

Water Availability Study for the  
Lake Erie North Shore Watershed Plan Area  
Niagara Peninsula Source Protection Area

Prepared for  
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NIAGARA PENINSULA  
**CONSERVATION**  
A U T H O R I T Y

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## **1. INTRODUCTION**

### **1.1 Background and Objectives**

The Niagara Peninsula Conservation Authority (NPCA) and AquaResource Inc. have completed this Water Availability Study (WAS) of the Lake Erie North Shore (LENS) as part of the Niagara Peninsula Source Protection Area (NPSP Area) Source Water Protection (SWP) Tier 1 Water Budget.

The NPSP Area Tier 1 Water Budget is funded by the province of Ontario. The study methodology was developed by NPCA in consultation with the Ministry of Natural Resources (MNR), based upon the March 2007 Draft Guidance Module – Water Budget and Water Quantity Risk Assessment.

The purpose of the WAS was to determine the water available for surface water flow, groundwater recharge and evapotranspiration on a monthly basis for the time period 1991 to 2005. This time period was chosen to best suit available datasets (e.g. Statistics Canada) and meet the minimum World Meteorological Organization climate normal criterion of fifteen (15) years.

This report documents the WAS for LENS Watershed Plan Area (WSPA). The study area is located in Haldimand County and the Regional Municipality of Niagara including portions of the local municipalities, Township of Wainfleet, City of Port Colborne and Town of Fort Erie (Figure 1.1).

### **1.2 Study Team and Approach**

AquaResource Inc. was awarded the contract to complete the Water Availability Study and has previously worked with NPCA on the Conceptual Water Budget Report (Franz Environmental Inc. et al., 2007). In late 2007 they also completed a technical memo on NPCA Water Survey of Canada stations regarding Baseflow Separation and Streamflow Recession. AquaResource Inc. is involved in the Source Water Protection Water Budget process at a number of levels, including the development of the Water Budget Guidance Module, ongoing technical support for the Ministry of Natural Resources and completing Tier 1, 2 and 3 Water Budget projects for conservation authorities and municipalities.

Peer review of the WAS project was provided primarily by Robert Muir of Dillon Consulting Limited. Mr. Muir previously assisted NPCA in development of the Tier 1 Water Budget work program. He is a Water Resources Engineer with almost two decades of experience and has provided peer review for the Lake Simcoe Region Conservation Authority Water Budget as well as surface water vulnerability studies for a number of conservation authorities.

NPCA staff from three (3) departments were involved throughout the study. These included Jeff Lee and Geoff Verkade from the Geographic Information Systems (GIS) group, Guangli Zhang from Engineering and Jayme Campbell and Brian Wright from Source Water Protection.

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The project approach was designed to take advantage of NPCA's GIS expertise and datasets (e.g. soils, land use and digital elevation model) and NPCA's Engineering Department's experience with HEC-HMS. HEC-HMS is the U.S. Army Corps of Engineers (USACE) Hydrologic Engineering Center Hydrologic Modelling System. This is the current software package utilized by the NPCA Engineering Department for its in-house floodplain mapping. HEC-GeoHMS was used by NPCA GIS specialists throughout the project to develop the hydrologic modelling inputs for HEC-HMS.

This report describes the work completed as part of the WAS of Lake Erie North Shore.

### **1.3 Project Tasks**

The principal objectives of the Water Availability Study are to derive monthly estimates of (i) evapotranspiration, (ii) groundwater recharge and (iii) water available for surface water flow for the time period 1991 to 2005.

The project tasks are:

- Initial parameterization of the numerical models to simulate watershed conditions;
- Calibration of the models to observed surface water flow data (where available) with an emphasis on volumes as opposed to peak flow rates; and
- Continuous model HEC-HMS hourly simulation runs for 1991-2005.

Reporting was completed on the model development, calibration uncertainty and outputs and recommendations for future work.

### **1.4 Relevant Reference Documents**

A variety of previous studies provide details regarding the hydrologic conceptual model; these include, most notably:

- Water Budget Conceptual Understanding for the Niagara Peninsula Source Protection Area (Franz et al, 2007); and
- Baseflow Separation and Streamflow Recession (AquaResource Inc., 2007).

These studies are referenced throughout this report.

### **1.5 Document Organization**

The sections within the report are organized as follows:

- Chapter 2 – Watershed Characteristics;
- Chapter 3 – Watershed Modelling; and
- Chapter 4 – References.

## **2. WATERSHED CHARACTERISTICS**

### **2.1 General Description of the Watershed**

The Lake Erie North Shore Watershed Planning Area is 117.7 km<sup>2</sup> in size. It contains sixteen (16) catchments which flow independent of one another into Lake Erie. These catchments include (from west to east as shown on Figure 3.2): Low Banks Drain (LB), Hoover Drain (HOD), Lake Erie 8 (LE8), Lake Erie 6G (LE6G), Casey Drain (CD), Wainfleet Marsh Drain (WM), Lake Erie 5 (LE5), Eagle Marsh Drain (EMD), Lake Erie 4 (LE4), Wignell Drain (WD), Beaver Dam Drain (BDD), Oil Mill Creek (OMC), Bearss Drain (BD), Port Albino Drain (PAD), Lake Erie 2B (LE2B) and Bay Beach Area drain (BBA).

These catchments cover 92.9 km<sup>2</sup> (79%) of the LENS WSPA. The remaining 24.8 km<sup>2</sup> include:

- 13.8 km<sup>2</sup> (12%) which has no (digital elevation model) mapped watercourse, these are generally located between a catchment and the Lake Erie shore;
- 4.6 km<sup>2</sup> (3%) which drain to Mill Race Creek within Big Forks Creek WSPA, located between Hoover Drain and Lake Erie 8; and
- 7.4 km<sup>2</sup> (6%) which drains to the Welland Canal within the Central Welland River WSPA, located at Port Colborne between Eagle Marsh Drain and Lake Erie 4.

These “missing” areas will be added to the final stress assessment through a combination of pro-rating and results from other WAS results.

### **2.2 Climate Setting**

The climate of Southern Ontario is characterized as having warm summers, mild winters, a long growing season, and usually reliable rainfall. The climate within southern Ontario differs somewhat from one location to another, and from one year to the next. Spatial variations are caused by the topography and varying exposure to the prevailing winds in relation to the Great Lakes (Schroeter et al, 1998).

According to Brown et al. (1980), Lake Erie North Shore is located in the Niagara Fruit Belt climatic region. Using the stations shown on Figure 2.1, Figures 2.2 and 2.3 show the 1991-2005 mean monthly precipitation and mean monthly temperature (Schroeter and Associates, 2007). Mean monthly precipitation ranged from a low of 49 mm at the Port Colborne Environment Canada station 6136606 in February to a high of 110 mm at Fort Erie Environment Canada station 6132470 in September. Generally the amount of precipitation increased from west to east across the WSPA. The mean annual range in temperature was 26.0 degrees Celsius (°C) (Figure 2.3).

Spatial variations in mean annual snowfall, air temperature and mean annual precipitation across Lake Erie North Shore in relation to the entire NPCA jurisdiction are illustrated in Figures 2.4, 2.5 and 2.6 and tabulated in Table 2.1. Annual precipitation and snow in Lake Erie North Shore appear to range from 1050-940 mm per year and 190 to 105 mm,

respectively, on mean across the WSPA. Mean annual temperatures range from 8.7 to 9.3 °C.

Figures 2.7, 2.8 and 2.9 show the annual precipitation, annual snow water equivalent and mean annual temperature for the 1991-2005 period respectively for the Dunnville, Port Colborne and Fort Erie stations. The total annual precipitation ranged from a 1998 low of 662 mm (Dunnville) to a high of 1380 mm in 1996 (Fort Erie). On average the annual precipitation was 991 mm (1991-2005). The amount of snow water equivalent ranged from a low of 48 mm in 2001 (Dunnville) to a high in 1995 of 292 mm (Fort Erie). Overall 144 mm (15%) of precipitation is delivered as snowfall. The amount of snow generally increased from west to east across the WSPA. The mean annual temperature was lowest in 1996 at 7.7°C and highest in 1998 at 10.6°C.

Brown et al. (1980) previously estimated the regional mean annual actual evapotranspiration between 533-559 mm and mean annual water surplus as about 279 mm.

### **2.2.1 Net Solar Radiation**

Six (6) solar radiation and two (2) sunshine station locations were located in and near NPCA ranging from Buffalo, New York to the Hamilton Royal Botanical Gardens (RBG), shown in Figure 2.1. Annual values of net radiation ranged from 26.29 KW/m<sup>2</sup> at Niagara Falls, New York in 2004 to 33.89 KW/m<sup>2</sup> at Hamilton RBG in 1991 (Figure 2.10). Overall all stations had their lowest annual net radiation results in 2004, for the period 1991-2005. A review of the results however indicates sunshine station results tend to be slightly higher than those measuring incoming radiation directly. The greatest monthly variation between station measurements occurs during the summer period (Figure 2.11) and shows an increase in net solar radiation going to the northwest from Buffalo to Hamilton RBG in the July means.

## **2.3 Topography**

The topography of LENS WSPA varies from 199 metres above sea level (m ASL) to 165 m ASL. Overall elevation decreases from north to south with some exceptions such as Lowbanks from the influence of the topographic high of the Mohawk moraine and the western Point Albino beach deposits.

On average the slope of the catchments was 2.1%, ranging from 0.5 to 4.4% (Table 3.1). The channel profiles of the largest drained areas are shown on Figures 2.12a and 2.12b. The profiles indicate that in a number of cases the channels are nearly flat over 3 km from the Lake Erie shore (e.g. Low Banks Drain, Eagle Marsh Drain, Wignell Drain and Port Albino Drain).

## **2.4 Physiography**

Major LENS WSPA physiographic features include the Haldimand Clay Plain and the Onondaga Escarpment (Chapman and Putnam, 1984). The Haldimand Clay Plain is characterized by levelness, local relief less than a few metres and poor drainage. The

clay plain is partially separated from Lake Erie by the east-northeast to west-southwest trending Onondaga Escarpment.

The Onondaga Escarpment (Figure 2.15) is a north-facing, east-west trending bedrock cuesta with approximately 10 metres of relief. A cuesta is a hill or ridge with a gentle slope on one side and a steep slope on the other (AGI, 1974). The Onondaga Escarpment (limestone plain) forms the northern drainage boundary for a number of catchments in the central portion of the WSPA (e.g. Eagle Marsh Drain, Wignell Drain, Beaver Dam Drain and Bearss Drain).

The Dunnville Sand Plain is shown as within LENS WSPA however this is not correct as shown by the more detailed soil and surficial geologic mapping (Figures 2.14 and 2.15).

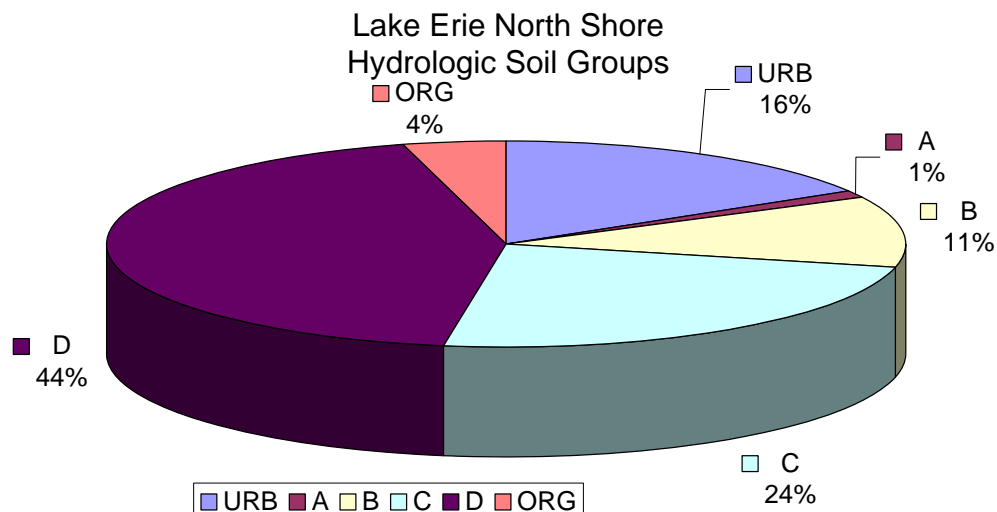
## 2.5 Soils

The mapped soils information was provided by the Ontario Ministry of Agriculture and Food and combines two (2) soil surveys, Haldimand County and Niagara Region (Figure 2.14).

The mapped soils (Figure 2.14) are classified into four hydrologic soil groups (A, B, C and D), organic or urban area. The grouping is according to the soil's minimum infiltration rate, obtained for bare soil after prolonged wetting. According to the United States Department of Agriculture the Hydrologic Soil Groups (HSG) are as follows:

- Group A: sand, loamy sand or sandy loam;
- Group B: silt loam or loam;
- Group C: sandy clay loam; and
- Group D: clay loam, silty clay loam, sandy clay, silty clay or clay.

The dominant soil groups are D and C. These amount to 44 and 24 percent of the area of LENS model respectively (as presented below). The remaining portion of the WSPA is mapped as 16% urban, 11% B, 4% organic and 1% A. The hydrologic soil group data inputs for the model are summarized on Table 2.2. Organic soils were assigned Group D HSG parameters. Urban soil polygons were not assigned HSG values.



## **2.6 Surficial geology**

The surficial geology of the LENS WSPA (Figure 2.15) is largely fine-textured glaciolacustrine deposits (e.g. clay and silt) or paleozoic bedrock at surface. In general the clay and silt is thin over bedrock highs and thickens considerably over topographically low areas (Jagger Hims Limited, 2005). The Paleozoic bedrock is primarily Onondaga and Bois Blanc formation limestone.

The Wainfleet and Crystal Beach moraines are the northern surface water divides for catchments CD\_W100 and BD\_W100/PAD\_W100/LE2B\_W100/BBA\_W100, respectively.

## **2.7 Land Cover**

Land cover information was provided by the Ministry of Natural Resources as part of its Southern Ontario Land Resource Information System (SOLRIS). Twenty-three (23) categories were provided as shown on Table 2.3 and simplified on Figure 2.16. The largest land use categories making up 89% of LENS were (i) monoculture 23%, (ii) rural land use 22%, (iii) swamp 16%, (iv) mixed agriculture 11%, (v) built-up impervious 6%, (vi) deciduous forest 6% and (vi) transportation 5%.

## **2.8 Streamflow**

There were no streamflow gauges located within the WSPA to provide measurements of surface water flow for this particular study.

The outflow from Eagle Marsh Drain and Wignell Drain can be reduced during elevated lake levels as control structures prevent backflow from Lake Erie. However this is not considered a concern for this study as results are assessed on a monthly basis and the outflows are pumped to Lake Erie when the control structures are in place.

### **3. WATERSHED MODELLING**

The following sections describe the construction and calibration of the Lake Erie North Shore (LENS) HEC-HMS model, and present the water balance estimates, as well as the Water Quantity Stress Assessment components.

#### **3.1 Model Description**

As outlined in the NPCA WAS Terms of Reference (TOR), HEC-HMS was chosen to model the hydrology of the fourteen (14) Watershed Protection Areas (WSPAs) within the NPCA official boundary. HEC-HMS is a numerical simulation model, supported by the U.S. Army Corps of Engineers, and is designed to simulate the precipitation-runoff processes of a watershed. The program is an integrated work environment, including a database management system, data entry utilities, a computation engine, results reporting tools, and a graphical user interface. A companion product, HEC-GeoHMS, is a software package for use with ArcView, and was employed to develop a significant portion of the required HEC-HMS inputs.

HEC-HMS can be run at a variety of time steps, from 1 minute to 1 day. For the LENS model, and other models created for this study, HEC-HMS was run on the hourly time step.

For complete documentation of the HEC-HMS program, as well as individual hydrologic processes included in HEC-HMS, please refer to the HEC-HMS User Manual and/or Technical Reference Manual (USACE, 2006, 2000).

##### **3.1.1 HEC-HMS Hydrologic Processes**

HEC-HMS includes a variety of algorithms for representing the dominant hydrologic processes. This allows the modelling approach to be tailored both to the available data and the overall goals of the study.

The modeller can specify the appropriate algorithm for the following processes:

- Evapotranspiration;
- Snowmelt;
- Loss (infiltration method);
- Baseflow Routing;
- Catchment Hydrograph Transform; and
- Channel Routing.

The algorithms used in the NPCA WAS (specified by the NPCA Water Availability Study Terms of Reference) for each of the six major hydrologic processes are described in the following sections. A conceptualization of the hydrologic processes simulated by HEC-HMS is included in Figure 3.1.



### **3.1.1.1 Evapotranspiration**

The Priestly-Taylor evapotranspiration routine was specified for use in this project. The Priestly-Taylor method relies upon solar radiation and temperature to generate estimates of potential evapotranspiration (PET).

The Priestley-Taylor equation is as follows:

$$PET = \alpha \frac{s(T_a)}{s(T_a) + \gamma} (K_n + L_n) \cdot \frac{1}{\rho_w \lambda_v}$$

Where;

$K_n$  = Short wave radiation

$L_n$  = Long wave radiation

$s(T_a)$  = Slope of the saturation-vapour pressure vs. temperature curve

$\alpha$  = Dryness coefficient

$\rho_w$  = Mass density of water

$\gamma$  = Psychrometric constant (ratio of the heat capacity of the air to the latent heat of vaporization)

$\lambda_v$  = Latent head of vaporization

Once the Priestley-Taylor PET estimate is generated, HEC-HMS applies crop coefficients to reflect cropping practices or vegetative cover. The crop coefficients are applied as multipliers to scale the Priestley-Taylor PET estimate for that time step.

Evapotranspiration rates are generated by applying the estimated potential evapotranspiration rates to the soil-water reservoir represented within HEC-HMS. Actual evapotranspiration is limited by the amount of water within the soil-water reservoir. When the soil-water reservoir is saturated, actual evapotranspiration is equal to potential evapotranspiration. When the soil-water reservoir is empty (water content is zero), evapotranspiration can no longer be supported bringing the actual evapotranspiration to zero. It remains at zero, until a precipitation event replenishes the soil-water reservoir.

### **3.1.1.2 Snowmelt**

The ability to simulate snow processes is critical to represent the hydrology of cold climate watersheds. The spring snowmelt period (March/April in Southern Ontario) is the season with the highest typical streamflow, and is also responsible for the majority of streamflow volume. This is also the period of time where saturated soil conditions are common producing groundwater recharge.

HEC-HMS considers snow processes by tracking changes to the snowpack. A snowpack is formed when precipitation occurs and the air temperature is below 0°C. HEC-HMS tracks the accumulation and melt of the snowpack through use of the Temperature Index Method. This method utilizes precipitation and temperature to simulate snow accumulation and melt processes. Water content of the snowpack can be increased by snow or rain falling on the snowpack.

Snowmelt is generated when temperatures rise to the point where there is sufficient energy to transform frozen water into liquid water. The amount of melt experienced by

the snowpack is dependent for each degree above the freezing point. Snowmelt is held within the snowpack until the snowpack's point of saturation is reached. When the snowpack becomes saturated (specified by the water capacity of the snowpack), liquid water is then provided to the soil surface as water available for infiltration or runoff.

Sublimation is the direct loss of water from the snowpack to the atmosphere. It is not represented within HEC-HMS. Over the winter season, sublimation can result in a significant loss of water content from the snowpack. Schroeter and Associates have estimated this loss to be 0.33 mm/day (Schroeter and Associates, 2004). This is considered a limitation of the HEC-HMS model, and may lead to an over-estimation of water content held within the snowpack.

For a detailed discussion on the snowmelt processes included in HEC-HMS, please refer to Appendix A for an unpublished presentation provided by the USACE.

### **3.1.1.3 Loss Method (Infiltration)**

The infiltration method, or as HEC-HMS terms it, the "loss method", is responsible for partitioning liquid precipitation into direct overland runoff, evapotranspiration, or percolation. The Deficit and Constant Loss method is utilized for this project, and is carried out on a catchment by catchment basis.

Liquid precipitation that falls as rainfall or snowmelt is input into a storage reservoir. This storage reservoir represents all storage elements within each catchment. This includes, but is not limited to, soil water storage, depression storage, and interception storage. The depth of water held within this element is specified by the user.

Water held within the storage reservoir can be removed by evaporation or by percolation. Evaporation, at the rate estimated by the Priestly-Taylor equation, can remove water held within the storage reservoir. If the storage reservoir is empty, actual evapotranspiration is zero for that time step. Water can also leave the reservoir via percolation, which is determined by the Constant Rate. Percolation can only occur when the storage reservoir is completely saturated, and ceases when the storage reservoir drops below the point of saturation. At this point, evapotranspiration is the sole process that is able to reduce the amount of water held in the storage reservoir. Direct overland runoff is only generated when the storage reservoir is full, and liquid precipitation falls at a rate faster than the Constant Rate.

A limitation of this method is the unlimited acceptance of precipitation into the storage element. Provided there is sufficient storage, the reservoir can accept all precipitation, and produce no runoff or recharge, regardless of the intensity of the event. This can result in an under-prediction of flow, particularly when the reservoir is near empty. The impact of this limitation would be most significant when comparing simulated and observed hydrographs for a particular event. Due to the modelling focus being on regional water budgeting, and not flood flow estimation, it is anticipated that this limitation will not be a major factor for the purposes of this exercise.

### **3.1.1.4 Baseflow Method**

Once the loss method generates estimates of percolation, this water is passed onto the Baseflow Method for a representation of the subsurface processes (see Figure 3.1). The Baseflow Method selected for this study is the Linear Reservoir Method.

Routing flows through a linear storage element is calculated by the following equations: (Schroeter and Watt, 1980)

$$Q_t = C \times Q_{t-1} + (1 - C) \times I_{t-1}$$
$$C = e^{\left(\frac{-dt}{KR}\right)}$$

Where:

$Q_{t-1}$ ,  $Q_t$  = Outflow

dt = time step

KR = recession constant (hr)

I = Inflow

The Linear Reservoir method uses two linear reservoirs to model the recession of baseflow after a precipitation event. The first linear reservoir is meant to represent a rapidly responding system, often termed “interflow”. Interflow is commonly understood to be subsurface stormflow moving through a shallow unsaturated soil horizon, towards a watercourse (Bedient and Huber, 2002).

The second linear reservoir is meant to represent a slower responding groundwater system, in comparison to the first reservoir. This is the system most commonly associated with baseflow and groundwater recharge.

Previous interpretations of the hydrologic/hydrogeologic system within the NPCA, carried out as part of the Conceptual Water Budget, have indicated that there is very little evidence of a regional groundwater flow system with strong interactions with the surface water system (Franz et al., 2007). The Conceptual Water Budget also stated there was minimal recharge to a deeper regional groundwater system, and that any groundwater discharge that did occur was “fed by localized groundwater recharge, which does not enter the regional aquifer system”. This localized groundwater discharge was termed, perhaps mistakenly, “interflow”. The term interflow, as it was used in the Conceptual Water Budget, meant to indicate discharge that was not sourced from a larger regional system, but rather from localized, near surface, aquifers. It was not meant to describe the shallow stormflow as described by Bedient and Huber (2002).

For the purposes of this study, flow from the first linear reservoir (interflow) will be considered to be part of the storm response, which travels laterally through the unsaturated soil horizon, before discharging into a watercourse. Flow that enters the second reservoir, and is discharged as baseflow, will represent the amount of water that percolates and reaches the saturated soil layer as groundwater recharge.

The percolation computed from the Deficit and Constant Loss method, is split evenly between both reservoirs. The proportion of water supplied to each reservoir, is specified by the program itself, and can not be modified.

Discharges from both of the linear reservoirs are added with any direct runoff, which create the catchment outflow hydrograph. As this method conserves mass within the catchment, there is no ability to route a portion of baseflow to a downstream catchment, or to remove water from the entirety of the watershed representing “deep recharge”. There are two sources of error associated with this limitation. The first source of error is that outflows of groundwater to downstream catchments cannot be represented. This may result in headwater catchments having too much groundwater discharge, with downstream catchments having too little. The error associated with this limitation is inversely proportional to the watershed area. This is due to net groundwater inflows/outflows becoming negligible as the area of interest increases.

The second source of error is that the loss of water to regional groundwater flow systems (removal of water from the watershed) is not able to be represented. By neglecting this loss, other water balance parameters could be over-estimated (ET, runoff, baseflow). Due to the conceptualization reported in the Conceptual Water Budget (Franz et al., 2007) of minimal interaction between the regional groundwater system and the surface water system, it is anticipated this will not be a significant source of error.

#### **3.1.1.5 Catchment Hydrograph Transform**

Whereas the Baseflow Method is responsible for the routing of percolated water, the Transform Method is responsible for the routing of overland runoff. For this study, the SCS Unit Hydrograph Method was specified for the transform method.

