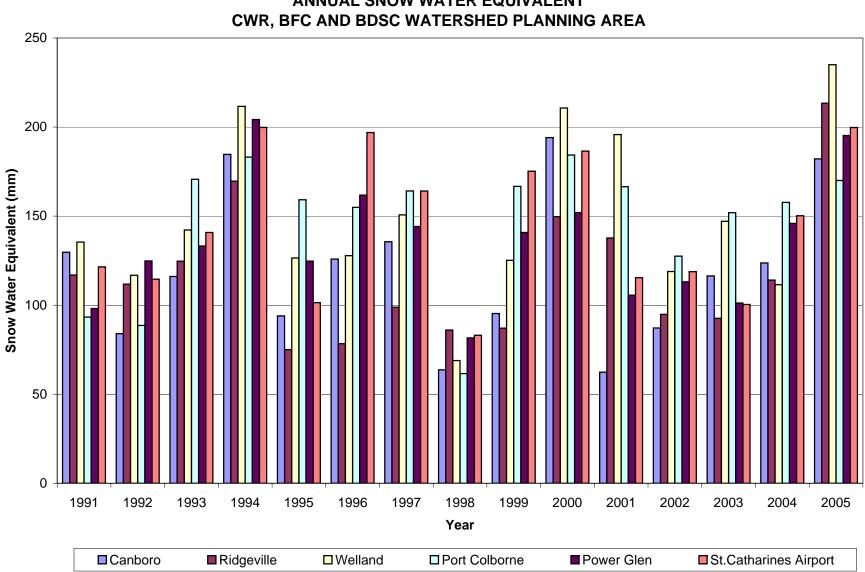


FIGURE 2.8 ANNUAL SNOW WATER EQUIVALENT CWR, BFC AND BDSC WATERSHED PLANNING AREA

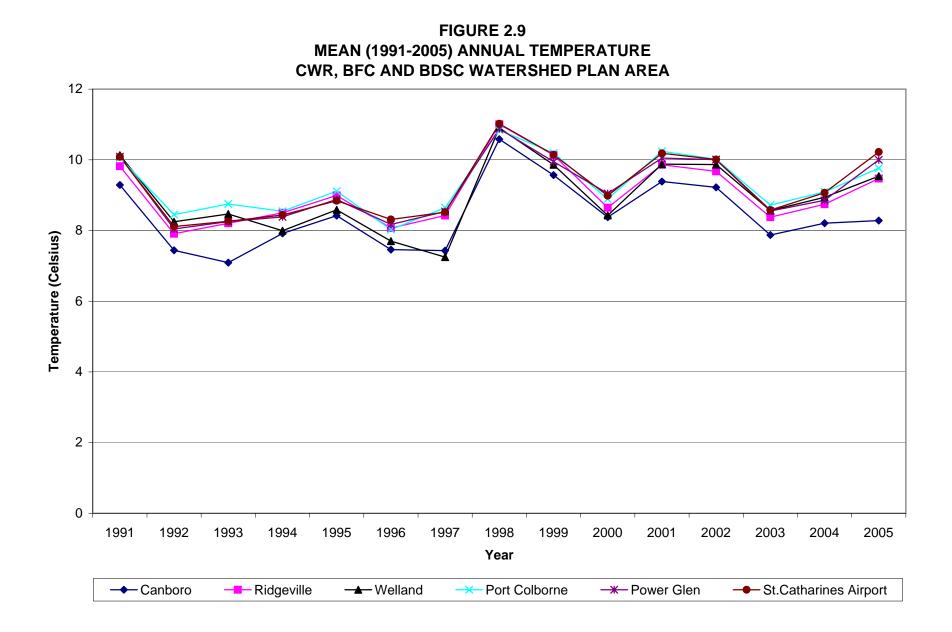


Figure 2.10 ANNUAL NET SOLAR RADIATION

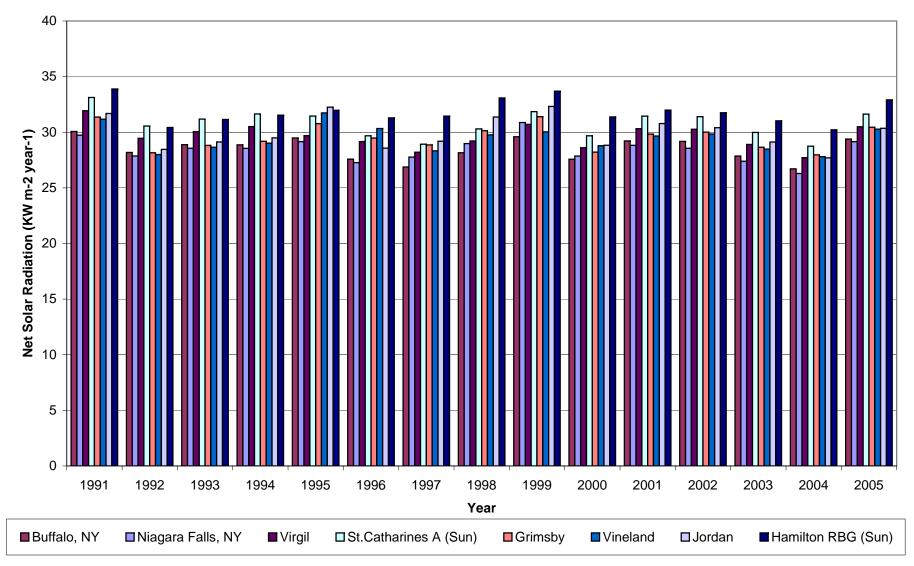


Figure 2.11 MEAN (1991-2005) MONTHLY NET SOLAR RADIATION

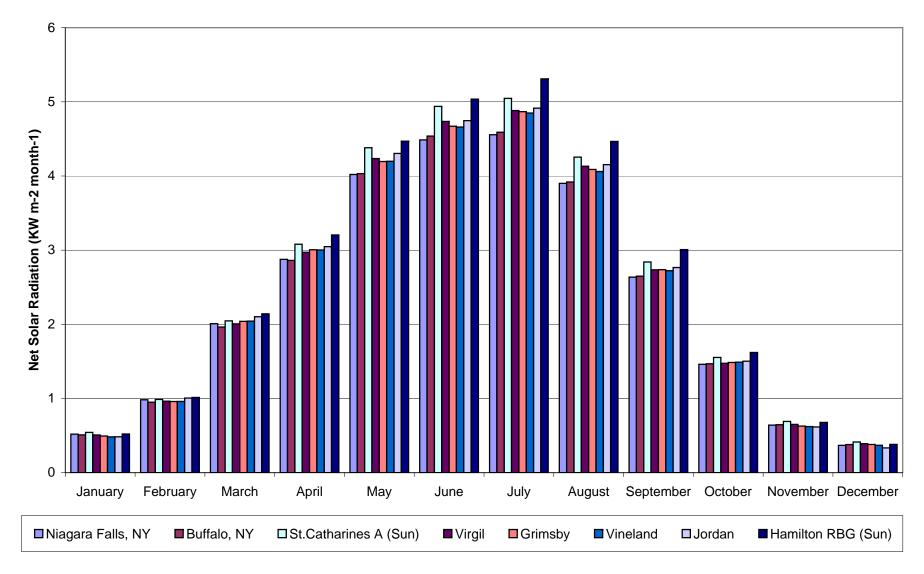


Figure 2.12a Channel Profile of Central Welland River

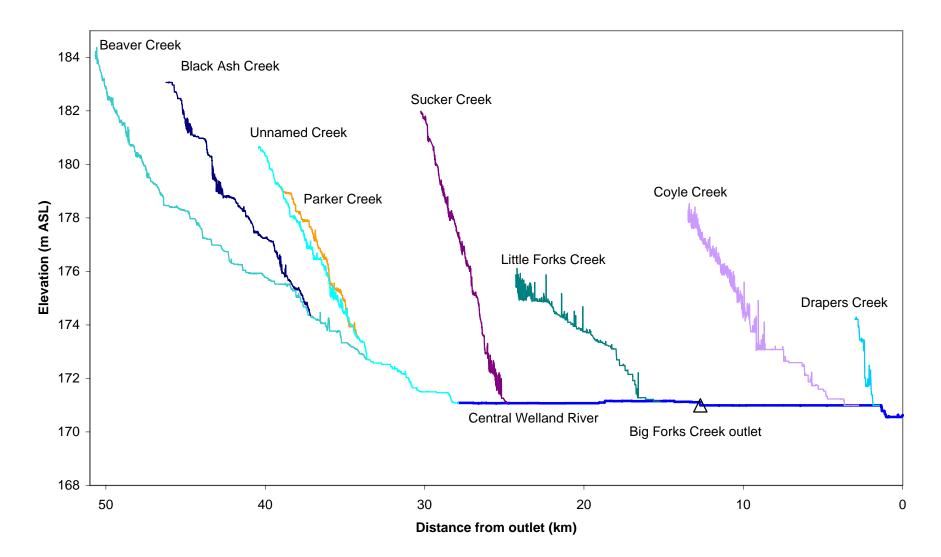


Figure 2.12b Channel Profile of Big Forks Creek

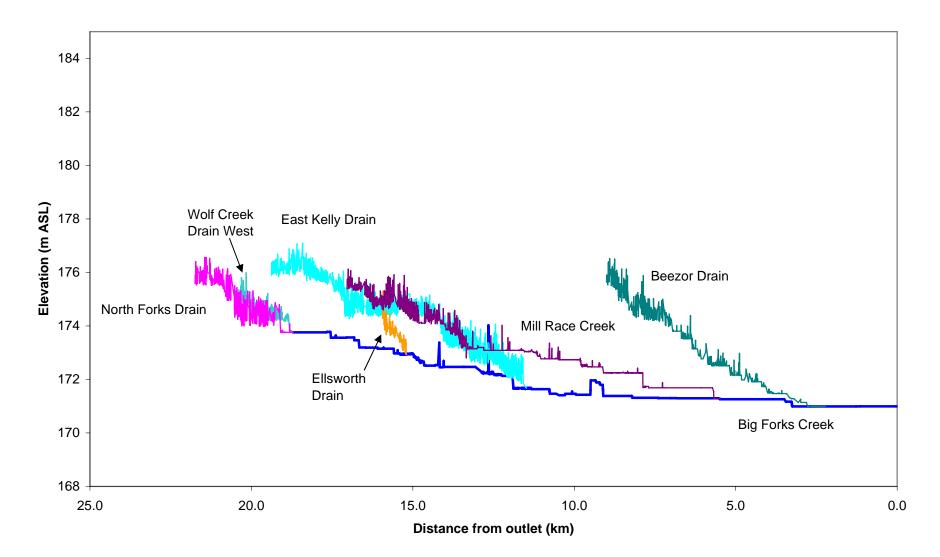


Figure 2.12c Channel Profile of Welland Canal South

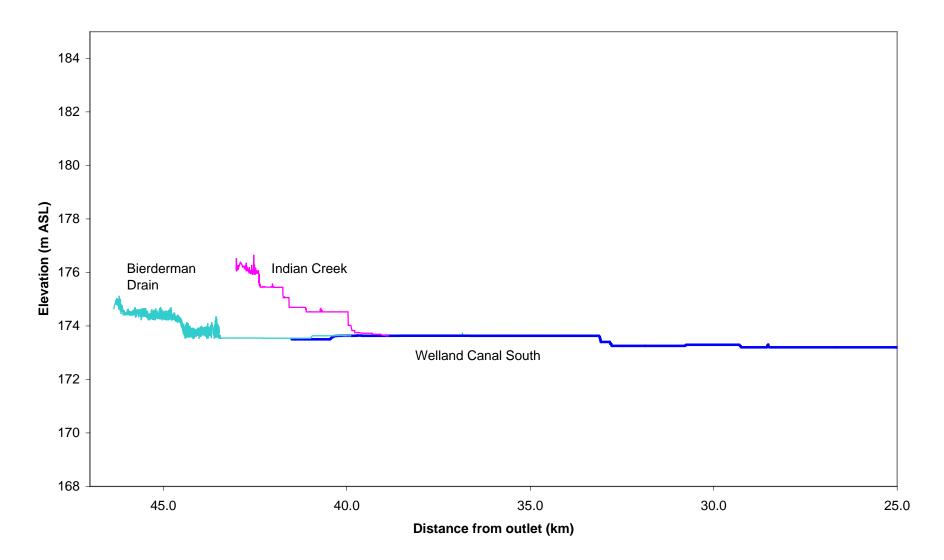
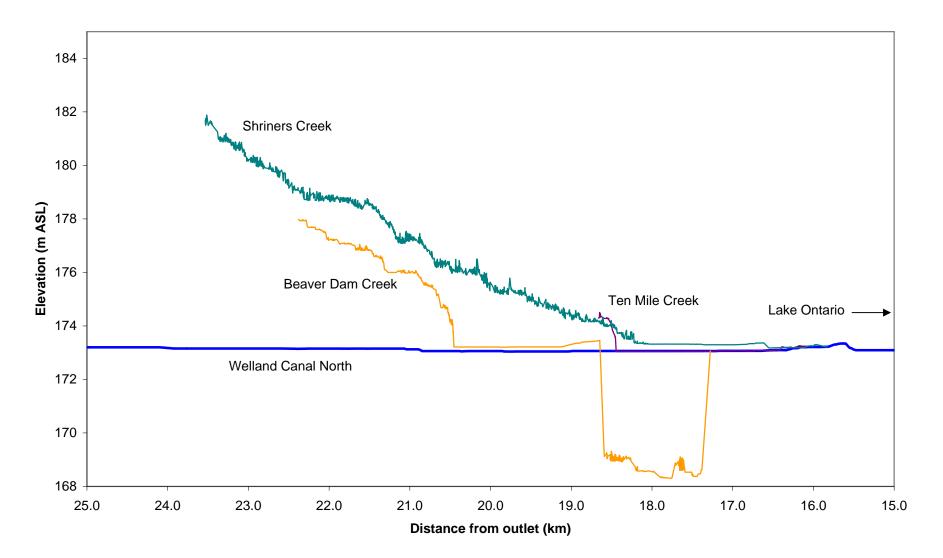
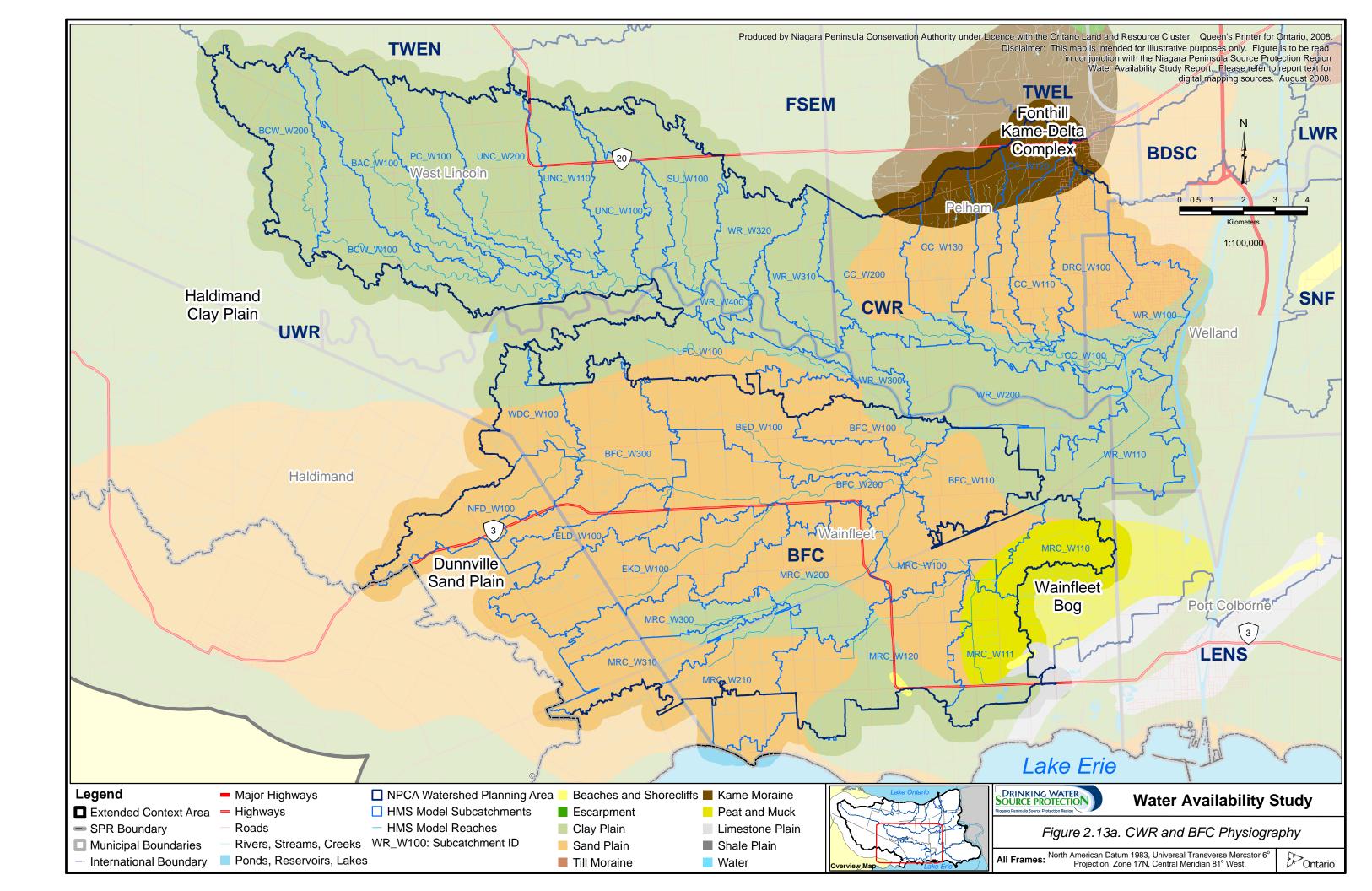
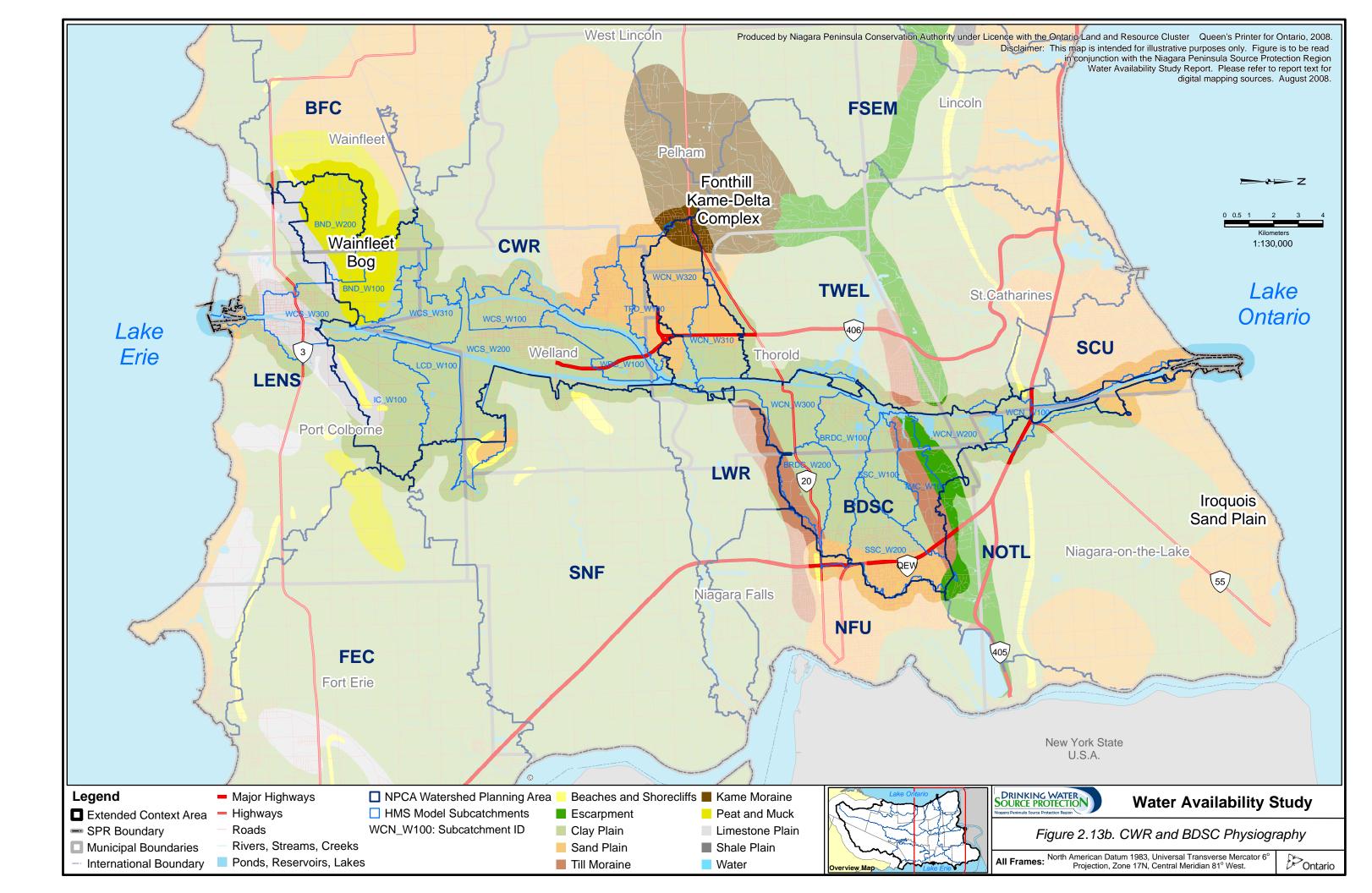
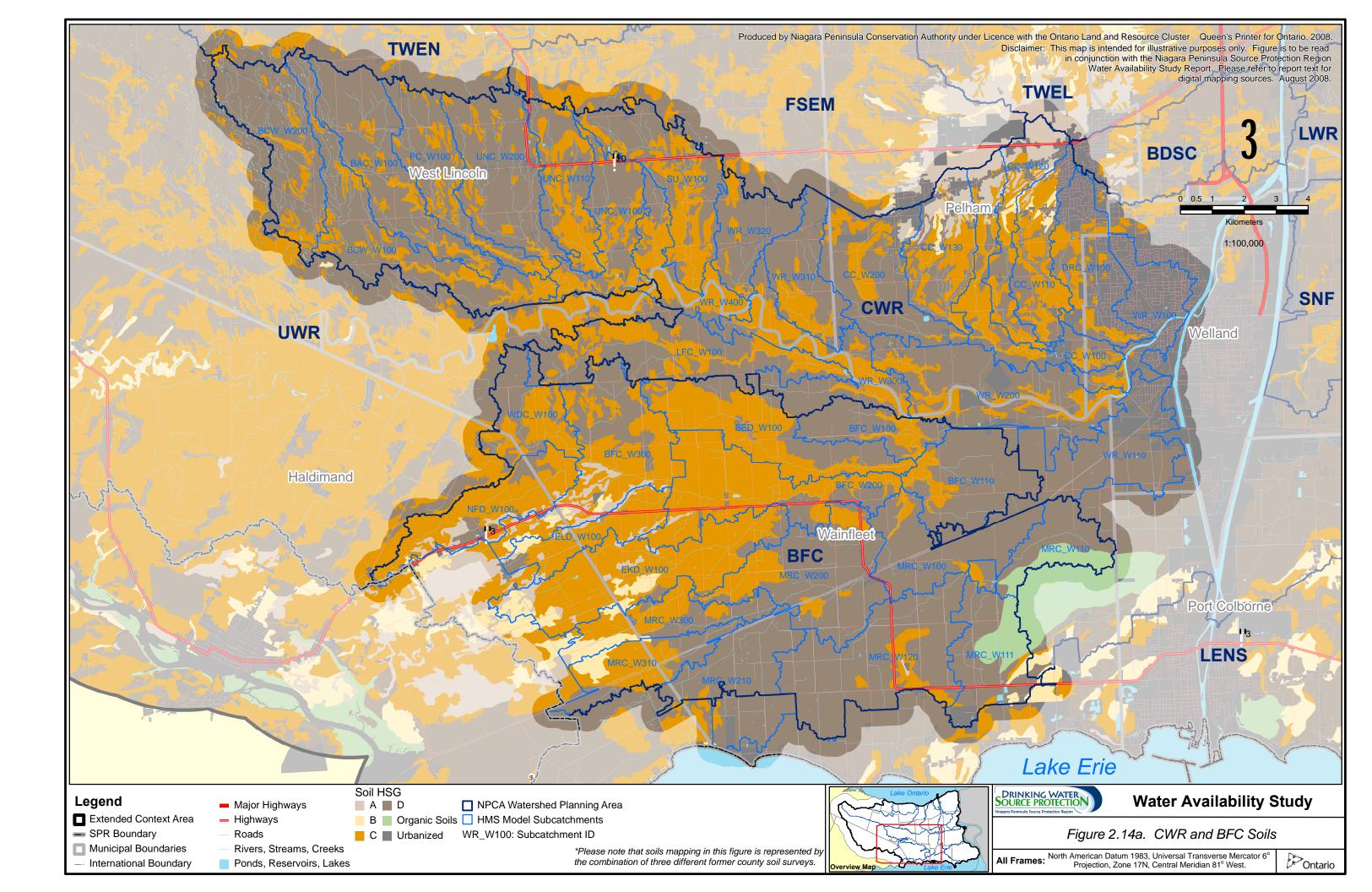