The SCS Unit Hydrograph Method was originally developed from observed data collected in small, agricultural watersheds. The observed data has been generalized as dimensionless hydrographs, and a best-approximate hydrograph was developed for general application. The SCS method scales the generalized hydrograph by a user specified time lag to produce the unit hydrograph. The time lag is approximated by taking 60% of the time of concentration.

#### **3.1.1.6 Channel Routing**

As catchment outflow hydrographs are generated and added to the main channel, the resulting hydrographs must be routed downstream. While HEC-HMS has a number of methods available for routing, the Muskingum-Cunge method has been specified for this study. The Muskingum-Cunge routing method is based on the combination of the conservation of mass and the diffusion representation of the conservation of momentum. It represents the attenuation of flood waves and can be used in river reaches with a small slope.

The attenuation of hydrographs is calculated by specifying the characteristics of the channel. These characteristics include length, slope, Manning’s n, and channel geometry (cross section).

## **3.2 Model Set-up**

HEC-HMS requires a number of datasets to represent the hydrology of a watershed. A large portion of the model set-up was completed by NPCA, utilizing both GIS and HEC-GeoHMS processing. The following sections summarize the methodology for the initial parameterization of the HEC-HMS model. For the detailed description on the implementation of HEC-GeoHMS, please refer to Appendix B.

### **3.2.1 Meteorological information**

To properly represent streamflow and significant hydrologic processes, climate and climate variability must be represented within the hydrologic model.

Climate data from three meteorological stations were considered when constructing the LENS HEC-HMS model (refer to Figure 2.1). All of the stations: Port Colborne, Dunnville Pumping Station and Fort Erie, are owned and operated by Environment Canada. As such, they are operated to a national standard, and undergo significant quality assurance/quality control procedures to ensure accurate data collection.

The majority of catchments in the LENS model were assigned climate data from the Port Colborne station. Catchments in the far western portion of the model were assigned climate data from the Dunnville station and catchments in the far eastern portion were assigned climate data from the Fort Erie station.

To ensure each climate station had a complete period of record, each dataset was cleaned up and filled-in by Schroeter and Associates (2007). The in-fill procedure was carried out on both the daily datasets (max/min temperatures, rainfall/snowfall totals), and the hourly rainfall datasets.

With the hourly modelling time interval, hourly data inputs were required. To produce hourly precipitation, daily snowfall depths were evenly distributed throughout the day and added to the hourly rainfall dataset. While it is unlikely that the reported daily snowfall depths are evenly distributed throughout a day, the fact that snowfall does not generate an immediate streamflow response, means the impact of this assumption is negligible.

Synthetic hourly temperatures were generated using the maximum and minimum daily temperatures and a generalized synoptic curve (Schroeter and Associates, 2004).

An hourly net solar radiation station was created for the LENS model using data from the Northeast Regional Climate Centre station at Buffalo, New York. The overall hourly net radiation was calculated using the methodology of Allen et al. (2005).

### **3.2.2 Streamflow Information**

The LENS WSPA consists of small creeks, and gullies, which drain directly to Lake Erie, without forming a typical dendritic drainage system. As described in Section 2.8, there are no active stream gauges in the LENS WSPA.

### **3.2.3 Catchment boundaries and characteristics**

General catchment parameters and specifically parameters for the transform and loss methods are shown on Table 3.1. Catchments were delineated by NPCA GIS specialists and AquaResource Inc., using the NPCA 2 m DEM. The catchments ranged in size from 0.03 to 9.5 km<sup>2</sup>. In other WAS (e.g. Twenty Mile Creek) smaller catchments were not possible without the reduction of the model time steps to less than one hour. This constraint is a modelling limitation related to the size of catchment and the model time step within the Transform algorithm (see Section 3.2.8). However, due to the nature of the multiple catchments routing to Lake Erie, no surface water gauges available for calibration and results being reported on a monthly basis, this was not a constraint for LENS. The model schematics and catchment boundaries are included in Figure 3.2.

As is shown in Figure 3.2 (and discussed in Section 2.1), there are areas of the LENS WSPA which were not included within the LENS model. These are very small areas that have no mapped watercourse, and are therefore unable to be included in the HEC-GeoHMS processing (unless already part of the WAS for Central Welland River and Big Forks Creek). To ensure volumes utilized for the Stress Assessment consider the full area of the LENS WSPA, modelled results will be area pro-rated upwards to include un-modelled areas. Given the small proportion (12%) of the LENS WSPA that is un-modelled, it is anticipated that the error introduced by this will be minimal.

### **3.2.4 Initial Parameterization – Loss Method**

The Loss Method relies on three parameters to determine the amount of water that infiltrates, or is available to become overland runoff. These parameters are the constant rate, the catchment storage capacity, and the percentage of impervious cover.

The Deficit and Constant Loss Method assumes that the soil has a constant infiltration rate approximated by the saturated soil hydraulic conductivity. Using the soil and water holding capacity information in Appendix C, average maximum infiltration rates were assigned to each polygon in the soil layer based on their soil type. The catchment average constant rate was determined by area weighting each of the soil polygons in the specific catchment.

HEC-HMS assumes that the soil has a fixed water holding capacity, based on the active rooting depth of vegetation and soil type. The soil water holding capacity layer was built by intersecting the SOLRIS land cover and the OMAF soils layers and by assigning soil water holding capacity values from Appendix C to each unique combination of land cover class and soil type. Like the methodology employed for the constant rate, the area weighted average for each polygon within the catchment, was used to calculate the catchment average.

HEC-HMS considers an impervious surface as an area in a watershed for which all contributing precipitation runs off, with no infiltration, no evaporation, and no other volume losses. This surface was built by assuming SOLRIS built-up impervious and transportation polygons were 100% impervious, with built-up pervious polygons being 50% impervious. All other polygons were assumed to have zero impervious cover.

### **3.2.5 Initial Parameterization – Evapotranspiration**

In the Deficit and Constant Loss Method, water is removed from the soil to simulate evapotranspiration. Potential evapotranspiration is calculated through use of the Priestly-Taylor method. This method uses a crop coefficient,  $K_c$ , indicating the ratio of crop potential and short grass reference evapotranspiration. Daily land use layers were created and assigned crop coefficients from Appendix D.

Combined with the solar radiation and temperature datasets, outlined in Section 3.2.1, and the crop coefficients, HEC-HMS calculates the potential evapotranspiration for each time step. This potential evapotranspiration value is then applied to the catchment storage reservoir to generate actual evapotranspiration.

### **3.2.6 Initial Parameterization – Snowmelt**

The following parameters are required to represent snowmelt. These generalized parameters are referenced from the HEC-HMS User Manual.

- Temperature at which precipitation falls as snow;
- Temperature at which the snowpack begins to melt;
- Water capacity of the snowpack;
- Amount of melt that occurs due to heat transfer from the underlying ground;
- Rate at which snow melts when rain occurs;
- Rate at which snow melts during rainfall-free periods.

### **3.2.7 Initial Parameterization – Baseflow**

Once water percolates through the soil column, HEC-HMS routes this water back to the stream as interflow or baseflow. The Linear Reservoir Method, specified for use by the WAS TOR, approximates the discharge by use of a linear reservoir. Groundwater recession constants, estimated via streamflow analysis, represent the reservoir response time and are used as the reservoir constant (also called the time constant) for the linear reservoir in each layer. There are two linear reservoirs that can be represented within HEC-HMS.

The first linear reservoir was parameterized with the intent to represent interflow. A groundwater coefficient of 18 hours was assigned to this reservoir for all catchments, as this was used in the final calibration parameters for Twenty Mile Creek and Upper Welland River.

The 2<sup>nd</sup> linear reservoir, meant to represent groundwater discharge to the watercourse, was parameterized based on streamflow recession analysis completed by AquaResource Inc. (2007). The streamflow recession analysis estimated the reservoir constant for nine streamflow gauges located within the NPCA. Due to the lack of a streamflow gauge within the WSPA, statistics from the, Oswego Creek at Canboro gauge, were used within

the LENS model. The median reservoir constant from the 1991-2005 period was assigned to the 2<sup>nd</sup> linear reservoir for each catchment in the model.

### **3.2.8 Initial Parameterization – Transform**

The lag time associated with the SCS transform method is a function of the Soil Conservation Service (SCS) Curve Number (Figure 3.3), the hydraulic length, and the catchment slope. This time lag is used to produce the unit hydrograph that allows precipitation excess (precipitation-infiltration) to be transformed into an overland runoff hydrograph. For adequate definition of the unit hydrograph ordinates, a modelling time step that is less than 29% of the time lag must be used. This constraint effectively places a minimum size requirement on the catchments represented within the model.

Curve Number (CN) values are used in the calculation of CN lag time for the SCS Unit Transform Method. The factors influencing CN values are land cover type, soil type and Antecedent Soil Moisture Condition (AMC). AMC is an estimate of soil water content prior to the beginning of the simulation period, and has 3 levels:

- AMC I reflects soils that are dry but with water content not below the wilting point.
- AMC II reflects soils having average soil water content, and
- AMC III reflects soils that have experienced rainfall in the five days previous to the simulation period.

CN values in the study area were assumed to reflect average soil water content (AMC II). The CN layer was built by intersecting the SOLRIS land cover and OMAF soil layer and by assigning CN values from Appendix C to each unique combination of land use class and soil type. Built-up impervious, built-up pervious and transportation SOLRIS polygons were considered under the impervious surface data field and not assigned CN values.

### **3.2.9 Initial Parameterization – Routing**

To simulate the effects of channel geometry on hydrograph shape, the traditional Muskingum-Cunge Routing Method was used assuming trapezoidal channel geometry. The following inputs are required:

- Channel Bottom Width. The channel width for each of the routing reaches was estimated by digitizing cross sections. This channel width estimation assumed that the water surface width on digital air photos approximated the width of the channel bed.
- Channel Side Slope. The channel side slope was approximated by digitizing two points at the end of each digitized channel width cross sections using a 2m resolution DEM as a guide. Slope values were extracted at the location where the points intersected a slope grid.



- Channel Manning's Roughness Coefficient. Appropriate Manning's roughness coefficients were assigned (Appendix C) to channel routing reaches based on a visual stream bed condition assessment of 10-20cm resolution digital air photos.

### **3.3 Model Calibration/Verification**

#### **3.3.1 Overview of Procedures**

Due to the lack of observed streamflow data for the LENS WSPA, the calibration and verification process undertaken for the Upper Welland River and Twenty Mile Creek models could not be completed. Instead, the parameters within the LENS model were modified using the parameter adjustments carried out for the Upper Welland River and Twenty Mile Creek models. This methodology assumes that adjustments to model parameters, which result in an acceptable calibration, are transferable between WSPAs. This assumption is validated by the fact that the adjustments required for both Upper Welland and Twenty Mile Creek were extremely similar, which suggests that these adjustments were regional in nature. Given the geologic homogeneity of the NPCA, this is to be expected.

To reduce the reliance on the user specified initial conditions, and to allow the model to "self-initialize", the modelling period was extended to 1990. No data from the 1990 extension was included in the final results.

#### **3.3.2 Model Parameters**

As described in the Twenty Mile and Upper Welland modelling reports, parameters related to the Constant Rate and Maximum Storage values in the Deficit and Constant Loss method were adjusted. These values affected the amount of overland runoff, baseflow and interflow, and the amount of evapotranspiration. The modelled values of the Maximum Storage and the Constant Rate for each catchment are shown in Table 3.2. These values provided the base case for the sensitivity analysis which is documented in Section 3.4.

**Table 3.2 - Calibrated Constant Rate and Maximum Storage Terms**

<b>Catchment LENS ID</b>	<b>Constant Rate (mm/h)</b>	<b>Max Storage (mm)</b>
BBA_W100	0.0363	125
BD_W100	0.2008	146
BD_W200	0.1962	149
BDD_W100	0.1568	154
BDD_W200	0.1795	160
CD_W100	0.1425	149
EMD_W100	0.1250	125
EMD_W200	0.2566	125
HOD_W100	0.2520	167
LB_W100	0.1421	150
LB_W200	0.1680	165
LE2B_W100	0.3838	181
LE4_W100	0.0205	125

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<b>Catchment LENS ID</b>	<b>Constant Rate (mm/h)</b>	<b>Max Storage (mm)</b>
LE5_W100	0.1572	125
LE6G_W100	0.1076	132
LE8_W100	0.1002	169
OMC_W100	0.2321	128
PAD_W100	0.3598	142
WD_W100	0.1887	125
WD_W200	0.2823	137
WM_W100	0.3161	186

As per previous models, crop coefficients were modified to adjust the intensity of evapotranspiration applied to the storage element within the Deficit and Constant Loss Method. These modifications are a means to adjust evapotranspiration to account for issues with temperature data, solar radiation, the potential evapotranspiration method itself, or the lack of a sublimation process. They are not due to the crop coefficients being non-representative of their respective catchments. All month's crop coefficients were increased, with the exception of October, which was slightly lowered from the original estimate. As the crop coefficients are direct multipliers to the potential evapotranspiration estimated by Priestley-Taylor Evapotranspiration Method, increasing the crop coefficients resulted in an increase in evapotranspiration, with a corresponding decrease in streamflow (runoff, interflow and baseflow). Table 3.3 displays the final adjustments applied to the original crop coefficients.

**Table 3.3 – Monthly Crop Coefficient Adjustments**

<b>Month</b>	<b>Crop Coefficient Adjustment</b>
January	12.42
February	7.45
March	2.48
April	1.38
May	1.38
June	1.21
July	1.38
August	1.38
September	1.38
October	0.92
November	4.97
December	6.21

The groundwater coefficients with the Linear Reservoir Baseflow Method were not adjusted, but are presented in Table 3.4 below. While these are simply routing parameters, and are not used in partitioning precipitation, they are important to properly represent how infiltrated water is returning to the watercourse.

**Table 3.4 - Groundwater Coefficients in Linear Reservoir Baseflow Model**

<b>Catchment ID</b>	<b>GW 1 Coefficient (hr)</b>	<b>GW 2 Coefficient (hr)</b>
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All Catchments	18	326
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### **3.4 Model Sensitivity**

A sensitivity analysis was carried out on both models to determine the model sensitivity to variations in hydrologic parameters.

Previous HEC-HMS studies have shown that simulated streamflow is most sensitive to two parameters. They are: (1) the maximum infiltration rate (equivalent to the Constant Loss in the Deficit and Constant Loss Method), and (2) the water content available for evapotranspiration (Deficit term in the Deficit and Constant Loss Method) (Fleming and Neary, 2004).

Based on this finding from Fleming and Neary, four scenarios were tested to judge the sensitivity of model outputs to variations in the Constant Rate and Maximum Storage terms, included in the Deficit and Constant Loss Method. It is recognized that many other parameters and inputs can have an impact on simulated streamflow (snowmelt parameters, temperature, crop coefficients, precipitation, baseflow recession constants, etc.). However, due to the constraints in the scope of this project, only a limited sensitivity analysis was possible.

Both the Constant Rate and the Maximum Storage were varied by  $\pm 25\%$  independently, resulting in the four scenarios. Changes in total outflow, mean evapotranspiration, runoff and recharge were calculated and tabulated in the following tables:

- Tables 3.5 lists the percent change in total outflow for each scenario, over the base case.
- Tables 3.6 displays the percent change in total outflow, evapotranspiration, runoff and recharge for each scenario, over the base case.

**Table 3.5 - Sensitivity Analysis Results – Change in Outflow**

Month	Constant Rate +25%	Constant Rate -25%	Max Storage +25%	Max Storage -25%
Jan	0.7%	-0.8%	-2.8%	1.6%
Feb	0.1%	-0.1%	-3.5%	2.6%
Mar	0.0%	0.0%	-0.2%	0.2%
Apr	1.2%	-1.5%	-0.6%	0.1%
May	1.0%	-1.1%	0.0%	0.0%
Jun	0.9%	-1.0%	0.0%	0.0%
Jul	0.6%	-0.6%	0.0%	1.8%
Aug	0.1%	-0.2%	0.0%	1.8%
Sep	-0.1%	0.1%	-13.4%	35.4%
Oct	-0.3%	0.4%	-27.7%	49.5%
Nov	-0.5%	0.6%	-27.4%	24.2%
Dec	0.2%	-0.2%	-11.5%	6.9%

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**Table 3.6- Sensitivity Analysis Results – Change in Water Balance Estimates**

Scenario	ET	Baseflow	Interflow	Runoff
1: Constant Rate +25%	0.0%	15.9%	15.9%	-5.1%
2: Constant Rate -25%	-0.2%	-18.7%	-18.7%	6.5%
3: Max Storage +25%	5.5%	-8.2%	-8.2%	-6.0%
4: Max Storage -25%	-6.2%	6.4%	6.4%	8.0%

As is shown by Tables 3.5 and 3.6, variations in the Constant Rate did not significantly affect overall streamflow volume, but did cause significant changes in water balance estimates. Given that the Constant Rate controls the drainage of the storage reservoir (when fully saturated) to the groundwater reservoirs, increasing the Constant Rate results in an increase in both baseflow and interflow, with a corresponding decrease in runoff. Alternatively, decreasing the Constant Rate, has the effect of increasing runoff, and decreasing baseflow and interflow. Due to the Constant Rate not impacting the amount of water that can be held in storage, evapotranspiration is not affected. Figure 3.4 illustrates the percent change in the mean monthly outflow of the model with a 25% increase and decrease in the Constant Rate. The dotted line at  $\pm 10\%$  represents the uncertainty associated with streamflow estimates (Winter, 1981). As shown in the figure, the percent change for both variations in the Constant Rate, is well within these boundaries, which suggests that estimated streamflow is insensitive to changes in the Constant Rate.

Since percolation and runoff only occur when the storage reservoir is full (i.e. when the soil is saturated), increasing the Maximum Storage results in decreases in baseflow, interflow, and runoff. Actual evapotranspiration increases due to a higher volume of water being held in the storage element. A decrease in the Maximum Storage has the reverse effect: increasing baseflow, interflow and runoff and decreasing evapotranspiration, as less water is required to reach the storage reservoir's point of saturation (refer to Table 3.6). As illustrated in Figure 3.5 and Table 3.5, the model outflow is highly sensitive to variations in Maximum Storage in the fall and early winter months, but insensitive to these variations in the spring and summer. This is due to the storage element either being completely empty (summer) or completely full (spring) during these seasons, regardless of the size of the storage element. Very large variations in Maximum Storage would be required to change streamflow during these seasons. Flows during the fall season do exhibit sensitivity to variations in the Maximum Storage term. This is due to the storage reservoir being filled during these months. A smaller storage reservoir would cause the storage reservoir to be filled quicker, resulting in more volume directed to percolation and direct overland runoff. Increases in the storage reservoir will yield the opposite effect: lowered flows, runoff, and percolation.

It is important to note that variations in the Constant Rate and Maximum Storage term, do not impact flows during the summer months. This suggests that uncertainty with these terms will not likely add significant levels of uncertainty to the Tier 1 Surface Water Stress Assessment.

The results of the sensitivity analysis suggest that the model solution for LENS is non-unique, particularly with respect to the Constant Rate. In a non-unique solution, it is

possible to calibrate the model to streamflow volumes and obtain a good fit with a number of differing sets of parameters. Frequently with non-unique solutions it is likely that compensating errors are present; whereby the model is simulating the correct streamflow, but incorrectly replicating the underlying physical processes.

In the case of LENS, the Constant Rate can vary by as much as 25%, with a negligible change in streamflow volume. While streamflow is not sensitive to the Constant Rate variation, there is a significant impact on the water balance parameters estimated by the model (+20% baseflow). Water balance estimates (runoff, baseflow) therefore have a greater degree of uncertainty than the streamflow estimates.

To reduce the level of uncertainty, it is recommended that a more detailed Loss Method, such as the Soil Moisture Accounting Method, be tested on a WSPA to validate the water balance estimates made via the Deficit and Constant Loss Method. Such a test would preferably be carried out within a WSPA with observed streamflow data. The modular approach of HEC-HMS would easily facilitate replacing the Deficit and Constant Loss method, with the Soil Moisture Accounting method. Should the more detailed Soil Moisture Accounting Method generate water balance estimates similar to the Deficit and Constant Loss, a higher level of certainty could be attached to estimates generated for other WSPAs (including LENS). Additionally, the Soil Moisture Accounting Loss method allows the modeller to account for the proportion of percolated water that is lost from the surface water system as “deep recharge”, a key limitation of the Deficit and Constant Loss method identified in Section 3.1.1.4.

## ***3.5 Results and Discussion***

### ***3.5.1 Water Balance Results***

HEC-HMS outputs a number of water balance parameters at the catchment level. These include, but are not limited to: total flow, runoff, percolated water, evapotranspiration, snow water equivalent, and hydrographs at catchment or reach junctions. These values are output to a HEC DSS file at an hourly time step.

Output from HEC-HMS is summarized in Table 3.7, presenting the mean annual water balance on a catchment basis and an overall WSPA basis for the 1991-2005 time period. The water balance terms are defined below:

- Precipitation – Climate data used to represent the precipitation over each of the catchments is summarized by HEC-HMS and is presented here.
- ET – Estimated actual evapotranspiration.
- Interflow – Outflow from 1<sup>st</sup> linear reservoir (half of percolated water); percolated water which moves laterally through the unsaturated soil horizon.
- Baseflow – Outflow from 2<sup>nd</sup> linear reservoir (half of percolated water); slow responding groundwater system. Consists of water which reaches the saturated soil zone.
- Overland Runoff – Depth of water that does not infiltrate, and reaches the surface water system via overland runoff.

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- Total Outflow – Total annual outflow from the catchment; is the sum of Baseflow, Interflow and Runoff.

**Table 3.7 - Summary of Water Balance Model Results**

<b>Catchment ID</b>	<b>Precipitation (mm)</b>	<b>ET (mm)</b>	<b>Interflow (mm)</b>	<b>Baseflow (mm)</b>	<b>Runoff (mm)</b>	<b>Outflow (mm)</b>
LENS_BBA_W100	1053	115	15	15	909	938
LENS_BD_W100	968	539	55	55	317	428
LENS_BD_W200	968	571	55	55	286	396
LENS_BDD_W100	968	581	47	47	293	386
LENS_BDD_W200	968	576	52	52	287	391
LENS_CD_W100	968	555	46	46	320	412
LENS_EMD_W100	968	309	37	37	585	659
LENS_EMD_W200	968	462	79	79	348	506
LENS_HOD_W100	939	531	60	60	289	409
LENS_LB_W100	939	593	35	35	277	346
LENS_LB_W200	939	602	37	37	263	337
LENS_LE2B_W100	1053	563	104	104	283	490
LENS_LE4_W100	968	359	8	8	592	608
LENS_LE5_W100	968	506	42	42	376	461
LENS_LE6G_W100	968	497	37	37	397	470
LENS_LE8_W100	968	561	32	32	342	407
LENS_OMC_W100	968	501	66	66	335	468
LENS_PAD_W100	968	538	80	80	269	429
LENS_WD_W100	968	537	58	58	315	431
LENS_WD_W200	968	554	74	74	266	414
LENS_WM_W100	968	619	71	71	205	348
BFC_MRC_W210	968	578	38	38	313	390
CWR_WCS_W300	968	410	17	17	523	557
<b>Overall WSPA</b>	<b>967</b>	<b>520</b>	<b>51</b>	<b>51</b>	<b>343</b>	<b>446</b>

The estimated values for baseflow and interflow are very similar for most of the catchments. The standard deviation for the range of these estimates is 22mm, which is equal to 2% of average annual precipitation, and is to be expected due to the homogeneity of geologic conditions found within LENS. The standard deviations for the range of direct overland runoff and evapotranspiration estimates are much higher, 153 mm and 113 mm, respectively. This level of variability is almost entirely related to the high impervious percentages in certain catchments.

The largest estimate for baseflow in LENS is in catchment LE2B\_W100. This baseflow estimate is double the WSPA average, and is attributable to near-shore perviousness, as well as higher rates of precipitation related to the Fort Erie climate station. Catchments with very low baseflow estimates (BBA\_W100 and LE4\_W100) have significant impervious areas, 49% and 28% respectively. Correspondingly, these catchments have high runoff rates associated with impervious areas. Evapotranspiration for these catchments is lower than the mean due to a smaller proportion of the catchment being pervious and able to support evapotranspiration.

### 3.5.2 Stress Assessment

As discussed in Section 1.0, the primary objective of this modelling is to determine water supply and reserve flows for use in the Tier 1 Water Quantity Stress Assessment. The Stress Assessment will be completed both for groundwater and surface water systems, and identifies those subwatersheds where there may be a potential for water taking related stress.

Following the methodology in Guidance Module 7, a subwatershed's stress is estimated by comparing the amount of water consumed with the amount of available water. This comparison is made by calculating "Percent Water Demand" as follows, with the terms defined in Table 3.8:

$$\% \text{ Water Demand} = \frac{Q_{DEMAND}}{Q_{SUPPLY} - Q_{RESERVE}} \times 100\%$$

**Table 3.8 – Percent Water Demand Components**

Term	Definition	Calculation
$Q_{DEMAND}$	Consumptive Demand	Mean annual or monthly consumptive demand is calculated as the estimated rate of locally consumptive takings. Water demands are grouped into surface and groundwater takings.  Estimates of consumptive demand will be made from PTTW analysis, agricultural water use coefficients and private well usage. This portion of the Stress Assessment is outside the scope of the Water Availability Study, and will be completed by NPCA staff.
$Q_{SUPPLY}$	Water Supply	For surface water, the supply is calculated as the monthly median outflow for the area to be assessed.  Groundwater supply is calculated as the estimated annual recharge rate plus the estimated groundwater inflow into a subwatershed.
$Q_{RESERVE}$	Water Reserve	Water Reserve is a specified amount of water that is not considered as part of the available water supply.  For surface water supplies, water reserve is estimated using the 90 <sup>th</sup> percentile monthly outflow, at a minimum. The 90 <sup>th</sup> percentile flow is defined as the flow that is equaled or exceeded 90% of the time.  Groundwater reserve is calculated as 10% of the total estimated groundwater discharge within a subwatershed.

It is noted that baseflow is considered in both the surface water supply (baseflow within the outflow hydrograph) and groundwater supply (recharge, which sustains baseflow) terms of the Water Quantity Stress Assessment. While this may seem to "double count" baseflow, one should keep in mind the original purpose of the Stress Assessment, which is only to identify areas that have a high proportion of consumptive water taking, in comparison to the water flowing through the system. Identified areas, particularly at the Tier 1 scale, may not necessarily be experiencing hydrologic or ecologic stress, but rather are identified as requiring additional study to better understand the cumulative impacts of water use. The Stress Assessment methodology should not be utilized as a design/allocation tool, in an attempt to determine the total amount of water available to



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be withdrawn within a subwatershed, as double counting of the baseflow term would then be a consideration.

For surface water systems, the Percent Water Demand equation is based on a mean monthly basis. The maximum percent water demand for all months is then used to estimate the Potential for Surface Water Stress as shown on Table 3.9.

**Table 3.9– Potential for Surface Water Stress Thresholds**

Surface Water Potential Stress Level Assignment	Maximum Monthly % Water Demand
Significant	> 50%
Moderate	20% - 50%
Low	<20 %

For groundwater systems, the stress assessment calculation is based on mean annual demand conditions, as well as for monthly maximum demand conditions. The stress level for groundwater systems is calculated according to the thresholds shown on Table 3.10.

**Table 3.10 – Potential for Groundwater Stress Thresholds**

Groundwater Potential Stress Level Assignment	Mean Annual	Monthly Maximum
Significant	> 25%	> 50%
Moderate	> 10%	> 25%
Low	0 – 10%	0 – 25%

### 3.5.2.1 Surface Water Supply Components

The sum of outflows for all modelled catchments within the LENS WSPA were used to calculate the monthly median and 90<sup>th</sup> percentile flows. This includes all modelled catchments in the LENS model, as well as the portions of the previously modelled catchments CWR\_WCS\_W300 and BFC\_MRC\_W210 that are located in the LENS WSPA. A total of 89% of the LENS WSPA area is included in these modelled catchments. The flows from the modelled catchments were then scaled upwards to include areas within LENS that were not modelled (described in Section 3.2.3). Table 3.11 includes the median and 90<sup>th</sup> percentile flows for the entirety of the LENS WSPA, (modelled and unmodelled areas). These flow estimates include the direct overland runoff calculated from the upstream drainage area, and the interflow and baseflow components.

**Table 3.11 – Surface Water Percent Water Demand Components**

LAKE ERIE NORTH SHORE WSPA		
Month	Water Supply (Median Flow) (m <sup>3</sup> /s)	Water Reserve (90 <sup>th</sup> % Flow) (m <sup>3</sup> /s)
Jan	0.69	0.27
Feb	0.87	0.28
Mar	1.35	0.43
Apr	0.71	0.21
May	0.26	0.05
Jun	0.09	0.01
Jul	0.04	0.00
Aug	0.04	0.00



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<b>Sep</b>	0.03	0.00
<b>Oct</b>	0.11	0.00
<b>Nov</b>	0.56	0.05
<b>Dec</b>	0.82	0.22

There is a greater amount of uncertainty with respect to the 90<sup>th</sup> percentile flows than with the median flows. The 90<sup>th</sup> percentile flow, being observed at the extreme low end of flows, may be affected by processes not considered by HEC-HMS. These processes may include, but are not limited to: water takings, evaporation from the stream channel, online ponds, and regional groundwater discharge. Due to the magnitude of these processes not being well known, the net effect of these processes (additional or less flow) is not able to be determined, but does introduce a level of uncertainty into the 90<sup>th</sup> percentile flows.

### **3.5.2.2 Groundwater Supply Components**

The determination of the groundwater supply term is slightly more complex, due to HEC-HMS producing estimates of both interflow and baseflow. As described in Section 3.1.1.4, interflow is the portion of stormflow that moves through a shallow, unsaturated soil horizon towards a watercourse. Based on this description, the portion of percolated water that is directed into the interflow array will not be considered as part of the available groundwater supply.

The portion of percolated water that is directed to the baseflow array within HEC-HMS, and is meant to represent a slower, deeper groundwater system (only relative to the interflow component), will be used to infer groundwater recharge. It is recognized that within the Haldimand Clay Plain, there is very little evidence of a regional aquifer that has strong interconnections with the surface water system (Franz et al., 2007).

Groundwater recharge estimates, inferred from HEC-HMS baseflow estimates, should not be considered recharge to deep, confined aquifers below the Haldimand Clay Plain, but rather recharge to shallow and localized aquifers near the surface. At the scale of a Tier 1 Water Quantity Stress Assessment, no distinction is made for recharge that supplies a specific aquifer unit; rather the stress assessment is carried out on the groundwater system as a whole. This may result in percent water demand being underestimated for a confined water source whose primary source of water is lateral groundwater inflow.

It is recognized that there is uncertainty associated with HEC-HMS's arbitrary proportioning of percolated water to half baseflow and half interflow. Actually this division would shift from year to year, and season to season, with possibly some periods experiencing all percolated water returning to the watercourse as either interflow or baseflow. Determining the exact proportion of percolated water that reaches the uppermost water table (groundwater recharge) is not an obtainable goal for the scope of this project. By considering half of percolated water that is directed to the baseflow array as available for groundwater taking, the Stress Assessment will be conservative in nature.

Included in Table 3.12 is the estimated 1991-2005 annual mean groundwater recharge rate for the LENS WSPA. Also included is the groundwater reserve value, which is equal to 10% of estimated groundwater discharge (baseflow).

**Table 3.12 – Groundwater Percent Water Demand Components**

<b>Water Supply (Groundwater Recharge) (mm)</b>	<b>Water Reserve (10% Discharge) (mm)</b>
51	5.1

To complete the groundwater Stress Assessment, groundwater inflow to LENS must be quantified. It is anticipated that NPCA staff will complete this portion of the Stress Assessment as part of a separate project.

### **3.6 Uncertainty**

Any model of a natural system is a simplification of reality, and as such, is inherently uncertain. Although the calibration and verification processes are performed in an attempt to reduce uncertainty, the model results and water budgets reflect the uncertainty in the input parameters.

The certainty of the water balance estimates is inexorably tied to the ability of the climate stations used in the model to accurately represent the average climatic conditions over the WSPA. The current density of climate stations with long term datasets is likely not sufficient to fully reflect spatial climate variability, particularly during the summer months where extremely localized precipitation events are common (thunderstorms).

Further climate-related uncertainty is introduced into the process by the measurement error in climate observations. Uncertainty with the precipitation measurement has been estimated by Cumming Coburn Limited (2000) to be approximately  $\pm 10\%$ , with uncertainty during winter months reaching  $\pm 20\%$ . Precipitation measurement in winter months has a higher degree of uncertainty due to the difficulty of measuring snowfall.

Snow accumulation, ablation, redistribution and melt are significant hydrologic processes in Canadian watersheds. The rates of these processes are determined by the inter-relation of many factors, including: land cover, albedo, solar radiation, wind speed/direction, cloud cover, temperature fluctuations, rainfall amount/temperature, and new snow density. Lack of available data and a complete understanding on the interrelations and implications of these factors introduces a level of uncertainty into hydrologic modelling. The simplified snow processes within HEC-HMS reflects this level of uncertainty.

With no streamflow data to calibrate to, and verify the performance of the model against, there is a greater degree of uncertainty associated with results from this model, than the Twenty Mile or Upper Welland models. Given the geologic homogeneity, and the resulting likelihood that adjustments for Twenty Mile or Upper Welland are transferable to the LENS model areas, this uncertainty is minimized.

Care should be taken when interpreting results from extreme events, such as the 1998-1999 drought. During extreme events certain processes that may be insignificant under “average” conditions, and therefore not considered in regional scale modelling, may become dominant and affect the hydrologic response of a watershed. This can lead to model estimates deviating from observed conditions for such periods.

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All modelling algorithms are simplified and are unable to accurately reflect the host of processes that can affect the hydrologic response of a catchment to a precipitation event. These limitations are not solely specific to HEC-HMS, nor the algorithms contained within HEC-HMS, but are rather a limitation of hydrologic science. With an insufficient ability to conceptualize and replicate all hydrologic processes, hydrologic model algorithms group or average, many processes into one. This averaging has the ability to introduce error into water balance estimates, and often the solution reached by an averaged approach results in a non-unique solution (as is shown to be the case in LENS). This limitation should be kept in mind for utilizing water balance estimates generated as part of this study.

The climate of southwestern Ontario significantly varies from season to season. As a result of these changing seasons, hydrologic parameters (e.g. infiltration, depression storage, overland runoff routing) also vary. In the case of the freezing and thawing of soils, this can have a significant impact on the ability of soil to infiltrate water. HEC-HMS does not have the ability to vary parameters with season, and as such, is a source of uncertainty. Due to the dominance of the Haldimand Clay Plain in the area, and its limited ability to infiltrate water, even under warm conditions, it is expected that the uncertainty associated with this limitation is less significant than for highly pervious watersheds.

While any modelling exercise contains inherent uncertainties, it should be noted that the constructed HEC-HMS model produces estimates of streamflow and water balance values that exceed the level of accuracy expected for a Tier 1 Water Quantity Stress Assessment.

## **4. Conclusions and Recommendations**

A HEC-HMS continuous hydrologic model has been constructed for the Lake Erie North Shore (LENS) WSPA. The model has been adjusted based on a successful calibration and verification exercise carried out for both Twenty Mile Creek and the Upper Welland River. Due to the lack of observed streamflow data, the performance of the LENS model in predicting streamflow is not able to be determined. However, based on the performance of Twenty Mile and Upper Welland, it is very likely that this model is predicting reasonable streamflow and water balance estimates.

The generated water balance and streamflow estimates reflect the most complete understanding of the hydrologic system that is available, and represent the best available estimates. Significant uncertainties do remain; however, there is insufficient information to quantify the net impact of these uncertainties on the water balance and streamflow estimates. These estimates will form the foundation of a future Tier 1 Stress Assessment.

To advance the Tier 1 Water Quantity Stress Assessment as well as the basic understanding of the significant hydrologic processes, the following recommendations are made:

1. That groundwater inflow volumes to LENS be approximated by use of regional groundwater mapping products. Groundwater inflows are required to fully quantify the water supply term of the Groundwater Stress Assessment.
2. That monthly consumptive surface and groundwater demand (non-Great Lakes sources only) be approximated from Permits To Take Water, Census of Agriculture, and Census of Population. These consumptive demands are required to complete the Water Quantity Stress Assessment.
3. Utilizing the estimated consumptive demands, the groundwater inflow volumes, and values presented in Section 3.5.2, that the Tier 1 Water Quantity Stress Assessment be carried out. This will identify WSPAs that have a potential for hydrologic stress related to water takings.
4. That the water balance estimates generated from the Deficit and Constant Loss Method for LENS, be validated against estimates generated from a more detailed loss Method (Soil Moisture Accounting Method). This validation exercise would preferably be carried out on a WSPA with observed streamflow data. Should the more detailed Soil Moisture Accounting Method generate water balance estimates similar to the Deficit and Constant Loss, a higher level of certainty could be attached to water balance estimates. The need for further model refinement could be re-evaluated following the subsequent stress assessment.

Despite the uncertainties inherent with any modelling exercise, the LENS HEC-HMS model is an excellent tool for estimating the water supply components of a Tier 1 Water Quantity Stress Assessment. In addition to exceeding the expectations of a Tier 1 Stress Assessment, it will greatly assist NPCA staff in characterizing and understanding the fundamental hydrologic processes occurring within LENS.

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# TABLES



**TABLE 2.1**  
**MEAN ANNUAL CLIMATE STATION VALUES**  
**WATER AVAILABILITY STUDY**

MSC ID	NAME	AVERAGE ANNUAL (1991-2005)		
		PRECIPITATION (mm)	SNOW WATER EQUIVALENT (mm)	TEMPERATURE (°C)
6132148	DUNNVILLE PUMPING STN	948	91.1	8.6
6132470	FORT ERIE	1053	193	8.7
6133055	GRIMSBY MOUNTAIN	923	184	9.3
6133120	HAGERSVILLE	947	102	8.4
6153194	HAMILTON A	898	160	8.0
6153298	HAMILTON PSYCH HOSPITAL	850	108	8.8
6153290	HAMILTON MUNICIPAL LAB	793	101	9.5
6135657	NIAGARA FALLS NPCSH	948	160	9.4
6155097	MIDDLEPORT TS	896	110	8.5
6135FF4	NIAGARA ON THE LAKE	838	110	9.3
6136606	PORT COLBORNE	971	147	9.3
6136626	PORT DALHOUSIE	895	140	9.5
6137161	RIDGEVILLE	918	117	9.1
6137287	ST CATHARINES A	897	145	9.3
6139141	VINELAND	875	137	9.5
6139143	VINELAND RITTENHOUSE	850	115	9.4
6137306	ST CATHARINES POWER GLEN	890	135	9.2
6139445	WELLAND	969	148	9.0
6139148	VINELAND STATION RCS	840	146	9.1
6131165	CANBORO	894	120	8.4

Table Notes:

MSC - Meteorological Survey of Canada

**TABLE 2.2**  
**HYDROLOGIC SOIL GROUPS BY CATCHMENT**  
**LAKE ERIE NORTH SHORE WATERSHED PLANNING AREA**  
**WATER AVAILABILITY STUDY**

Catchment LENS ID	Area (km <sup>2</sup> )	Hydrologic Soil Groups (%)					
		A	B	C	D	ORG	URB
LENS_BBA_W100	1.6	0.1		12.0	1.2		86.6
LENS_BD_W100	5.8	0.7	9.2	27.7	36.6	13.5	12.3
LENS_BD_W200	5.3		3.9	38.4	57.7		
LENS_BDD_W100	3.9	0.6	14.5	0.9	79.4		4.6
LENS_BDD_W200	8.0	1.3	5.5	28.6	52.2	8.3	4.1
LENS_CD_W100	6.8		7.3	11.3	76.4		5.0
LENS_EMD_W100	5.9	1.2	10.9	13.4	4.1		70.2
LENS_EMD_W200	4.9		25.3	19.1	30.7		24.9
LENS_HOD_W100	1.5	2.6	1.9	67.7	27.4		0.1
LENS_LB_W100	5.6	0.3		29.4	70.2		
LENS_LB_W200	9.5	0.1		24.9	75.0		
LENS_LE2B_W100	2.0	2.8	27.6	57.4		6.4	5.8
LENS_LE4_W100	2.2		0.2		3.9	21.5	74.4
LENS_LE5_W100	1.8	2.2	16.0	6.0	24.3		51.5
LENS_LE6G_W100	1.1	0.5	2.0	0.1	84.4		13.0
LENS_LE8_W100	1.4	0.7			90.4		8.9
LENS_OMC_W100	2.9		17.0	24.7	39.7		18.3
LENS_PAD_W100	9.0	9.2	18.5	40.7	5.1	9.7	16.9
LENS_WD_W100	6.7		17.7	15.5	25.7	15.0	26.1
LENS_WD_W200	5.3		28.4	20.8	50.7		
LENS_WM_W100	1.5		15.8	45.2	39.0		0.1
Average %		1.7	13.0	25.5	43.7	12.4	24.9
% of LENS model		1.3%	10.9%	23.9%	43.4%	4.2%	16.3%
Area (km <sup>2</sup> )		1.24	10.11	22.17	40.32	3.94	15.13

Table Notes:

ORG - Organic, URB - Urban, LENS - Lake Erie North Shore

BBA - Bay Beach Area Drain, BD - Bearss Drain, BDD - Beaver Dam Drain, CD - Casey Drain

EMD - Eagle Marsh Drain, HOD - Hoover Drain, LB - Lowbanks Drain, LE2B - Lake Erie 2B, LE4 - Lake Erie 4

LE5 - Lake Erie 5, LE6G - Lake Erie 6G, LE8 - Lake Erie 8, OMC - Oil Mill Creek, PAD - Port Albino Drain

WD - Wignell Drain, WM - Wainfleet Marsh Drain

TABLE 2.3  
LAND COVER BY CATCHMENT  
LAKE ERIE NORTH SHORE WATERSHED PLANNING AREA  
WATER AVAILABILITY STUDY

Catchment LENS ID	Annual Crop	Mixed Agriculture	Mixed Crop	Monoculture	Orchards	Perennial Crop	Plantations	Vineyards	Coniferous Forest	Deciduous Forest	Forest	Hedge Rows	Mixed Forest	Built Up Pervious	Idle Land	Rural Land Use	Marsh	Swamp	Bog	Open Water	Built Up Impervious	Extraction	Transportation
LENS_BBA_W100		5.7		2.8		0.7				2.3	0.4	0.6		6.8		2.0		5.5		0.0	47.8		25.4
LENS_BD_W100	0.3	11.9		12.2						7.8	1.8	1.0		5.9	0.4	26.3	0.1	20.7			7.2		4.4
LENS_BD_W200		15.0		23.4		0.2	0.6			3.9	1.0	1.0		7.3	0.4	37.6		4.6		0.0			4.9
LENS_BDD_W100		22.5		16.5		0.1	1.3			8.2	2.4	0.5	0.1	5.8	0.0	17.1	0.4	21.6			0.2		3.3
LENS_BDD_W200		17.4		22.3						4.3	1.1	0.7		1.5	0.0	29.5	0.1	18.7		0.4	0.3		3.6
LENS_CD_W100		17.3		26.3			1.0		0.1	7.8	1.3	0.2		0.6	0.1	24.9		14.9		0.0	3.0		2.7
LENS_EMD_W100	0.0	5.8		15.3		0.0	0.2			4.4	0.2	0.0		4.1	0.4	4.3	6.8	11.3		1.4	31.1		14.7
LENS_EMD_W200		9.0		27.5						7.9	0.6	1.5		2.5	0.3	15.4		5.0			4.9	20.8	4.6
LENS_HOD_W100		2.7		49.7						2.9	2.3	0.2		1.1	0.2	24.3		11.6			4.1		0.9
LENS_LB_W100	0.1	3.1		57.9			0.9			1.3	1.5	0.1		1.8		18.0	0.3	13.2		0.0	0.2		1.6
LENS_LB_W200		5.4		40.2			0.6			1.1	0.9	0.1		1.2		22.1	0.1	25.4			0.4		2.6
LENS_LE2B_W100	0.6	20.6		3.0		2.8				3.5	1.1	2.6		0.3	2.0	15.9		37.2		0.0	7.3		3.1
LENS_LE4_W100	0.1	1.3		6.1		0.5				3.8	0.2			20.3	2.1	2.8	9.6	10.0		1.0	27.1		15.0
LENS_LE5_W100		6.2		25.9		0.1				7.9	0.5			0.0	1.1	8.4	1.7	13.4		18.1	10.4		6.3
LENS_LE6G_W100		14.0		18.7						5.2	1.7			0.4		35.5	1.9	9.4		0.0	6.3		6.8
LENS_LE8_W100		7.7		22.0						11.0	2.5				0.3	20.7		27.3		0.0	4.6		4.0
LENS_OMC_W100		10.7		15.6		0.1	0.7			10.7	2.5	3.2	0.5	3.6	0.2	27.2		11.8		0.3	7.6		5.6
LENS_PAD_W100	0.5	11.6		11.0						10.5	2.2	0.4		5.7	1.4	13.2		30.3		1.4	7.7		3.9
LENS_WD_W100	7.6	12.7		17.6			0.5		0.5	5.5	0.6	0.2	1.1	3.5	2.0	19.8	1.5	2.8		0.0	1.5	18.6	4.0
LENS_WD_W200	2.1	11.8		23.8			0.1		0.0	3.3	0.3	0.4		3.6		46.1	0.0	3.1			1.0	0.0	4.5
LENS_WM_W100		9.6		26.7						25.2	0.8	0.5				26.7		7.3		0.0	2.4		0.9
Average %	1.4	10.6	NA	22.1	NA	0.6	0.7	NA	0.2	6.6	1.2	0.8	0.6	4.0	0.7	20.8	2.1	14.5	NA	1.5	8.7	13.1	5.8
% of LENS model	0.8%	11.1%		23.4%		0.1%	0.4%		0.0%	5.9%	1.2%	0.6%	0.1%	3.6%	0.5%	21.7%	0.9%	15.5%		0.6%	6.2%	2.4%	5.0%
Area/Land Cover (km <sup>2</sup> )	0.7	10.3		21.8		0.1	0.3		0.0	5.4	1.1	0.5	0.1	3.3	0.5	20.1	0.8	14.4		0.6	5.8	2.3	4.6

Table Notes:  
LENS - Lake Erie North Shore  
BBA - Bay Beach Area Drain, BD - Bearss Drain, BDD - Beaver Dam Drain, CD - Casey Drain, EMD - Eagle Marsh Drain, HOD - Hoover Drain, LB - Lowbanks Drain, LE2B - Lake Erie 2B, LE4 - Lake Erie 4, LE5 - Lake Erie 5, LE6G - Lake Erie 6G, LE8 - Lake Erie 8  
OMC - Oil Mill Creek, PAD - Port Albino Drain, WD - Wignell Drain, WM - Wainfleet Marsh Drain

**TABLE 3.1**  
**CATCHMENT PARAMETERS**  
**LAKE ERIE NORTH SHORE WATERSHED PLANNING AREA**  
**WATER AVAILABILITY STUDY**

Catchment LENS ID	Area (km <sup>2</sup> )	Slope (%)	Impervious Area (%)	Curve Number (CN)	Basin Time Lag (hours)	Maximum storage (mm)	Infiltration Rate (mm/hour)
LENS_BBA_W100	1.63	2.4	49.0	84	1.2	28.1	0.3
LENS_BD_W200	5.30	1.4	5.0	85	1.8	198.2	1.8
LENS_BD_W100	5.78	1.9	8.0	80	2.0	195.1	1.8
LENS_BDD_W100	3.91	2.2	3.2	82	1.6	205.4	1.4
LENS_BDD_W200	8.04	1.5	3.8	82	2.6	212.7	1.6
LENS_CD_W100	6.81	1.5	4.2	83	2.2	198.8	1.3
LENS_EMD_W100	5.90	2.4	30.2	81	2.0	65.0	1.1
LENS_EMD_W200	4.91	3.7	7.0	84	1.2	137.0	2.3
LENS_HOD_W100	1.55	1.4	3.1	82	1.5	222.5	2.3
LENS_LE10A_W100	0.44	0.5	1.4	86	1.3	187.7	2.1
LENS_LE10_W100	0.29	2.3	9.2	85	0.6	111.2	0.8
LENS_LE11_W100	0.46	2.2	3.9	84	0.6	202.6	2.0
LENS_LE2_W100	0.15	1.7	6.3	90	0.5	20.0	0.0
LENS_LE2A_W100	0.82	2.4	59.1	96	0.7	20.0	0.0
LENS_LE2B_W100	1.97	1.5	6.9	75	2.2	240.7	3.5
LENS_LE4_W100	2.24	2.9	28.3	87	1.0	45.7	0.2
LENS_LE5_W100	1.79	3.2	11.6	84	1.0	104.7	1.4
LENS_LE5A_W100	0.83	1.3	8.5	79	1.6	144.2	3.3
LENS_LE6_W100	0.35	1.5	9.3	84	0.8	188.5	2.4
LENS_LE6A_W100	0.88	2.2	6.4	81	0.9	184.9	2.8
LENS_LE6B_W100	0.33	2.3	18.9	72	0.9	178.1	4.0
LENS_LE6C_W100	0.23	2.1	12.2	74	0.6	218.4	2.9
LENS_LE6D_W100	0.63	2.0	8.4	77	0.9	178.7	3.8
LENS_LE6E_W100	0.10	2.2	11.5	79	0.5	144.4	4.2
LENS_LE6F_W100	0.31	1.8	25.4	82	0.8	107.0	2.2
LENS_LE6G_W100	1.13	1.3	10.3	84	1.3	176.1	1.0
LENS_LE7_W100	0.34	1.4	11.6	84	0.8	145.9	0.9
LENS_LE8_W100	1.45	1.4	6.5	81	1.4	225.7	0.9
LENS_LE8A_W100	0.05	2.8	32.6	86	0.3	102.3	0.5
LENS_LE9_W100	0.03	2.8	40.5	87	0.2	80.4	0.6
LENS_LE9A_W100	0.12	2.3	13.1	94	0.3	1.9	0.2
LENS_LB_W200	9.47	1.4	2.8	83	3.3	219.5	1.5
LENS_LB_W100	5.63	1.4	1.7	85	2.5	200.1	1.3
LENS_OMC_W100	2.94	1.9	9.3	80	1.8	171.0	2.1
LENS_PAD_W100	9.00	3.1	7.6	77	2.4	189.9	3.3
LENS_WM_W100	1.46	1.7	2.4	79	1.5	247.6	2.9
LENS_WD_W200	5.29	1.5	4.9	83	3.1	182.7	2.6
LENS_WD_W100	6.74	4.4	4.8	86	1.7	133.7	1.7