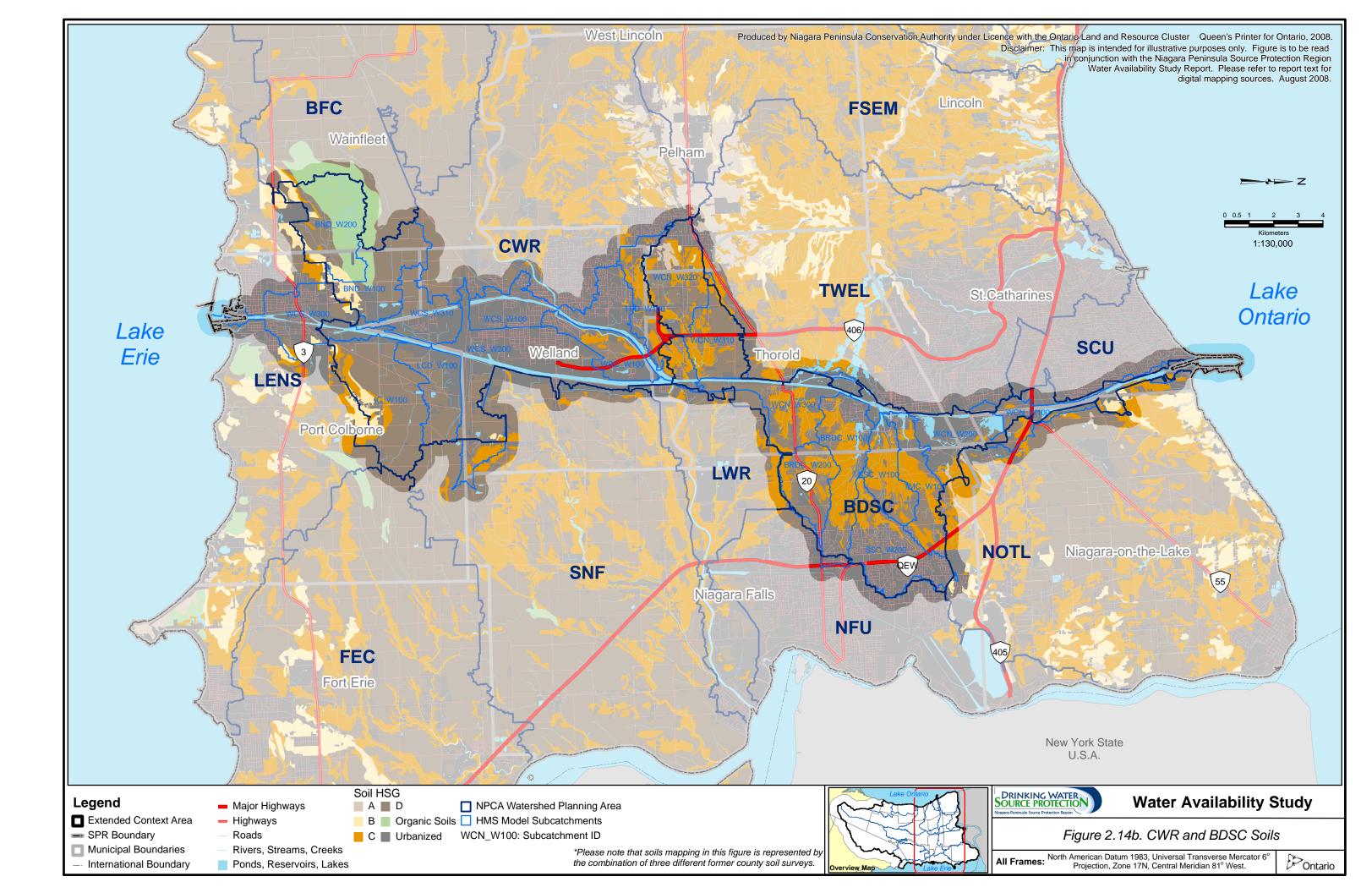
Figure 2.12d Channel Profile of Beaverdams Shriners Creek (Welland Canal North)