**TABLE 3.1****CATCHMENT PARAMETERS****LAKE ERIE NORTH SHORE WATERSHED PLANNING AREA****WATER AVAILABILITY STUDY**

Minimum	0.03	0.5	1.4	72.1	0.2	1.9	0.0
Maximum	9.47	4.4	59.1	95.7	3.3	247.6	4.2
Average	2.61	2.1	12.9	82.8	1.4	153.1	1.8

## Table Notes:

LENS - Lake Erie North Shore, BBA - Bay Beach Area Drain, BD - Bearss Drain

BDD - Beaver Dam Drain, CD - Casey Drain, EMD - Eagle Marsh Drain, HOD - Hoover Drain

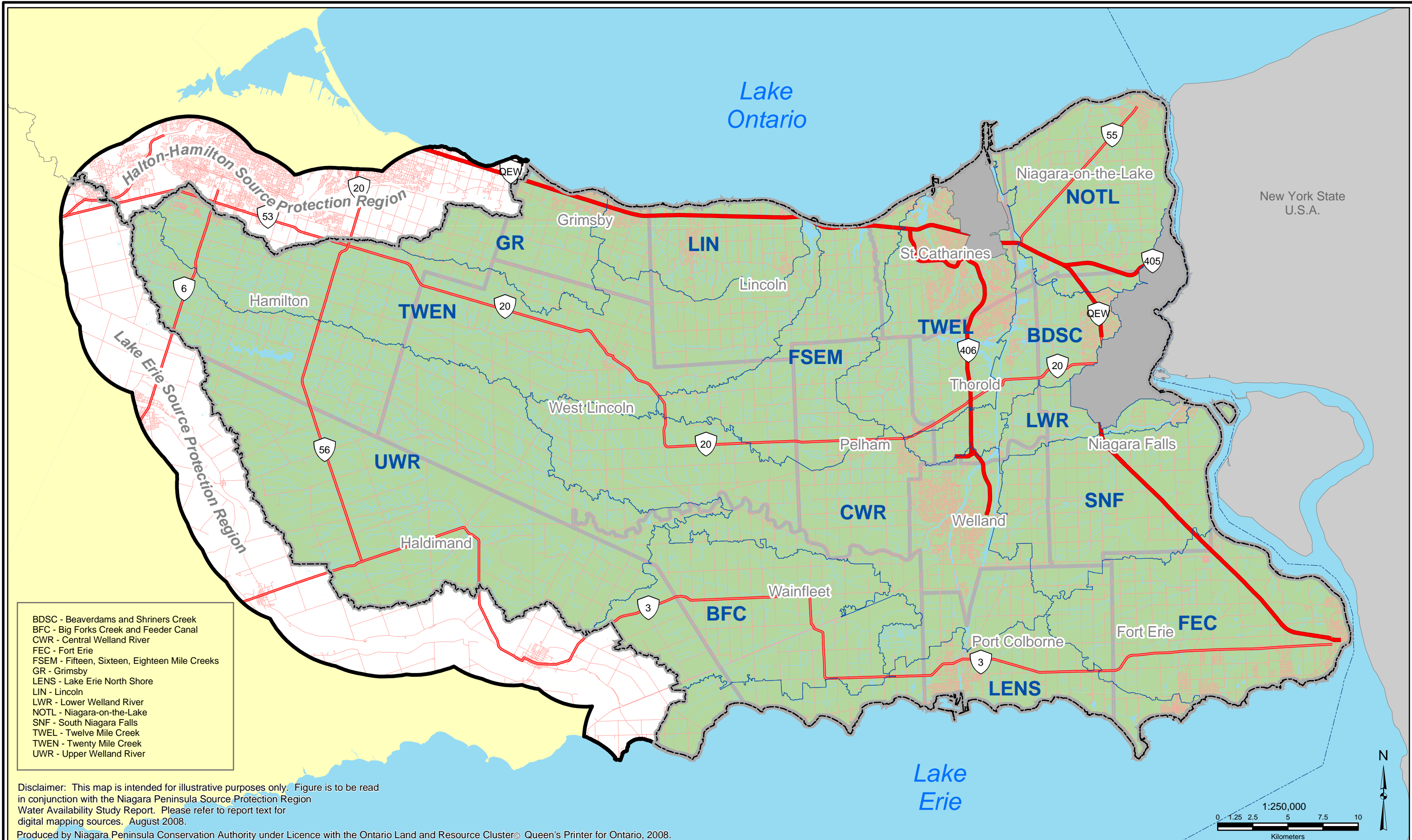
LB - Lowbanks Drain, LE2B - Lake Erie 2B, LE4 - Lake Erie 4, LE5 - Lake Erie 5, LE6G - Lake Erie 6G

LE8 - Lake Erie 8, OMC - Oil Mill Creek, PAD - Port Albino Drain, WD - Wignell Drain

WM - Wainfleet Marsh Drain

# FIGURES





- BDSC - Beaverdams and Shriners Creek  
BFC - Big Forks Creek and Feeder Canal  
CWR - Central Welland River  
FEC - Fort Erie  
FSEM - Fifteen, Sixteen, Eighteen Mile Creeks  
GR - Grimsby  
LENS - Lake Erie North Shore  
LIN - Lincoln  
LWR - Lower Welland River  
NOTL - Niagara-on-the-Lake  
SNF - South Niagara Falls  
TWEL - Twelve Mile Creek  
TWEN - Twenty Mile Creek  
UWR - Upper Welland River

Disclaimer: This map is intended for illustrative purposes only. Figure is to be read in conjunction with the Niagara Peninsula Source Protection Region Water Availability Study Report. Please refer to report text for digital mapping sources. August 2008.

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**Legend**

- |                        |                         |  |
|------------------------|-------------------------|--|
| Extended Context Area  | Major Highways          | Niagara Peninsula Source Protection Region |
| SPR Boundary           | Highways                | NPCA Watershed Planning Areas              |
| Municipal Boundaries   | Roads                   | Urban Areas                                |
| International Boundary | Rivers, Streams, Creeks |  |



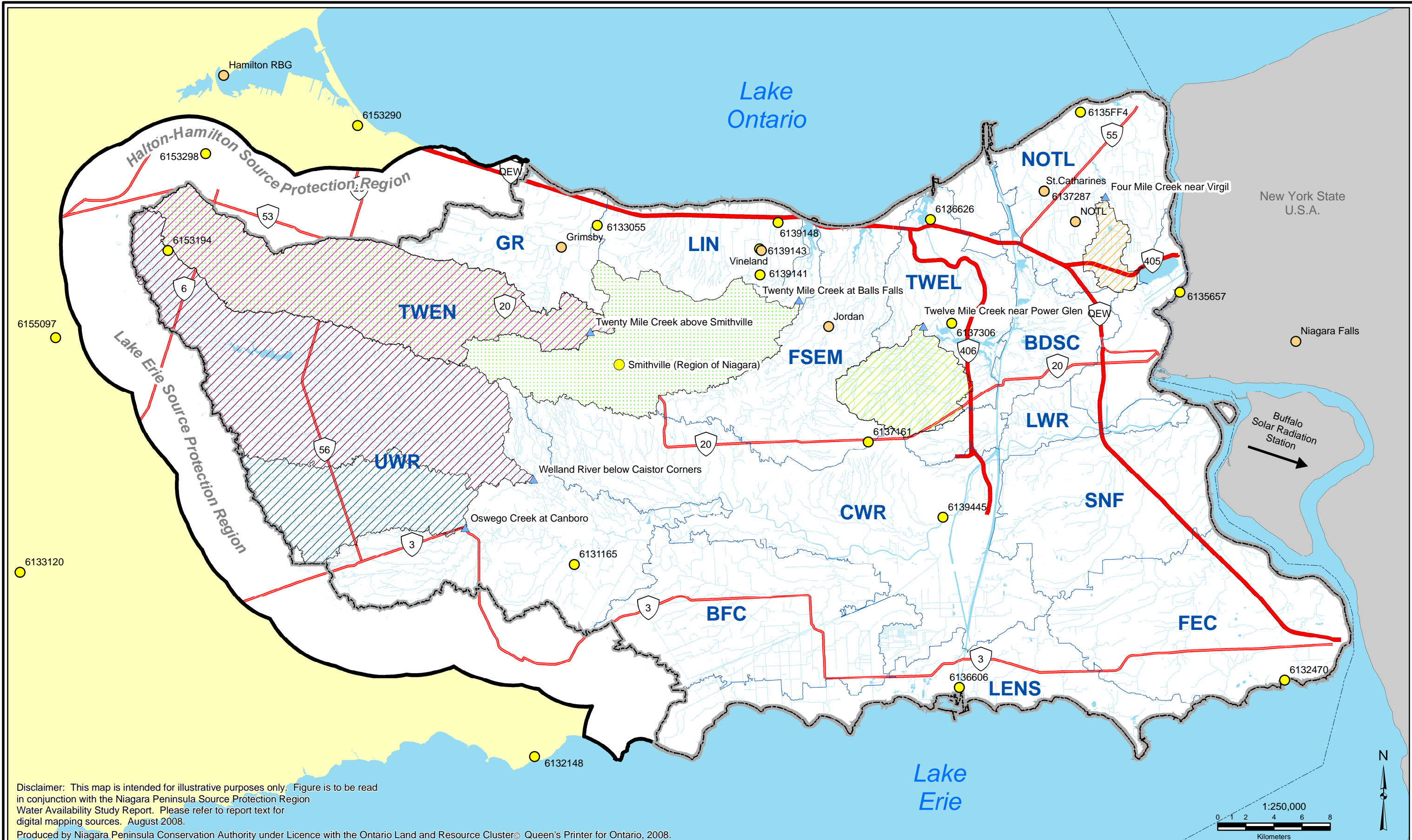
**Water Availability Study**

*Figure 1.1. Study Area*

All Frames: North American Datum 1983, Universal Transverse Mercator 6° Projection, Zone 17N, Central Meridian 81° West.







Disclaimer: This map is intended for illustrative purposes only. Figure is to be read in conjunction with the Niagara Peninsula Source Protection Region Water Availability Study Report. Please refer to report text for digital mapping sources. August 2008.

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**Legend**

- Extended Context Area
- SPR Boundary
- Municipal Boundaries
- International Boundary
- Major Highways
- Highways
- Rivers, Streams, Creeks
- NPCA Watershed Planning Area
- MSC Climate Station
- WSC Stream Gauge Station

**Water Survey Canada Stream Gauge Drainage Area**

- Four Mile Near Virgil
- Twelve Mile Creek
- Twenty Mile Creek (Smithville)
- Welland River (Caistor Corners)
- Oswego Creek
- Twenty Mile Creek (Balls Falls)

**Overview Map**

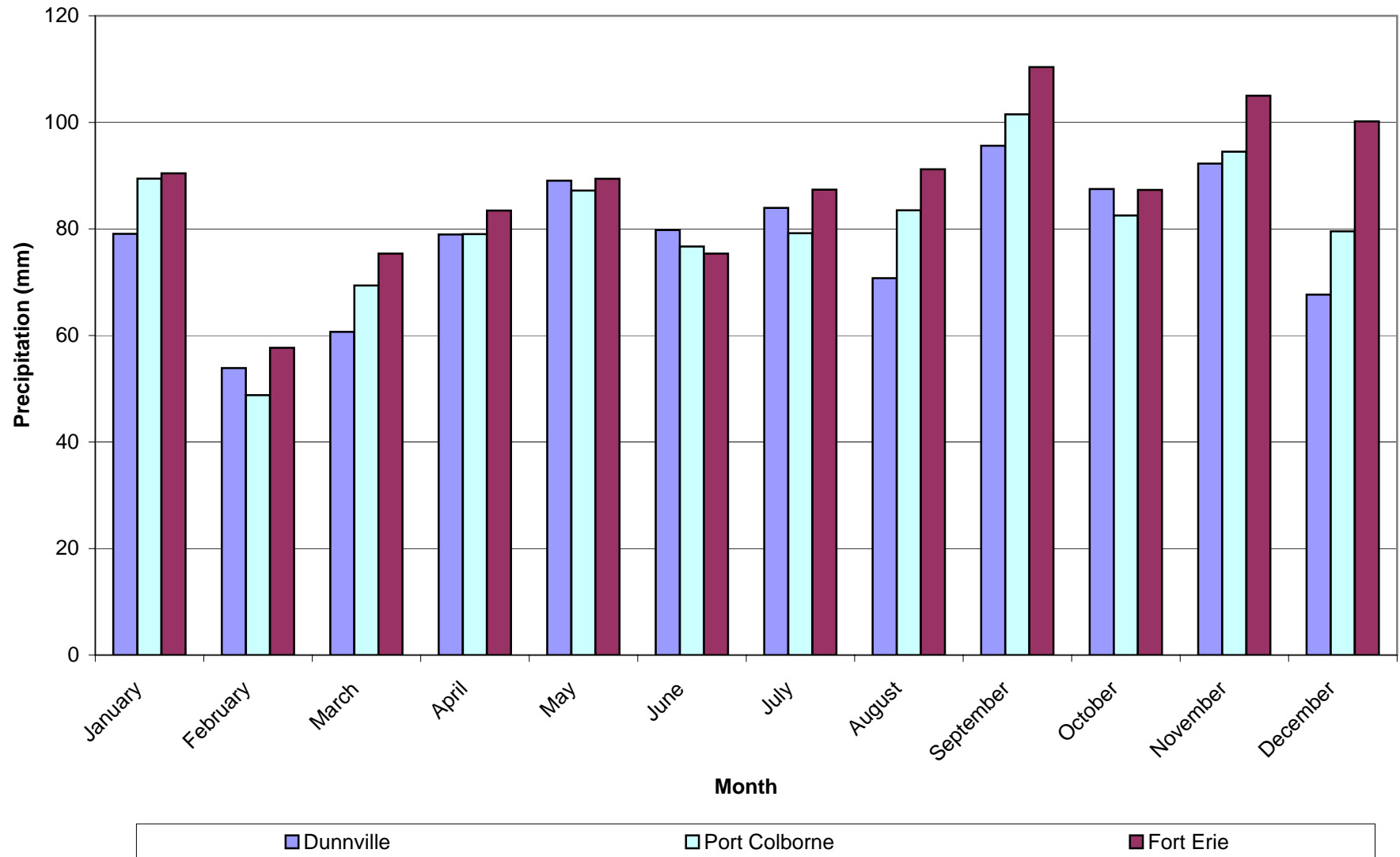
**Water Availability Study**

**Figure 2.1. Study Meteorological Stations and Stream Gauges**

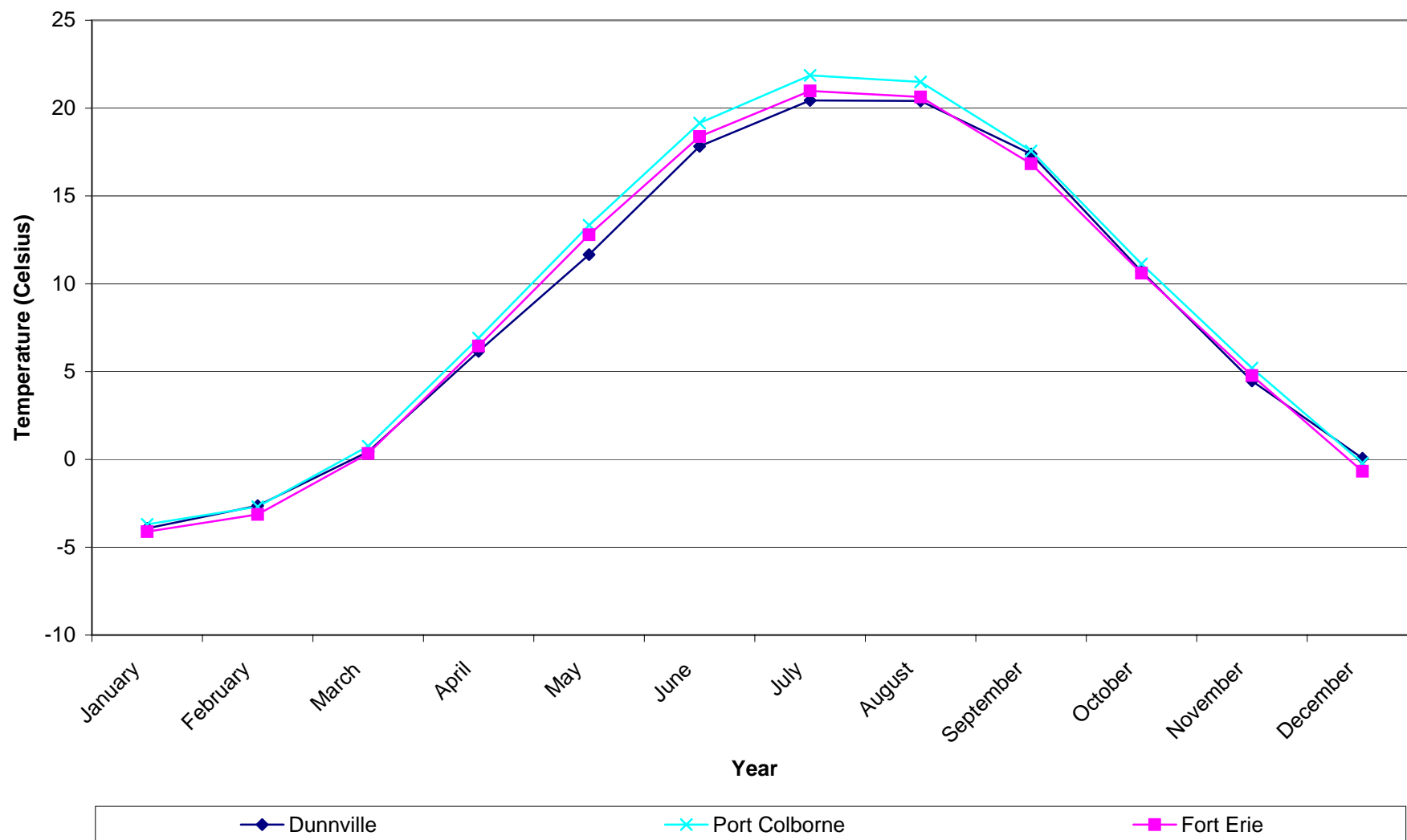
All Frames: North American Datum 1983, Universal Transverse Mercator 6° Projection, Zone 17N, Central Meridian 81° West.