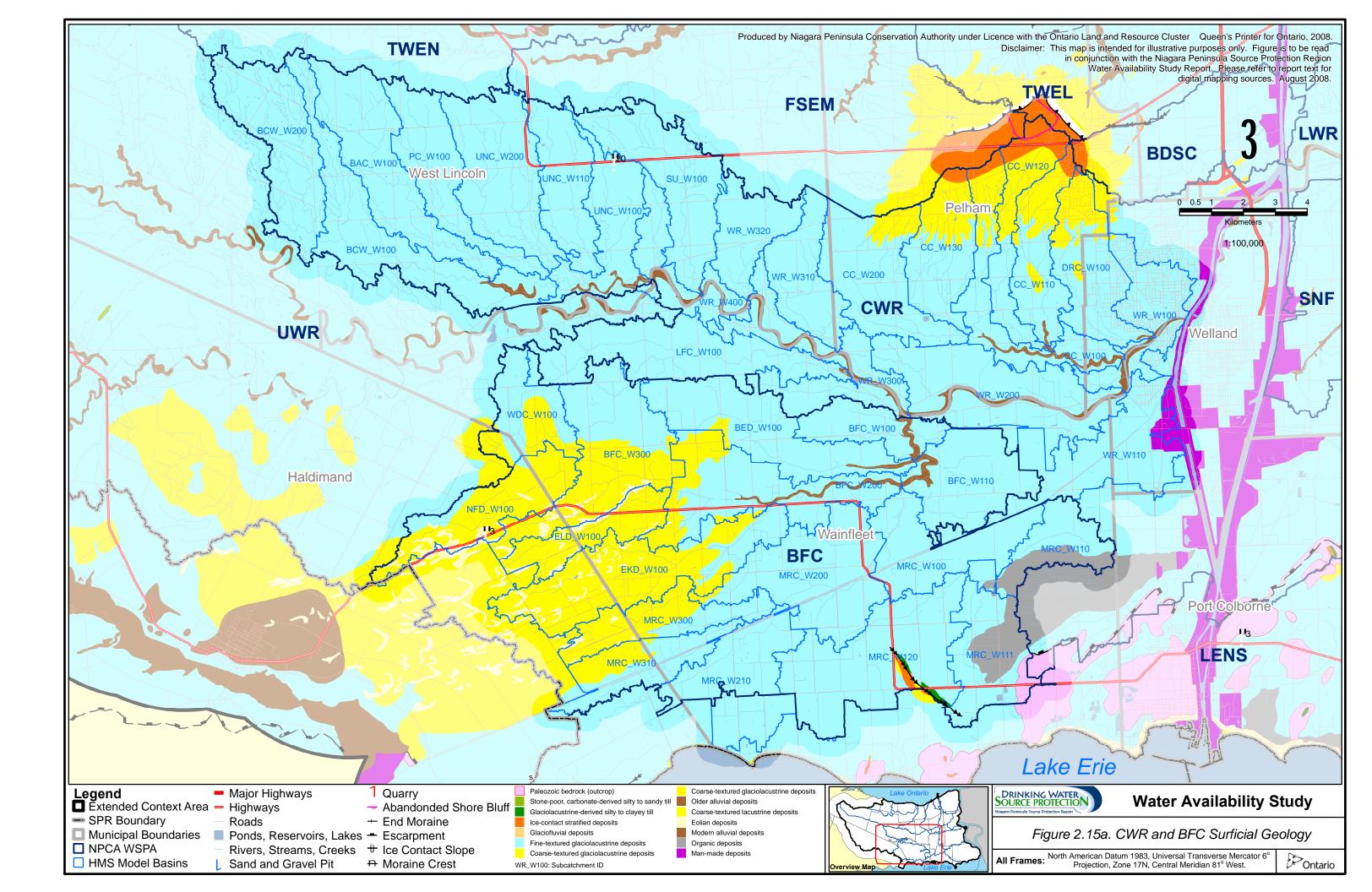


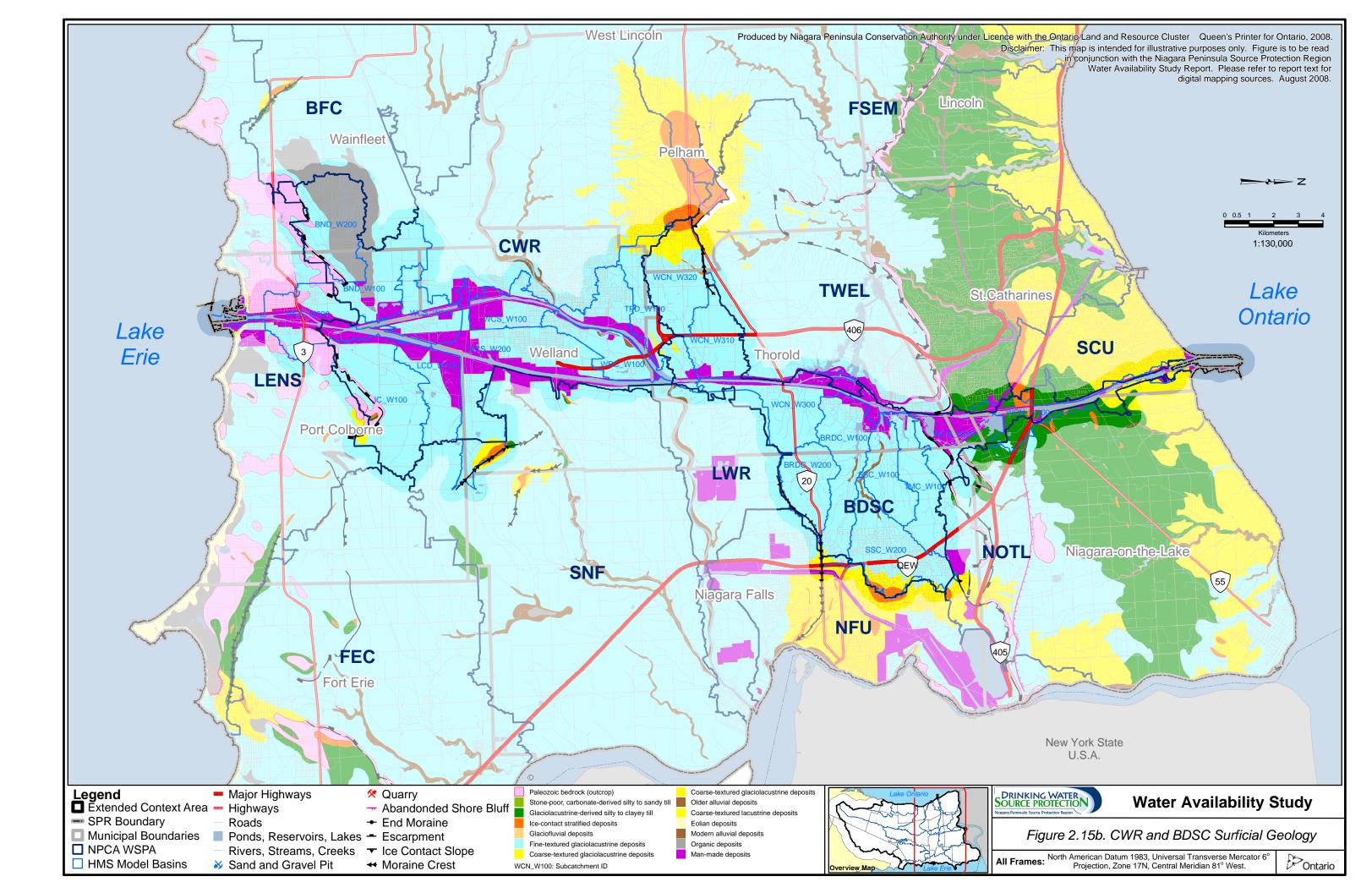


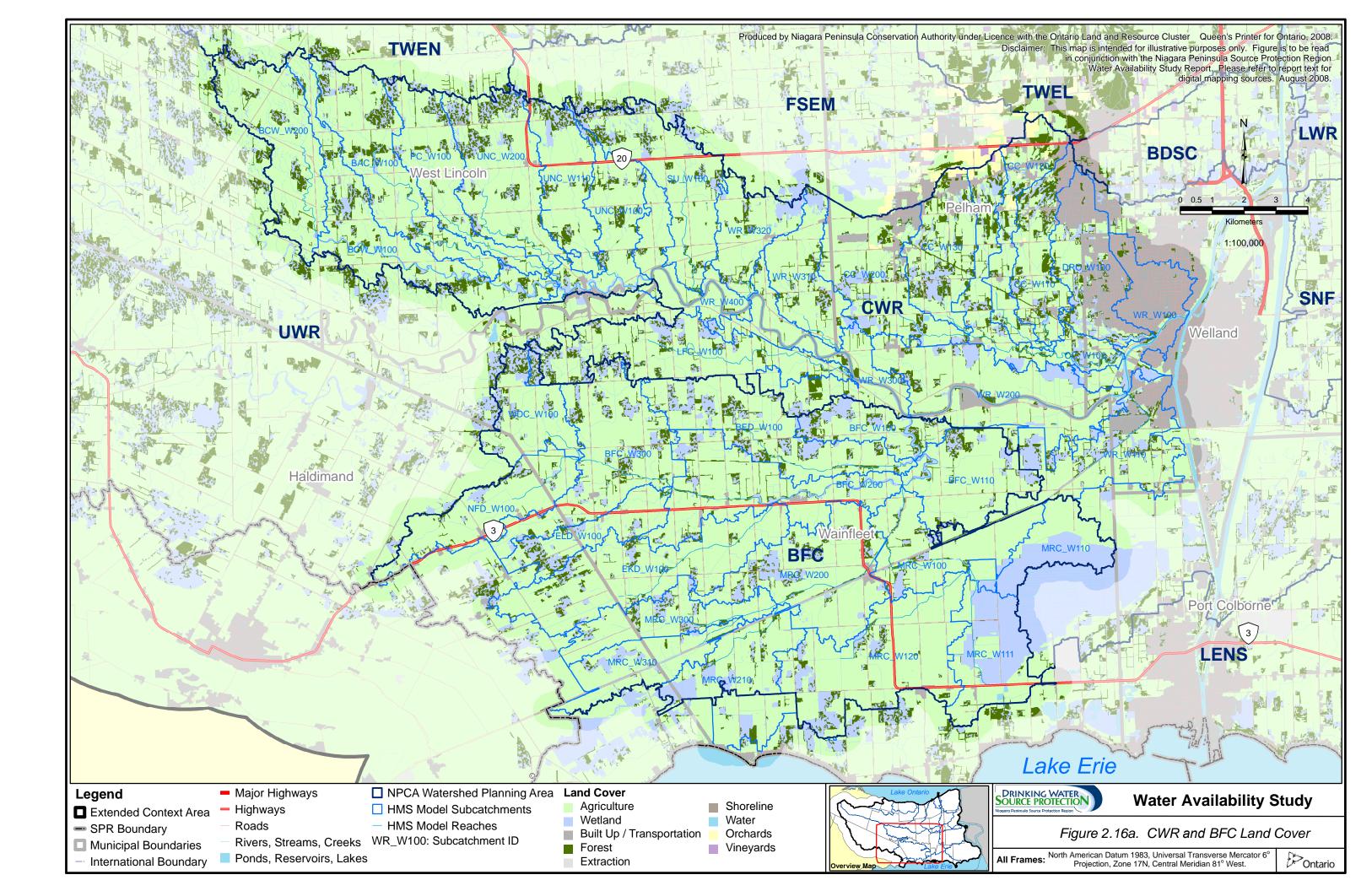


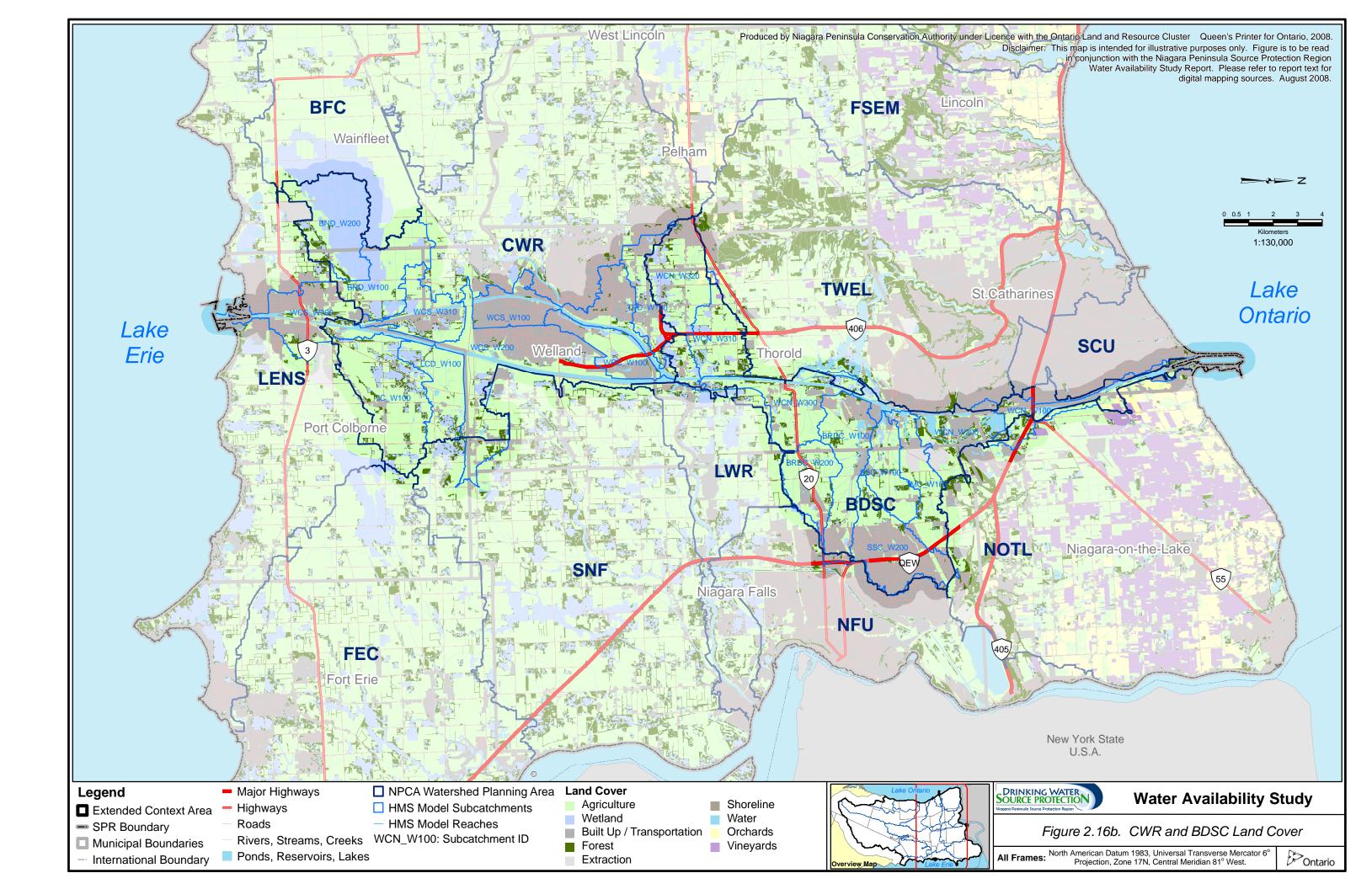




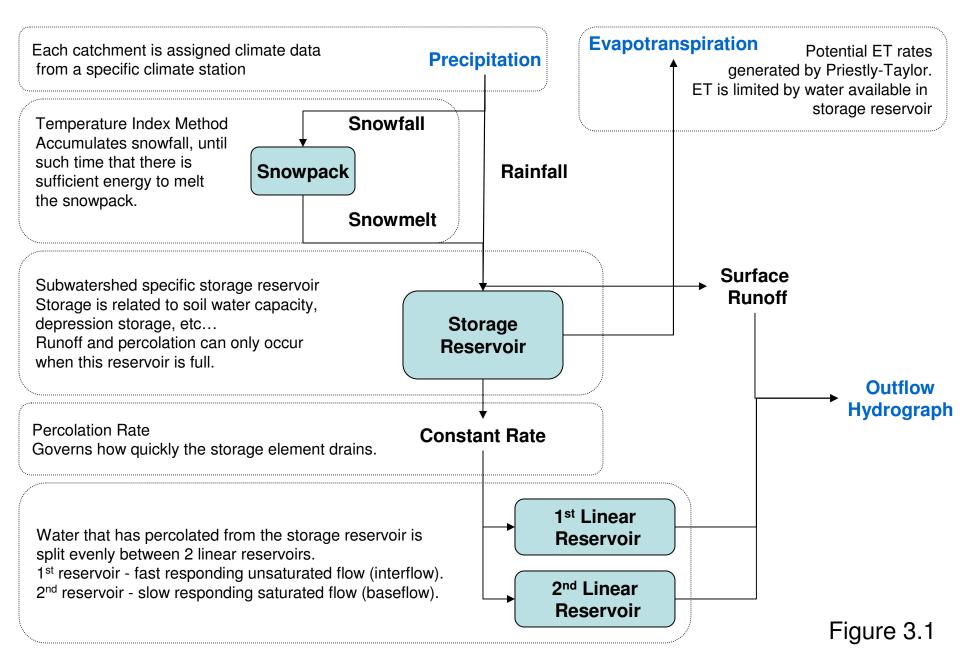


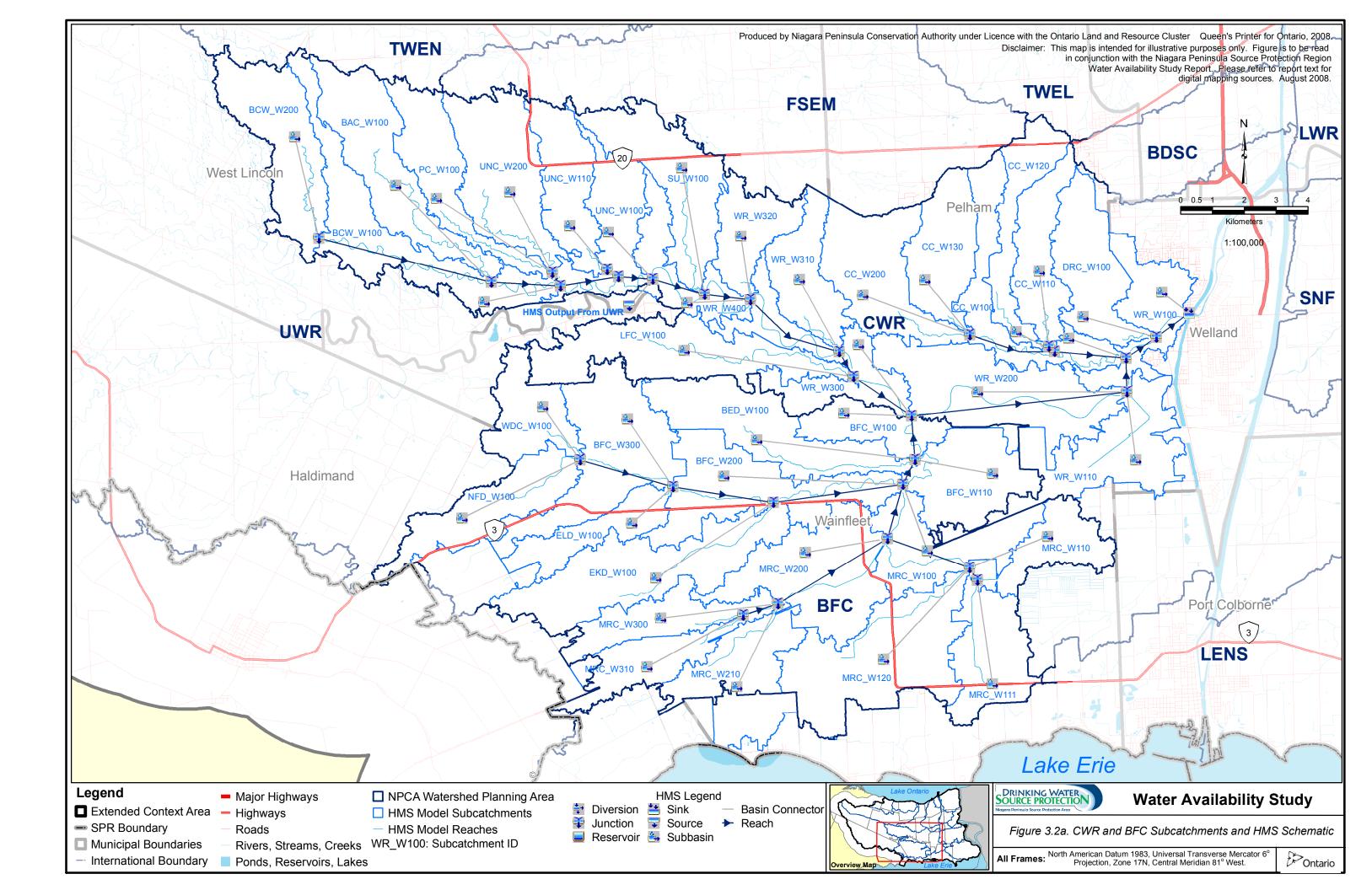


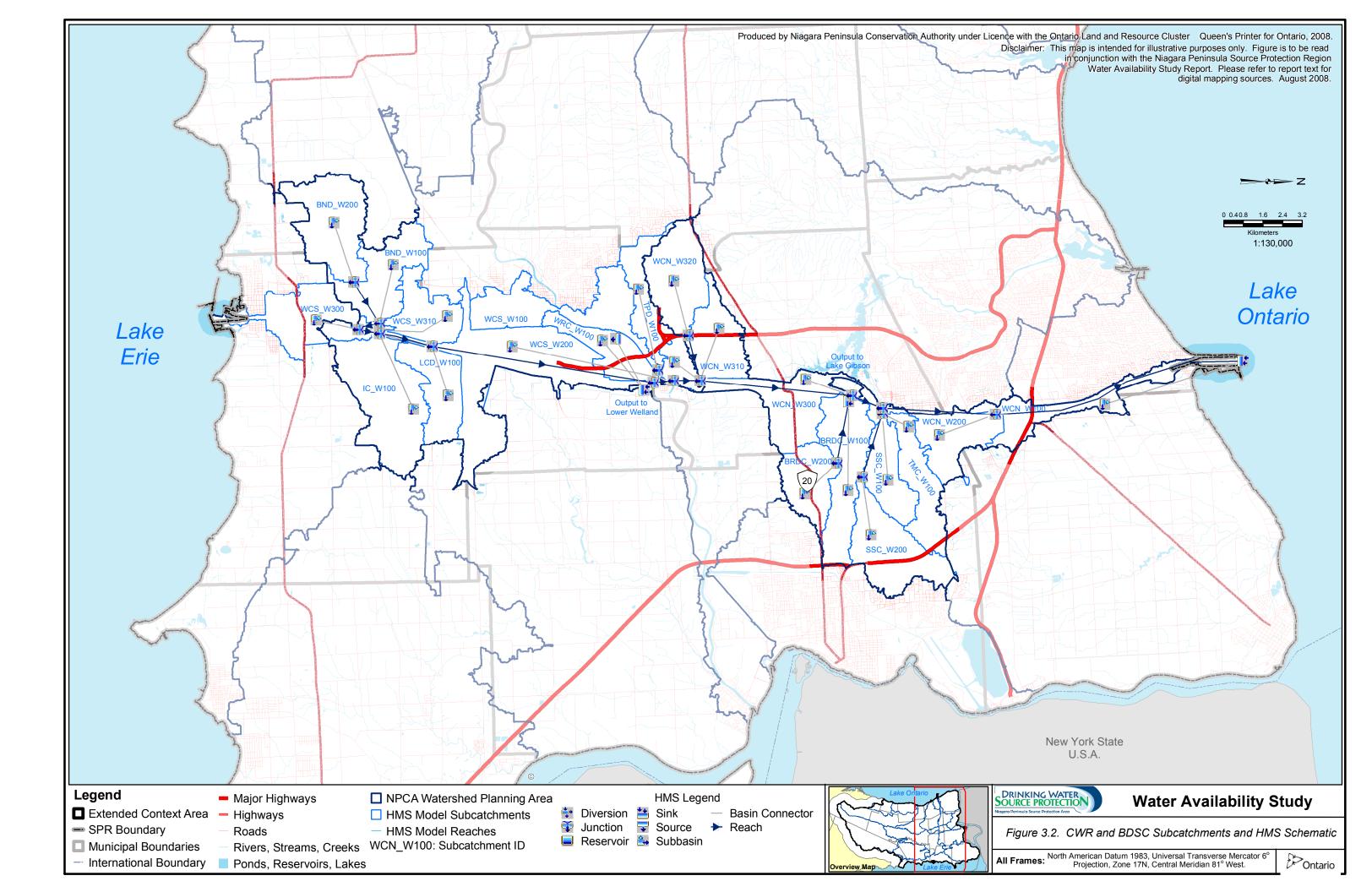


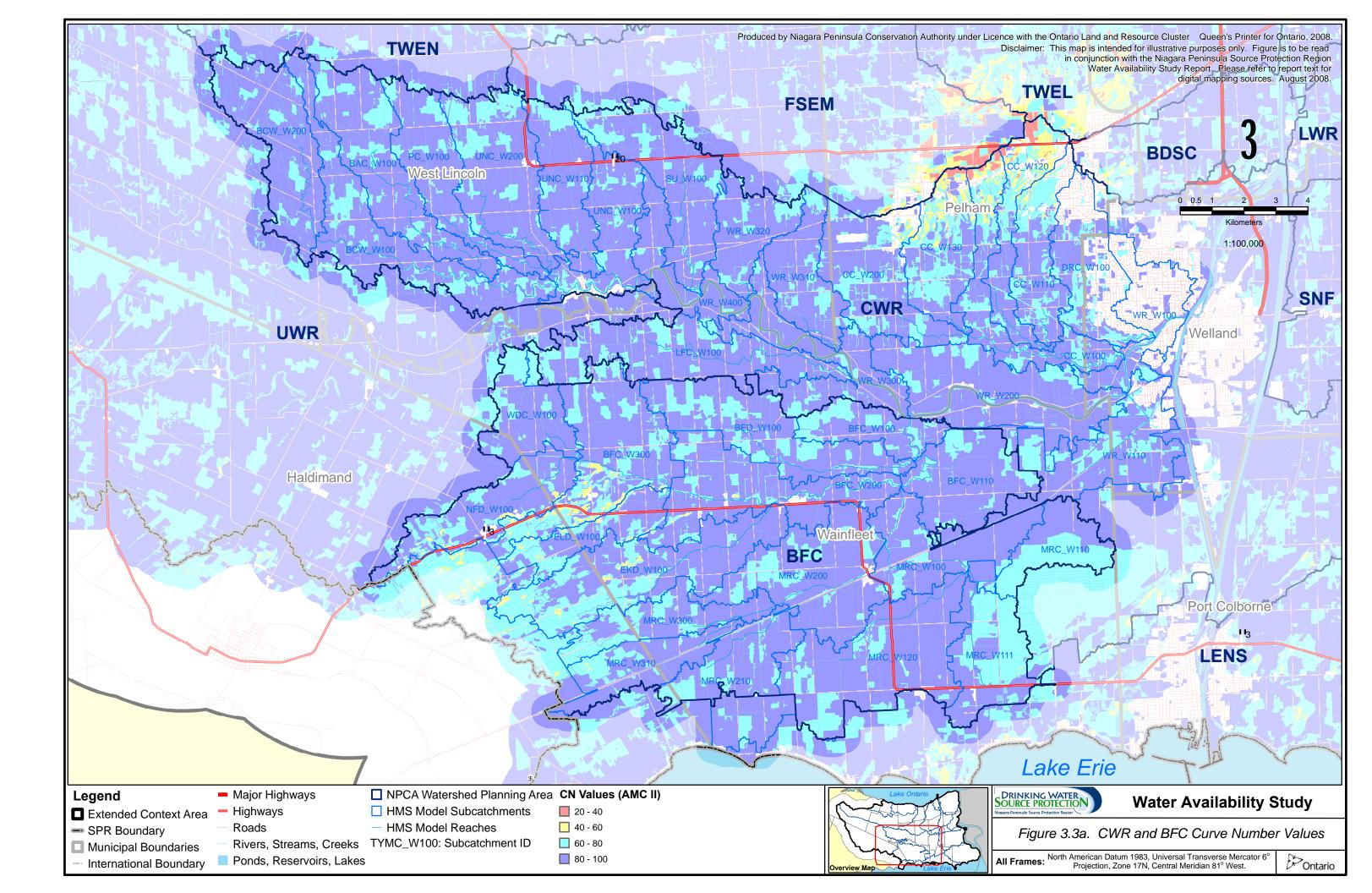


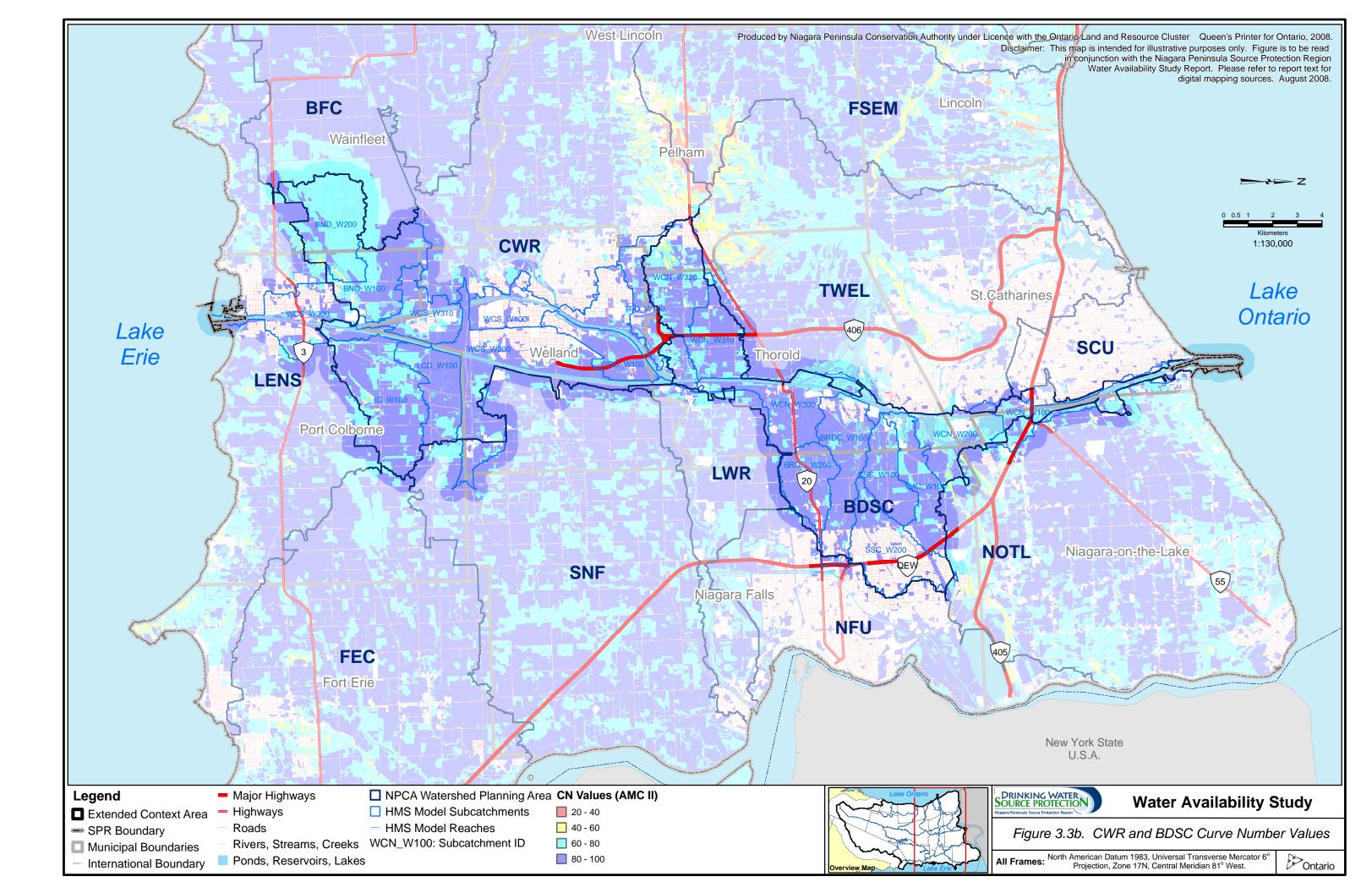
Conceptualization of Hydrologic Processes in HEC-HMS











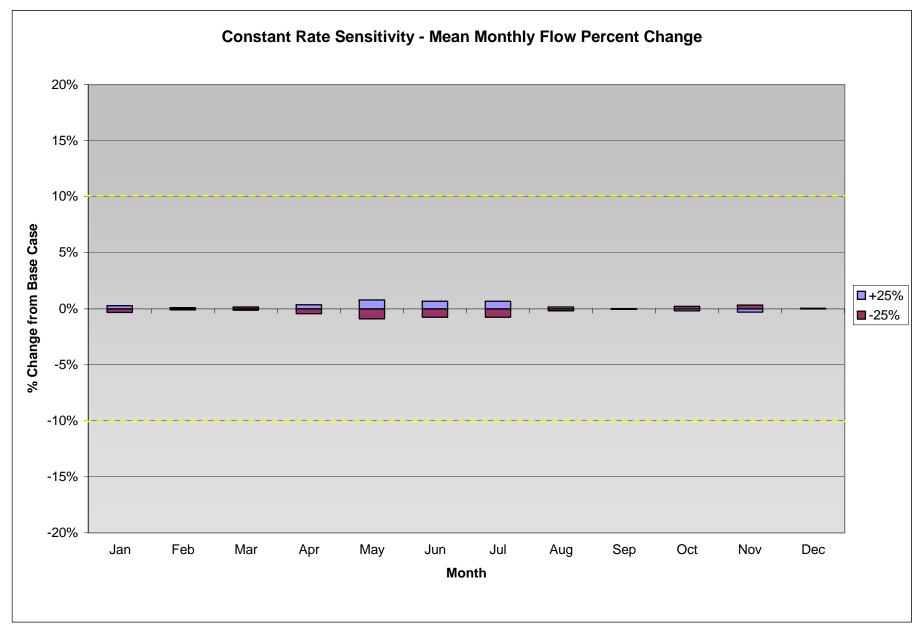


Figure 3.4 Central Welland River and Big Forks Creek

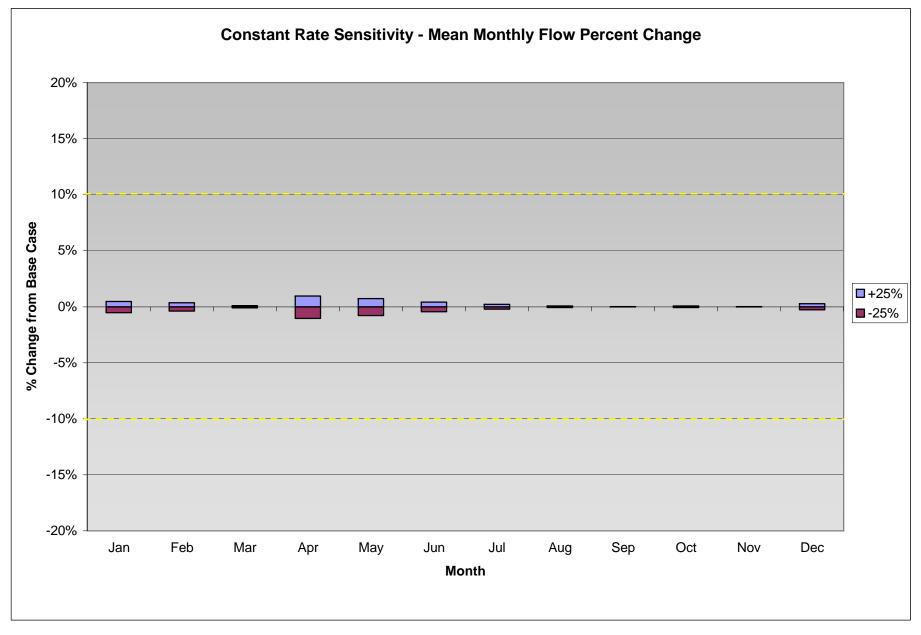


Figure 3.5 Beaverdams and Shriners Creeks

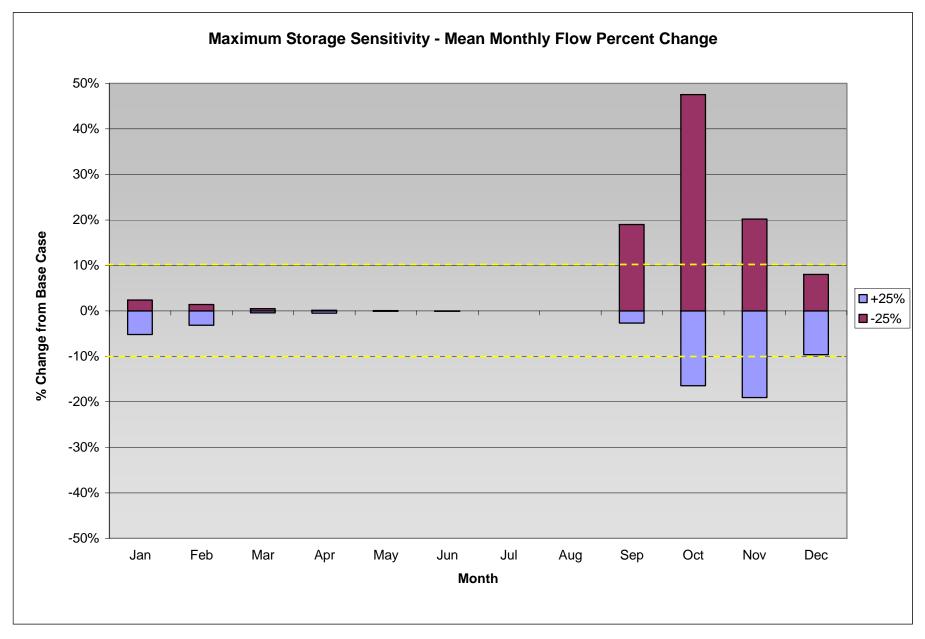


Figure 3.6 Central Welland River and Big Forks Creek

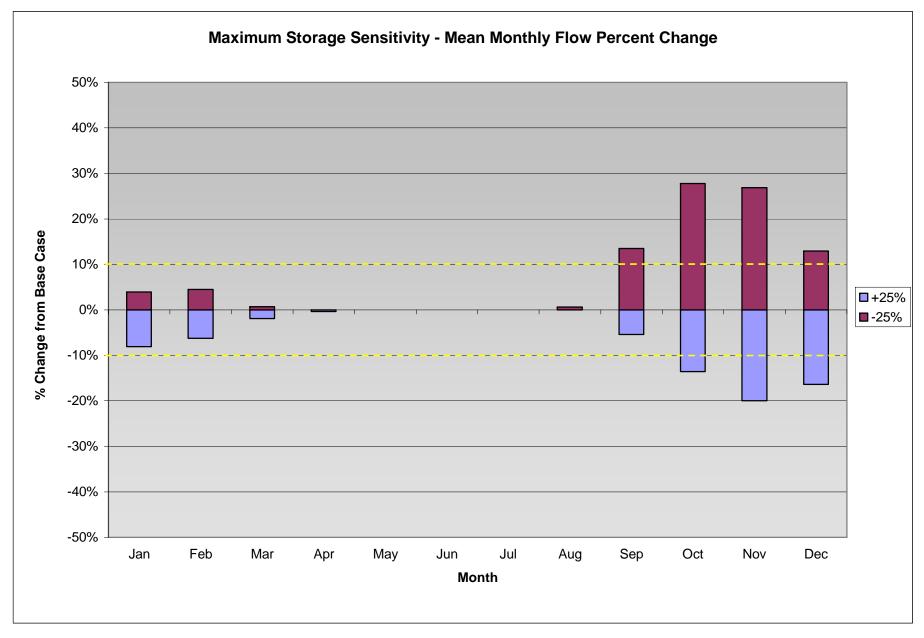


Figure 3.7 Beaverdams and Shriners Creeks

Appendix A

Snow Modeling Overview

Dr. Steven F. Daly

USACE ERDC/CRREL Hanover, NH 03755



US Army Corps of Engineers

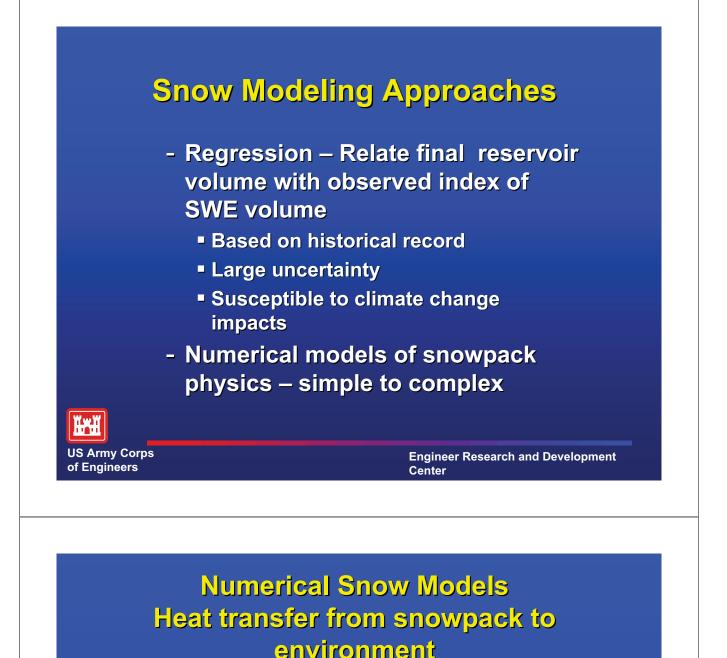
Engineer Research and Development Center

Snow Modeling can support our Snow Hydrology Goals

- Snow Accumulation
 - Estimation of the distribution of watershed snow water equivalent (SWE)
- Snow melt (Ablation)

Timing and magnitude of snowmelt





Simulate each heat transfer mode

- (Complete energy balance)
 - Data intensive
 - Varies widely due to slope, aspect, vegetation , elevation, etc.
- Simplify heat transfer by considering only key meteorological parameters (temperature index)
 - Air temperature
 - precipitation

US Army Corps of Engineers

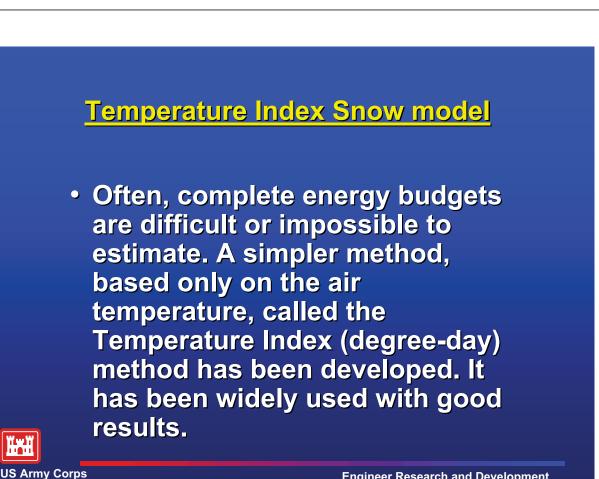
Numerical Snow Models Representing snowpack physical properties

- Multi-layer snow packs
 - Each layer with separate properties
 - Temperature, Density, Liquid water
- Single Snow Layer
 - Average snow properties
 - SWE, Cold Content, Liquid Water, etc

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Engineer Research and Development

Numerical Snow Models

Energy Balance

- Wind
- Temperature
- Water Vapor
- Radiation (Net)
- Precipitation
- Advection (rain)
- Ground
- Detailed (layered) snow pack

- Temperature Index
 - Temperature
 - Precipitation
- Single layer snow
 - SWE
 - Cold Content
 - Liquid water
- Calibration
 required

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Snowmelt Modeling in HMS

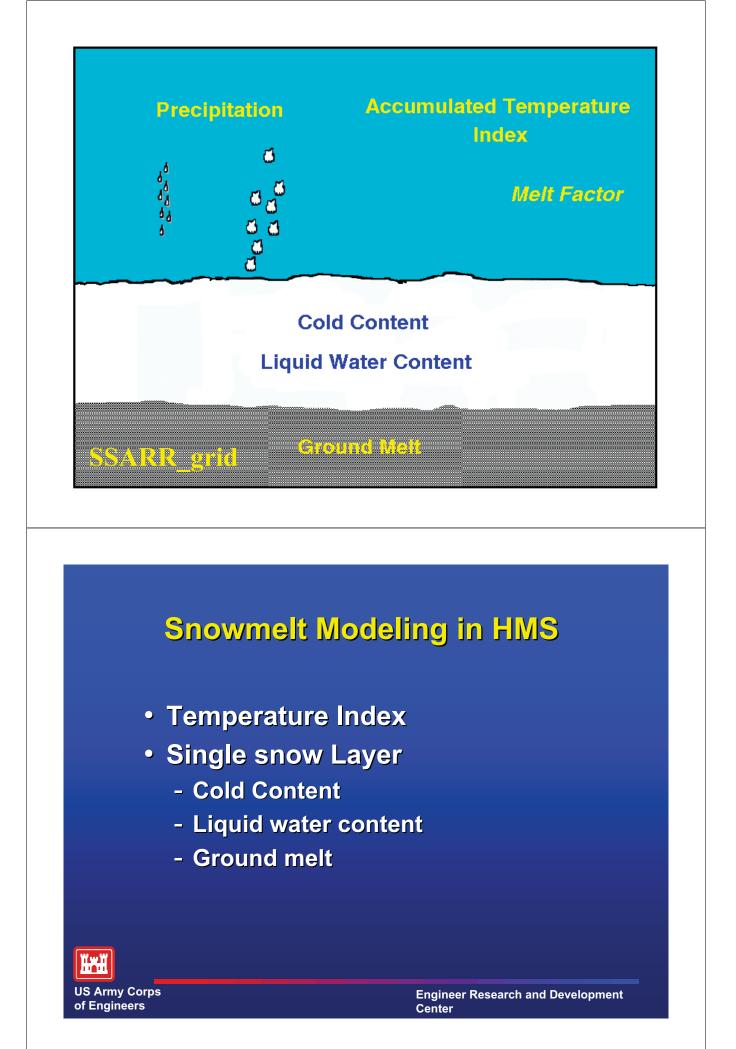
- Streamflow Synthesis And Reservoir Regulation – SSARR – North Pacific Division – NWD
 - Snow model, hydrology model, reservoir model for Pacific NW
- HEC-1 simple snow model
- SSARR snow model was made stand alone SSARR_grid
- SSARR_grid made into a distributed model Distributed Snow Process Model – DSPM
- SSARR_grid added to HEC-HMS

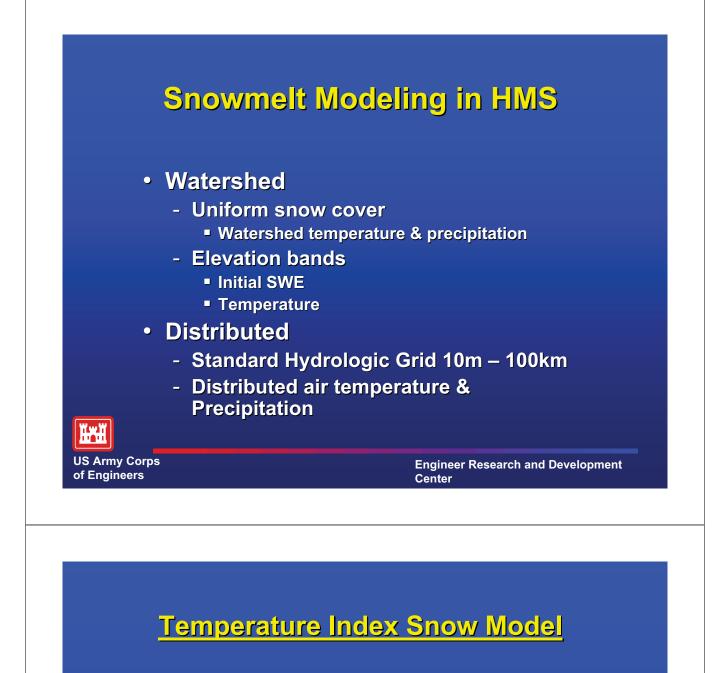


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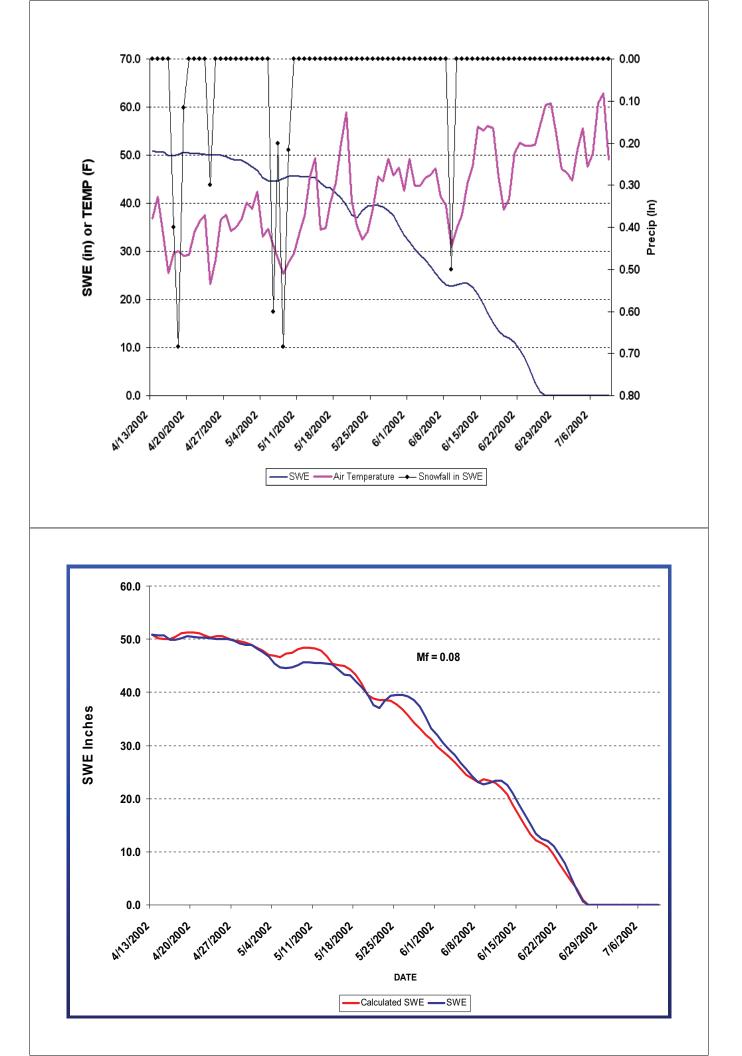
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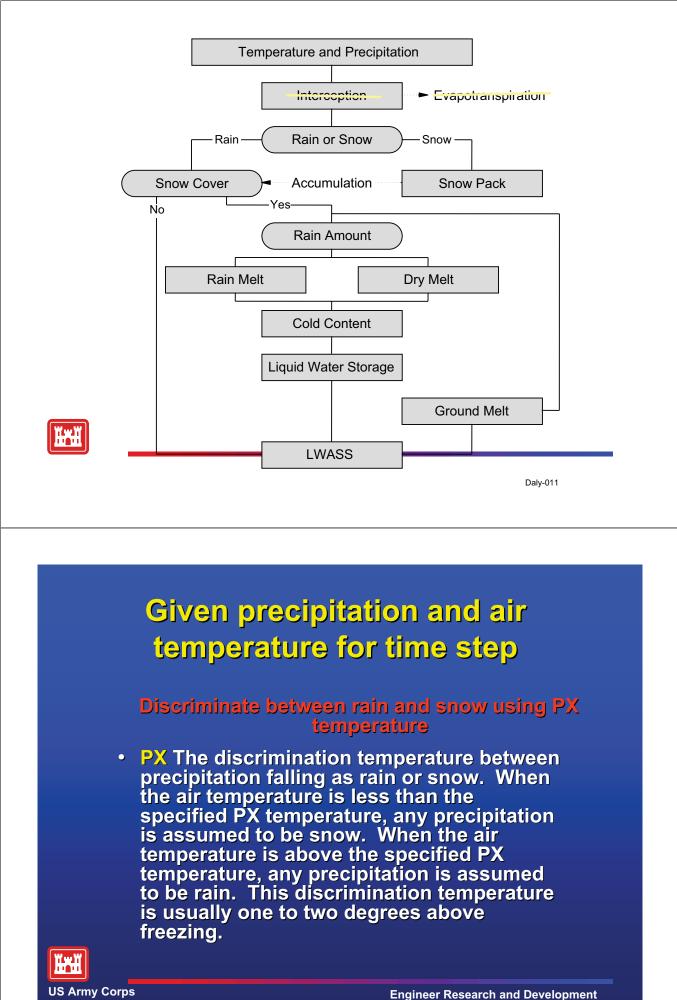




- Snow Melt is estimated as
 - $L_s = M_f (Ta T_{base})$
 - Mf serves as an index of the total heat transfer at the snow surface which includes long wave, short wave, latent heat, and sensible heat transfer. Also strongly influenced by the wind speed, aspect, slope, vegetation, etc.
 - Mf can be a constant, or set as a function of the accumulated thawing days or set as a function of the month of year.







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Given precipitation and air temperature for time step

Discriminate between melt and non-melt using Base temperature

• Base Temperature. The difference between the base temperature and the air temperature defines the temperature index used in calculating snowmelt. The meltrate is multiplied by the difference between the air temperature and the base temperature to estimate the snowmelt amount. If the air temperature is less than the base temperature, then the amount of melt is assumed to be zero. Typically, the base temperature should be 32F (0C) or close to it.



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Given precipitation and air temperature for time step

Discriminate between melt and non-melt using Base temperature

T <T_{base} - No Melt
 SWE accumulates T< PX; P>0

• T >T_{base} - Melt



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T >T_{base} - Melt

Discriminate between rain melt and dry-melt using rain rate limit

 The rain rate limit. The discrimination rain rate in inches/day between dry melt and wet melt. The wet meltrate is applied as the meltrate when it is raining at rates greater than the rain rate limit. If the rain rate is less than the rain rate limit, the meltrate is computed as if there were no precipitation.



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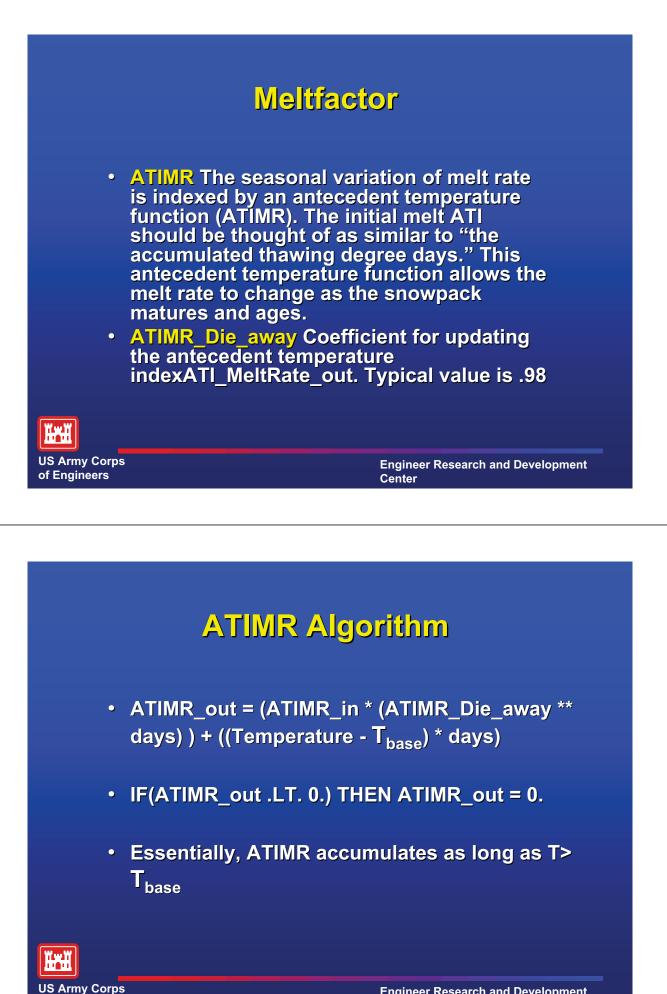
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Meltfactor

L_s = M_f (Ta – T_{base})

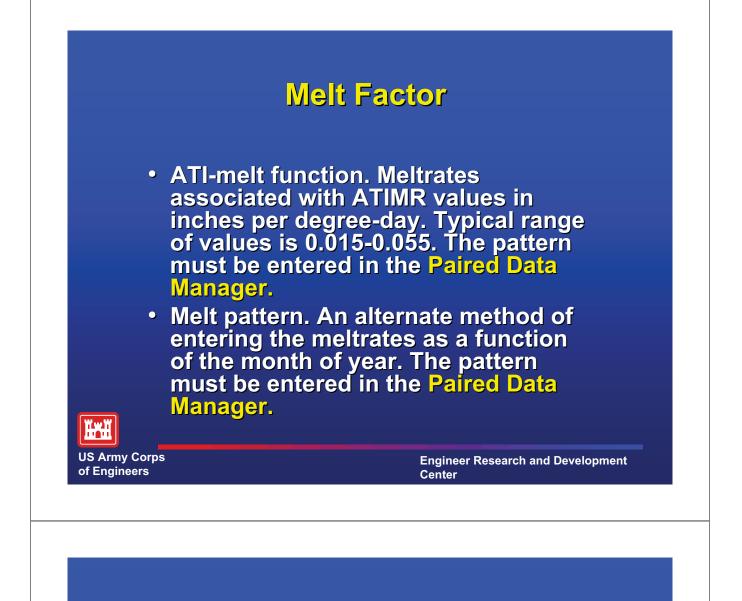
 Meltfactor can be constant or variable with time. A variable meltfactor recognizes that as snowpack matures the heat transfer rates change AND/OR different components of heat transfer change in importance.





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Typical Melt Factors

- 0.04-0.08 inches/F-day
- 1.8-3.7 mm/C-day
 - Average daily temperature
 - No rain
 - Not heavily forested
 - No extreme conditions high winds, etc
- Thin ephemeral snowpacks that melt out in a very short time may have a constant melt factor



Temperature Index Snow Model

• Rain melt

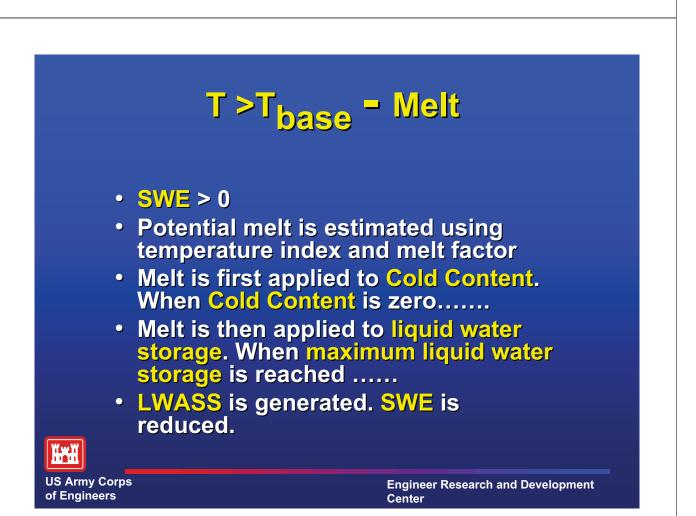
- Snow melt that occurs when the air temperature is above the snow/rain temperature and the precipitation rate is significant. Rain is assumed to fall at the air temperature
- Melt from rain and condensation of water vapor in the snowpack.

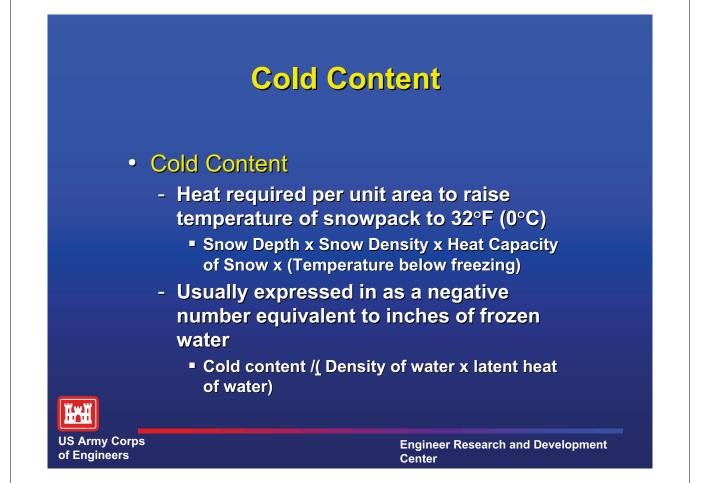
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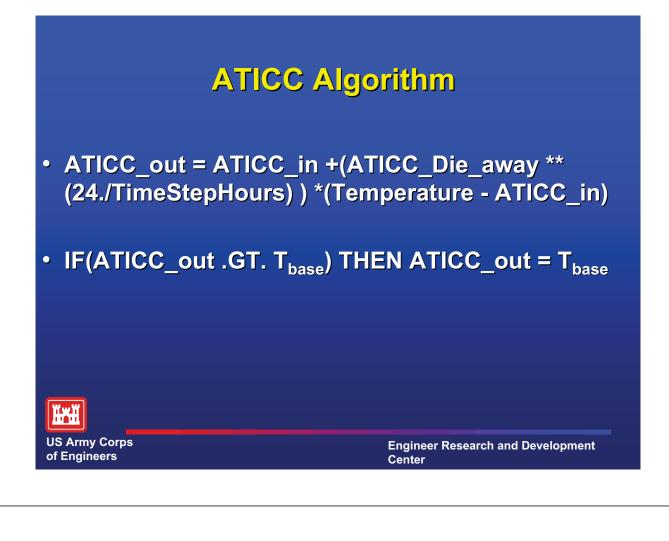


Cold Content -Parameters

- ATICC. The ATICC is an index to the snow temperature near the surface. It is calculated assuming an approximation to the transient heat flow equations. This value is used to estimate the cold content of the snow. It should be set to the approximate snowpack temperature if known. If not known, it can be set to 32F (0C).
- ATICC_die_away. Coefficient for updating the antecedent temperature index <u>ATI_ColdContent_out.</u> Typical value .84



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Cold Content -Parameters

- ATI cold function. Cold rates associated with the ATIMR values, in inches per degree-day. Typical range of values is 0.010-0.025. The pattern must be entered in the Paired Data Manager.
- interval_Cold = (Temperature ATICC_out)
 *coldRate / 24.
- ColdContent_out = ColdContent_in + interval_Cold *TimeStepHours



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Cold Content-Parameters

 ATICC_Snow_MAX. Precipitation rate, inches per hour. If the precipitation rate exceeds ATICC_Snow_MAX, the antecedent coldness index ATICC is set to the temperature of the precipitation (or the base temperature, which ever is lower) If the precipitation rate is less than ATICC_Snow_MAX, ATICC is computed as an antecedent index. Typical value is (.8 inches/day)



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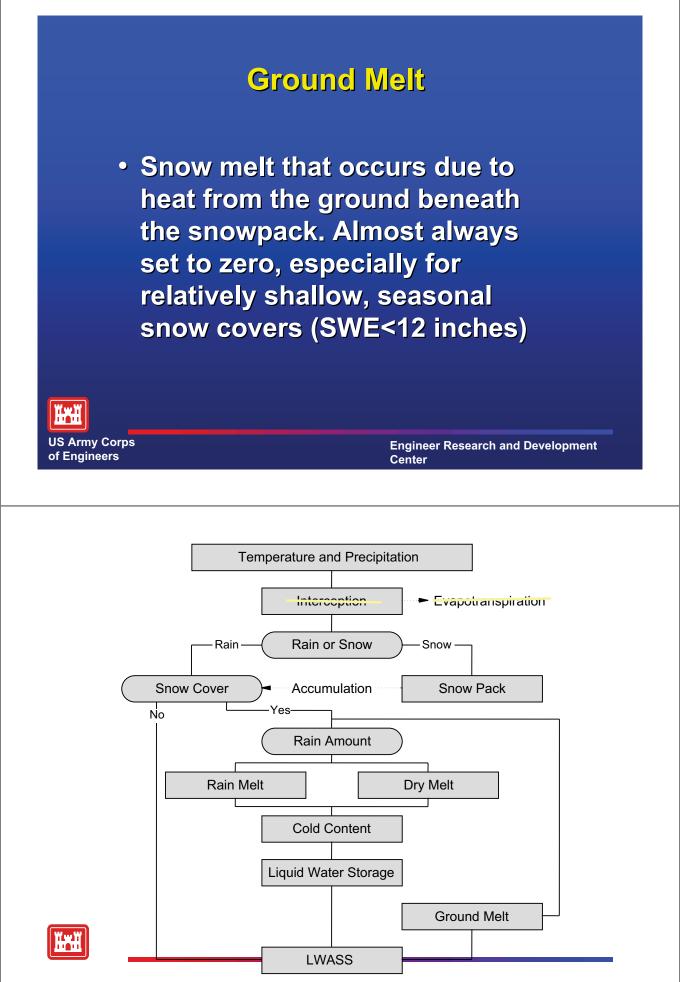
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Liquid Water Storage -Parameters

 Maximum liquid water capacity. The maximum liquid water capacity specifies the amount of melted water that must accumulate in the snowpack before liquid water becomes available at the soil surface for infiltration or runoff. Typically, the maximum liquid water held in the snowpack is on the order of 3%-5% of the SWE, although it can be higher. Liquid water can persist in the snow only if the snowpack temperature is at 32F (0C); at which point the cold content is zero.

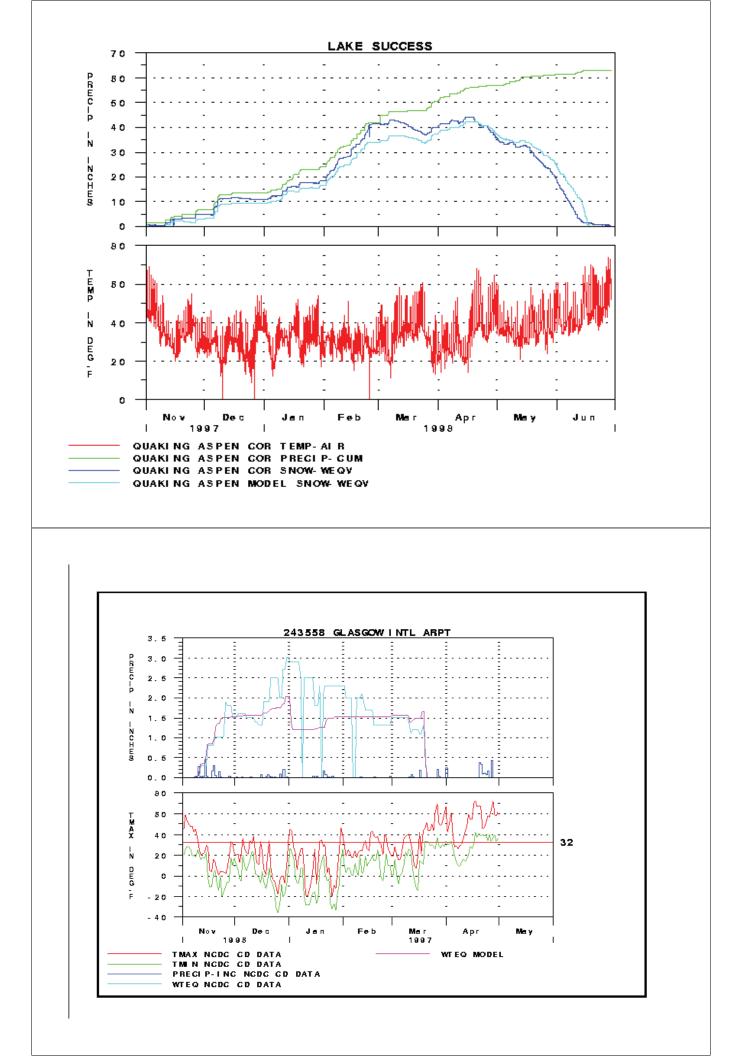


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Entering Snow Parameters in HMS

	PX Temperature (DEG F)	35.69		
	Wet Meltrate (IN/DEG F-DAY)			
	Rain Rate Limit (IN/DAY)			
	ATI-Meltrate Coefficient:			
		snowmelt_atimelt_9697		
	Meltrate Pattern:			
	Cold Limit (IN/DAY)			
	ATI-Coldrate Coefficient:	0.90		
		snowmelt_aticold_9697		
	Water Capacity (%) Groundmelt Method:			
			×	
	Groundmelt (IN/DAY)	0		
S Army Corps Engineers		Center		nd Development
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Summary -Terms

- Temperature Index Snow Model
- Complete Energy Budget Snow Model
- Snow Water Equivalent (SWE)
- Cold Content
- Snow Liquid Water Storage
- PX
- Base Temperature
- Rain Rate Limit
- Melt factor
- ATIMR
- Rain melt
- LWASS
- ATICC
- Maximum Liquid Water Capacity
- Ground melt



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Appendix B

GeoHMS Processing

With the wide availability of Geographic Information System (GIS) data layers for the study area, a good proportion of the model set up and parameterization was achieved using ArcHydro, and HEC-GeoHMS, which are publicly available ArcGIS extensions. The primary GIS data sources for the study consisted of a high resolution (3m) Digital Elevation Model (DEM) and stream layer produced by the Niagara Peninsula Conservation Authority; Southern Ontario Land Resource Information System (SOLRIS) land cover layer (Version 2) produced by the Ontario Ministry of Natural Resources (2006) and soils layer obtained from the Ontario Ministry of Agriculture and Food (Niagara Region 1989, Haldimand County 1984, Hamilton-Wentworth 1965). These layers were further processed in a GIS to produce the required HEC HMS model inputs.

Primary GIS layers processing

The digital soil layer from OMAF aggregates county-based soils maps onto a seamless and standardized product. Included in the product are three tables (soil component, soil name file, soil layer file) that can be relationally linked to the spatial data. The soils component attribute table maintains 17 variables for each soil component, which includes a multifaceted variety of soil and soil-related data. Based on their ability to drain precipitation inputs, soils can be categorized into 4 Hydrologic Soil Groups (HSG) ranging from A to D (Appendix C). HSG A soils have the highest infiltration rates, while HSG D soils have the lowest infiltration rates. Each map unit polygon in the component table can be comprised of up to 2 different (HSG) that contribute to the total area of the map unit. The area occupied by each HSG is represented as a percent of the map unit. However, there is no specific information on the location of the individual HSB within any polygon. For example, a BRADY soil series map unit can be comprised of 70 % HSG (A) and 30 % HSG (B). To avail ourselves with the most detailed information for the study, all HEC-HMS model parameter calculations dependent upon HSG were percent-weighted based on HSG.

SOLRIS is a primary data layer that provides a comprehensive landscape level inventory of land use for the study area. The SOLRIS land use classes for the study area are provided in Appendix II. The version in use in this study groups agricultural crop such as corn, grains, wheat, alfalfa, and soybeans into broad agricultural land use classes such as monoculture and annual crops. The level of detail was insufficient for the application of crop specific coefficients required as input for the evapotranspiration calculations in the Priestly-Taylor method used in HEC-HMS. Statistic Canada field crop data (percent by type) at the census consolidated subdivision level was used as a weight to calculate composite crop coefficient values for the SOLRIS agricultural land use classes.

Derivative GIS layer processing

Setting up the model required the user to select methods to simulate infiltration losses, transform excess precipitation into runoff, represent baseflow contribution to subbasin outflow, and simulate flow in open channels. Each method requires one or more parameters that describe the state of each catchment and allow the model to simulate the hydrologic processes. Table 1 shows the simulation methods used and the required GIS derived parameters. With the exception of the crop coefficient, channel bottom width, and channel side slope all vector products were converted to a 15m raster-based product.

Category	Method	Required GIS Parameter			
Loss	Deficit and Constant	Soil Water Holding Capacity			
		Soil maximum infiltration rate			
		Crop Coefficient (Priestly Taylor)			
Transform	SCS Unit Hydrograph	CN, Initial Abstraction, Lag Time			
Baseflow	Linear Reservoir	N/A			
Routing	Muskingum-Cunge	Channel Side Slope, Channel Bottom Width			

Table 1. HEC-HMS simulation methods used

Constant infiltration rates

The deficit and constant loss method assumes that the soil has a set maximum infiltration rate approximated by the saturated soil hydraulic conductivity. Using the information in Appendix C, average maximum infiltration rates were assigned to each polygon in the soil layer based on their HSG.

Soil water holding capacity

In the simulation, the soil is also assumed to have a fixed water holding capacity, typically affected by the active rooting depth of vegetation and HSG. The soil water holding capacity layer was built by intersecting the SOLRIS land cover and the OMAF soils layers and by assigning soil water holding capacity values from Appendix C to each unique combination of land cover class and soil HSG. OMAF polygons mapped as urban were not included in the procedure; imperviousness is addressed later.

Crop Coefficient

In the deficit constant method, water is removed from the soil to simulate evapotranspiration. In the model, evapotranspiration was calculated through the Priestly-Taylor. This method requires the use of crop coefficient K_c , which indicates the ratio of crop potential and grass reference evapotranspiration. Land use layers were created for each day of the year and daily crop coefficients from Appendix D were assigned to the land use classes.

CN grid

CN values are used in the calculation of CN lag time for the SCS Unit Hydrograph transform method. The factors influencing CN values are land cover type, HSG, and Antecedent Soil Moisture Condition (AMC). AMC is an estimate of soil water content prior to the beginning of the simulation period. AMC I reflect soils that are dry but with water content not below the wilting point. AMC II reflects soils having average soil water content, and AMC III reflects soils that have experienced rainfall in the five days previous to the beginning of the simulation period. CN values in the study area were assumed to reflect average soil water content. The CN layer was built by intersecting the SOLRIS land cover and OMAF soil layer and by assigning CN values from Appendix II to each unique combination of land use class and HSG. CN values were not assigned for built-up impervious, built-up pervious, transportation or open water SOLRIS land cover types.

Impervious

HEC-HMS considers an impervious surface as an area in a watershed for which all contributing precipitation runs off, with no infiltration, evaporation, or other volume losses. This surface was built by assigning percentages of 100, 100 and 50 to the transportation, built-up impervious and built-up pervious polygons respectively. All other polygons were assigned a value of 0.

Channel width and side slope

In the model the traditional Muskingum-Cunge routing method was used assuming trapezoidal channel geometry. The method requires the input of channel bottom width, channel side slope, and channel manning roughness coefficient. Channel width for each of the routing reaches was estimated by digitizing cross sections across the channel assuming that the extent (i.e. width) of the water surface on the digital air photos roughly approximates the width of the channel bed. Channel side slope was approximated by digitizing two points at the end of each digitized channel width cross sections using a 2m resolution DEM as a guide. Appropriate channel Manning roughness coefficients from Appendix C were assigned to channel routing reaches following visual channel stream bed condition assessment from 10 and 20 cm resolution digital air photos.

The GIS approach to building a HEC HMS model is generally done in two phases: the terrain processing phase and the model parameterization phase. These are described below.

Terrain Processing

The terrain processing phase requires a terrain model that is hydrologically correct. The terrain is created by integrating a fully connected dendritic stream network into a DEM. This process can be summarized as follows: 1) rasterization of the vector stream network to the same resolution as the DEM, 2) reclassifying the rasterized stream

network by assigning an arbitrary elevation (i.e. 50) value to the cells of the stream network. 3) Subtracting the reclassified grid from the DEM. This has the effect of decreasing the elevation of all DEM cells underlying the stream network by the aforementioned elevation value (50 m). 4) Filling the DEM sinks, thus ensuring that no water is trapped in DEM depressions and that all DEM cells drain to the outlet.

The next steps are the creation of two terrain derivatives from the filled DEM and a series of processing steps to delineate the watershed subbasins. These are performed using the ArcHydro Tools and are briefly outlined below:

- 1) Flow direction grid: Shows the orientation of the DEM cell's to its neighbour steepest down slope.
- 2) Flow accumulation grid: Indicates the number of upstream cells draining to each DEM cell.
- 3) Stream definition: The flow accumulation grid was then used to produce a synthetic stream network by applying a suitable area threshold value. The area threshold value indicates the minimum upstream area required to initiate a synthetic stream network. A 500 ha threshold value was selected so that average catchment size in the study was between 5 and 10 km² and lag time for most of these catchments greater than 2 hrs.
- 4) Stream segmentation: The synthetic stream network is divided at the synthetic stream network confluences. All cells belonging to each stream segment are assigned a unique value.
- 5) Catchment Grid Delineation: This step generates a grid representation of a subbasin for each stream segment. All cells belonging to a subbasin are assigned a unique number.
- 6) Catchment Polygon Processing: This step converts the grid representation of the subbasin to a vector representation
- 7) Drainage line processing: This step converts the grid representation of the segmented synthetic stream network into a vector representation.
- 8) Adjoint catchment processing: This step aggregates the upstream subbasins at every stream confluence. This step has no hydrological significance and is done to increase the performance of the point delineation process.