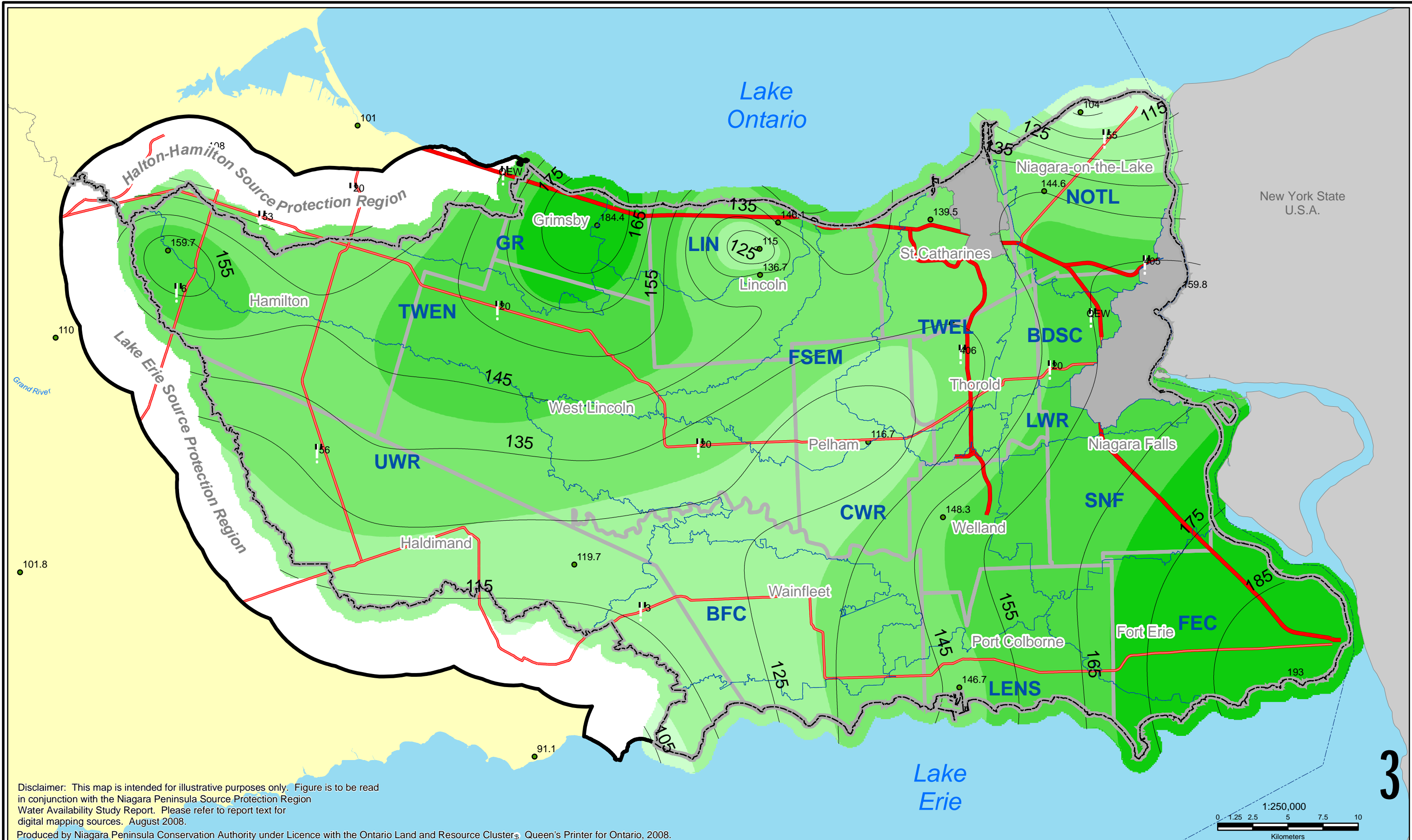


**FIGURE 2.2**  
**MEAN (1991-2005) MONTHLY PRECIPITATION**  
**LENS WATERSHED PLANNING AREAS**

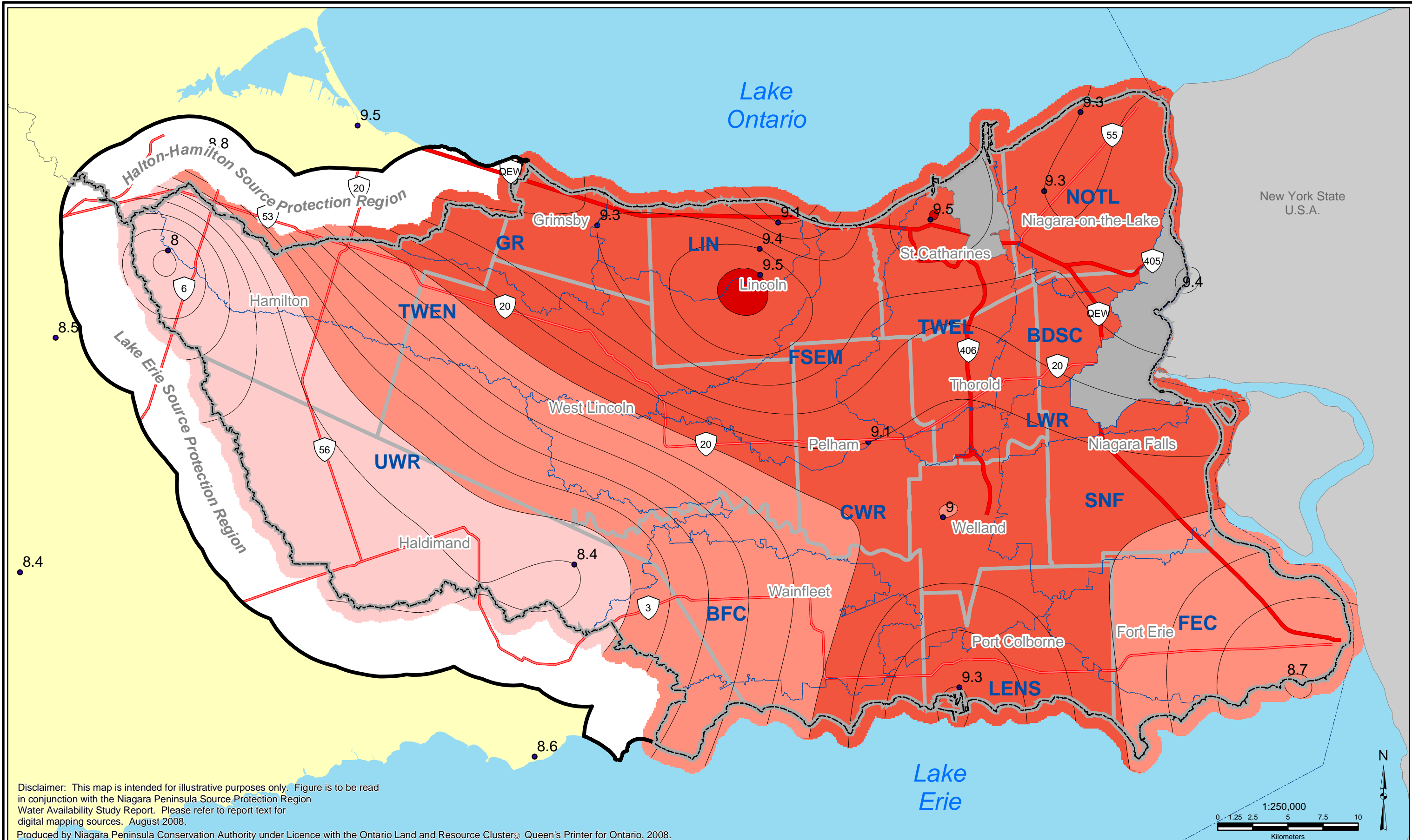


**FIGURE 2.3**  
**MEAN (1991-2005) MONTHLY TEMPERATURE**  
**LENS WATERSHED PLAN AREA**





<b>Legend</b> <div> <div>Extended Context Area</div> <div>SPR Boundary</div> <div>Municipal Boundaries</div> <div>International Boundary</div> </div>		<div> <div>Major Highways</div> <div>Highways</div> <div>Roads</div> <div>Rivers, Streams, Creeks</div> <div>Ponds, Reservoirs, Lakes</div> </div>		<div> <div>Niagara Peninsula Source Protection Region</div> <div>NPCA Watershed Planning Areas</div> <div>Urban Areas</div> <div>MSC Snow Data</div> </div>		<b>1991 - 2005 (mm)</b> <div> <div>90 - 110</div> <div>110 - 130</div> <div>130 - 150</div> <div>150 - 170</div> <div>170 - 190</div> </div>	<div> <div> <div>Lake Huron</div> <div>Lake Ontario</div> <div>Lake Erie</div> </div> <div>Study Area</div> <div>Overview Map</div> </div>	<div> <div> <div>DRINKING WATER SOURCE PROTECTION</div> <div>Niagara Peninsula Source Protection Region</div> </div> <div> <div>Water Availability Study</div> <div>Figure 2.4. Mean Annual Snowfall Water Equivalent</div> <div> <div>All Frames:</div> <div>North American Datum 1983, Universal Transverse Mercator 6° Projection, Zone 17N, Central Meridian 81° West.</div> </div> </div> </div>	<div> <div>Ontario</div> </div>
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**Legend**

- Extended Context Area
- SPR Boundary
- Municipal Boundaries
- International Boundary
- Major Highways
- Highways
- Roads
- Rivers, Streams, Creeks
- Ponds, Reservoirs, Lakes
- Niagara Peninsula Source Protection Region 1991 - 2005 (oC)
- NPCA Watershed Planning Areas
- Urban Areas
- MSC Temperature Stations
- 8.0 - 8.5
- 8.5 - 9.0
- 9.0 - 9.5
- 9.5 - 10.0

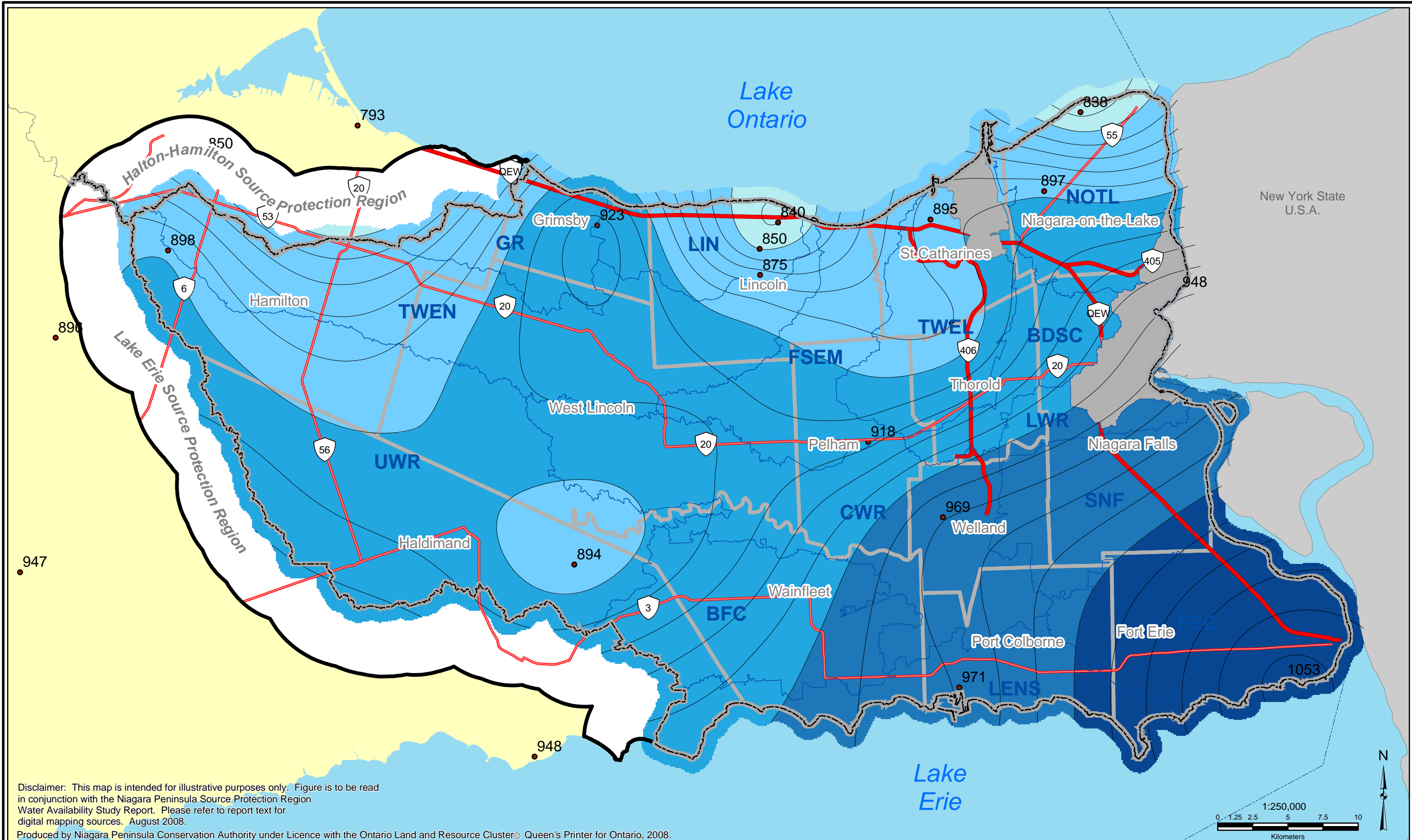
Overview Map

**Water Availability Study**

Figure 2.5. Mean Annual Temperature

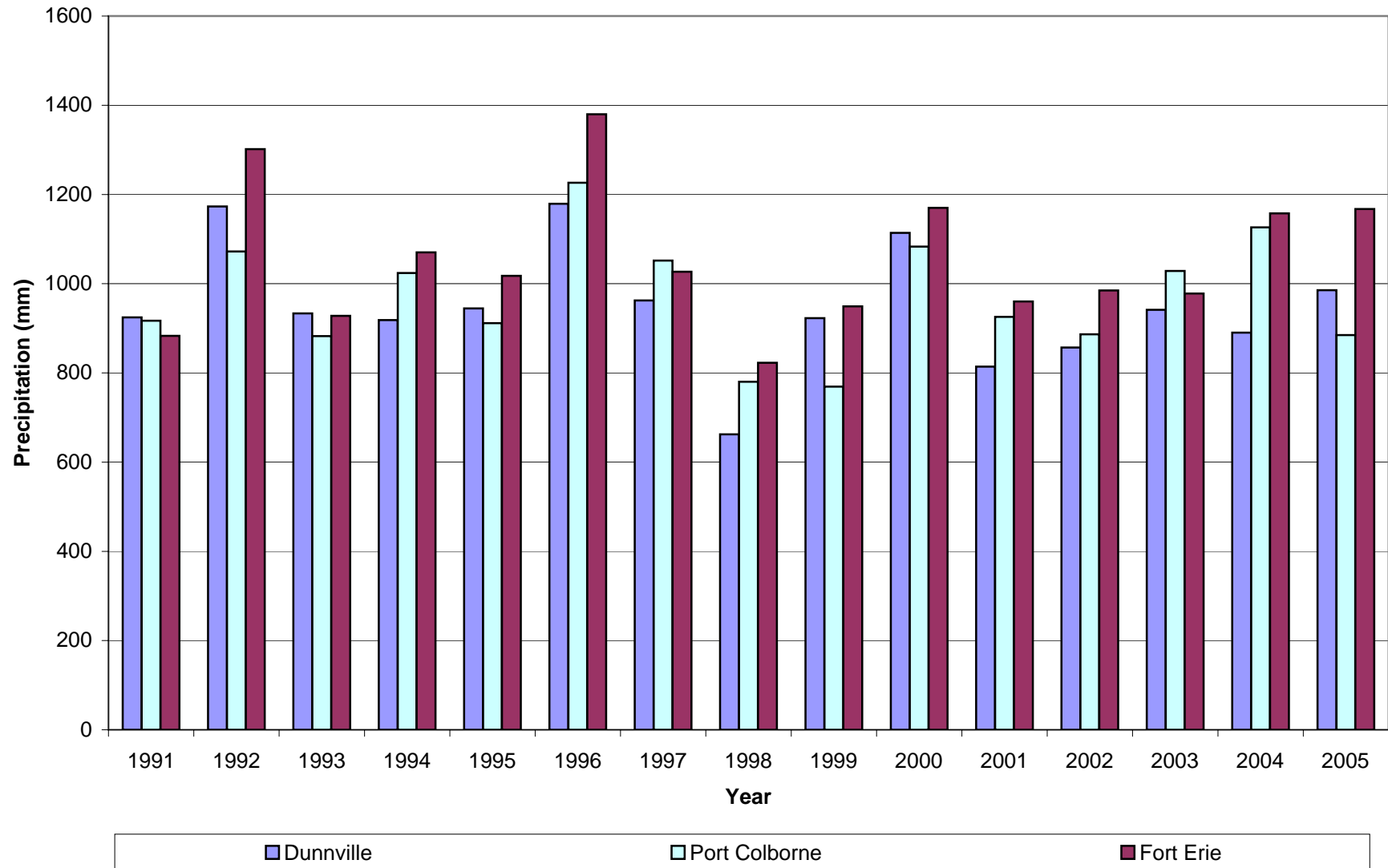
All Frames: North American Datum 1983, Universal Transverse Mercator 6° Projection, Zone 17N, Central Meridian 81° West.



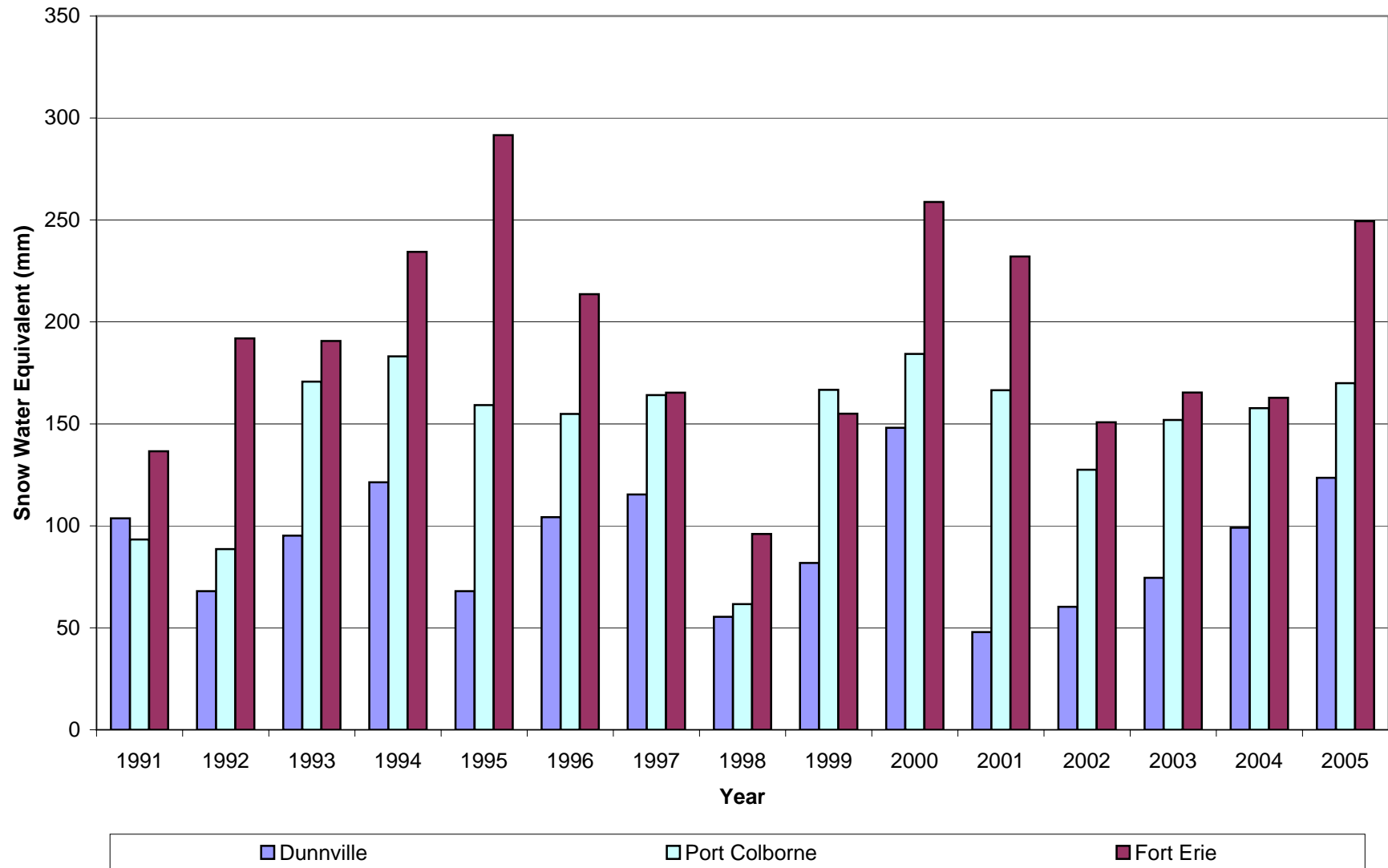


<b>Legend</b> <div> <div>Extended Context Area</div> <div>SPR Boundary</div> <div>Municipal Boundaries</div> <div>International Boundary</div> </div>		<div> <div>Major Highways</div> <div>Highways</div> <div>Roads</div> <div>Rivers, Streams, Creeks</div> <div>Ponds, Reservoirs, Lakes</div> </div>		<div> <div>Niagara Peninsula Source Protection Region 1991 - 2005 (mm)</div> <div> <div>NPCA Watershed Planning Areas</div> <div>Urban Areas</div> <div>MSC Temperature Stations</div> </div> </div>		<div> <div>800 - 850</div> <div>850 - 900</div> <div>900 - 950</div> <div>950 - 1000</div> <div>1000 - 1050</div> </div>		<div> <div> <div>Lake Huron</div> <div>Lake Ontario</div> <div>Lake Erie</div> <div>Study Area</div> </div> <div>Overview Map</div> </div>		<div> <div> <div>DRINKING WATER SOURCE PROTECTION</div> <div>Niagara Peninsula Source Protection Area</div> </div> <div>Water Availability Study</div> </div>		<div> <div>Figure 2.6. Mean Annual Precipitation</div> <div> <div>All Frames:</div> <div>North American Datum 1983, Universal Transverse Mercator 6° Projection, Zone 17N, Central Meridian 81° West.</div> </div> </div>		<div> <div>Ontario</div> </div>	
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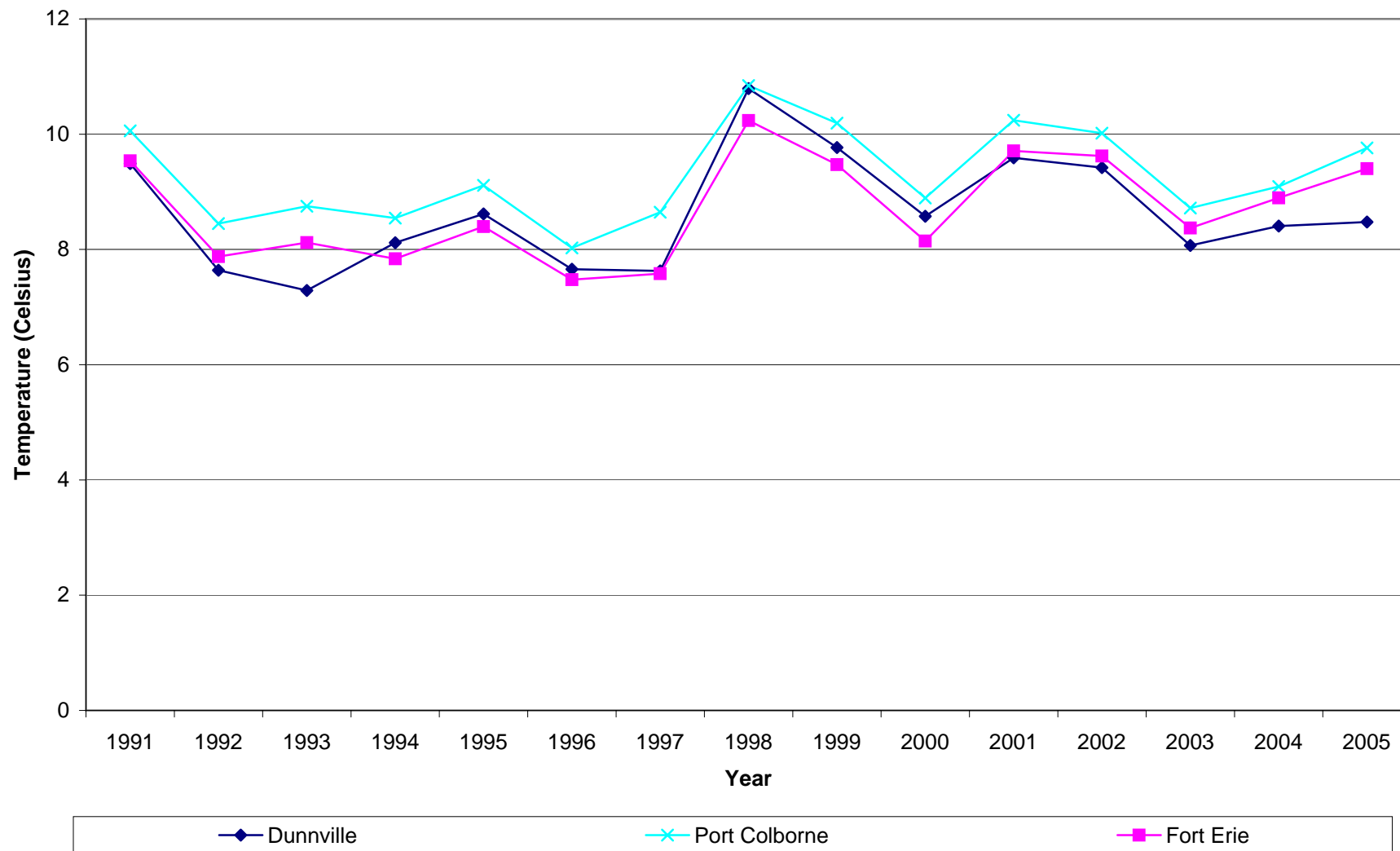
**FIGURE 2.7**  
**ANNUAL PRECIPITATION**  
**LENS WATERSHED PLANNING AREA**



**FIGURE 2.8**  
**ANNUAL SNOW WATER EQUIVALENT**  
**LENS WATERSHED PLANNING AREA**

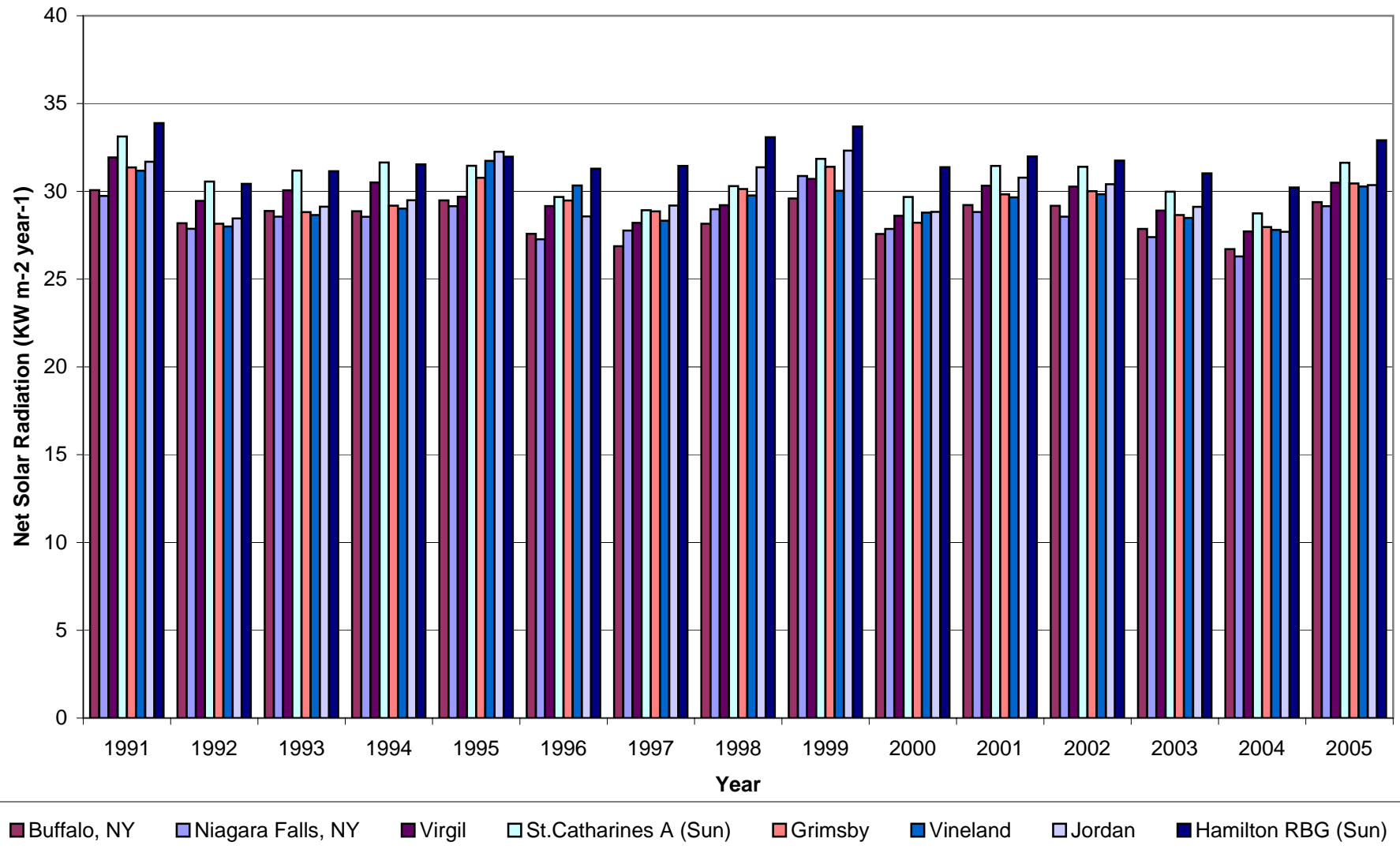


**FIGURE 2.9**  
**MEAN (1991-2005) ANNUAL TEMPERATURE**  
**LENS WATERSHED PLAN AREA**

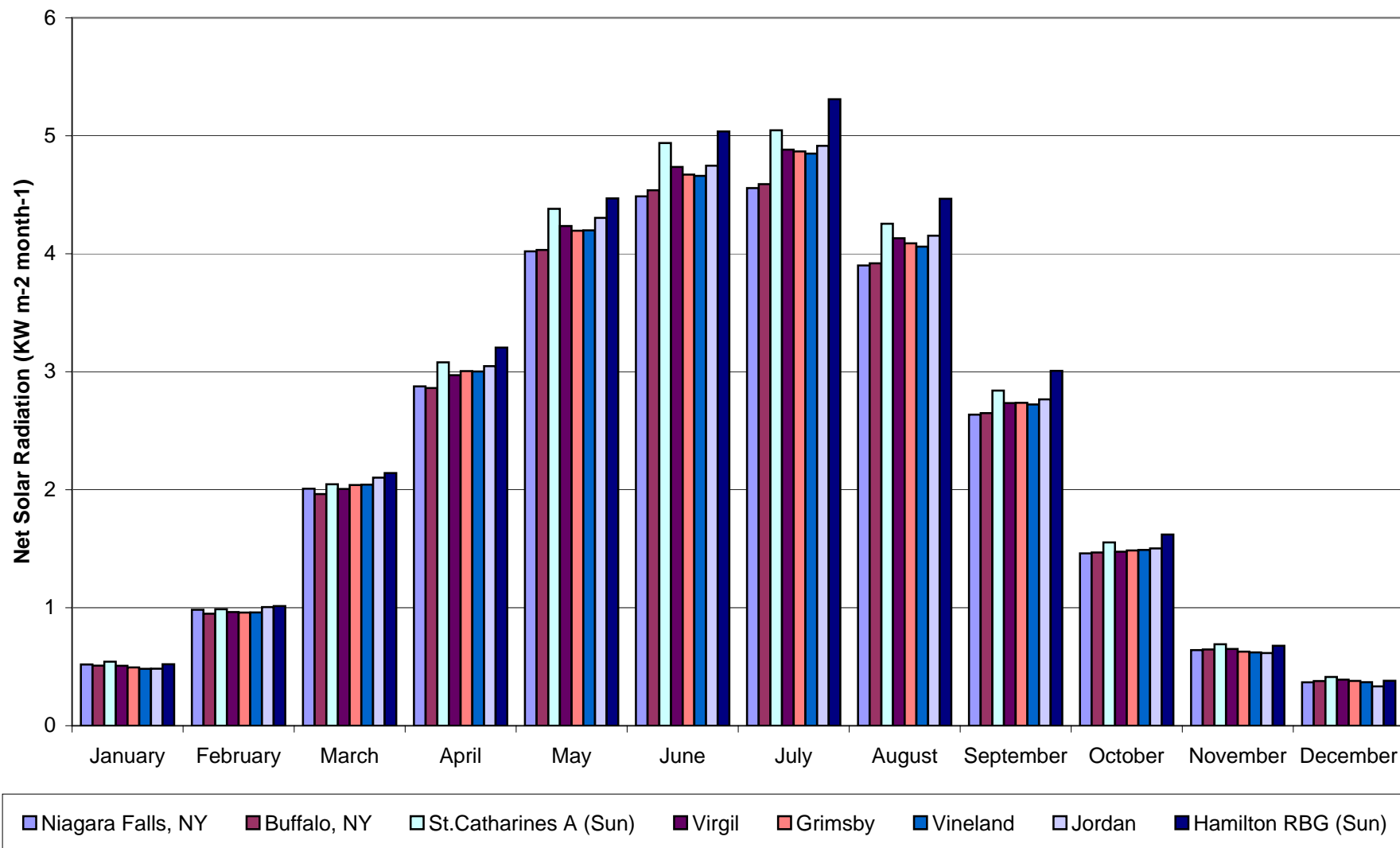




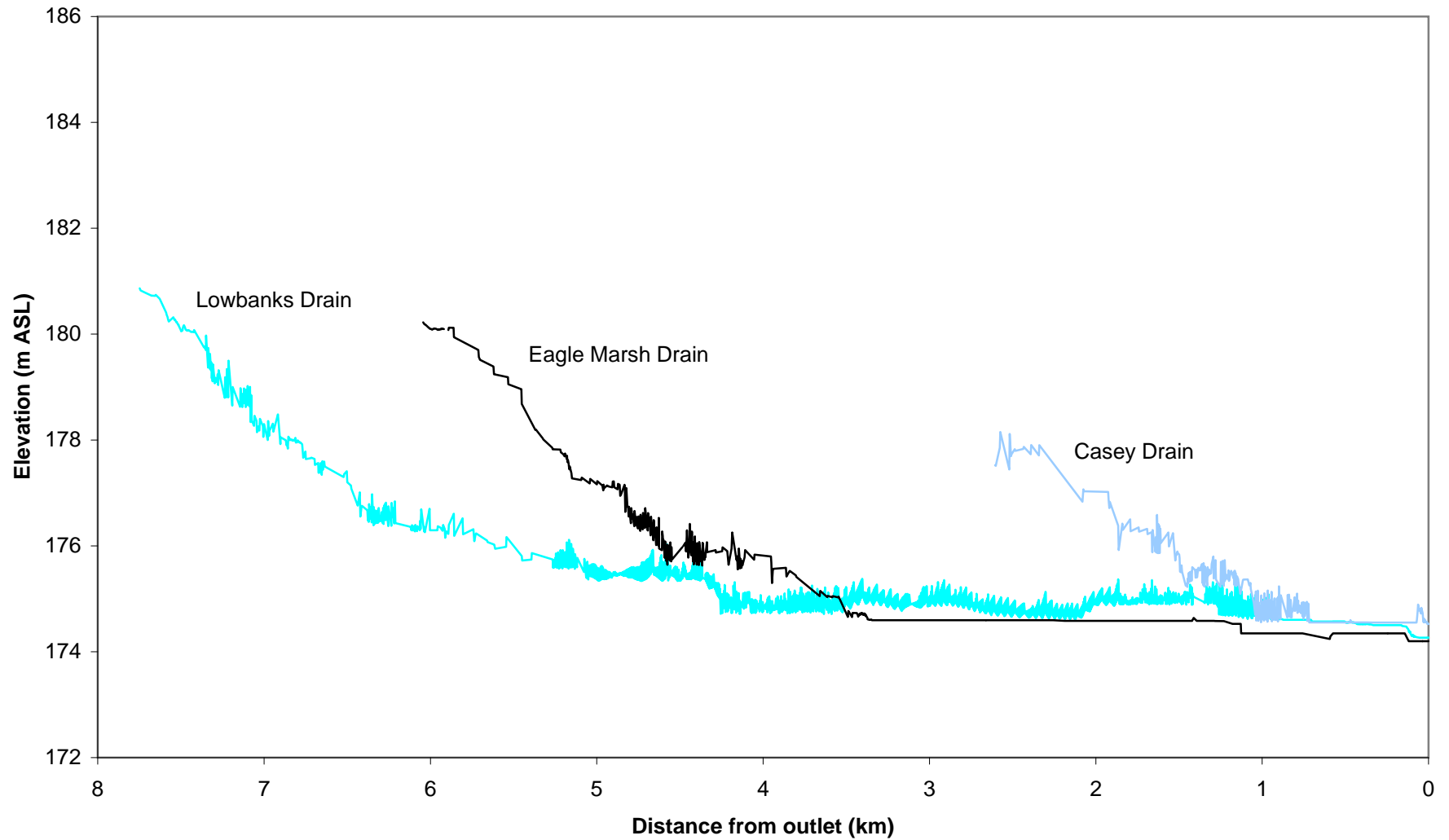
**Figure 2.10**  
**ANNUAL NET SOLAR RADIATION**



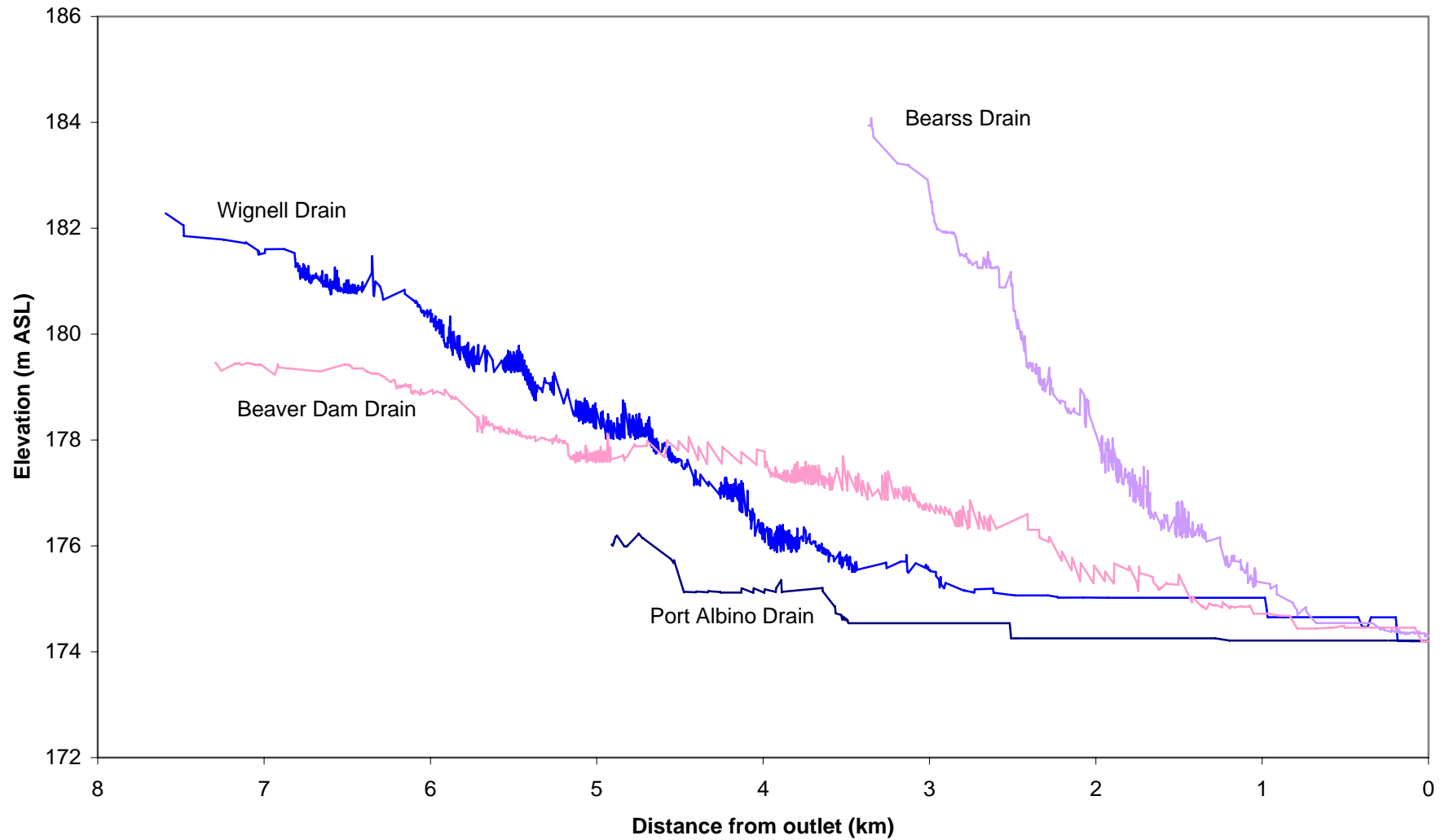
**Figure 2.11**  
**MEAN (1991-2005) MONTHLY NET SOLAR RADIATION**

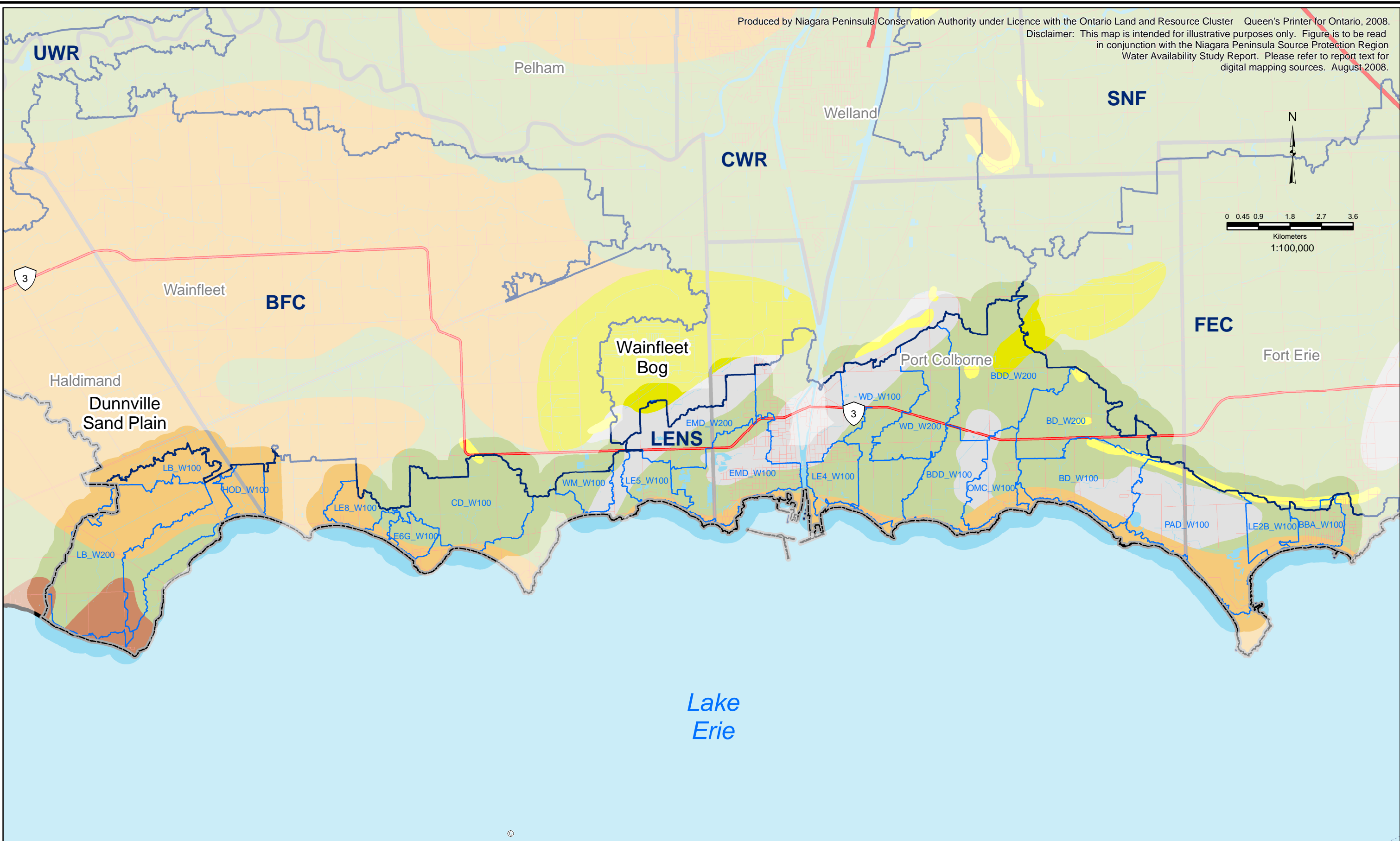


**Figure 2.12a**  
**Major Channel Profiles for Western Lake Erie North Shore Watershed Planning Area**



**Figure 2.12b**  
**Channel Profiles for Eastern Lake Erie North Shore Watershed Planning Area**

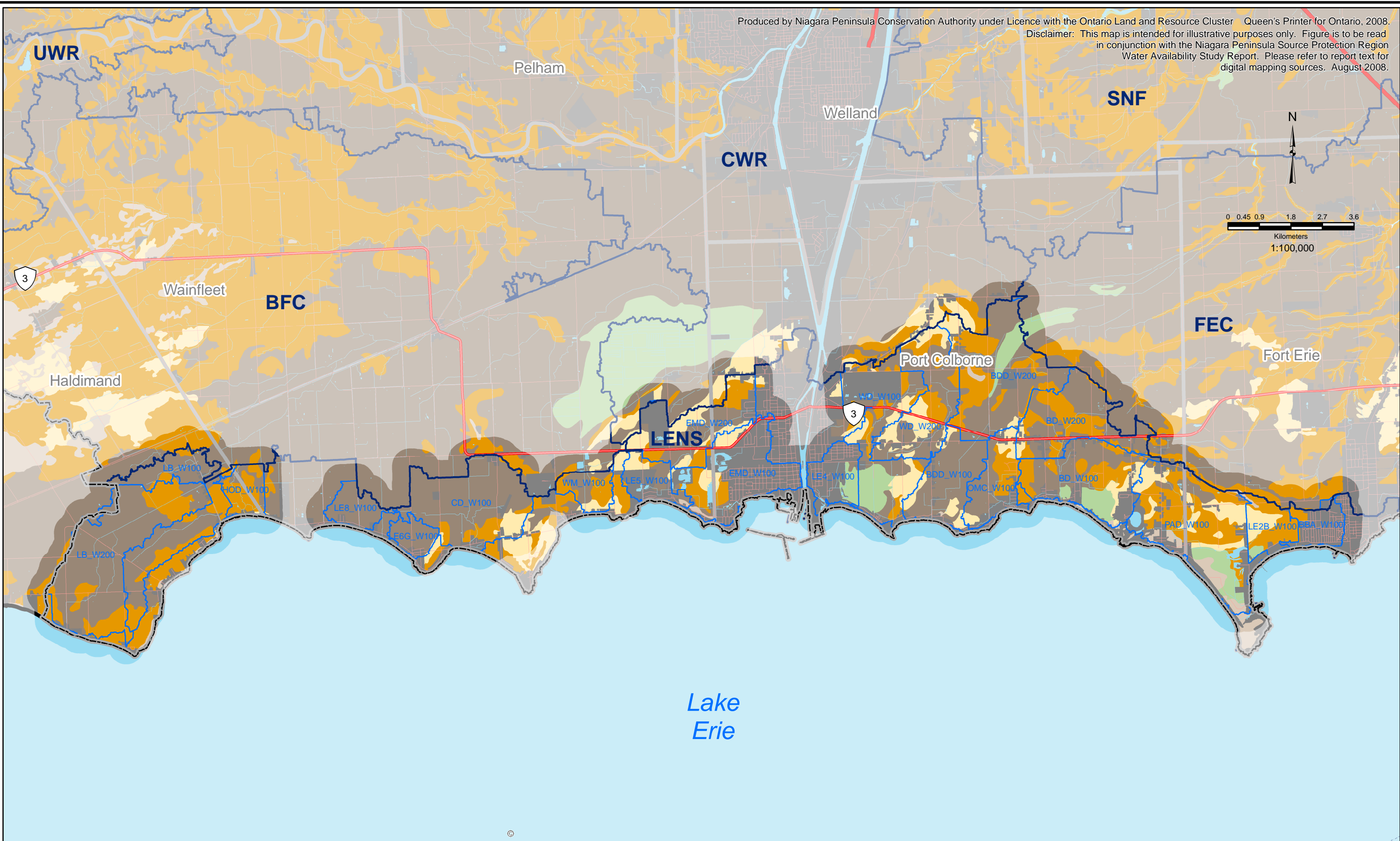




<b>Legend</b> Extended Context Area SPR Boundary Municipal Boundaries International Boundary		Major Highways Highways Roads Rivers, Streams, Creeks Ponds, Reservoirs, Lakes		NPCA Watershed Planning Area HMS Model Subcatchments BBA_W100: Subcatchment ID		Beaches and Shorecliffs Escarpment Clay Plain Sand Plain Till Moraine		Kame Moraine Peat and Muck Limestone Plain Shale Plain Water	
--	--	--	--	--	--	---	--	--	--

Overview Map

**Water Availability Study**  
*Figure 2.13. LENS Physiography*  
 All Frames: North American Datum 1983, Universal Transverse Mercator 6° Projection, Zone 17N, Central Meridian 81° West.



**Legend**

- Extended Context Area
- SPR Boundary
- Municipal Boundaries
- International Boundary
- Major Highways
- Highways
- Roads
- Rivers, Streams, Creeks
- Ponds, Reservoirs, Lakes

**Soil HSG**

- A
- B
- C
- D
- Organic Soils
- Urbanized

**NPCA Watershed Planning Area**

- HMS Model Subcatchments
- BBA\_W100: Subcatchment ID

*\*Please note that soils mapping in this figure is represented by the combination of three different former county soil surveys.*

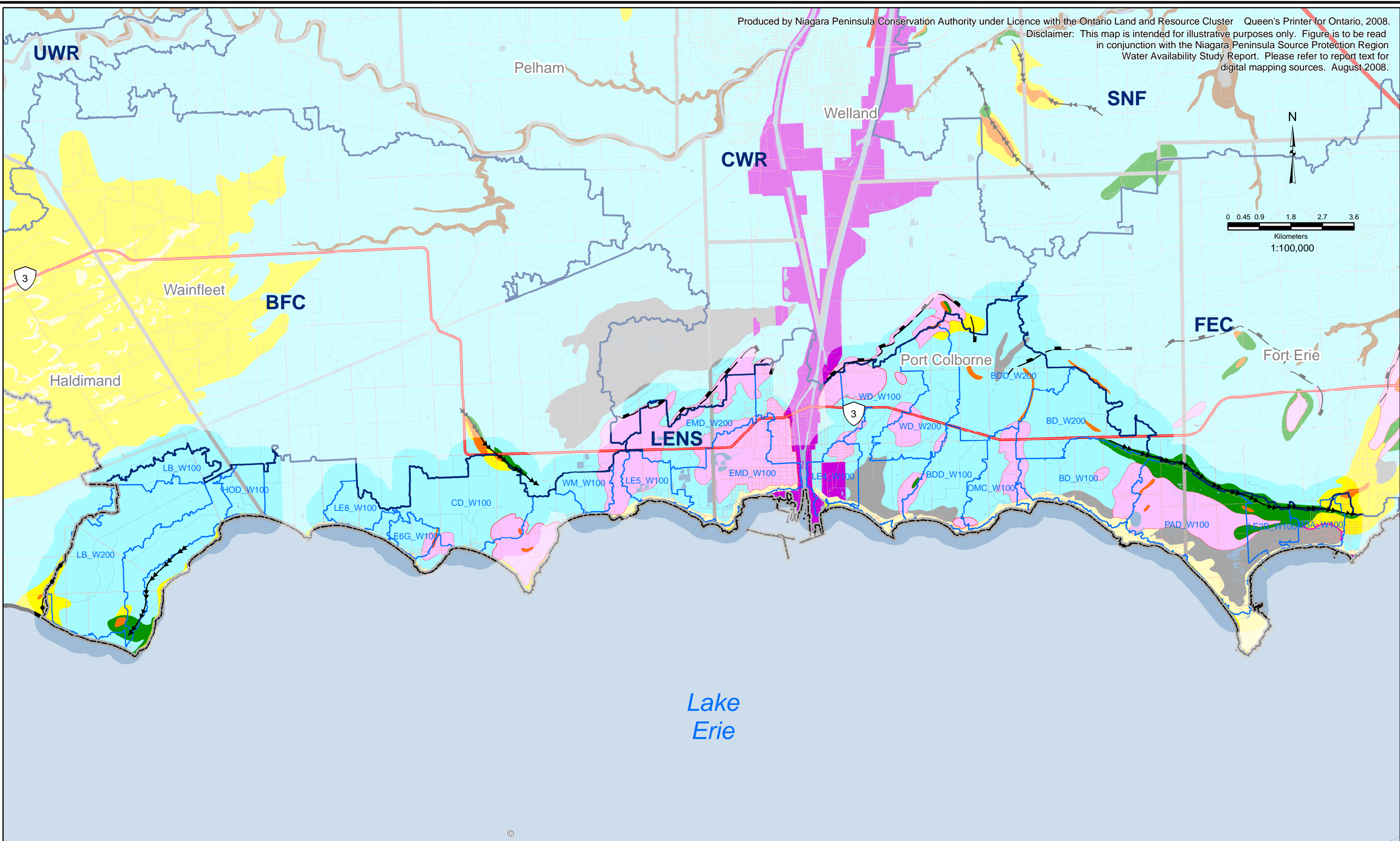
Overview Map

**Water Availability Study**

*Figure 2.14. LENS Soils*

All Frames: North American Datum 1983, Universal Transverse Mercator 6° Projection, Zone 17N, Central Meridian 81° West.