Hydrologic Model Creation

Once the terrain processing is completed, the data required to support model creation and model parameterization can be extracted for the study area using the HEC-GeoHMS tools. The main steps are HEC-HMS model set up, Watershed subbasin

boundary refinement and model parameterization. These steps are briefly outlined below.

HEC-HMS model set up

An HEC-HMS project is created by specifying the outlet point of the study area. During the project generation, the following datasets are created.

- 1) Filled DEM: Hydrologically corrected DEM.
- 2) Raw DEM: Original DEM.
- 3) Flow direction grid
- 4) Flow accumulation grid
- 5) Stream grid: Synthetic stream network in grid representation.
- 6) Stream link grid : Synthetic stream network segmented at confluences.
- 7) Catchment grid: Subbasin extents in grid representation.
- 8) Catchment polygon: Represent the extracted subbasin extent in vector format.
- 9) Rivers: Represents the synthetic stream network in vector format.
- 10) Project point: Represent the watershed outlet.

Watershed subbasin boundary refinement

Once a HMS project has been set up, the watershed subbasin boundaries can be revised. This was done mainly by combining and by subdividing subbasins. Subbasins larger than the 10 km^2 threshold were subdivided at hydrologic control points such as road crossings where changes in flow regime were most likely to occur. Subbasins smaller than the 5 km² threshold were merged to adjacent subbasins.

HEC-HMS model parameterization

Once a satisfactory watershed subbasin layout was defined, the next step was the parameterization of the model. Model parameterization is done in two phases. These are the extraction of the watershed physical parameters and the extraction of the model hydrologic parameters. These steps are outlined below.

Extraction of the watershed physical parameters

The HEC-GeoHMS toolbar can compute several topographic related characteristics of streams and subbasins. These include the following:

- 1) River length: Computes the length of each river features.
- 2) River slope: Extracts the upstream and downstream elevation of each stream segment and calculates the river slope.
- 3) Basin slope: Computes the average slope for each subbasin based on an input slope grid that was generated from a 2m DEM.
- 4) Longest flow path: Computes the longest flow path for each subbasin.
- 5) Basin centroid: Calculates the centroid as the center of the longest flow path within the subbasin.
- 6) Centroidal flow path: Calculates the flow path from the projected point of the subbasin centroid on the longest flow path to the subbasin outlet, along the longest flow path route.

Extraction of the watershed hydrologic parameters

In addition to extracting watershed physical parameters the HEC-GeoHMS tool can also extract a number of hydrologic inputs for the HEC-HMS model. The steps involved are briefly outlined below:

- 1) Selection of HMS processes: In this step, the loss, transform, and baseflow type methods for the subbasins and the routing method for the rivers as outlined in Table 1 were selected.
- 2) Subbasin curve number: Computation of the average subbasin curve number.
- 3) Muskingum-Cunge parameters: Allows the selection of the Muskingum-Cunge channel shape. A trapezoidal channel shape was selected for this study.
- 4) CN lag: Computation of the CN lag for each subbasin.

A number of required hydrologic parameters could not be extracted using the HEC-GeoHMS tools. These model parameters were manually calculated or extracted through the development and application of Python scripts. These parameters are listed below:

5) Basin imperviousness: Computation of the average subbasin imperviousness (%).

- 6) Basin maximum infiltration rate: Computation of the average subbasin maximum infiltration rate (mm/hr).
- 7) Basin water holding capacity: Computation of the average subbasin maximum water holding capacity (mm).
- 8) Basin Initial Abstraction: The initial abstraction defines the amount of precipitation that must fall before runoff is observed. This was calculated using the following formula:

I = 0.2*((25400/[CN])-254)

Where:

I = initial abstraction (mm) CN = Subbasin curve number

- 9) Initial Deficit: Initial Deficit represents the empty storage depth (mm) at the beginning of the simulation period. This quantity was set at half of the basin water holding capacity implying average soil moisture content in the soil at the beginning of the simulation period.
- 10) Routing channel bottom width: Computation of the average channel bottom width (m) for each routing reach.
- 11) Routing channel side slope: Computation of the average channel side slope (dimensionless).
- 12) Crop coefficient: Composite crop coefficient values were calculated for each day of the simulation period by calculating an area-weighted value for each catchment.

Export Model to HMS

Before exporting the developed hydrological modelling inputs to an HEC-HMS model input file, the HEC-GeoHMS tools were used to check the GIS layers for stream and basin connectivity, generate HMS schematic, legend, and a background map file.

Appendix C

Hydrologic Soil Group	Description	Average Infiltration Rate (mm/hr)
Â	Soils having high infiltration rates	9.51
	even when thoroughly wet.	
	These soils consist mainly of deep,	
	well-drained to excessively drained	
	sands or gravels.	
	These soils have a high rate of water	
	transmission and therefore a low	
	runoff potential.	
В	Soils having moderate infiltration	5.72
	rates when thoroughly wet,	
	consisting mainly of moderately	
	deep to deep, moderately well to	
	well-drained soils with moderately	
	fine to moderately coarse textures.	
	These soils have a moderate rate of	
	water transmission.	
С	Soils having slow infiltration rates	2.54
	when thoroughly wet,	
	consisting mainly of either soils with	
	a layer that impedes the downward	
	movement of water	
	or soils with moderately fine or fine	
	textures and slow infiltration rates.	
	These soils have a	
	slow rate of water transmission.	
D	Soils having very slow infiltration	0.64
	rates when thoroughly wet.	
	These are mainly comprised of	
	either clayey soil with high swelling	
	capacity or potential,	
	soils with a high permanent water	
	table, soils with a clay layer at or	
	near the surface, and/or	
	shallow soils over nearly impervious	
	materials. These soils have a very	
	slow rate of water	
	transmission and therefore a high	
	runoff potential.	

Average soil infiltration rates based on Hydrologic Soil Group (Haan et al., 1982)

	Hydrologic Soil Group						
Land use class	Α	B	C	D			
Annual Crop	67	78	85	89			
Bog	72	72	72	72			
Coniferous Forest	48	58	70	77			
Deciduous Forest	48	58	70	77			
Extraction	98	98	98	98			
Forest	48	58	70	77			
Hedge Rows	48	58	70	77			
Idle Land	50	61	74	80			
Marsh	85	85	85	85			
Mixed Agriculture	64	74	80	85			
Mixed Crop	67	78	85	89			
Mixed Forest	48	58	70	77			
Monoculture	40	62	76	81			
Open Shoreline	72	72	72	72			
Open Water	100	100	100	100			
Orchards	40	62	76	81			
Perennial Crop	59	74	83	86			
Plantations	38	60	74	80			
Rural Land Use	56	70	80	82			
Shoreline	72	72	72	72			
Swamp	72	72	72	72			

Runoff CN number for SOLRIS land use classes and HSG groups.

Soil water holding capacity (mm) for SOLRIS land use classes and HSG groups.

	Hydrologic Soil Group						
Land use class	Α	B	C	D			
Annual Crop	75	150	200	150			
Bog	250	300	400	350			
Built Up Impervious	0	0	0	0			
Built Up Pervious	50	75	113	75			
Coniferous Forest	250	300	400	350			
Deciduous Forest	250	300	400	350			
Extraction	0	0	0	0			
Forest	250	300	400	350			
Hedge Rows	250	300	400	350			
Idle Land	100	150	250	200			
Marsh	0	0	0	0			
Mixed Agriculture	75	150	200	150			
Mixed Crop	75	150	200	150			
Mixed Forest	250	300	400	350			
Monoculture	75	150	200	150			
Open Shoreline	0	0	0	0			
Open Water	0	0	0	0			
Orchards	250	300	400	350			
Perennial Crop	100	150	250	200			
Plantations	100	150	250	100			
Rural Land Use	100	150	250	200			
Shoreline	0	0	0	0			
Swamp	250	300	400	350			
Transportation	0	0	0	0			

Channel Stream bed condition	Minimum	Average	Maximum
a. Clean, straight, full, no rifts or deep pools	0.025	0.030	0.033
b. Same as above, but more stones and weeds	0.030	0.035	0.040
c. Clean, winding, some pools and shoals	0.033	0.040	0.045
d. Same as above, but some stones and	0.035	0.045	0.050
weeds.			
e. Same as above, lower stages, more	0.040	0.048	0.050
ineffective slopes and sections			
f. Same as "d" but more stones	0.045	0.050	0.060
g. Sluggish reaches, weedy, deep pools	0.050	0.070	0.080
h. Very weedy reaches, deep pools, or	0.070	0.100	0.150
floodways with heavy stands of timber and			
brush.			

Channel manning n coefficients under various channel stream bed conditions.

Appendix D



Weather INnovations Incorporated 7159 Queen's Line RR 5, Chatham, ON N7M 5J5 Ph:(519)352-5334 Fax:(519)352-7630

Land-Use Evapotranspiration Coefficient Study

Weather INnovations Incorporated (WIN) was contracted by the Niagara Peninsula Conservation Authority (NPCA) to determine coefficients relating evapotranspiration rates to a reference potential evapotranspiration for the highest density land uses in the Niagara Peninsula.

The land uses of greatest interest, the majority of which were identified to be agricultural, were outlined by the NPCA. The crops of greatest density were used to further define the agricultural land use.

[able]	I: Identified Land Use
Agric	cultural Land
- S	oybeans
- H	lay
- G	rain Corn
- W	Vinter Wheat
- V	ineyards
	each
- P	ear/Apple/Cherry/Plum
- G	razing Land
Non-	Agricultural Land Use
- Ic	lle Land (more than 10
years	out of agricultural
-	iction)
- D	eciduous Forest
- Sv	wamp
	uilt-Up Pervious land
	grass)
	ler Coverage Land
Uses	
	pen/Shallow Water
	Iarsh
- C	oniferous Forest
- T	allgrass
- F	en
- B	og

The results of many researchers have been used to develop the evapotranspiration coefficients suggested by the Food and Agricultural Organization of the United Nations (FAO), and are accepted internationally. Due to the complexity of measuring evapotranspiration and the variability in the values year to year, there are very few (if any) results indicating a daily estimate for evapotranspiration values. WIN has determined,

where possible, the growth stages for the various agricultural crops listed, based on suggestions by the FAO and other researchers. The step-wise values have been determined where a linear progression occurs from one coefficient to another.

Very few crop coefficients for Ontario are available, and a majority of the land uses identified by the NPCA could not be found in Ontario documentation. As a result, values from the FAO were substituted, except in the situation of grapes and deciduous fruit trees. These values, although based on FAO findings, were more refined by the OMAFRA Best Management - Irrigation Management guide, and the British Columbia Ministry of Agriculture, Food and Fisheries. These two sources were used in combination, to determine the evapotranspiration coefficient for grapes and deciduous fruit trees.

The FAO suggests modifying the mid-growth stage values depending on the minimum daily relative humidity value, and the wind speed (at 2m) based on the region of interest, and the crop height. They suggest this adjustment to both Kc_{mid} and Kc_{end} . The equation indicated is identified for Kc_{mid} , however the equation for Kc_{end} is the same, just with the table value for Kc_{end} substituted in place of $Kc_{mid(tab)}$.

$$Kc_{mid} = Kc_{mid(tab)} + [0.04(u_2-2)-0.004[RH_{min}-45]](h/3)^{0.3},$$

where $Kc_{mid(tab)}$ is the published FAO crop coefficient, u_2 is the wind speed at 2m (in m/s), RH_{min} is the minimum daily RH value (%), and h is the crop height (m).

In order to conduct this calculation, wind speed and RH data from a station in the Niagara Peninsula was used for 2006. It is important to note that these values will change yearly depending on the season. A yearly average of the 'adjusted' Kc value will be used for the purpose of this project.

The following changes to Kc_{mid} and Kc_{end} were made to the following crops.

Crop	Crop Height	Kc _{mid(tab)}	Adjusted Kc _{mid}	Kc _{end(tab)}	Adjusted Kc _{end}
~ 1	(m)				0.51
Soybeans	0.7	1.15	1.15	0.50	0.51
Winter	1.0	1.15	1.17	0.32	0.34
Wheat					
Maize	2.0	1.20	1.25	0.48	0.65
Rye Grass	0.3	1.05	1.05	1.00	1.00
Hay					
Clover Hay	0.6	0.90	0.90	0.85	0.85
Pasture	0.4	0.95	0.95	0.85	0.85
Wetlands	1.5	1.20	1.27	0.30	0.37

Table 2: Adjusted Kcmid and Kcend values

FAO provides estimated duration for each crop coefficient. These were taken into consideration when determining the change from one coefficient to another. However, in some instances, alterations were made to better suit the Ontario growing season. Many of the planting dates set for crops are close to the earliest planting date for the crop. These values should be adjusted, if required, to reflect a variety of situations based on planting dates.

The evapotranspiration coefficients for boreal deciduous and coniferous forests were determined from a study by Komastsu (2005). The results, from various research projects around the world, were examined for a comparison of coefficients for the Priestly-Taylor model. In order to determine the values for this study, the average of the findings for both the boreal deciduous and the boreal coniferous forests were used. The values for boreal coniferous forest range from 0.38 to 0.69, with an average of 0.55. Values estimating the winter evapotranspiration coefficients for conifer trees could not be determined in the time allotted for this project. As such, evaporative losses of 0.2 were substituted, the value which is currently used for deciduous trees.

The determination for a coefficient for idle land becomes more complex. As the land has been out of agricultural production for 10 or more years, it is assumed that grasses, weeds and native vegetation are now established. The FAO indicates the use of the following equations to determine the mid-season evaporation rates. Due to the lack of information regarding leaf area index (LAI) values, it was indicated by the FAO that full coverage vegetation would have an LAI value of 3. For the scenario of tallgrass, a similar methodology was used. However, a study by Verma and Berry (1997) indicates that the LAI from a tallgrass prairie was 0.2 from mid-March to early May, and ranged from 0.3 to 1.8 in the later part of May (average of 1.05 will be used for this study). At the peak of the season, the range was 2.5 to 2.8 (average of 2.65 used for this study).

 $K_{cb, h} = 1.0 + 0.1h$, for $h \le 2m$

 $K_{cb full} = K_{cb}, h + [0.04(u_2-2)-(0.004(RH_{min}-45)](h/3)^{0.3}$

 $K_{cbmid} = K_{cmin} + (K_{cfull} - K_{cmin})(1 - exp[-0.7LAI]),$

where Kc_{min} is the minimum Kc value for bare soil (ranging from 0.15-0.2). OMAFRA indicates a bare soil coefficient of 0.2. This value was used.

Due to a lack of information regarding evapotranspiration values for swamps, bogs, marshes and fens individually, the value for wetlands was used in all four situations.

Many models used to calculate potential evapotranspiration (ETo) utilize a well-watered turf surface as the reference point. The coefficient for the built-up pervious area (e.g. sod/grass) will be 1.

Open water, especially water at a depth greater than 5m, creates a complex situation. The FAO indicates that deep bodies of water experience fluctuating temperatures, but this may not be true of frozen surfaces. During periods of peak evapotranspirative losses, radiation is being absorbed into the water. Therefore, the evaporative losses are less than ETo. During cooler temperature periods, the energy exchange is reversed. This causes the evaporation rates to be higher than those for grass, during the same period.

The following tables are daily evapotranspiration coefficients for the identified land use classes.

corn.						
	Soybeans	Kc	Winter Wheat	Kc	Grain Corn	Kc
Jan. 1	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 2	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 3	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 4	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 5	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 6	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 7	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 8	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 9	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 10	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 11	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 12	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 13	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 14	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 15	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 16	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 17	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 18	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 19	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 20	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 21	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 22	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 23	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 24	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 25	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 26	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 27	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 28	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 29	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 30	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 31	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 1	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 2	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 3	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 4	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 5	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 6	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 7	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 8	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 9	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 10	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 11	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 12	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 13	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 14	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 15	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 16	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 17	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 18	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 19	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 20	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 21	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 22	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
	Bare Soil	0.20	Kini			0.20
Feb. 23	Bare Noti	(12)(1	K 1n1	0.40	Bare Soil	

Table 3: Annual evapotranspiration coefficients for soybeans, winter wheat and grain corn.

Feb. 25	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 26	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 27	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 28	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 1	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 2	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 3	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 4	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 5	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 6	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 7	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 8	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
		0.20	Kini	0.40	+	
Mar. 9	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 10	Bare Soil				Bare Soil	0.20
Mar. 11	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 12	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 13	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 14	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 15	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 16	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 17	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 18	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 19	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 20	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 21	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 22	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 23	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 24	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 25	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 26	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 27	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 28	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 29	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 30	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Mar. 31	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Apr. 1	Bare Soil	0.20	Kdev	0.42	Bare Soil	0.20
Apr. 2	Bare Soil	0.20	Kdev	0.43	Bare Soil	0.20
Apr. 3	Bare Soil	0.20	Kdev	0.45	Bare Soil	0.20
Apr. 4	Bare Soil	0.20	Kdev	0.47	Bare Soil	0.20
Apr. 5	Bare Soil	0.20	Kdev	0.48	Bare Soil	0.20
Apr. 6	Bare Soil	0.20	Kdev	0.50	Bare Soil	0.20
Apr. 7	Bare Soil	0.20	Kdev	0.52	Bare Soil	0.20
Apr. 8	Bare Soil	0.20	Kdev	0.53	Bare Soil	0.20
Apr. 9	Bare Soil	0.20	Kdev	0.55	Bare Soil	0.20
Apr. 10	Bare Soil	0.20	Kdev	0.57	Bare Soil	0.20
Apr. 10	Bare Soil	0.20	Kdev	0.57	Bare Soil	0.20
Apr. 12	Bare Soil	0.20	Kdev	0.60	Bare Soil	0.20
Apr. 12 Apr. 13	Bare Soil	0.20	Kdev	0.62	Bare Soil	0.20
Apr. 13 Apr. 14	Bare Soil	0.20	Kdev	0.62	Bare Soil	0.20
Apr. 14 Apr. 15	Bare Soil	0.20	Kdev	0.65	Kini	0.20
*	Bare Soil	0.20		0.63	Kini	0.30
Apr. 16			Kdev			
Apr. 17	Bare Soil	0.20	Kdev	0.68	Kini	0.30
Apr. 18	Bare Soil	0.20	Kdev	0.70	Kini	0.30
Apr. 19	Bare Soil	0.20	Kdev	0.72	Kini	0.30
Apr. 20	Bare Soil	0.20	Kdev	0.73	Kini	0.30
Apr. 21	Bare Soil	0.20	Kdev	0.75	Kini	0.30
Apr. 22	Bare Soil	0.20	Kdev	0.77	Kini	0.30
Apr. 23	Bare Soil	0.20	Kdev	0.78	Kini	0.30
Apr. 24	Bare Soil	0.20	Kdev	0.80	Kini	0.30

Apr. 25	Bare Soil	0.20	Kdev	0.82	Kini	0.30
Apr. 26	Bare Soil	0.20	Kdev	0.84	Kini	0.30
Apr. 27	Bare Soil	0.20	Kdev	0.85	Kini	0.30
Apr. 28	Bare Soil	0.20	Kdev	0.87	Kini	0.30
Apr. 29	Bare Soil	0.20	Kdev	0.89	Kini	0.30
Apr. 30	Bare Soil	0.20	Kdev	0.90	Kini	0.30
May 1	Bare Soil	0.20	Kdev	0.92	Kini	0.30
May 2	Bare Soil	0.20	Kdev	0.92	Kini	0.30
May 3	Bare Soil	0.20	Kdev	0.95	Kini	0.30
May 4	Bare Soil	0.20	Kdev	0.97	Kini	0.30
May 5	Bare Soil	0.20	Kdev	0.99	Kini	0.30
	Bare Soil	0.20	Kdev	1.00	Kini	0.30
May 6						
May 7	Bare Soil	0.20	Kdev	1.02	Kini	0.30
May 8	Bare Soil	0.20	Kdev	1.04	Kini	0.30
May 9	Bare Soil	0.20	Kdev	1.05	Kini	0.30
May 10	Bare Soil	0.20	Kdev	1.07	Kini	0.30
May 11	Bare Soil	0.20	Kdev	1.09	Kini	0.30
May 12	Bare Soil	0.20	Kdev	1.10	Kini	0.30
May 13	Bare Soil	0.20	Kdev	1.12	Kini	0.30
May 14	Bare Soil	0.20	Kdev	1.14	Kini	0.30
May 15	Kini	0.40	Kdev	1.15	Kini	0.30
May 16	Kini	0.40	Kmid	1.17	Kdev	0.32
May 17	Kini	0.40	Kmid	1.17	Kdev	0.35
May 18	Kini	0.40	Kmid	1.17	Kdev	0.37
May 19	Kini	0.40	Kmid	1.17	Kdev	0.39
May 20	Kini	0.40	Kmid	1.17	Kdev	0.42
May 21	Kini	0.40	Kmid	1.17	Kdev	0.44
May 22	Kini	0.40	Kmid	1.17	Kdev	0.46
May 23	Kini	0.40	Kmid	1.17	Kdev	0.49
May 24	Kini	0.40	Kmid	1.17	Kdev	0.51
May 25	Kini	0.40	Kmid	1.17	Kdev	0.53
May 26	Kini	0.40	Kmid	1.17	Kdev	0.55
May 27	Kini	0.40	Kmid	1.17	Kdev	0.58
May 28	Kini	0.40	Kmid	1.17	Kdev	0.60
May 29	Kini	0.40	Kmid	1.17	Kdev	0.62
May 30	Kini	0.40	Kmid	1.17	Kdev	0.65
May 31	Kini	0.40	Kmid	1.17	Kdev	0.67
June 1	Kini	0.40	Kmid	1.17	Kdev	0.69
June 2	Kini	0.40	Kmid	1.17	Kdev	0.09
June 3	Kini	0.40	Kmid	1.17	Kdev	0.72
June 4	Kdev	0.40	Kmid	1.17	Kdev	0.74
June 5	Kdev	0.42	Kmid	1.17	Kdev	0.70
June 6	Kdev	0.43	Kmid	1.17	Kdev	0.79
June 7	Kdev	0.49	Kmid	1.17	Kdev	0.83
June 8	Kdev	0.51	Kmid	1.17	Kdev	0.86
June 9	Kdev	0.54	Kmid	1.17	Kdev	0.88
June 10	Kdev	0.56	Kmid	1.17	Kdev	0.90
June 11	Kdev	0.58	Kmid	1.17	Kdev	0.93
June 12	Kdev	0.60	Kmid	1.17	Kdev	0.95
June 13	Kdev	0.63	Kmid	1.17	Kdev	0.97
June 14	Kdev	0.65	Kmid	1.17	Kdev	1.00
June 15	Kdev	0.67	Kmid	1.17	Kdev	1.02
June 16	Kdev	0.70	Kmid	1.17	Kdev	1.04
June 17	Kdev	0.72	Kmid	1.17	Kdev	1.06
June 18	Kdev	0.74	Kmid	1.17	Kdev	1.09
June 19	Kdev	0.76	Kmid	1.17	Kdev	1.11
June 20	Kdev	0.79	Kmid	1.17	Kdev	1.13
June 21	Kdev	0.81	Kmid	1.17	Kdev	1.16
	Kdev	0.83	Kmid	1.17	Kdev	1.18

June 23	Kdev	0.85	Kmid	1.17	Kdev	1.20
June 24	Kdev	0.88	Kmid	1.17	Kdev	1.23
June 25	Kdev	0.90	Kmid	1.17	Kmid	1.25
June 26	Kdev	0.92	Kmid	1.17	Kmid	1.25
June 27	Kdev	0.95	Kmid	1.17	Kmid	1.25
June 28	Kdev	0.93	Kmid	1.17	Kmid	1.25
June 29	Kdev	0.97	Kmid	1.17	Kmid	1.25
June 30	Kdev	1.01	Kmid	1.17	Kmid	1.25
July 1	Kdev	1.04	Kmid	1.17	Kmid	1.25
July 2	Kdev	1.06	Kmid	1.17	Kmid	1.25
July 3	Kdev	1.08	Kmid	1.17	Kmid	1.25
July 4	Kdev	1.10	Kmid	1.17	Kmid	1.25
July 5	Kdev	1.13	Kmid	1.17	Kmid	1.25
July 6	Kmid	1.15	Kmid	1.17	Kmid	1.25
July 7	Kmid	1.15	Kmid	1.17	Kmid	1.25
July 8	Kmid	1.15	Kmid	1.17	Kmid	1.25
July 9	Kmid	1.15	Kmid	1.17	Kmid	1.25
,					Kmid	
July 10	Kmid	1.15	Kmid	1.17		1.25
July 11	Kmid	1.15	Klate	1.14	Kmid	1.25
July 12	Kmid	1.15	Klate	1.12	Kmid	1.25
July 13	Kmid	1.15	Klate	1.09	Kmid	1.25
July 14	Kmid	1.15	Klate	1.06	Kmid	1.25
July 15	Kmid	1.15	Klate	1.03	Kmid	1.25
July 16	Kmid	1.15	Klate	1.01	Kmid	1.25
July 17	Kmid	1.15	Klate	0.98	Kmid	1.25
July 18	Kmid	1.15	Klate	0.95	Kmid	1.25
July 19	Kmid	1.15	Klate	0.92	Kmid	1.25
July 20	Kmid	1.15	Klate	0.92	Kmid	1.25
,						
July 21	Kmid	1.15	Klate	0.87	Kmid	1.25
July 22	Kmid	1.15	Klate	0.84	Kmid	1.25
July 23	Kmid	1.15	Klate	0.82	Kmid	1.25
July 24	Kmid	1.15	Klate	0.79	Kmid	1.25
July 25	Kmid	1.15	Klate	0.76	Kmid	1.25
July 26	Kmid	1.15	Klate	0.73	Kmid	1.25
July 27	Kmid	1.15	Klate	0.71	Kmid	1.25
July 28	Kmid	1.15	Klate	0.68	Kmid	1.25
July 29	Kmid	1.15	Klate	0.65	Kmid	1.25
July 30	Kmid	1.15	Klate	0.62	Kmid	1.25
July 31	Kmid	1.15	Klate	0.60	Kmid	1.25
	Kmid	1.15	Klate	0.57	Kmid	1.25
Aug. 1						
Aug. 2	Kmid	1.15	Klate	0.54	Kmid	1.25
Aug. 3	Kmid	1.15	Klate	0.52	Kmid	1.25
Aug. 4	Kmid	1.15	Klate	0.49	Kmid	1.25
Aug. 5	Kmid	1.15	Klate	0.46	Kmid	1.25
Aug. 6	Kmid	1.15	Klate	0.43	Kmid	1.25
Aug. 7	Kmid	1.15	Klate	0.41	Kmid	1.25
Aug. 8	Kmid	1.15	Klate	0.38	Kmid	1.25
Aug. 9	Kmid	1.15	Klate	0.35	Kmid	1.25
Aug. 10	Kmid	1.15	Klate	0.33	Kmid	1.25
Aug. 10	Kmid	1.15	Bare Soil	0.20	Kmid	1.25
Aug. 11 Aug. 12	Kmid	1.15	Bare Soil	0.20	Kmid	1.25
U						
Aug. 13	Kmid	1.15	Bare Soil	0.20	Kmid	1.25
Aug. 14	Kmid	1.15	Bare Soil	0.20	Klate	1.23
Aug. 15	Kmid	1.15	Bare Soil	0.20	Klate	1.22
Aug. 16	Kmid	1.15	Bare Soil	0.20	Klate	1.20
Aug. 17	Kmid	1.15	Bare Soil	0.20	Klate	1.19
Aug. 18	Kmid	1.15	Bare Soil	0.20	Klate	1.17
-	Kmid	1.15	Bare Soil	0.20	Klate	1.16
Aug. 19	NIIIQ	1.1./	Date Son	0.20	INIAU	

Aug. 21	Kmid	1.15	Bare Soil	0.20	Klate	1.13
Aug. 22	Kmid	1.15	Bare Soil	0.20	Klate	1.11
Aug. 23	Kmid	1.15	Bare Soil	0.20	Klate	1.10
Aug. 24	Kmid	1.15	Bare Soil	0.20	Klate	1.08
Aug. 25	Kmid	1.15	Bare Soil	0.20	Klate	1.07
Aug. 26	Kmid	1.15	Bare Soil	0.20	Klate	1.05
Aug. 27	Kmid	1.15	Bare Soil	0.20	Klate	1.03
Aug. 27	Kmid	1.15	Bare Soil	0.20	Klate	1.04
Aug. 29	Kmid	1.15	Bare Soil	0.20	Klate	1.01
Aug. 30	Kmid	1.15	Bare Soil	0.20	Klate	0.99
Aug. 31	Kmid	1.15	Bare Soil	0.20	Klate	0.98
Sept. 1	Kmid	1.15	Bare Soil	0.20	Klate	0.96
Sept. 2	Kmid	1.15	Bare Soil	0.20	Klate	0.95
Sept. 3	Kmid	1.15	Bare Soil	0.20	Klate	0.93
Sept. 4	Klate	1.13	Bare Soil	0.20	Klate	0.92
Sept. 5	Klate	1.10	Bare Soil	0.20	Klate	0.90
Sept. 6	Klate	1.08	Bare Soil	0.20	Klate	0.89
Sept. 7	Klate	1.05	Bare Soil	0.20	Klate	0.87
Sept. 8	Klate	1.03	Bare Soil	0.20	Klate	0.85
Sept. 9	Klate	1.01	Bare Soil	0.20	Klate	0.84
Sept. 10	Klate	0.98	Bare Soil	0.20	Klate	0.82
Sept. 11	Klate	0.96	Bare Soil	0.20	Klate	0.81
Sept. 12	Klate	0.93	Bare Soil	0.20	Klate	0.79
Sept. 12 Sept. 13	Klate	0.93	Bare Soil	0.20	Klate	0.79
-	Klate	0.91	Bare Soil	0.20		0.78
Sept. 14					Klate	
Sept. 15	Klate	0.86	Kini	0.40	Klate	0.75
Sept. 16	Klate	0.84	Kini	0.40	Klate	0.73
Sept. 17	Klate	0.81	Kini	0.40	Klate	0.72
Sept. 18	Klate	0.79	Kini	0.40	Klate	0.70
Sept. 19	Klate	0.76	Kini	0.40	Klate	0.69
Sept. 20	Klate	0.74	Kini	0.40	Klate	0.67
Sept. 21	Klate	0.72	Kini	0.40	Klate	0.66
Sept. 22	Klate	0.69	Kini	0.40	Klate	0.64
Sept. 23	Klate	0.67	Kini	0.40	Klate	0.63
Sept. 24	Klate	0.64	Kini	0.40	Klate	0.61
Sept. 25	Klate	0.62	Kini	0.40	Klate	0.60
Sept. 26	Klate	0.60	Kini	0.40	Klate	0.58
Sept. 27	Klate	0.57	Kini	0.40	Klate	0.57
Sept. 28	Klate	0.55	Kini	0.40	Klate	0.55
Sept. 29	Klate	0.52	Kini	0.40	Klate	0.54
Sept. 30	Klate	0.50	Kini	0.40	Klate	0.54
Oct. 1	Bare Soil	0.20	Kini	0.40	Klate	0.52
Oct. 1 Oct. 2	Bare Soil Bare Soil	0.20	Kini	0.40	Klate	0.31
	Bare Soil Bare Soil					
Oct. 3		0.20	Kini	0.40	Bare Soil	0.20
Oct. 4	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 5	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 6	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 7	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 8	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 9	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 10	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 11	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 12	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 13	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 14	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 15	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 16	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 17	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 18	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
001.10	Date Soli	0.20	KIIII	0.40	Dare Soll	0.20

Oct. 19	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 20	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 21	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 22	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 23	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 24	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 25	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 26	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 27	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 28	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 29	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 30	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 31	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 1	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 2	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 3	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 4	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 5	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 6	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 7	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 8	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 9	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 10	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 11	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 12	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 12	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 13	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 14	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 15	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
		0.20	Kini	0.40	+	0.20
Nov. 17	Bare Soil			0.40	Bare Soil	
Nov. 18 Nov. 19	Bare Soil Bare Soil	0.20	Kini Kini	0.40	Bare Soil Bare Soil	0.20
Nov. 19	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 20		0.20	Kini	0.40	+	0.20
Nov. 21	Bare Soil Bare Soil	0.20	Kini	0.40	Bare Soil Bare Soil	0.20
		0.20	Kini	0.40		0.20
Nov. 23	Bare Soil				Bare Soil	
Nov. 24	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 25	Bare Soil		Kini	0.40	Bare Soil	0.20
Nov. 26	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 27	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 28	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 29	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Nov. 30	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 1	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 2	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 3	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 4	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 5	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 6	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 7	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 8	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 9	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 10	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 11	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 12	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 13	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 14	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 15	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 16	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20

Dec. 17	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 18	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 19	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 20	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 21	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 22	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 23	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 24	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 25	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 26	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 27	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 28	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 29	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 30	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 31	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20

Table 4: Annual evapotranspiration coefficients for hay, pasture, and wetlands.

		<u> </u>		, F		
	Hay ⁺	Kc	Pasture	Kc	Wetlands	Kc
Jan. 1	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 2	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 3	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 4	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 5	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 6	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 7	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 8	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 9	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 10	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 11	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 12	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 13	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 14	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 15	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 16	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 17	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 18	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 19	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 20	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 21	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 22	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 23	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 24	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 25	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 26	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 27	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 28	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 29	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 30	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 31	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 1	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 2	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 3	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 4	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 5	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 6	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 7	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 8	Bare Soil	0.20	Dormant	0.40	Dormant	0.37

Feb. 9	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 10	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 11	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 12	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 13	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 14	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 15	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 16	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 17	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 18	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 19	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 20	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 21	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 22	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 23	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 24	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 25	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 26	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 27	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 28	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 1	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 2	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 3	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 4	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 5	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 6	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 7	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 8	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 9	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 10	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 11	Bare Soil	0.20	+	0.40	_	0.37
Mar. 12	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 12 Mar. 13	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
			Dormant		Dormant	
Mar. 14	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 15	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 16	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 17	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 18	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 19	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 20	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 21	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 22	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 23	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 24	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 25	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 26	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 27	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 28	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 29	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 30	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Mar. 31	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 1	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 2	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 3	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 4	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 5	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
		0.00	Demast	0.40	Dormant	0.37
Apr. 6	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 6 Apr. 7	Bare Soil Bare Soil	0.20	Dormant	0.40	Dormant	0.37

Apr. 9	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 10	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 11	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 12	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 13	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 14	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 15	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 16	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 17	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 18	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 19	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 20	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 21	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 22	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 23	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 24	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 25	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 26	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 27	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 28	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 29	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 30	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
May 1	Kini	0.95	Kini	0.40	Dormant	0.37
May 2	Kini	0.95	Kini	0.40	Dormant	0.37
May 3	Kini	0.95	Kini	0.40	Dormant	0.37
May 4	Kini	0.95	Kini	0.40	Dormant	0.37
May 5	Kini	0.95	Kini	0.40	Dormant	0.37
May 6	Kini	0.95	Kini	0.40	Dormant	0.37
May 7	Kini	0.95	Kini	0.40	Dormant	0.37
May 8	Kini	0.95	Kini	0.40	Dormant	0.37
May 9	Kini	0.95	Kini	0.40	Dormant	0.37
May 10	Kini	0.95	Kini	0.40	Dormant	0.37
May 11	Kini	0.95	Kdev	0.43	Dormant	0.37
May 12	Kini	0.95	Kdev	0.45	Dormant	0.37
May 13	Kini	0.95	Kdev	0.48	Dormant	0.37
May 14	Kini	0.95	Kdev	0.50	Dormant	0.37
May 15	Kini	0.95	Kdev	0.53	Kini	0.30
May 16	Kini	0.95	Kdev	0.56	Kini	0.30
May 17	Kini	0.95	Kdev	0.58	Kini	0.30
May 18	Kini	0.95	Kdev	0.61	Kini	0.30
May 19	Kini	0.95	Kdev	0.64	Kini	0.30
May 20	Kini	0.95	Kdev	0.66	Kini	0.30
May 21	Kdev	0.95	Kdev	0.69	Kini	0.30
May 22	Kdev	0.96	Kdev	0.71	Kini	0.30
May 23	Kdev	0.96	Kdev	0.74	Kini	0.30
May 24	Kdev	0.97	Kdev	0.77	Kini	0.30
May 25	Kdev	0.97	Kdev	0.79	Kdev	0.33
May 26	Kdev	0.98	Kdev	0.82	Kdev	0.36
May 27	Kdev	0.98	Kdev	0.85	Kdev	0.39
May 28	Kdev	0.98	Kdev	0.87	Kdev	0.43
		0.99	Kdev	0.90	Kdev	0.46
May 29	Kdev	0.77				
May 29	Kdev Kdev	0.99	Kdev	0.92	Kdev	0.49
May 29 May 30	Kdev	0.99				
May 29 May 30 May 31	Kdev Kdev	0.99	Kmid	0.95	Kdev	0.52
May 29 May 30 May 31 June 1	Kdev Kdev Kdev	0.99 1.00 1.00	Kmid Kmid	0.95 0.95	Kdev Kdev	0.52 0.55
May 29 May 30 May 31 June 1 June 2	Kdev Kdev Kdev Kdev	0.99 1.00 1.00 1.00	Kmid Kmid Kmid	0.95 0.95 0.95	Kdev Kdev Kdev	0.52 0.55 0.58
May 29 May 30 May 31 June 1 June 2 June 3	Kdev Kdev Kdev Kdev Kdev	0.99 1.00 1.00 1.00 1.01	Kmid Kmid Kmid Kmid	0.95 0.95 0.95 0.95	Kdev Kdev Kdev Kdev	0.52 0.55 0.58 0.61
May 29 May 30 May 31 June 1 June 2	Kdev Kdev Kdev Kdev	0.99 1.00 1.00 1.00	Kmid Kmid Kmid	0.95 0.95 0.95	Kdev Kdev Kdev	0.52 0.55 0.58

June 7	Kdev	1.03	Kmid	0.95	Kdev	0.74
June 8	Kdev	1.03	Kmid	0.95	Kdev	0.77
June 9	Kdev	1.03	Kmid	0.95	Kdev	0.80
June 10	Kdev	1.04	Kmid	0.95	Kdev	0.83
June 11	Kdev	1.04	Kmid	0.95	Kdev	0.86
June 12	Kdev	1.04	Kmid	0.95	Kdev	0.89
June 13	Kdev	1.05	Kmid	0.95	Kdev	0.93
June 14	Kdev	1.06	Kmid	0.95	Kdev	0.96
June 15	Kmid	1.06	Kmid	0.95	Kdev	0.99
June 16	Kmid	1.06	Kmid	0.95	Kdev	1.02
June 17	Kmid	1.06	Kmid	0.95	Kdev	1.05
June 18	Kmid	1.06	Kmid	0.95	Kdev	1.08
June 19	Kmid	1.06	Kmid	0.95	Kdev	1.11
June 20	Kmid	1.06	Kmid	0.95	Kdev	1.14
June 21	Kmid	1.06	Kmid	0.95	Kdev	1.18
June 22	Kmid	1.06	Kmid	0.95	Kdev	1.21
June 23	Kmid	1.06	Kmid	0.95	Kdev	1.24
June 24	Kmid	1.06	Kmid	0.95	Kdev	1.27
June 25	Kmid	1.06	Kmid	0.95	Kmid	1.27
June 26	Kmid	1.06	Kmid	0.95	Kmid	1.27
June 27	Kmid	1.06	Kmid	0.95	Kmid	1.27
June 28	Kmid	1.06	Kmid	0.95	Kmid	1.27
June 29	Kmid	1.06	Kmid	0.95	Kmid	1.27
June 30	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 1	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 2	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 3	Kmid	1.06	Kmid	0.95	Kmid	1.27
		1.06	Kmid	0.95	Kmid	1.27
July 4	Kmid					
July 5	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 6	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 7	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 8	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 9	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 10	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 11	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 12	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 13	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 14	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 15	Klate	1.05	Kmid	0.95	Kmid	1.27
July 15 July 16			Kmid	0.95	Kmid	
	Klate	1.04				1.27
July 17	Klate	1.02	Kmid	0.95	Kmid	1.27
July 18	Klate	1.01	Kmid	0.95	Kmid	1.27
July 19	Klate	1.00	Kmid	0.95	Kmid	1.27
July 20	Kini	0.95	Kmid	0.95	Kmid	1.27
July 21	Kini	0.95	Kmid	0.95	Kmid	1.27
July 22	Kini	0.95	Kmid	0.95	Kmid	1.27
July 23	Kini	0.95	Kmid	0.95	Kmid	1.27
July 24	Kini	0.95	Kmid	0.95	Kmid	1.27
July 25	Kini	0.95	Kmid	0.95	Kmid	1.27
July 26	Kdev	0.95	Kmid	0.95	Kmid	1.27
July 20 July 27	Kdev	0.96	Kmid	0.95	Kmid	1.27
				0.95		
July 28	Kdev	0.97	Kmid		Kmid	1.27
July 29	Kdev	0.98	Kmid	0.95	Kmid	1.27
July 30	Kdev	0.98	Kmid	0.95	Kmid	1.27
July 31	Kdev	0.99	Kmid	0.95	Kmid	1.27
	17.1	1.00	Kmid	0.95	Kmid	1.27
Aug. 1	Kdev	1.00				
	Kdev	1.00	Kmid	0.95	Kmid	1.27
Aug. 1				0.95 0.95	Kmid Kmid	1.27 1.27

Aug. 5	Kdev	1.03	Kmid	0.95	Kmid	1.27
Aug. 6	Kdev	1.03	Kmid	0.95	Kmid	1.27
Aug. 7	Kdev	1.04	Kmid	0.95	Kmid	1.27
Aug. 8	Kdev	1.05	Kmid	0.95	Kmid	1.27
Aug. 9	Kdev	1.05	Kmid	0.95	Kmid	1.27
Aug. 10	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 11	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 12	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 12 Aug. 13	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 13 Aug. 14	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 14 Aug. 15	Kmid	1.06	Kmid	0.95	Kmid	1.27
	Kmid	1.00	Kmid	0.95	Kmid	1.27
Aug. 16						
Aug. 17	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 18	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 19	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 20	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 21	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 22	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 23	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 24	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 25	Klate	1.07	Kmid	0.95	Kmid	1.27
Aug. 26	Klate	1.08	Kmid	0.95	Kmid	1.27
Aug. 27	Klate	1.10	Kmid	0.95	Kmid	1.27
Aug. 28	Klate	1.11	Kmid	0.95	Kmid	1.27
Aug. 29	Klate	1.00	Kmid	0.95	Kmid	1.27
Aug. 30	Kini	0.95	Kmid	0.95	Kmid	1.27
Aug. 31	Kini	0.95	Kmid	0.95	Kmid	1.27
Sept. 1	Kini	0.95	Kmid	0.95	Kmid	1.27
Sept. 2	Kini	0.95	Kmid	0.95	Kmid	1.27
Sept. 2 Sept. 3	Kini	0.95	Kmid	0.95	Kmid	1.27
Sept. 3 Sept. 4	Kini	0.95	Kmid	0.95	Kmid	1.27
Sept. 4 Sept. 5	Kini	0.95	Kmid	0.95	Kmid	1.27
1						
Sept. 6	Kdev	0.96	Kmid	0.95	Kmid	1.27
Sept. 7	Kdev	0.96	Kmid	0.95	Kmid	1.27
Sept. 8	Kdev	0.97	Kmid	0.95	Kmid	1.27
Sept. 9	Kdev	0.98	Kmid	0.95	Kmid	1.27
Sept. 10	Kdev	0.98	Kmid	0.95	Kmid	1.27
Sept. 11	Kdev	0.99	Kmid	0.95	Kmid	1.27
Sept. 12	Kdev	1.00	Kmid	0.95	Kmid	1.27
Sept. 13	Kdev	1.01	Kmid	0.95	Kmid	1.27
Sept. 14	Kdev	1.01	Kmid	0.95	Kmid	1.27
Sept. 15	Kdev	1.02	Kmid	0.95	Kmid	1.27
Sept. 16	Kdev	1.03	Kmid	0.95	Kmid	1.27
Sept. 17	Kdev	1.03	Kmid	0.95	Kmid	1.27
Sept. 18	Kdev	1.04	Kmid	0.95	Kmid	1.27
Sept. 19	Kdev	1.05	Kmid	0.95	Kmid	1.27
Sept. 20	Kdev	1.05	Kmid	0.95	Klate	1.23
Sept. 20 Sept. 21	Kmid	1.06	Kmid	0.95	Klate	1.20
Sept. 21 Sept. 22	Kmid	1.06	Kmid	0.95	Klate	1.16
Sept. 22 Sept. 23	Kmid	1.06	Kmid	0.95	Klate	1.12
Sept. 23 Sept. 24	Kmid	1.06	Kmid	0.95	Klate	1.08
Sept. 24 Sept. 25	Kmid	1.06	Kmid	0.95	Klate	1.05
Sept. 25 Sept. 26	Kmid	1.06	Kmid	0.95	Klate	1.03
1			Kmid	0.95		0.97
Sept. 27	Kmid	1.06			Klate	
Sept. 28	Kmid	1.06 1.06	Kmid	0.95	Klate	0.93
0		1 1 1 6	Kmid	0.95	Klate	0.90
Sept. 29	Kmid			0.05	T71 /	0.07
Sept. 30	Kmid	1.06	Kmid	0.95	Klate	0.86
1				0.95 0.95 0.94	Klate Klate Klate	0.86 0.82 0.79

Oct. 3	Kmid	1.06	Klate	0.94	Klate	0.75
Oct. 4	Kmid	1.06	Klate	0.93	Klate	0.71
Oct. 5	Kmid	1.06	Klate	0.92	Klate	0.67
Oct. 6	Klate	1.05	Klate	0.92	Klate	0.64
Oct. 7	Klate	1.05	Klate	0.91	Klate	0.60
Oct. 8	Klate	1.04	Klate	0.90	Klate	0.56
Oct. 9	Klate	1.04	Klate	0.90	Klate	0.52
Oct. 10	Klate	1.03	Klate	0.89	Klate	0.49
Oct. 11	Klate	1.02	Klate	0.88	Klate	0.45
Oct. 12	Klate	1.02	Klate	0.88	Klate	0.43
Oct. 12 Oct. 13	Klate	1.02		0.87	Klate	0.41
			Klate			
Oct. 14	Klate	1.01	Klate	0.86	Klate	0.34
Oct. 15	Klate	1.00	Klate	0.86	Klate	0.30
Oct. 16	Bare Soil	0.20	Klate	0.85	Dormant	0.37
Oct. 17	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 18	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 19	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 20	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 21	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 22	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 23	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 24	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 25	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 26	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 27	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 27	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 28 Oct. 29	Bare Soil	0.20		0.40	+ +	0.37
			Dormant		Dormant	
Oct. 30	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 31	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 1	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 2	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 3	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 4	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 5	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 6	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 7	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 8	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 9	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 10	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 11	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 12	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 13	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 14	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 14	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 15	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 10	Bare Soil	0.20		0.40		0.37
			Dormant		Dormant	
Nov. 18	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 19	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 20	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 21	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 22	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 23	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 24	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 25	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 26	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 27	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 28	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 29	Bare Soil	0.20	Dormant	0.40	Dormant	0.37

Bara Soil	0.20	Dormant	0.40	Dormant	0.37
					0.37
		_		_	0.37
					0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Bare Soil	0.20	Dormant	0.40	Dormant	0.37
	Bare SoilBare Soil	Bare Soil 0.20 Bare Soil 0.20	Bare Soil0.20DormantBare Soil	Bare Soil0.20Dormant0.40Bare Soil0.20 <t< td=""><td>Bare Soil0.20Dormant0.40DormantBare Soil0.20Dormant0.40Dormant</td></t<>	Bare Soil0.20Dormant0.40DormantBare Soil0.20Dormant0.40Dormant

+3 cuttings were assumed for Hay

Table 5 : Annual evapotranspiration coefficients
for peaches, apples, cherries, pears, and grapes.

Tor pedenes, appres, enerries, pears, and grapes.								
	Peaches*	Apples, Cherries, Pears [‡]	Grapes ^a					
Jan. 1	0.2	0.2	0.2					
Jan. 2	0.2	0.2	0.2					
Jan. 3	0.2	0.2	0.2					
Jan. 4	0.2	0.2	0.2					
Jan. 5	0.2	0.2	0.2					
Jan. 6	0.2	0.2	0.2					
Jan. 7	0.2	0.2	0.2					
Jan. 8	0.2	0.2	0.2					
Jan. 9	0.2	0.2	0.2					
Jan. 10	0.2	0.2	0.2					
Jan. 11	0.2	0.2	0.2					
Jan. 12	0.2	0.2	0.2					
Jan. 13	0.2	0.2	0.2					
Jan. 14	0.2	0.2	0.2					
Jan. 15	0.2	0.2	0.2					
Jan. 16	0.2	0.2	0.2					
Jan. 17	0.2	0.2	0.2					
Jan. 18	0.2	0.2	0.2					
Jan. 19	0.2	0.2	0.2					
Jan. 20	0.2	0.2	0.2					

Jan. 21	0.2	0.2	0.2
Jan. 21	0.2	0.2	0.2
Jan. 22	0.2	0.2	0.2
Jan. 24	0.2	0.2	0.2
Jan. 24	0.2	0.2	0.2
Jan. 25	0.2	0.2	0.2
Jan. 20	0.2	0.2	0.2
Jan. 27	0.2	0.2	0.2
Jan. 29	0.2	0.2	0.2
Jan. 30	0.2	0.2	0.2
Jan. 30	0.2	0.2	0.2
Feb. 1	0.2	0.2	0.2
Feb. 2	0.2	0.2	0.2
Feb. 3	0.2	0.2	0.2
Feb. 4	0.2	0.2	0.2
Feb. 5	0.2	0.2	0.2
Feb. 6	0.2	0.2	0.2
Feb. 7	0.2	0.2	0.2
Feb. 7 Feb. 8	0.2	0.2	0.2
Feb. 8 Feb. 9	0.2	0.2	0.2
Feb. 9	0.2	0.2	0.2
Feb. 11	0.2	0.2	0.2
Feb. 12	0.2	0.2	0.2
Feb. 13	0.2	0.2	0.2
Feb. 14	0.2	0.2	0.2
Feb. 15	0.2	0.2	0.2
Feb. 16	0.2	0.2	0.2
Feb. 17	0.2	0.2	0.2
Feb. 18	0.2	0.2	0.2
Feb. 19	0.2	0.2	0.2
Feb. 20	0.2	0.2	0.2
Feb. 21	0.2	0.2	0.2
Feb. 22	0.2	0.2	0.2
Feb. 23	0.2	0.2	0.2
Feb. 24	0.2	0.2	0.2
Feb. 25	0.2	0.2	0.2
Feb. 26	0.2	0.2	0.2
Feb. 27	0.2	0.2	0.2
Feb. 28	0.2	0.2	0.2
Mar. 1	0.2	0.2	0.2
Mar. 2	0.2	0.2	0.2
Mar. 3	0.2	0.2	0.2
Mar. 4	0.2	0.2	0.2
Mar. 5	0.2	0.2	0.2
Mar. 6	0.2	0.2	0.2
Mar. 7	0.2	0.2	0.2
Mar. 8	0.2	0.2	0.2
Mar. 9	0.2	0.2	0.2
Mar. 10	0.2	0.2	0.2
Mar. 11	0.2	0.2	0.2
Mar. 12	0.2	0.2	0.2
Mar. 13	0.2	0.2	0.2
Mar. 14	0.2	0.2	0.2
Mar. 15	0.2	0.2	0.2
Mar. 16	0.2	0.2	0.2
Mar. 17	0.2	0.2	0.2
Mar. 18	0.2	0.2	0.2
Mar. 19	0.2	0.2	0.2
Mar. 20	0.2	0.2	0.2