**Legend**

- Extended Context Area
- SPR Boundary
- Municipal Boundaries
- NPCA WSPA
- HMS Model Basins
- Major Highways
- Highways
- Roads
- Ponds, Reservoirs, Lakes
- Rivers, Streams, Creeks
- Sand and Gravel Pit
- Quarry
- Abandoned Shore Bluff
- End Moraine
- Escarpment
- Ice Contact Slope
- Moraine Crest

- Paleozoic bedrock (outcrop)
- Stone-poor, carbonate-derived silty to sandy till
- Glaciolacustrine-derived silty to clayey till
- Ice-contact stratified deposits
- Glaciofluvial deposits
- Fine-textured glaciolacustrine deposits
- Coarse-textured glaciolacustrine deposits
- Coarse-textured glaciolacustrine deposits
- Older alluvial deposits
- Coarse-textured lacustrine deposits
- Eolian deposits
- Modern alluvial deposits
- Organic deposits
- Man-made deposits

BBA\_W100: Subcatchment ID

Overview Map

**DRINKING WATER SOURCE PROTECTION**

Niagara Peninsula Source Protection Region

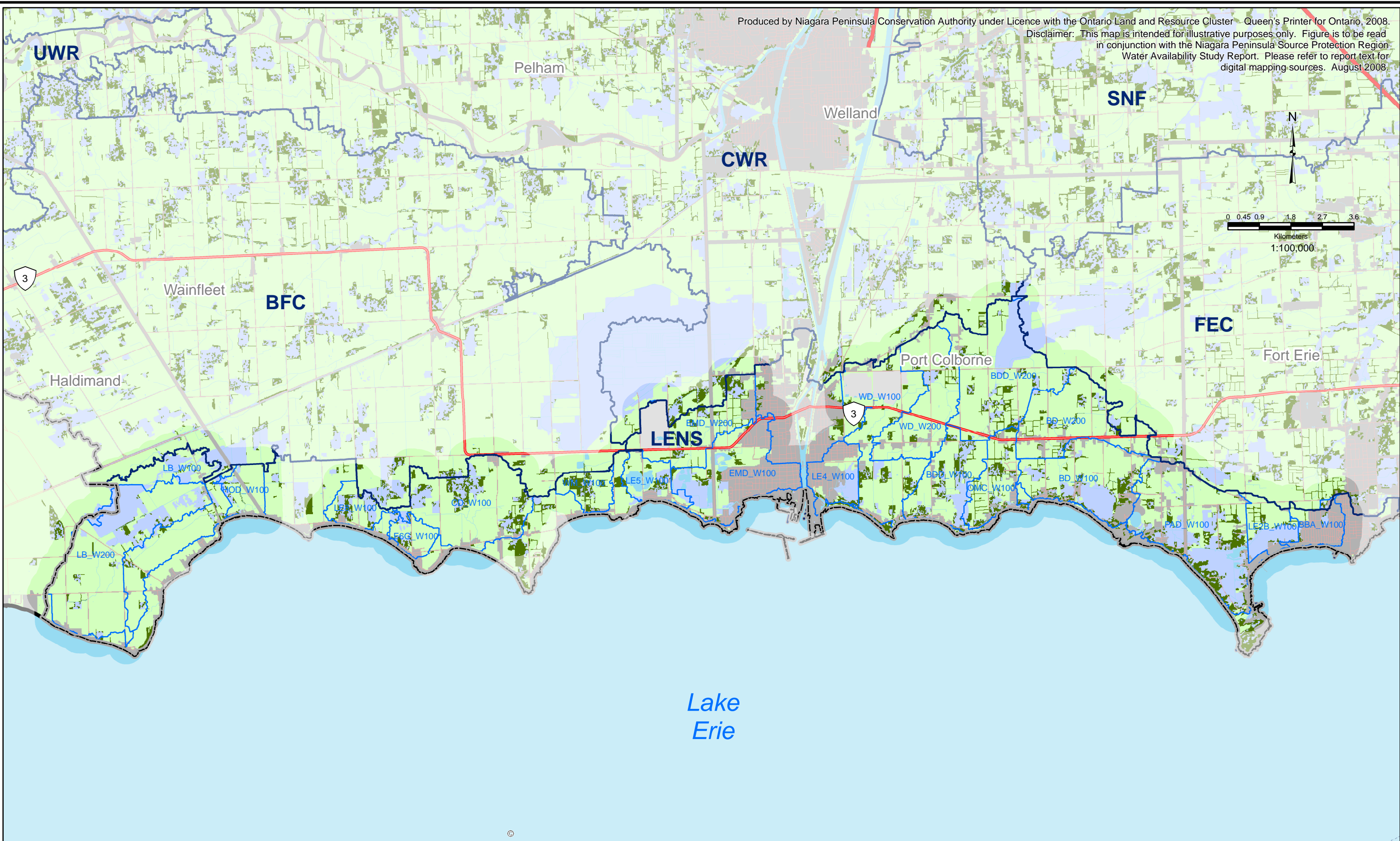
**Water Availability Study**

*Figure 2.15. LENS Surficial Geology*

All Frames: North American Datum 1983, Universal Transverse Mercator 6° Projection, Zone 17N, Central Meridian 81° West.

Ontario





**Legend**

- Extended Context Area
- SPR Boundary
- Municipal Boundaries
- International Boundary
- Major Highways
- Highways
- Roads
- Rivers, Streams, Creeks
- Ponds, Reservoirs, Lakes
- NPCA Watershed Planning Area
- HMS Model Subcatchments
- HMS Model Reaches
- BBA\_W100: Subcatchment ID

**Land Cover**

- Agriculture
- Wetland
- Built Up / Transportation
- Forest
- Extraction
- Shoreline
- Water
- Orchards
- Vineyards

**DRINKING WATER SOURCE PROTECTION**  
Niagara Peninsula Source Protection Region

**Water Availability Study**

*Figure 2.16. LENS Land Cover*

All Frames: North American Datum 1983, Universal Transverse Mercator 6° Projection, Zone 17N, Central Meridian 81° West.

Ontario

Overview Map



# Conceptualization of Hydrologic Processes in HEC-HMS

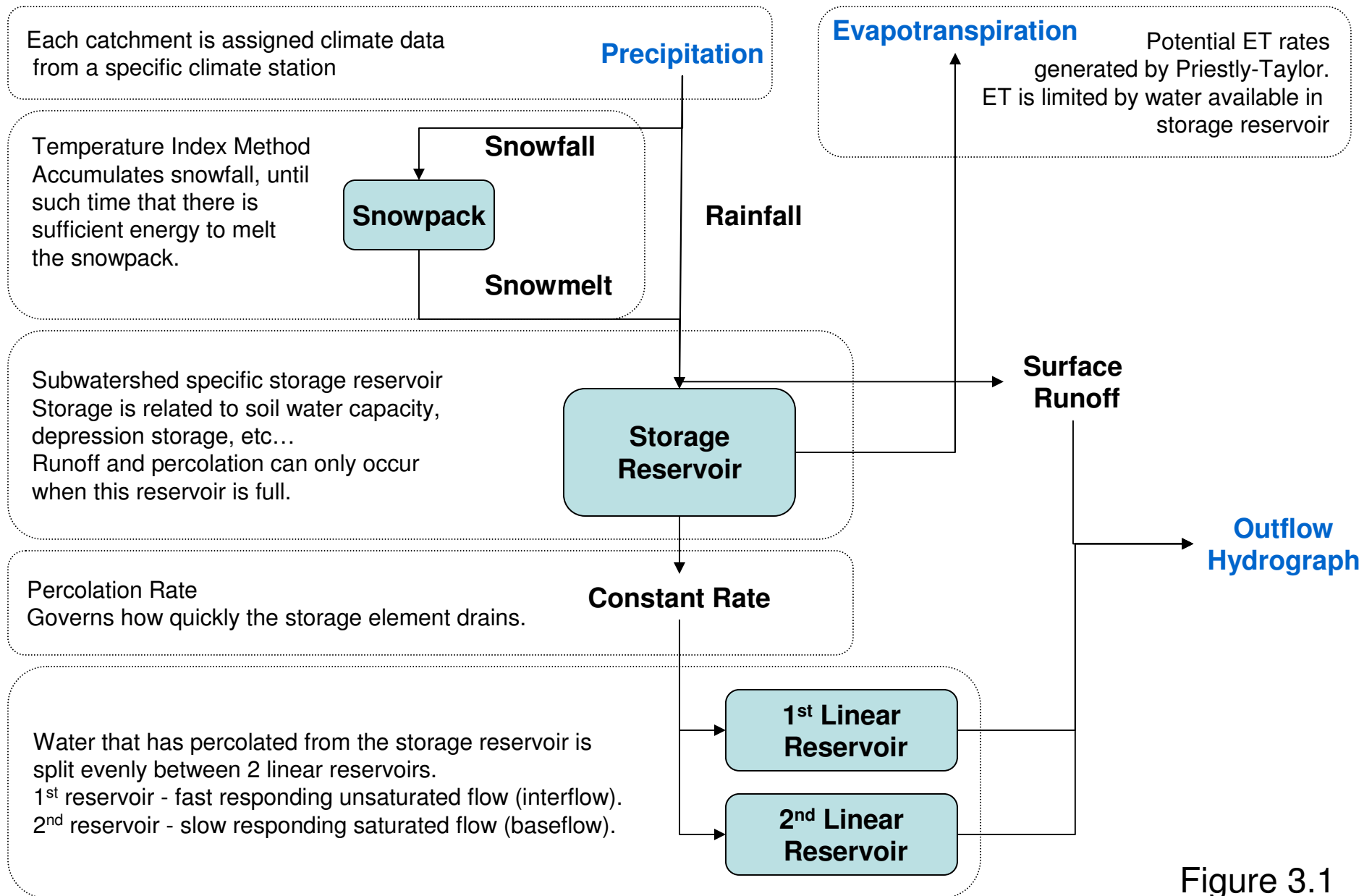
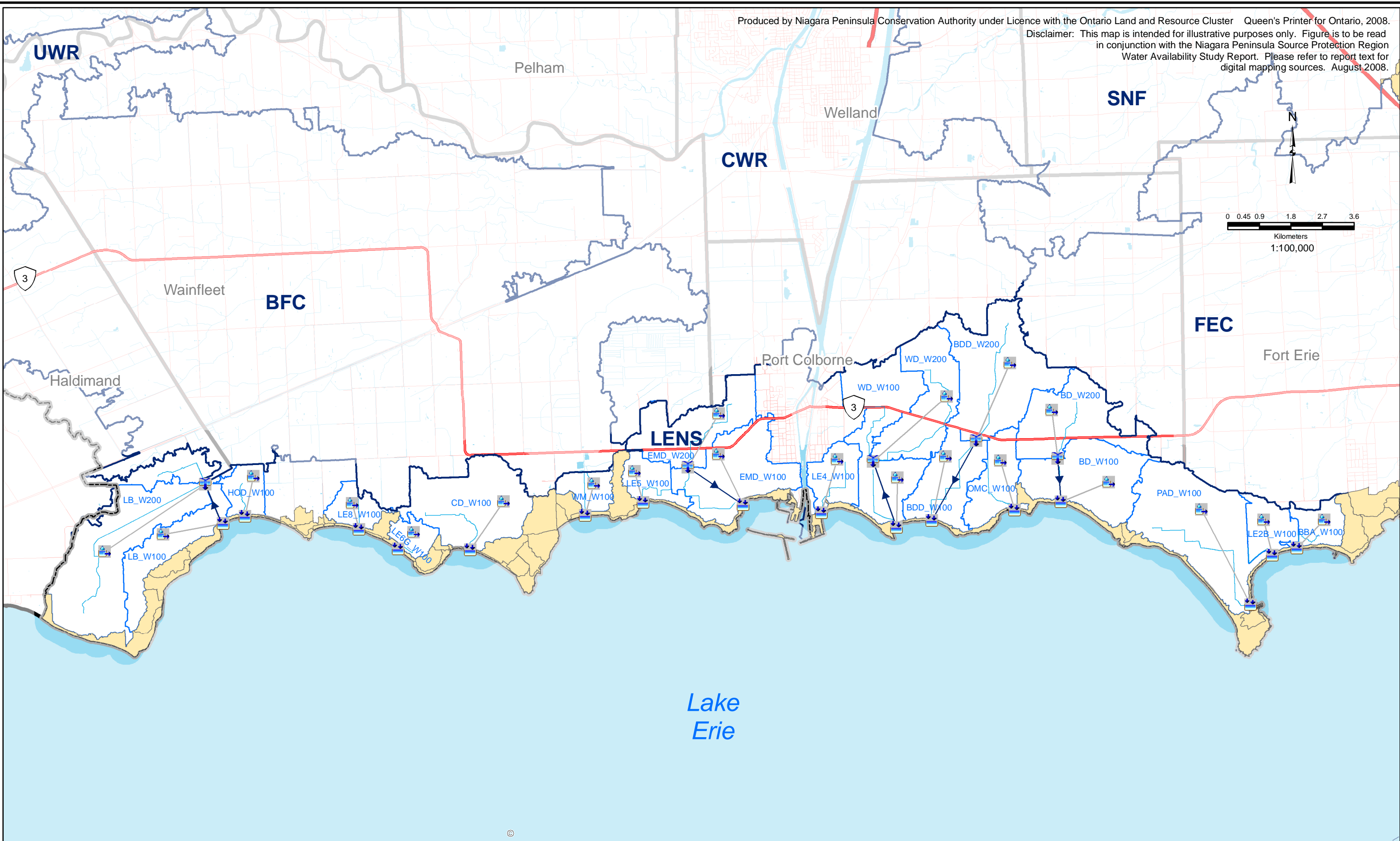


Figure 3.1



**Legend**

- Extended Context Area
- SPR Boundary
- Municipal Boundaries
- International Boundary

- Major Highways
- Highways
- Roads
- Rivers, Streams, Creeks
- Ponds, Reservoirs, Lakes

- NPCA Watershed Planning Area
- HMS Model Subcatchments
- HMS Model Reaches
- Non Modeled Areas
- BBA\_W100: Subcatchment ID

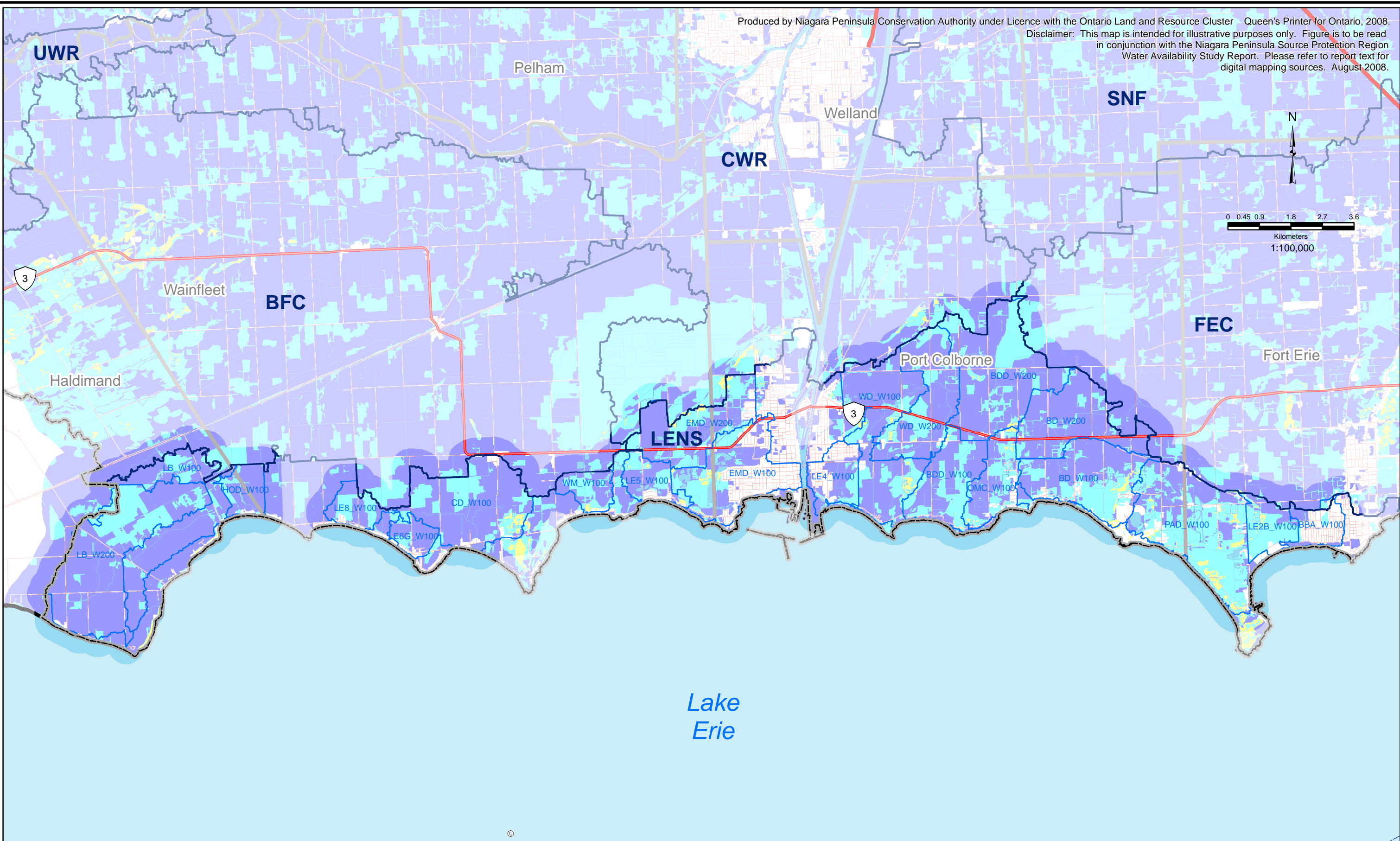
**HMS Legend**

- Diversion
- Junction
- Reservoir
- Sink
- Source
- Subbasin
- Basin Connector
- Reach

**Water Availability Study**

**Figure 3.2. LENS Subcatchments and HMS Schematic**

All Frames: North American Datum 1983, Universal Transverse Mercator 6° Projection, Zone 17N, Central Meridian 81° West.



**Legend**

- Extended Context Area
- SPR Boundary
- Municipal Boundaries
- International Boundary
- Major Highways
- Highways
- Roads
- Rivers, Streams, Creeks
- Ponds, Reservoirs, Lakes
- NPCA Watershed Planning Area
- HMS Model Subcatchments
- HMS Model Reaches
- BBA\_W100: Subcatchment ID

**CN Values (AMC II)**

- 20 - 40
- 40 - 60
- 60 - 80
- 80 - 100

**Overview Map**

**DRINKING WATER SOURCE PROTECTION**  
Niagara Peninsula Source Protection Region

**Water Availability Study**

**Figure 3.3. LENS Curve Number Values**

**All Frames:** North American Datum 1983, Universal Transverse Mercator 6° Projection, Zone 17N, Central Meridian 81° West.

Ontario

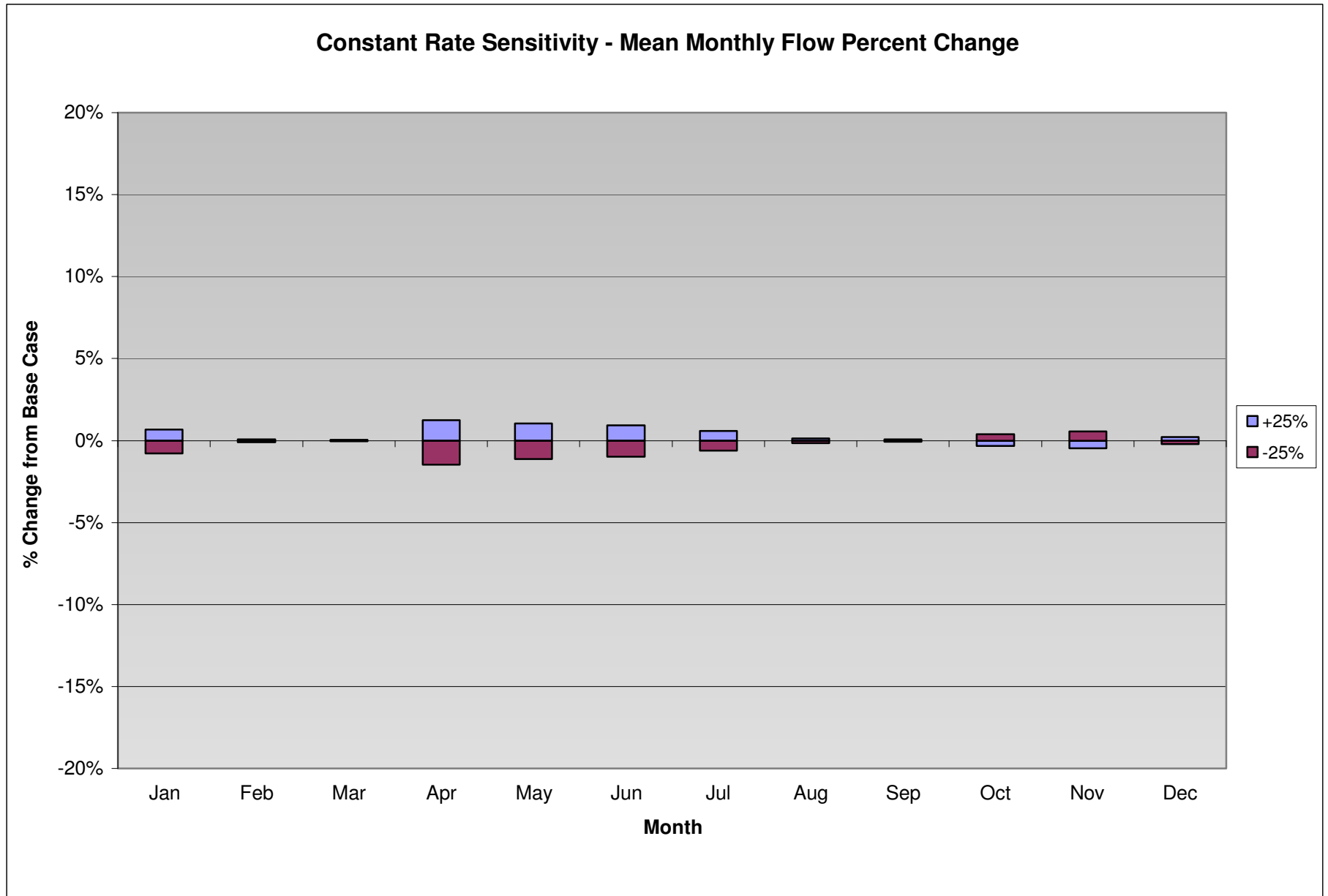


Figure 3.4  
Lake Erie North Shore

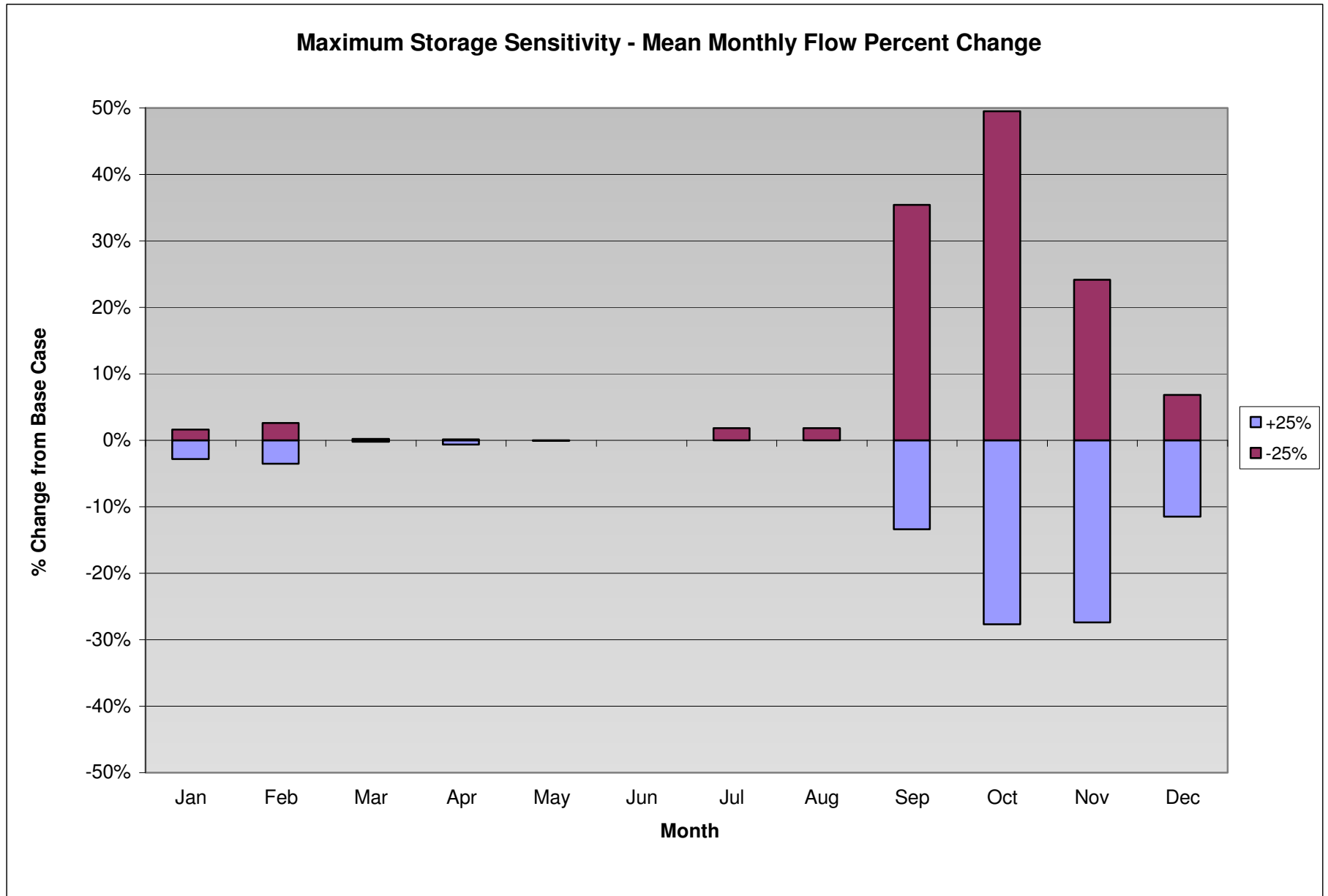


Figure 3.5  
Lake Erie North Shore

## **Appendix A**



# Snow Modeling Overview

Dr. Steven F. Daly

**USACE ERDC/CRREL**

Hanover, NH 03755



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## Snow Modeling can support our Snow Hydrology Goals