M 21	0.2	0.2	0.0
Mar. 21	0.2	0.2	0.2
Mar. 22	0.2	0.2	0.2
Mar. 23	0.2	0.2	0.2
Mar. 24	0.2	0.2	0.2
Mar. 25	0.2	0.2	0.2
Mar. 26	0.2	0.2	0.2
Mar. 27	0.2	0.2	0.2
Mar. 28	0.2	0.2	0.2
Mar. 29	0.2	0.2	0.2
Mar. 30	0.2	0.2	0.2
Mar. 31	0.2	0.2	0.2
Apr. 1	0.2	0.2	0.2
Apr. 2	0.2	0.2	0.2
Apr. 3	0.2	0.2	0.2
Apr. 4	0.2	0.2	0.2
Apr. 5	0.2	0.2	0.2
Apr. 6	0.2	0.2	0.2
Apr. 7	0.2	0.2	0.2
Apr. 8	0.2	0.2	0.2
Apr. 9	0.2	0.2	0.2
Apr. 10	0.2	0.2	0.2
Apr. 11	0.2	0.2	0.2
Apr. 12	0.2	0.2	0.2
Apr. 13	0.2	0.2	0.2
Apr. 14	0.2	0.2	0.2
Apr. 15	0.2	0.2	0.2
Apr. 16	0.2	0.2	0.2
Apr. 17	0.2	0.2	0.2
Apr. 18	0.2	0.2	0.2
Apr. 19	0.2	0.2	0.2
Apr. 20	0.2	0.2	0.2
Apr. 21	0.2	0.2	0.2
Apr. 22	0.2	0.2	0.2
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Apr. 23	0.2	0.2	0.2
Apr. 24	0.2	0.2	0.2
Apr. 25	0.2	0.2	0.2
Apr. 26	0.2	0.2	0.2
Apr. 27	0.2	0.2	0.2
Apr. 28	0.2	0.2	0.2
Apr. 29	0.2	0.2	0.2
Apr. 30	0.2	0.2	0.2
May 1	0.3	0.3	0.5
May 2	0.3	0.3	0.5
May 3	0.3	0.3	0.5
May 4	0.3	0.3	0.5
May 5	0.3	0.3	0.5
May 6	0.3	0.3	0.5
	0.3	0.3	0.5
May 7			
May 8	0.3	0.3	0.5
May 9	0.3	0.3	0.5
May 10	0.3	0.3	0.5
May 11	0.3	0.3	0.5
May 12	0.3	0.3	0.5
May 13	0.3	0.3	0.5
May 14	0.3	0.3	0.5
May 14 May 15	0.3	0.3	0.5
	0.3	0.3	0.5
May 16			
May 17	0.3	0.3	0.5
May 18	0.3	0.3	0.5

May 19	0.3	0.3	0.5
May 20	0.3	0.3	0.5
May 21	0.3	0.3	0.5
May 22	0.3	0.3	0.5
May 23	0.3	0.3	0.5
May 24	0.3	0.3	0.5
May 25	0.3	0.3	0.5
May 26	0.3	0.3	0.5
May 27	0.3	0.3	0.5
May 28	0.3	0.3	0.5
May 29	0.3	0.3	0.5
May 30	0.3	0.3	0.5
May 31	0.3	0.3	0.5
June 1	0.4	0.4	0.68
June 2	0.4	0.4	0.68
June 3	0.4	0.4	0.68
June 4	0.4	0.4	0.68
June 5	0.4	0.4	0.68
June 6	0.4	0.4	0.68
June 7	0.4	0.4	0.68
June 8	0.4	0.4	0.68
June 9	0.4	0.4	0.68
June 10	0.4	0.4	0.68
June 11	0.4	0.4	0.68
June 12	0.4	0.4	0.68
June 13	0.4	0.4	0.68
June 14	0.4	0.4	0.68
June 15	0.4	0.4	0.68
June 16	0.6	0.6	0.68
June 17	0.6	0.6	0.68
June 18	0.6	0.6	0.68
June 19	0.6	0.6	0.68
June 20 June 21	0.6	0.6	0.68
June 21 June 22	0.6	0.6	0.68
June 23	0.6	0.6	0.68
June 24	0.6	0.6	0.68
June 25	0.6	0.6	0.68
June 26	0.6	0.6	0.68
June 27	0.6	0.6	0.68
June 28	0.6	0.6	0.68
June 29	0.6	0.6	0.68
June 30	0.6	0.6	0.68
July 1	1.0	1.0	0.77
July 2	1.0	1.0	0.77
July 3	1.0	1.0	0.77
July 4	1.0	1.0	0.77
July 5	1.0	1.0	0.77
July 6	1.0	1.0	0.77
July 7	1.0	1.0	0.77
July 8	1.0	1.0	0.77
July 9	1.0	1.0	0.77
July 10	1.0	1.0	0.77
July 11	1.0	1.0	0.77
July 12	1.0	1.0	0.77
July 13	1.0	1.0	0.77
July 14	1.0	1.0	0.77
July 15	1.0	1.0	0.77
July 16	1.0	1.0	0.77

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July 17	1.0	1.0	0.77
July 18	1.0	1.0	0.77
July 19	1.0	1.0	0.77
July 20	1.0	1.0	0.77
July 21	1.0	1.0	0.77
July 21			
July 22	1.0	1.0	0.77
July 23	1.0	1.0	0.77
July 24	1.0	1.0	0.77
July 25	1.0	1.0	0.77
July 26	1.0	1.0	0.77
July 27	1.0	1.0	0.77
July 28	1.0	1.0	0.77
July 29	1.0	1.0	0.77
July 30	1.0	1.0	0.77
July 31	1.0	1.0	0.77
Aug. 1	1.0	1.0	0.80
Aug. 2	1.0	1.0	0.80
Aug. 3	1.0	1.0	0.80
Aug. 4	1.0	1.0	0.80
Aug. 5	1.0	1.0	0.80
Aug. 6	1.0	1.0	0.80
Aug. 7	1.0	1.0	0.80
	1.0	1.0	0.80
Aug. 8			
Aug. 9	1.0	1.0	0.80
Aug. 10	1.0	1.0	0.80
Aug. 11	1.0	1.0	0.80
Aug. 12	1.0	1.0	0.80
Aug. 13	1.0	1.0	0.80
Aug. 14	1.0	1.0	0.80
Aug. 15	1.0	1.0	0.80
Aug. 16	1.0	1.0	0.80
Aug. 17	1.0	1.0	0.80
Aug. 18	1.0	1.0	0.80
Aug. 19	1.0	1.0	0.80
Aug. 20	1.0	1.0	0.80
Aug. 21	1.0	1.0	0.80
Aug. 22	1.0	1.0	0.80
Aug. 23	1.0	1.0	0.80
Aug. 24	1.0	1.0	0.80
Aug. 25	1.0	1.0	0.80
Aug. 25	1.0	1.0	0.80
	1.0	1.0	
Aug. 27			0.80
Aug. 28	1.0	1.0	0.80
Aug. 29	1.0	1.0	0.80
Aug. 30	1.0	1.0	0.80
Aug. 31	1.0	1.0	0.80
Sept. 1	0.95	0.95	0.75
Sept. 2	0.95	0.95	0.75
Sept. 3	0.95	0.95	0.75
Sept. 3	0.95	0.95	0.75
Sept. 4 Sept. 5	0.95	0.95	0.75
Sept. 6	0.95	0.95	0.75
Sept. 7	0.95	0.95	0.75
Sept. 8	0.95	0.95	0.75
Sept. 9	0.95	0.95	0.75
Sept. 10	0.95	0.95	0.75
Sept. 11	0.95	0.95	0.75
Sept. 12	0.95	0.95	0.75
Sept. 12 Sept. 13	0.95	0.95	0.75
50pt. 15	0.75	0.75	0.75

Sept. 14	0.95	0.95	0.75
Sept. 15	0.95	0.95	0.75
Sept. 16	0.95	0.95	0.75
Sept. 17	0.95	0.95	0.75
Sept. 18	0.95	0.95	0.75
Sept. 19	0.95	0.95	0.75
Sept. 20	0.95	0.95	0.75
Sept. 20	0.95	0.95	0.75
Sept. 21 Sept. 22	0.95	0.95	0.75
Sept. 22 Sept. 23	0.95	0.95	0.75
Sept. 23	0.95	0.95	0.75
Sept. 24 Sept. 25	0.95	0.95	0.75
			0.75
Sept. 26	0.95	0.95	
Sept. 27	0.95	0.95	0.75
Sept. 28	0.95	0.95	0.75
Sept. 29	0.95	0.95	0.75
Sept. 30	0.95	0.95	0.75
Oct. 1	0.83	0.80	0.63
Oct. 2	0.83	0.80	0.63
Oct. 3	0.83	0.80	0.63
Oct. 4	0.83	0.80	0.63
Oct. 5	0.83	0.80	0.63
Oct. 6	0.83	0.80	0.63
Oct. 7	0.83	0.80	0.63
Oct. 8	0.83	0.80	0.63
Oct. 9	0.83	0.80	0.63
Oct. 10	0.83	0.80	0.63
Oct. 11	0.83	0.80	0.63
Oct. 12	0.83	0.80	0.63
Oct. 13	0.83	0.80	0.63
Oct. 14	0.83	0.80	0.63
Oct. 15	0.83	0.80	0.63
Oct. 16	0.83	0.80	0.63
Oct. 17	0.83	0.80	0.63
Oct. 18	0.83	0.80	0.63
Oct. 18 Oct. 19	0.83	0.80	0.63
Oct. 19 Oct. 20	0.83	0.80	0.63
Oct. 21	0.83	0.80	0.63
Oct. 22	0.83	0.80	0.63
Oct. 23	0.83	0.80	0.63
Oct. 24	0.83	0.80	0.63
Oct. 25	0.83	0.80	0.63
Oct. 26	0.83	0.80	0.63
Oct. 27	0.83	0.80	0.63
Oct. 28	0.83	0.80	0.63
Oct. 29	0.83	0.80	0.63
Oct. 30	0.83	0.80	0.63
Oct. 31	0.83	0.80	0.63
Nov. 1	0.2	0.2	0.2
Nov. 2	0.2	0.2	0.2
Nov. 3	0.2	0.2	0.2
Nov. 4	0.2	0.2	0.2
Nov. 5	0.2	0.2	0.2
Nov. 6	0.2	0.2	0.2
Nov. 7	0.2	0.2	0.2
Nov. 8	0.2	0.2	0.2
Nov. 9	0.2	0.2	0.2
Nov. 10	0.2	0.2	0.2
Nov. 11	0.2	0.2	0.2
1,07,11	0.2	0.2	0.2

Nov. 12	0.2	0.2	0.2
Nov. 13	0.2	0.2	0.2
Nov. 14	0.2	0.2	0.2
Nov. 15	0.2	0.2	0.2
Nov. 16	0.2	0.2	0.2
Nov. 17	0.2	0.2	0.2
Nov. 18	0.2	0.2	0.2
Nov. 19	0.2	0.2	0.2
Nov. 20	0.2	0.2	0.2
Nov. 21	0.2	0.2	0.2
Nov. 22	0.2	0.2	0.2
Nov. 23	0.2	0.2	0.2
Nov. 24	0.2	0.2	0.2
Nov. 24	0.2	0.2	0.2
Nov. 26	0.2	0.2	0.2
Nov. 20	0.2	0.2	0.2
		0.2	0.2
Nov. 28	0.2	0.2	0.2
Nov. 29	0.2		
Nov. 30	0.2	0.2	0.2
Dec. 1	0.2	0.2	0.2
Dec. 2	0.2	0.2	0.2
Dec. 3	0.2	0.2	0.2
Dec. 4	0.2	0.2	0.2
Dec. 5	0.2	0.2	0.2
Dec. 6	0.2	0.2	0.2
Dec. 7	0.2	0.2	0.2
Dec. 8	0.2	0.2	0.2
Dec. 9	0.2	0.2	0.2
Dec. 10	0.2	0.2	0.2
Dec. 11	0.2	0.2	0.2
Dec. 12	0.2	0.2	0.2
Dec. 13	0.2	0.2	0.2
Dec. 14	0.2	0.2	0.2
Dec. 15	0.2	0.2	0.2
Dec. 16	0.2	0.2	0.2
Dec. 17	0.2	0.2	0.2
Dec. 18	0.2	0.2	0.2
Dec. 19	0.2	0.2	0.2
Dec. 20	0.2	0.2	0.2
Dec. 21	0.2	0.2	0.2
Dec. 22	0.2	0.2	0.2
Dec. 23	0.2	0.2	0.2
Dec. 24	0.2	0.2	0.2
Dec. 25	0.2	0.2	0.2
Dec. 26	0.2	0.2	0.2
Dec. 27	0.2	0.2	0.2
Dec. 28	0.2	0.2	0.2
Dec. 29	0.2	0.2	0.2
Dec. 30	0.2	0.2	0.2
Dec. 30	0.2	0.2	0.2
* Values for	D 1		

* Values for Peaches were determined using the OMAFRA Best Management Practices -Irrigation Management guide for mature fruit trees with permanent sod and herbicide strip. Values for January to April, and November to December were based on the April coefficient, which is equal to that of bare soil. The values from the British Columbia Ministry of Agriculture, Food and Fisheries were used for October. The values used are an average of the coefficients indicated for the 3 regions in British Columbia. [‡]The values for all months, excluding October, were determined from OMAFRA Best Management Practices - Irrigation Management guide for mature fruit trees with permanent sod and herbicide strip. The values for October were determined based on the average values indicated by the 3 regions in British Columbia by the Ministry of Agriculture, Food and Fisheries.

^aOntario does not have published values for grapes. The BC Ministry of Agriculture, Food and Fisheries was used to determine these values from May to the end of October. The values are an average of the 3 identified regions in BC. The coefficient for bare soil was used for the off season months.

Tule Tullu.						
	Deciduous Forest	Kc ^b	Coniferous Forest	Kc ^b	Idle Land	Kc ^b
Jan. 1	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 2	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 3	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 4	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 5	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 6	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 7	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 8	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 9	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 10	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 11	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 12	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 13	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 14	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 15	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 16	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 17	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 18	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 19	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 20	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 21	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 22	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 23	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 24	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 25	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 26	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 27	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 28	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 29	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 30	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 31	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 1	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 2	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 3	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 4	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 5	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 6	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 7	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 8	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 9	Dormant	0.2	Off Season	0.2	Dormant	0.2 0.2
Feb. 10	Dormant	0.2	Off Season	0.2	Dormant	
Feb. 11 Feb. 12	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 12	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 13	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 14	Dormant	0.2	Off Season	0.2	Dormant	0.2

Table 6: Annual evapotranspiration coefficients for deciduous and coniferous forests, and idle land.

Feb. 15	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 16	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 17	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 18	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 19	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 20	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 21	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 22	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 23	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 24	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 25	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 26	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 27	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 28	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 1	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 2	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 3	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 4	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 5	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 6	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 7	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 8	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 9	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 10	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 11	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 12	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 13	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 14	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 15	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 16	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 17	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 18	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 19	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 20	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 21	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 22	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 23	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 24	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 25	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 26	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 27	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 28	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 29	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 30	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 31	Dormant	0.2	Off Season	0.2	Dormant	0.2
Apr. 1	Kmid	1.09	Kmid	0.55	Kmid	1.00
Apr. 2	Kmid	1.09	Kmid	0.55	Kmid	1.01
Apr. 3	Kmid	1.09	Kmid	0.55	Kmid	1.07
Apr. 4	Kmid	1.09	Kmid	0.55	Kmid	1.06
Apr. 5	Kmid	1.09	Kmid	0.55	Kmid	0.96
Apr. 6	Kmid	1.09	Kmid	0.55	Kmid	0.98
Apr. 7	Kmid	1.09	Kmid	0.55	Kmid	0.93
Apr. 8	Kmid	1.09	Kmid	0.55	Kmid	1.01
Apr. 9	Kmid	1.09	Kmid	0.55	Kmid	1.00
Apr. 10	Kmid	1.09	Kmid	0.55	Kmid	1.03
Apr. 11	Kmid	1.09	Kmid	0.55	Kmid	1.00
<i>i</i> ipi, i i						
-	Kmid	1.09	Kmid	0.55	Kmid	1.05
Apr. 12 Apr. 13	Kmid Kmid	1.09 1.09	Kmid Kmid	0.55	Kmid Kmid	1.03

Apr. 15	Kmid	1.09	Kmid	0.55	Kmid	0.99
Apr. 16	Kmid	1.09	Kmid	0.55	Kmid	0.99
Apr. 17	Kmid	1.09	Kmid	0.55	Kmid	0.99
Apr. 18	Kmid	1.09	Kmid	0.55	Kmid	1.00
Apr. 19	Kmid	1.09	Kmid	0.55	Kmid	1.02
Apr. 20	Kmid	1.09	Kmid	0.55	Kmid	1.00
Apr. 20	Kmid	1.09	Kmid	0.55	Kmid	0.90
Apr. 22	Kmid	1.09	Kmid	0.55	Kmid	0.90
1		1.09				0.83
Apr. 23	Kmid		Kmid	0.55	Kmid	
Apr. 24	Kmid	1.09	Kmid	0.55	Kmid	1.04
Apr. 25	Kmid	1.09	Kmid	0.55	Kmid	1.06
Apr. 26	Kmid	1.09	Kmid	0.55	Kmid	1.03
Apr. 27	Kmid	1.09	Kmid	0.55	Kmid	1.03
Apr. 28	Kmid	1.09	Kmid	0.55	Kmid	1.07
Apr. 29	Kmid	1.09	Kmid	0.55	Kmid	1.08
Apr. 30	Kmid	1.09	Kmid	0.55	Kmid	1.07
May 1	Kmid	1.09	Kmid	0.55	Kmid	1.04
May 2	Kmid	1.09	Kmid	0.55	Kmid	1.04
May 3	Kmid	1.09	Kmid	0.55	Kmid	1.00
May 4	Kmid	1.09	Kmid	0.55	Kmid	0.98
May 5	Kmid	1.09	Kmid	0.55	Kmid	1.03
May 6	Kmid	1.09	Kmid	0.55	Kmid	1.00
May 7	Kmid	1.09	Kmid	0.55	Kmid	1.00
	Kmid	1.09	Kmid	0.55	Kmid	1.04
May 8						
May 9	Kmid	1.09	Kmid	0.55	Kmid	1.04
May 10	Kmid	1.09	Kmid	0.55	Kmid	1.02
May 11	Kmid	1.09	Kmid	0.55	Kmid	0.97
May 12	Kmid	1.09	Kmid	0.55	Kmid	0.96
May 13	Kmid	1.09	Kmid	0.55	Kmid	0.97
May 14	Kmid	1.09	Kmid	0.55	Kmid	0.94
May 15	Kmid	1.09	Kmid	0.55	Kmid	0.91
May 16	Kmid	1.09	Kmid	0.55	Kmid	0.96
May 17	Kmid	1.09	Kmid	0.55	Kmid	0.94
May 18	Kmid	1.09	Kmid	0.55	Kmid	0.97
May 19	Kmid	1.09	Kmid	0.55	Kmid	1.05
May 20	Kmid	1.09	Kmid	0.55	Kmid	1.04
May 21	Kmid	1.09	Kmid	0.55	Kmid	1.02
May 22	Kmid	1.09	Kmid	0.55	Kmid	1.03
May 22 May 23	Kmid	1.09	Kmid	0.55	Kmid	1.03
May 24	Kmid	1.09	Kmid	0.55	Kmid	0.99
May 24 May 25	Kmid	1.09	Kmid	0.55	Kmid	0.99
		1.09				
May 26	Kmid		Kmid	0.55	Kmid	0.98
May 27	Kmid	1.09	Kmid	0.55	Kmid	0.97
May 28	Kmid	1.09	Kmid	0.55	Kmid	1.01
May 29	Kmid	1.09	Kmid	0.55	Kmid	0.99
May 30	Kmid	1.09	Kmid	0.55	Kmid	0.97
May 31	Kmid	1.09	Kmid	0.55	Kmid	0.88
June 1	Kmid	1.09	Kmid	0.55	Kmid	0.91
June 2	Kmid	1.09	Kmid	0.55	Kmid	0.87
June 3	Kmid	1.09	Kmid	0.55	Kmid	1.01
June 4	Kmid	1.09	Kmid	0.55	Kmid	1.01
		1.09	Kmid	0.55	Kmid	1.00
June 5	Kmid	1.09				
June 5 June 6	Kmid Kmid		Kmid	0.55	Kmid	1.00
June 6	Kmid	1.09	Kmid Kmid	0.55	Kmid Kmid	1.00
June 6 June 7	Kmid Kmid	1.09 1.09	Kmid	0.55	Kmid	0.99
June 6 June 7 June 8	Kmid Kmid Kmid	1.09 1.09 1.09	Kmid Kmid	0.55 0.55	Kmid Kmid	0.99 0.98
June 6 June 7 June 8 June 9	Kmid Kmid Kmid Kmid	1.09 1.09 1.09 1.09	Kmid Kmid Kmid	0.55 0.55 0.55	Kmid Kmid Kmid	0.99 0.98 1.08
June 6 June 7 June 8 June 9 June 10	Kmid Kmid Kmid Kmid Kmid	1.09 1.09 1.09 1.09 1.09	Kmid Kmid Kmid Kmid	0.55 0.55 0.55 0.55	Kmid Kmid Kmid Kmid	0.99 0.98 1.08 1.00
June 6 June 7 June 8 June 9	Kmid Kmid Kmid Kmid	1.09 1.09 1.09 1.09	Kmid Kmid Kmid	0.55 0.55 0.55	Kmid Kmid Kmid	0.99 0.98 1.08

June 13 Kmid 1.09 Kmid	0.55 Kmid 0.98
June 14 Kmid 1.09 Kmid	0.55 Kmid 1.03
June 15 Kmid 1.09 Kmid	0.55 Kmid 1.04
June 16 Kmid 1.09 Kmid	0.55 Kmid 1.06
June 17 Kmid 1.09 Kmid	0.55 Kmid 1.03
June 18 Kmid 1.09 Kmid	0.55 Kmid 0.97
June 19 Kmid 1.09 Kmid	0.55 Kmid 1.00
June 20 Kmid 1.09 Kmid	0.55 Kmid 0.97
June 21 Kmid 1.09 Kmid	
June 22 Kmid 1.09 Kmid	0.55 Kmid 0.95
June 23 Kmid 1.09 Kmid	0.55 Kmid 0.93
June 24 Kmid 1.09 Kmid	0.55 Kmid 0.97
June 25 Kmid 1.09 Kmid	0.55 Kmid 0.93
June 26 Kmid 1.09 Kmid	0.55 Kmid 0.89
June 27 Kmid 1.09 Kmid	0.55 Kmid 0.95
June 28 Kmid 1.09 Kmid	0.55 Kmid 0.96
June 29 Kmid 1.09 Kmid	0.55 Kmid 0.98
June 30 Kmid 1.09 Kmid	0.55 Kmid 1.04
July 1 Kmid 1.09 Kmid	0.55 Kmid 1.03
July 2 Kmid 1.09 Kmid	0.55 Kmid 0.97
July 3 Kmid 1.09 Kmid	0.55 Kmid 1.00
July 4 Kmid 1.09 Kmid	0.55 Kmid 1.00
July 5 Kmid 1.09 Kmid	0.55 Kmid 0.98
	0.55 Kmid 1.01
July 7 Kmid 1.09 Kmid	0.55 Kmid 1.02
July 8 Kmid 1.09 Kmid	0.55 Kmid 1.03
July 9 Kmid 1.09 Kmid	0.55 Kmid 0.94
July 10 Kmid 1.09 Kmid	0.55 Kmid 0.92
July 11 Kmid 1.09 Kmid	0.55 Kmid 0.88
July 12 Kmid 1.09 Kmid	0.55 Kmid 0.96
July 13 Kmid 1.09 Kmid	0.55 Kmid 0.94
July 14 Kmid 1.09 Kmid	0.55 Kmid 0.97
July 15 Kmid 1.09 Kmid	0.55 Kmid 0.99
July 16 Kmid 1.09 Kmid	0.55 Kmid 1.01
July 17 Kmid 1.09 Kmid	0.55 Kmid 1.01
July 18 Kmid 1.09 Kmid	0.55 Kmid 0.92
July 19 Kmid 1.09 Kmid	0.55 Kmid 0.95
July 20 Kmid 1.09 Kmid	0.55 Kmid 0.92
July 20 Kmid 1.09 Kmid	0.55 Kmid 0.90
July 22 Kmid 1.09 Kmid	0.55 Kmid 0.94
July 23 Kmid 1.09 Kmid	0.55 Kmid 1.01
	0.55 Kmid 0.99
July 25 Kmid 1.09 Kmid	0.55 Kmid 1.01
July 26 Kmid 1.09 Kmid	0.55 Kmid 0.96
July 27 Kmid 1.09 Kmid	0.55 Kmid 0.94
July 28 Kmid 1.09 Kmid	0.55 Kmid 0.95
July 29 Kmid 1.09 Kmid	0.55 Kmid 0.91
July 30 Kmid 1.09 Kmid	0.55 Kmid 0.95
July 31 Kmid 1.09 Kmid	0.55 Kmid 0.99
Aug. 1 Kmid 1.09 Kmid	0.55 Kmid 1.01
Aug. 2 Kmid 1.09 Kmid	0.55 Kmid 0.90
ě	0.55 Kmid 1.00
Aug. 3 Kmid 1.09 Kmid	
Ę	0.55 Kmid 0.95
Aug. 4 Kmid 1.09 Kmid	0.55 Kmid 0.95
Aug. 4Kmid1.09KmidAug. 5Kmid1.09Kmid	0.55 Kmid 0.98
Aug. 4Kmid1.09KmidAug. 5Kmid1.09KmidAug. 6Kmid1.09Kmid	0.55 Kmid 0.98 0.55 Kmid 1.00
Aug. 4 Kmid 1.09 Kmid Aug. 5 Kmid 1.09 Kmid Aug. 6 Kmid 1.09 Kmid Aug. 7 Kmid 1.09 Kmid	0.55 Kmid 0.98 0.55 Kmid 1.00 0.55 Kmid 1.02
Aug. 4 Kmid 1.09 Kmid Aug. 5 Kmid 1.09 Kmid Aug. 6 Kmid 1.09 Kmid Aug. 7 Kmid 1.09 Kmid Aug. 8 Kmid 1.09 Kmid	0.55 Kmid 0.98 0.55 Kmid 1.00 0.55 Kmid 1.02 0.55 Kmid 1.01
Aug. 4Kmid1.09KmidAug. 5Kmid1.09KmidAug. 6Kmid1.09KmidAug. 7Kmid1.09Kmid	0.55 Kmid 0.98 0.55 Kmid 1.00 0.55 Kmid 1.02

Aug. 11	Kmid	1.09	Kmid	0.55	Kmid	1.01
Aug. 12	Kmid	1.09	Kmid	0.55	Kmid	1.01
Aug. 13	Kmid	1.09	Kmid	0.55	Kmid	0.97
Aug. 14	Kmid	1.09	Kmid	0.55	Kmid	1.02
Aug. 15	Kmid	1.09	Kmid	0.55	Kmid	0.98
Aug. 16	Kmid	1.09	Kmid	0.55	Kmid	1.00
Aug. 17	Kmid	1.09	Kmid	0.55	Kmid	0.96
Aug. 17	Kmid	1.09	Kmid	0.55	Kmid	0.88
Aug. 18 Aug. 19	Kmid	1.09	Kmid	0.55	Kmid	0.93
Aug. 19 Aug. 20	Kmid	1.09	Kmid	0.55	Kmid	1.00
U						
Aug. 21	Kmid	1.09	Kmid	0.55	Kmid	0.99
Aug. 22	Kmid	1.09	Kmid	0.55	Kmid	1.01
Aug. 23	Kmid	1.09	Kmid	0.55	Kmid	0.95
Aug. 24	Kmid	1.09	Kmid	0.55	Kmid	0.91
Aug. 25	Kmid	1.09	Kmid	0.55	Kmid	0.92
Aug. 26	Kmid	1.09	Kmid	0.55	Kmid	0.92
Aug. 27	Kmid	1.09	Kmid	0.55	Kmid	0.91
Aug. 28	Kmid	1.09	Kmid	0.55	Kmid	0.93
Aug. 29	Kmid	1.09	Kmid	0.55	Kmid	0.98
Aug. 30	Kmid	1.09	Kmid	0.55	Kmid	0.97
Aug. 31	Kmid	1.09	Kmid	0.55	Kmid	0.97
Sept. 1	Kmid	1.09	Kmid	0.55	Kmid	0.93
Sept. 2	Kmid	1.09	Kmid	0.55	Kmid	0.90
Sept. 2 Sept. 3	Kmid	1.09	Kmid	0.55	Kmid	0.92
Sept. 3	Kmid	1.09	Kmid	0.55	Kmid	0.88
Sept. 4 Sept. 5	Kmid	1.09	Kmid	0.55	Kmid	0.88
-						0.92
Sept. 6	Kmid	1.09	Kmid	0.55	Kmid	
Sept. 7	Kmid	1.09	Kmid	0.55	Kmid	0.99
Sept. 8	Kmid	1.09	Kmid	0.55	Kmid	0.91
Sept. 9	Kmid	1.09	Kmid	0.55	Kmid	0.98
Sept. 10	Kmid	1.09	Kmid	0.55	Kmid	0.98
Sept. 11	Kmid	1.09	Kmid	0.55	Kmid	0.90
Sept. 12	Kmid	1.09	Kmid	0.55	Kmid	0.91
Sept. 13	Kmid	1.09	Kmid	0.55	Kmid	0.88
Sept. 14	Kmid	1.09	Kmid	0.55	Kmid	0.88
Sept. 15	Kmid	1.09	Kmid	0.55	Kmid	0.87
Sept. 16	Kmid	1.09	Kmid	0.55	Kmid	0.94
Sept. 17	Kmid	1.09	Kmid	0.55	Kmid	1.00
Sept. 18	Kmid	1.09	Kmid	0.55	Kmid	0.98
Sept. 19	Kmid	1.09	Kmid	0.55	Kmid	1.00
Sept. 20	Kmid	1.09	Kmid	0.55	Kmid	0.97
Sept. 20 Sept. 21	Kmid	1.09	Kmid	0.55	Kmid	0.97
Sept. 21 Sept. 22	Kmid	1.09	Kmid	0.55	Kmid	0.92
1						
Sept. 23	Kmid	1.09	Kmid	0.55	Kmid	1.01
Sept. 24	Kmid	1.09	Kmid	0.55	Kmid	1.00
Sept. 25	Kmid	1.09	Kmid	0.55	Kmid	0.96
Sept. 26	Kmid	1.09	Kmid	0.55	Kmid	1.01
Sept. 27	Kmid	1.09	Kmid	0.55	Kmid	0.92
Sept. 28	Kmid	1.09	Kmid	0.55	Kmid	0.94
Sept. 29	Kmid	1.09	Kmid	0.55	Kmid	0.95
Sept. 30	Kmid	1.09	Kmid	0.55	Kmid	0.95
Oct. 1	Kmid	1.09	Kmid	0.55	Kmid	0.95
Oct. 2	Kmid	1.09	Kmid	0.55	Kmid	0.91
Oct. 3	Kmid	1.09	Kmid	0.55	Kmid	0.98
Oct. 4	Kmid	1.09	Kmid	0.55	Kmid	0.98
Oct. 5	Kmid	1.09	Kmid	0.55	Kmid	0.93
Oct. 6	Kmid	1.09	Kmid	0.55	Kmid	0.95
Oct. 7	Kmid	1.09	Kmid	0.55	Kmid	0.92
Oct. 8	Kmid	1.09	Kmid	0.55	Kmid	0.92
001.0	KIIIU	1.09	KIIIU	0.35	Killiu	0.95

Oct. 9 Oct. 10 Oct. 11	V. 1					
	Kmid	1.09	Kmid	0.55	Kmid	0.94
Oct. 11	Kmid	1.09	Kmid	0.55	Kmid	0.93
	Kmid	1.09	Kmid	0.55	Kmid	0.98
Oct. 12	Kmid	1.09	Kmid	0.55	Kmid	1.01
Oct. 13	Kmid	1.09	Kmid	0.55	Kmid	1.00
Oct. 14	Kmid	1.09	Kmid	0.55	Kmid	1.01
Oct. 15	Kmid	1.09	Kmid	0.55	Kmid	0.99
Oct. 16	Kmid	1.09	Kmid	0.55	Kmid	0.98
Oct. 17	Kmid	1.09	Kmid	0.55	Kmid	0.98
Oct. 17	Kmid	1.09	Kmid	0.55	Kmid	0.94
		1.09				0.90
Oct. 19	Kmid		Kmid	0.55	Kmid	
Oct. 20	Kmid	1.09	Kmid	0.55	Kmid	0.91
Oct. 21	Kmid	1.09	Kmid	0.55	Kmid	0.92
Oct. 22	Kmid	1.09	Kmid	0.55	Kmid	1.00
Oct. 23	Kmid	1.09	Kmid	0.55	Kmid	0.97
Oct. 24	Kmid	1.09	Kmid	0.55	Kmid	0.99
Oct. 25	Kmid	1.09	Kmid	0.55	Kmid	0.95
Oct. 26	Kmid	1.09	Kmid	0.55	Kmid	0.92
Oct. 27	Kmid	1.09	Kmid	0.55	Kmid	0.95
Oct. 28	Kmid	1.09	Kmid	0.55	Kmid	1.07
Oct. 29	Kmid	1.09	Kmid	0.55	Kmid	1.01
Oct. 30	Kmid	1.09	Kmid	0.55	Kmid	1.03
Oct. 31	Kmid	1.09	Kmid	0.55	Kmid	0.99
Nov. 1	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 1 Nov. 2	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 2 Nov. 3	Dormant	0.2	Off Season	0.2	Dormant	0.2
		0.2				0.2
Nov. 4	Dormant		Off Season	0.2	Dormant	
Nov. 5	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 6	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 7	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 8	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 9	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 10	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 11	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 12	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 13	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 14	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 15	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 16	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 17	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 18	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 19	Dormant	0.2	Off Season	0.2	Dormant	0.2
		0.2	Off Season		+	
Nov. 20	Dormant			0.2	Dormant	0.2
Nov. 21	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 22	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 23	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 24	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 25	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 26	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 27	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 28	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 29	Dormant	0.2	Off Season	0.2	Dormant	0.2
INUV. 29	Dormant	0.2	Off Season	0.2	Dormant	0.2
	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 30		0.2	Off Season	0.2	Dormant	0.2
Nov. 30 Dec. 1	Dormant			0.2	2 on main	
Nov. 30 Dec. 1 Dec. 2	Dormant Dormant		Off Season	0.2	Dormant	0.2
Nov. 30 Dec. 1 Dec. 2 Dec. 3	Dormant	0.2	Off Season	0.2	Dormant Dormant	0.2
Nov. 30 Dec. 1 Dec. 2			Off Season Off Season Off Season	0.2 0.2 0.2	Dormant Dormant Dormant	

	D	0.0	0.00	0.0	D	0.0
Dec. 7	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 8	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 9	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 10	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 11	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 12	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 13	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 14	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 15	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 16	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 17	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 18	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 19	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 20	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 21	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 22	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 23	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 24	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 25	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 26	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 27	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 28	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 29	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 30	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 31	Dormant	0.2	Off Season	0.2	Dormant	0.2
	• • • • •	.1 1 1	11 /1	•		1

^bBased on the information available regarding the evapotranspiration of deciduous and coniferous forest, and idle land, step-wise changes in values were used. A growing season of April 1 to Oct. 31 was estimated in order to include bud development as well as killing frost at the end of the season. Should these values not be appropriate for the region of concern or vary seasonally, they should be adjusted accordingly.

Table 7: Annual evapotranspiration coefficients for tallgrass, built-up pervious areas, and open water.

	Tallgrass ^c	Kc ^d	Built-Up Pervious Area	Kc ^d	Open Water (Shallow)	Open Water (>5m depth) ^e
Jan. 1	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 2	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 3	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 4	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 5	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 6	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 7	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 8	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 9	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 10	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 11	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 12	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 13	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 14	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 15	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 16	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 17	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 18	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 19	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 20	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 21	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 22	Dormant	0.20	Reference	1.0	1.05	1.25

Jan. 23	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 24	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 25	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 26	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 27	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 28	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 20	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 30	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 30	Dormant	0.20	Reference	1.0	1.05	1.25
		0.20	Reference	1.0	1.05	1.25
Feb. 1	Dormant					
Feb. 2	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 3	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 4	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 5	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 6	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 7	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 8	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 9	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 10	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 11	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 12	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 13	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 14	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 15	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 16	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 17	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 18	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 19	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 20	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 21	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 22	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 22 Feb. 23	_	0.20	Reference	1.0	1.05	1.25
	Dormant	0.20				
Feb. 24	Dormant		Reference	1.0	1.05	1.25
Feb. 25	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 26	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 27	Dormant	0.20	Reference	1.0	1.05	1.25
Feb. 28	Dormant	0.20	Reference	1.0	1.05	1.25
Mar. 1	Dormant	0.20	Reference	1.0	1.05	1.25
Mar. 2	Dormant	0.20	Reference	1.0	1.05	1.25
Mar. 3	Dormant	0.20	Reference	1.0	1.05	1.25
Mar. 4	Dormant	0.20	Reference	1.0	1.05	1.25
Mar. 5	Dormant	0.20	Reference	1.0	1.05	1.25
Mar. 6	Dormant	0.20	Reference	1.0	1.05	1.25
Mar. 7	Dormant	0.20	Reference	1.0	1.05	1.25
Mar. 8	Dormant	0.20	Reference	1.0	1.05	1.25
Mar. 9	Dormant	0.20	Reference	1.0	1.05	1.25
Mar. 10	Dormant	0.20	Reference	1.0	1.05	1.25
Mar. 11	Dormant	0.20	Reference	1.0	1.05	1.25
Mar. 12	Dormant	0.20	Reference	1.0	1.05	1.25
Mar. 13	Dormant	0.20	Reference	1.0	1.05	1.25
Mar. 14	Dormant	0.20	Reference	1.0	1.05	1.25
Mar. 15	Kini	0.20	Reference	1.0	1.05	1.25
Mar. 16	Kini	0.38	Reference	1.0	1.05	1.25
Mar. 17	Kini	0.38	Reference	1.0	1.05	1.25
Mar. 18	Kini	0.38	Reference	1.0	1.05	1.25
	Kini	0.38	Reference	1.0	1.05	1.25
Mar. 19						
Mar. 20	Kini	0.38	Reference	1.0	1.05	1.25
Mar. 21	Kini	0.37	Reference	1.0	1.05	1.25
Mar. 22	Kini	0.37	Reference	1.0	1.05	1.25

Mar. 23	Kini	0.36	Reference	1.0	1.05	1.25
Mar. 24	Kini	0.36	Reference	1.0	1.05	1.25
Mar. 25	Kini	0.36	Reference	1.0	1.05	1.25
Mar. 26	Kini	0.37	Reference	1.0	1.05	1.25
Mar. 27	Kini	0.38	Reference	1.0	1.05	1.25
Mar. 28	Kini	0.37	Reference	1.0	1.05	1.25
Mar. 29	Kini	0.37	Reference	1.0	1.05	1.25
Mar. 30	Kini	0.37	Reference	1.0	1.05	1.25
Mar. 31	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 1	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 2	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 3	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 4	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 5	Kini	0.36	Reference	1.0	1.05	1.25
Apr. 6	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 7	Kini	0.36	Reference	1.0	1.05	1.25
Apr. 8	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 9	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 10	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 11	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 12	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 13	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 14	Kini	0.35	Reference	1.0	1.05	1.25
Apr. 15	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 16	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 17	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 18	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 19	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 20	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 21	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 22	Kini	0.36	Reference	1.0	1.05	1.25
Apr. 23	Kini	0.35	Reference	1.0	1.05	1.25
Apr. 24	Kini	0.36	Reference	1.0	1.05	1.25
Apr. 25	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 26	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 27	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 28	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 29	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 30	Kini	0.38	Reference	1.0	1.05	1.25
May 1	Kini	0.38	Reference	1.0	1.05	0.65
May 2	Kini	0.38	Reference	1.0	1.05	0.65
May 3	Kini	0.38	Reference	1.0	1.05	0.65
May 4	Kini	0.37	Reference	1.0	1.05	0.65
May 5	Kini	0.37	Reference	1.0	1.05	0.65
May 6	Kini	0.38	Reference	1.0	1.05	0.65
May 7	Kini	0.37	Reference	1.0	1.05	0.65
May 8	Kini	0.38	Reference	1.0	1.05	0.65
May 9	Kini	0.38	Reference	1.0	1.05	0.65
May 10	Kini	0.38	Reference	1.0	1.05	0.65
May 11	Kini	0.37	Reference	1.0	1.05	0.65
May 12	Kini	0.37	Reference	1.0	1.05	0.65
May 13	Kini	0.36	Reference	1.0	1.05	0.65
May 14	Kini	0.37	Reference	1.0	1.05	0.65
May 15	Kdev	0.85	Reference	1.0	1.05	0.65
May 16	Kdev	0.83	Reference	1.0	1.05	0.65
May 17	Kdev	0.86	Reference	1.0	1.05	0.65
May 18	Kdev	0.85	Reference	1.0	1.05	0.65
May 19	Kdev	0.86	Reference	1.0	1.05	0.65
May 20	Kdev	0.91	Reference	1.0	1.05	0.65

May 21	Kdev	0.91	Reference	1.0	1.05	0.65
May 22	Kdev	0.90	Reference	1.0	1.05	0.65
May 23	Kdev	0.90	Reference	1.0	1.05	0.65
May 24	Kdev	0.90	Reference	1.0	1.05	0.65
May 25	Kdev	0.87	Reference	1.0	1.05	0.65
May 26	Kdev	0.81	Reference	1.0	1.05	0.65
May 27	Kdev	0.87	Reference	1.0	1.05	0.65
,				1.0	1.05	
May 28	Kdev	0.87	Reference			0.65
May 29	Kdev	0.89	Reference	1.0	1.05	0.65
May 30	Kdev	0.88	Reference	1.0	1.05	0.65
May 31	Kdev	0.87	Reference	1.0	1.05	0.65
June 1	Kmid	1.19	Reference	1.0	1.05	0.65
June 2	Kmid	1.22	Reference	1.0	1.05	0.65
June 3	Kmid	1.18	Reference	1.0	1.05	0.65
June 4	Kmid	1.32	Reference	1.0	1.05	0.65
June 5	Kmid	1.32	Reference	1.0	1.05	0.65
June 6	Kmid	1.32	Reference	1.0	1.05	0.65
June 7	Kmid	1.31	Reference	1.0	1.05	0.65
June 8	Kmid	1.30	Reference	1.0	1.05	0.65
June 9	Kmid	1.29	Reference	1.0	1.05	0.65
June 10	Kmid	1.38	Reference	1.0	1.05	0.65
June 11	Kmid	1.31	Reference	1.0	1.05	0.65
June 12	Kmid	1.28	Reference	1.0	1.05	0.65
June 13	Kmid	1.29	Reference	1.0	1.05	0.65
June 14	Kmid	1.29	Reference	1.0	1.05	0.65
June 15	Kmid	1.33	Reference	1.0	1.05	0.65
June 16	Kmid	1.34	Reference	1.0	1.05	0.65
June 17	Kmid	1.36	Reference	1.0	1.05	0.65
June 18	Kmid	1.33	Reference	1.0	1.05	0.65
June 19	Kmid	1.28	Reference	1.0	1.05	0.65
June 20	Kmid	1.31	Reference	1.0	1.05	0.65
June 21	Kmid	1.28	Reference	1.0	1.05	0.65
June 22	Kmid	1.33	Reference	1.0	1.05	0.65
June 23	Kmid	1.26	Reference	1.0	1.05	0.65
June 24	Kmid	1.23	Reference	1.0	1.05	0.65
June 25	Kmid	1.27	Reference	1.0	1.05	0.65
June 26	Kmid	1.24	Reference	1.0	1.05	0.65
June 27	Kmid	1.20	Reference	1.0	1.05	0.65
June 28	Kmid	1.26	Reference	1.0	1.05	0.65
June 29	Kmid	1.26	Reference	1.0	1.05	0.65
June 30	Kmid	1.29	Reference	1.0	1.05	0.65
July 1	Kmid	1.34	Reference	1.0	1.05	0.65
July 2	Kmid	1.34	Reference	1.0	1.05	0.65
July 3	Kmid	1.28	Reference	1.0	1.05	0.65
July 4	Kmid	1.30	Reference	1.0	1.05	0.65
July 5	Kmid	1.31	Reference	1.0	1.05	0.65
July 6	Kmid	1.29	Reference	1.0	1.05	0.65
July 7	Kmid	1.32	Reference	1.0	1.05	0.65
July 8	Kmid	1.32	Reference	1.0	1.05	0.65
July 9	Kmid	1.34	Reference	1.0	1.05	0.65
July 10	Kmid	1.25	Reference	1.0	1.05	0.65
July 10	Kmid	1.23	Reference	1.0	1.05	0.65
,						
July 12	Kmid	1.19	Reference	1.0	1.05	0.65
July 13	Kmid	1.26	Reference	1.0	1.05	0.65
July 14	Kmid	1.25	Reference	1.0	1.05	0.65
July 15	Kmid	1.27	Reference	1.0	1.05	0.65
July 16	Kmid	1.29	Reference	1.0	1.05	0.65
,	Kmid	1.32	Reference	1.0	1.05	0.65
July 17	IXIIIIU	1.52	Reference	1.0	1.05	0.05

July 19	Kmid	1.23	Reference	1.0	1.05	0.65
July 20	Kmid	1.26	Reference	1.0	1.05	0.65
July 21	Kmid	1.23	Reference	1.0	1.05	0.65
July 22	Kmid	1.21	Reference	1.0	1.05	0.65
July 23	Kmid	1.25	Reference	1.0	1.05	0.65
July 24	Kmid	1.32	Reference	1.0	1.05	0.65
July 25	Kmid	1.29	Reference	1.0	1.05	0.65
July 25	Kmid	1.32	Reference	1.0	1.05	0.65
				1.0		0.65
July 27	Kmid	1.27	Reference		1.05	
July 28	Kmid	1.25	Reference	1.0	1.05	0.65
July 29	Kmid	1.26	Reference	1.0	1.05	0.65
July 30	Kmid	1.22	Reference	1.0	1.05	0.65
July 31	Kmid	1.26	Reference	1.0	1.05	0.65
Aug. 1	Kmid	1.29	Reference	1.0	1.05	0.65
Aug. 2	Kmid	1.32	Reference	1.0	1.05	0.65
Aug. 3	Kmid	1.21	Reference	1.0	1.05	0.65
Aug. 4	Kmid	1.30	Reference	1.0	1.05	0.65
Aug. 5	Kmid	1.26	Reference	1.0	1.05	0.65
Aug. 6	Kmid	1.29	Reference	1.0	1.05	0.65
Aug. 7	Kmid	1.31	Reference	1.0	1.05	0.65
Aug. 8	Kmid	1.33	Reference	1.0	1.05	0.65
Aug. 9	Kmid	1.32	Reference	1.0	1.05	0.65
Aug. 10	Kmid	1.29	Reference	1.0	1.05	0.65
Aug. 10 Aug. 11	Kmid	1.31	Reference	1.0	1.05	0.65
Aug. 12	Kmid	1.32	Reference	1.0	1.05	0.65
Aug. 13	Kmid	1.31	Reference	1.0	1.05	0.65
Aug. 14	Kmid	1.27	Reference	1.0	1.05	0.65
Aug. 15	Kmid	1.32	Reference	1.0	1.05	0.65
Aug. 16	Kmid	1.29	Reference	1.0	1.05	0.65
Aug. 17	Kmid	1.30	Reference	1.0	1.05	0.65
Aug. 18	Kmid	1.26	Reference	1.0	1.05	0.65
Aug. 19	Kmid	1.19	Reference	1.0	1.05	0.65
Aug. 20	Kmid	1.24	Reference	1.0	1.05	0.65
Aug. 21	Kmid	1.30	Reference	1.0	1.05	0.65
Aug. 22	Kmid	1.30	Reference	1.0	1.05	0.65
Aug. 23	Kmid	1.31	Reference	1.0	1.05	0.65
Aug. 24	Kmid	1.26	Reference	1.0	1.05	0.65
Aug. 25	Kmid	1.22	Reference	1.0	1.05	0.65
Aug. 26	Kmid	1.23	Reference	1.0	1.05	0.65
Aug. 20	Kmid	1.23	Reference	1.0	1.05	0.65
Aug. 27 Aug. 28	Kmid	1.22	Reference	1.0	1.05	0.65
		1.22				
Aug. 29	Kmid		Reference	1.0	1.05	0.65
Aug. 30	Kmid	1.28	Reference	1.0	1.05	0.65
Aug. 31	Kmid	1.27	Reference	1.0	1.05	0.65
Sept. 1	Klate	0.86	Reference	1.0	1.05	0.65
Sept. 2	Klate	0.84	Reference	1.0	1.05	0.65
Sept. 3	Klate	0.82	Reference	1.0	1.05	0.65
Sept. 4	Klate	0.84	Reference	1.0	1.05	0.65
Sept. 5	Klate	0.81	Reference	1.0	1.05	0.65
Sept. 6	Klate	0.84	Reference	1.0	1.05	0.65
Sept. 7	Klate	0.85	Reference	1.0	1.05	0.65
Sept. 8	Klate	0.87	Reference	1.0	1.05	0.65
Sept. 9	Klate	0.83	Reference	1.0	1.05	0.65
Sept. 10	Klate	0.87	Reference	1.0	1.05	0.65
Sept. 10	Klate	0.87	Reference	1.0	1.05	0.65
-		0.87		1.0		
Sept. 12	Klate		Reference		1.05	0.65
Sept. 13	Klate	0.83	Reference	1.0	1.05	0.65
Sept. 14	Klate	0.81	Reference Reference	1.0	1.05 1.05	0.65
Sept. 15	Klate					

Sept. 16	Klate	0.80	Reference	1.0	1.05	0.65
Sept. 17	Klate	0.85	Reference	1.0	1.05	0.65
Sept. 18	Klate	0.88	Reference	1.0	1.05	0.65
Sept. 19	Klate	0.87	Reference	1.0	1.05	0.65
Sept. 20	Klate	0.88	Reference	1.0	1.05	0.65
Sept. 21	Klate	0.87	Reference	1.0	1.05	0.65
Sept. 22	Klate	0.84	Reference	1.0	1.05	0.65
Sept. 23	Klate	0.84	Reference	1.0	1.05	0.65
Sept. 23 Sept. 24	Klate	0.89	Reference	1.0	1.05	0.65
Sept. 21 Sept. 25	Klate	0.88	Reference	1.0	1.05	0.65
Sept. 25	Klate	0.86	Reference	1.0	1.05	0.65
Sept. 20	Klate	0.89	Reference	1.0	1.05	0.65
Sept. 27 Sept. 28	Klate	0.83	Reference	1.0	1.05	0.65
<u> </u>	Klate	0.85			1.05	0.65
Sept. 29			Reference	1.0		
Sept. 30	Klate	0.85	Reference	1.0	1.05	0.65
Oct. 1	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 2	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 3	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 4	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 5	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 6	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 7	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 8	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 9	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 10	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 11	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 12	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 13	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 14	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 15	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 16	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 10 Oct. 17	Klate	0.37	Reference	1.0	1.05	1.25
		0.37				
Oct. 18	Klate		Reference	1.0	1.05	1.25
Oct. 19	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 20	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 21	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 22	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 23	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 24	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 25	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 26	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 27	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 28	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 29	Klate	0.38	Reference	1.0	1.05	1.25
Oct. 30	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 31	Klate	0.38	Reference	1.0	1.05	1.25
Nov. 1	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 2	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 3	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 4	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 4	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 5 Nov. 6		0.20	Reference	1.0	1.05	1.23
	Dormant					
Nov. 7	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 8	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 9	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 10	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 11	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 12	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 13	Dormant	0.20	Reference	1.0	1.05	1.25

Nov. 14	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 15	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 16	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 17	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 18	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 19	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 20	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 21	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 22	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 23	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 24	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 25	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 26	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 27	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 28	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 29	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 30	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 1	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 2	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 2 Dec. 3	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 4	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 4 Dec. 5	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 6	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 0 Dec. 7		0.20	Reference	1.0	1.05	1.25
	Dormant					1.25
Dec. 8	Dormant	0.20	Reference	1.0	1.05	
Dec. 9	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 10	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 11	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 12	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 13	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 14	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 15	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 16	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 17	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 18	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 19	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 20	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 21	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 22	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 23	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 24	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 25	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 26	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 27	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 28	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 29	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 30	Dormant	0.20	Reference	1.0	1.05	1.25

^cA LAI of 0 was used from January 1 to March 14, 0.2 from March 15 to May 15, 1.05 from May 15 to May 31, and 2.65 for June, July and August. An LAI of 1.05 was used from Sept. 1 to Sept. 30, and 0.2 from Oct. 1 to Oct. 31; 0 was used for the remainder of the year. The crop height used was 1.5m.

^dBased on the information available regarding the evapotranspiration of deciduous and coniferous forest, and idle land, step-wise changes in values were used.

^eThe dates chosen for the change in values for open water >5m are an estimated time as to when the average daily temperature exceeds water temperature (May 1) and when the

average daily temperature is below water temperature (Oct. 1). These values should be adjusted should the conditions vary by the season or by region.

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Disclaimer: The evapotranspiration coefficients indicated in this report are estimated from published literature produced in various climatic regions. These regions do not necessarily reflect the situations found in the Niagara Peninsula. The reported values are based on "well-watered" soil conditions and dry plant canopies. Information on dormant season (winter) ET is very limited. Any errors in the published literature may be reflected in the values presented in this report. The evapotranspiration coefficients reported are the best estimates available, but they should be used with the full recognition of these limitations.