- **Snow Accumulation**
  - Estimation of the distribution of watershed snow water equivalent (SWE)
- **Snow melt (Ablation)**
  - Timing and magnitude of snowmelt



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# Snow Modeling Approaches

- Regression – Relate final reservoir volume with observed index of SWE volume
  - Based on historical record
  - Large uncertainty
  - Susceptible to climate change impacts
- Numerical models of snowpack physics – simple to complex



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## Numerical Snow Models

### Heat transfer from snowpack to environment

- Simulate each heat transfer mode (Complete energy balance)
  - Data intensive
  - Varies widely due to slope, aspect, vegetation, elevation, etc.
- Simplify heat transfer by considering only key meteorological parameters (temperature index)
  - Air temperature
  - precipitation



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## **Numerical Snow Models**

### **Representing snowpack physical properties**

- **Multi-layer snow packs**
  - Each layer with separate properties
    - Temperature, Density, Liquid water
- **Single Snow Layer**
  - Average snow properties
    - SWE, Cold Content, Liquid Water, etc



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### **Temperature Index Snow model**

- Often, complete energy budgets are difficult or impossible to estimate. A simpler method, based only on the air temperature, called the Temperature Index (degree-day) method has been developed. It has been widely used with good results.



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## Numerical Snow Models

- **Energy Balance**
  - Wind
  - Temperature
  - Water Vapor
  - Radiation (Net)
  - Precipitation
  - Advection (rain)
  - Ground
- **Temperature Index**
  - Temperature
  - Precipitation
- **Single layer snow**
  - SWE
  - Cold Content
  - Liquid water
- **Calibration required**
- **Detailed (layered) snow pack**



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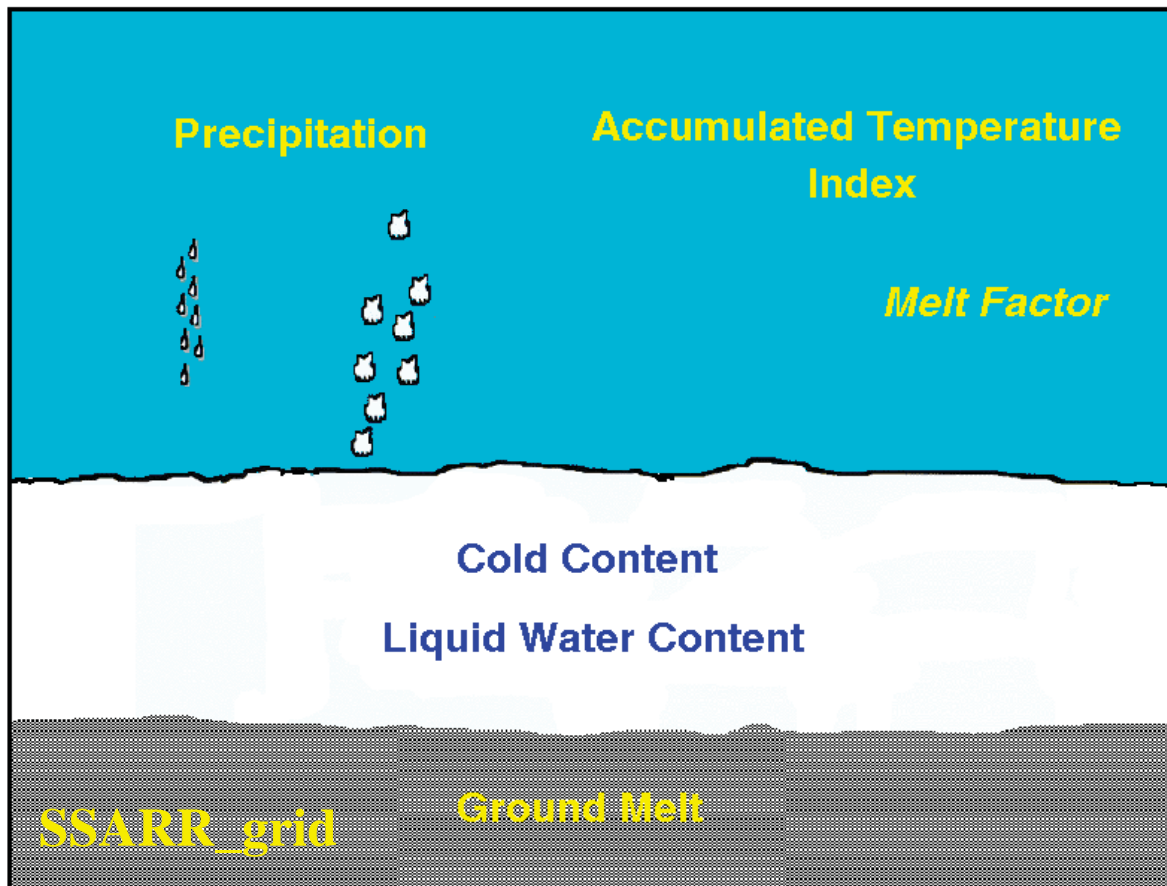
## Snowmelt Modeling in HMS

- **Streamflow Synthesis And Reservoir Regulation – SSARR – North Pacific Division – NWD**
  - Snow model, hydrology model, reservoir model for Pacific NW
- HEC-1 simple snow model
- SSARR snow model was made stand alone **SSARR\_grid**
- SSARR\_grid made into a distributed model **Distributed Snow Process Model – DSPM**
- **SSARR\_grid** added to HEC-HMS



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## Snowmelt Modeling in HMS

- Temperature Index
- Single snow Layer
  - Cold Content
  - Liquid water content
  - Ground melt



# Snowmelt Modeling in HMS

- **Watershed**
  - Uniform snow cover
    - Watershed temperature & precipitation
  - Elevation bands
    - Initial SWE
    - Temperature
- **Distributed**
  - Standard Hydrologic Grid 10m – 100km
  - Distributed air temperature & Precipitation



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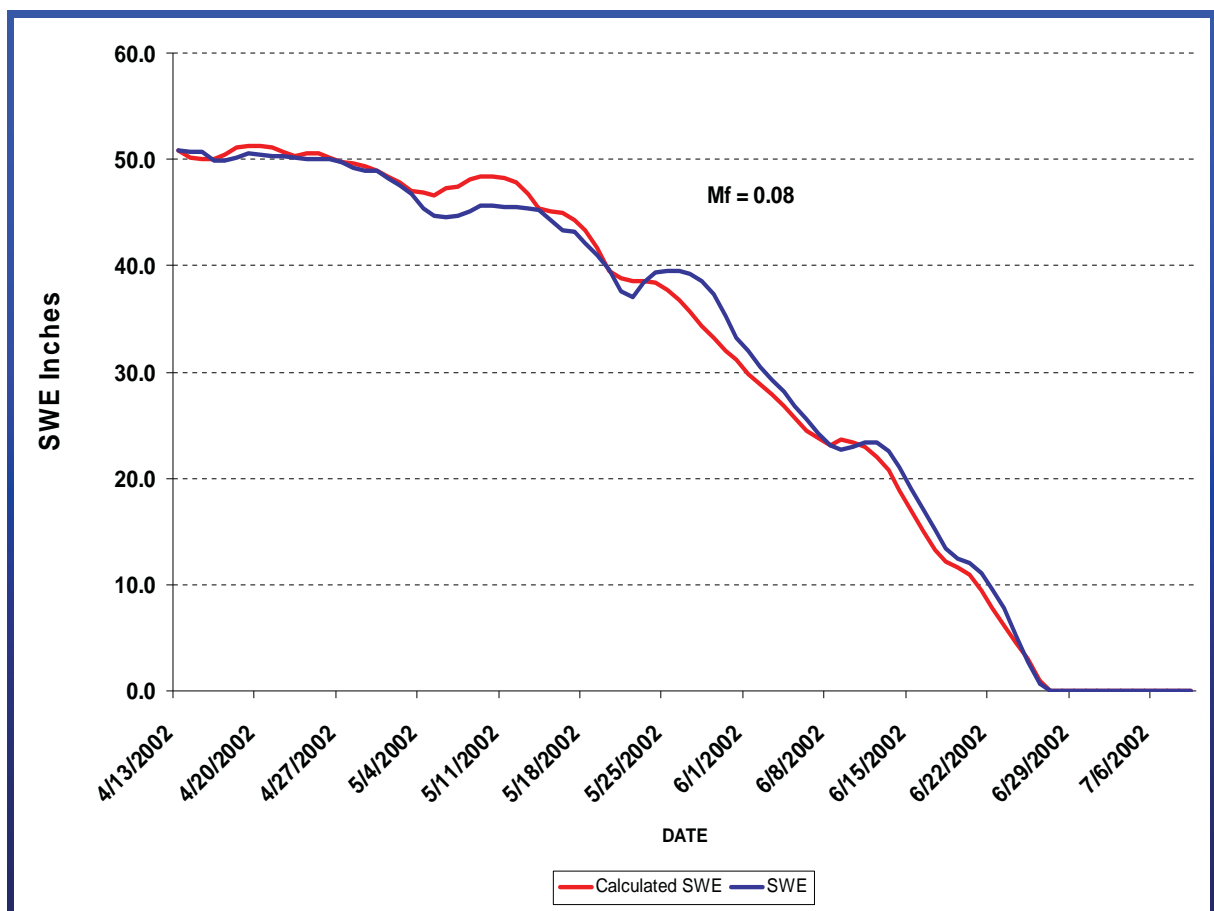
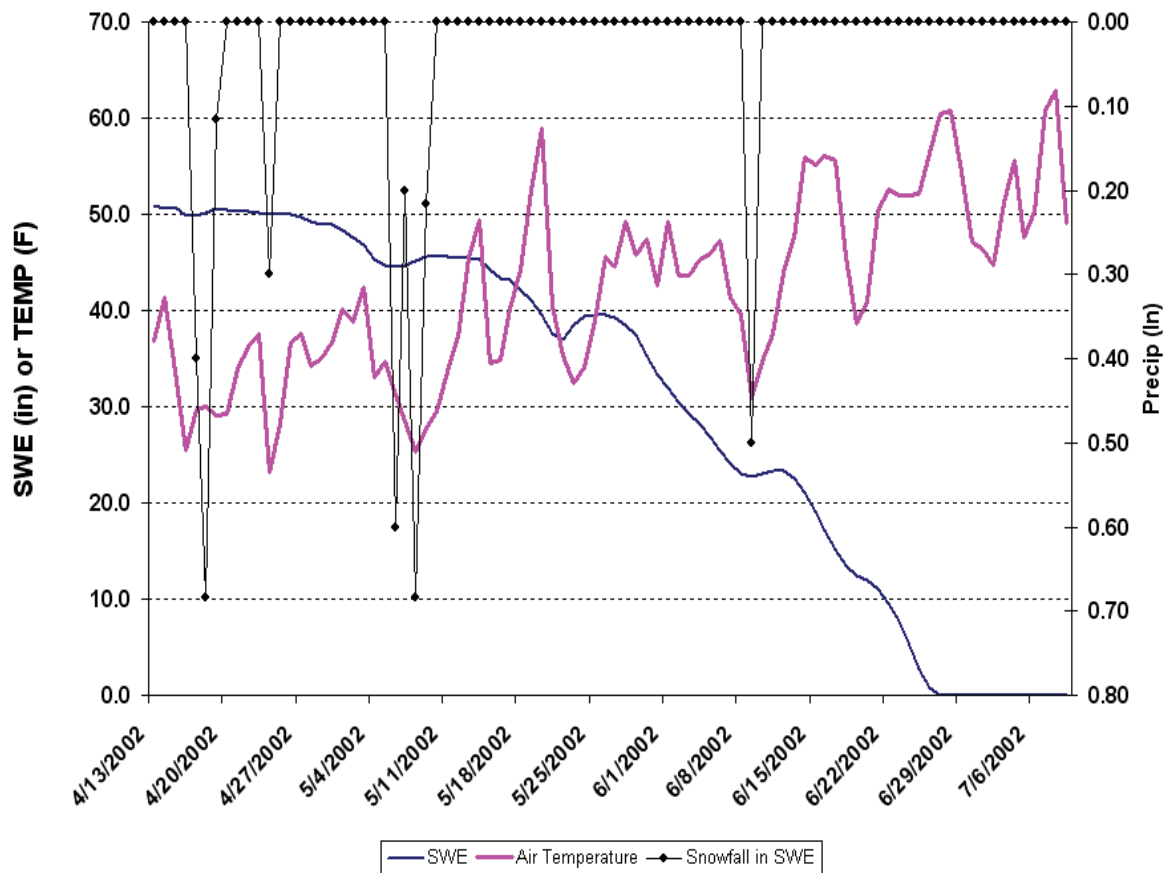
## Temperature Index Snow Model

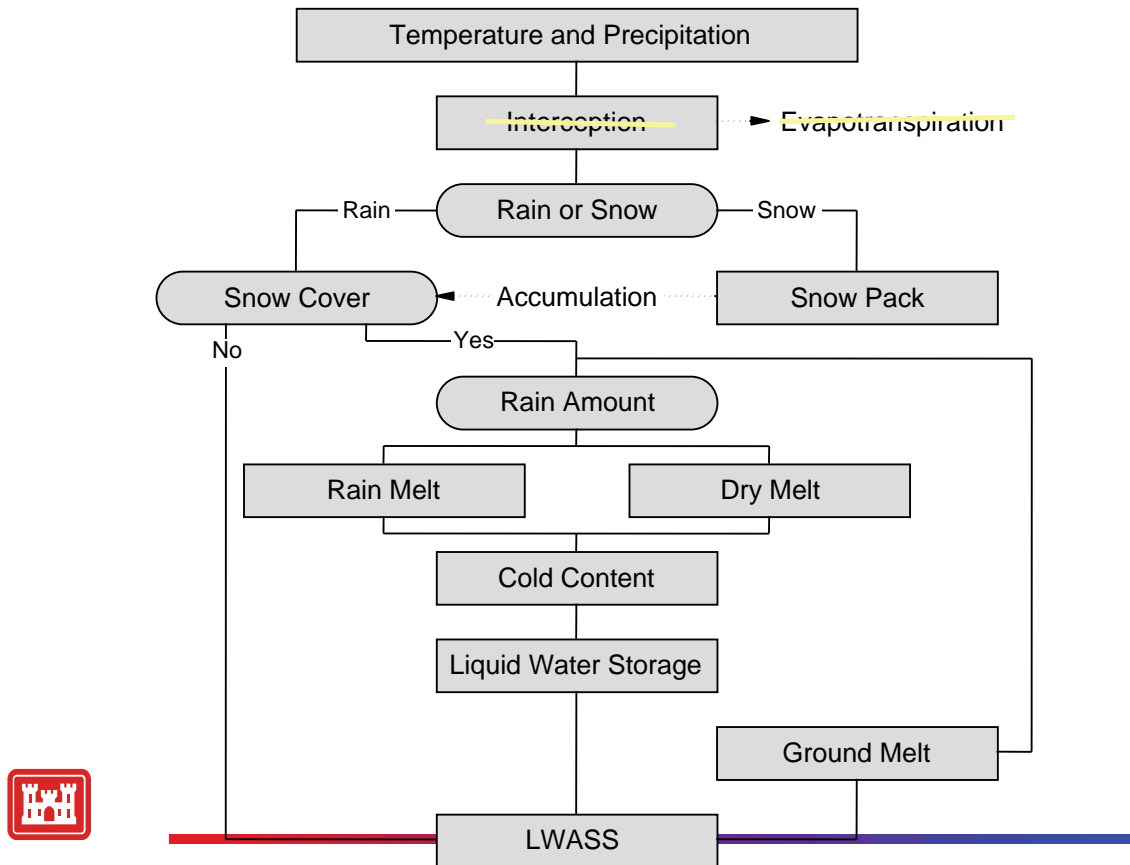
- Snow Melt is estimated as
  - $L_s = M_f (T_a - T_{base})$
  - $M_f$  serves as an index of the total heat transfer at the snow surface which includes long wave, short wave, latent heat, and sensible heat transfer. Also strongly influenced by the wind speed, aspect, slope, vegetation, etc.
  - $M_f$  can be a constant, or set as a function of the accumulated thawing days or set as a function of the month of year.



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Daly-011

## Given precipitation and air temperature for time step

### Discriminate between rain and snow using PX temperature

- **PX** The discrimination temperature between precipitation falling as rain or snow. When the air temperature is less than the specified PX temperature, any precipitation is assumed to be snow. When the air temperature is above the specified PX temperature, any precipitation is assumed to be rain. This discrimination temperature is usually one to two degrees above freezing.



## Given precipitation and air temperature for time step

Discriminate between melt and non-melt using  
Base temperature

- **Base Temperature.** The difference between the base temperature and the air temperature defines the temperature index used in calculating snowmelt. The meltrate is multiplied by the difference between the air temperature and the base temperature to estimate the snowmelt amount. If the air temperature is less than the base temperature, then the amount of melt is assumed to be zero. Typically, the base temperature should be 32F (0C) or close to it.



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## Given precipitation and air temperature for time step

Discriminate between melt and non-melt using  
Base temperature

- $T < T_{\text{base}}$  - No Melt  
- SWE accumulates  $T < P_X$ ;  $P > 0$
- $T > T_{\text{base}}$  - Melt



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$$T > T_{\text{base}} - \text{Melt}$$

Discriminate between rain melt and dry-melt using rain rate limit

- **The rain rate limit.** The discrimination rain rate in inches/day between dry melt and wet melt. The wet meltrate is applied as the meltrate when it is raining at rates greater than the rain rate limit. If the rain rate is less than the rain rate limit, the meltrate is computed as if there were no precipitation.



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## Meltfactor

$$L_s = M_f (T_a - T_{\text{base}})$$

- **Meltfactor** can be constant or variable with time. A variable **meltfactor** recognizes that as snowpack matures the heat transfer rates change AND/OR different components of heat transfer change in importance.



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# Meltfactor

- **ATIMR** The seasonal variation of melt rate is indexed by an antecedent temperature function (ATIMR). The initial melt ATI should be thought of as similar to “the accumulated thawing degree days.” This antecedent temperature function allows the melt rate to change as the snowpack matures and ages.
- **ATIMR\_Die\_away** Coefficient for updating the antecedent temperature indexATI\_MeltRate\_out. Typical value is .98



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## ATIMR Algorithm

- $ATIMR\_out = (ATIMR\_in * (ATIMR\_Die\_away^{**} days) ) + ((Temperature - T_{base}) * days)$
- IF(ATIMR\_out .LT. 0.) THEN ATIMR\_out = 0.
- Essentially, ATIMR accumulates as long as  $T > T_{base}$



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# Melt Factor

- ATI-melt function. Melrates associated with ATIMR values in inches per degree-day. Typical range of values is 0.015-0.055. The pattern must be entered in the **Paired Data Manager**.
- Melt pattern. An alternate method of entering the melrates as a function of the month of year. The pattern must be entered in the **Paired Data Manager**.



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## Typical Melt Factors

- 0.04-0.08 inches/F-day
- 1.8-3.7 mm/C-day
  - Average daily temperature
  - No rain
  - Not heavily forested
  - No extreme conditions – high winds, etc
- Thin ephemeral snowpacks that melt out in a very short time may have a constant melt factor



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## Temperature Index Snow Model

- Rain melt
  - Snow melt that occurs when the air temperature is above the snow/rain temperature and the precipitation rate is significant. Rain is assumed to fall at the air temperature
  - Melt from rain and condensation of water vapor in the snowpack.



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$$T > T_{\text{base}} - \text{Melt}$$

- **SWE** > 0
- Potential melt is estimated using temperature index and melt factor
- Melt is first applied to **Cold Content**. When **Cold Content** is zero.....
- Melt is then applied to **liquid water storage**. When **maximum liquid water storage** is reached .....
- **LWASS** is generated. **SWE** is reduced.



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# Cold Content

- **Cold Content**
  - Heat required per unit area to raise temperature of snowpack to 32°F (0°C)
    - $\text{Snow Depth} \times \text{Snow Density} \times \text{Heat Capacity of Snow} \times (\text{Temperature below freezing})$
  - Usually expressed in as a negative number equivalent to inches of frozen water
    - $\text{Cold content} / (\text{Density of water} \times \text{latent heat of water})$



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## Cold Content -Parameters

- **ATICC**. The **ATICC** is an index to the snow temperature near the surface. It is calculated assuming an approximation to the transient heat flow equations. This value is used to estimate the cold content of the snow. It should be set to the approximate snowpack temperature if known. If not known, it can be set to 32F (0C).
- **ATICC\_die\_away**. Coefficient for updating the antecedent temperature index **ATI\_ColdContent\_out**. Typical value .84



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## ATICC Algorithm

- $ATICC\_out = ATICC\_in + (ATICC\_Die\_away^{**} (24./TimeStepHours) ) * (Temperature - ATICC\_in)$
- IF(ATICC\_out .GT. T<sub>base</sub>) THEN ATICC\_out = T<sub>base</sub>



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## Cold Content -Parameters

- ATI cold function. Cold rates associated with the ATIMR values, in inches per degree-day. Typical range of values is 0.010-0.025. The pattern must be entered in the Paired Data Manager.
- $interval\_Cold = (Temperature - ATICC\_out) * coldRate / 24.$
- $ColdContent\_out = ColdContent\_in + interval\_Cold * TimeStepHours$



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## Cold Content-Parameters

- **ATICC\_Snow\_MAX.** Precipitation rate, inches per hour. If the precipitation rate exceeds **ATICC\_Snow\_MAX**, the antecedent coldness index ATICC is set to the temperature of the precipitation (or the base temperature, which ever is lower) If the precipitation rate is less than **ATICC\_Snow\_MAX**, ATICC is computed as an antecedent index. Typical value is (.8 inches/day)



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## Liquid Water Storage -Parameters

- **Maximum liquid water capacity.** The maximum liquid water capacity specifies the amount of melted water that must accumulate in the snowpack before liquid water becomes available at the soil surface for infiltration or runoff. Typically, the maximum liquid water held in the snowpack is on the order of 3%-5% of the SWE, although it can be higher. Liquid water can persist in the snow only if the snowpack temperature is at 32F (0C); at which point the cold content is zero.



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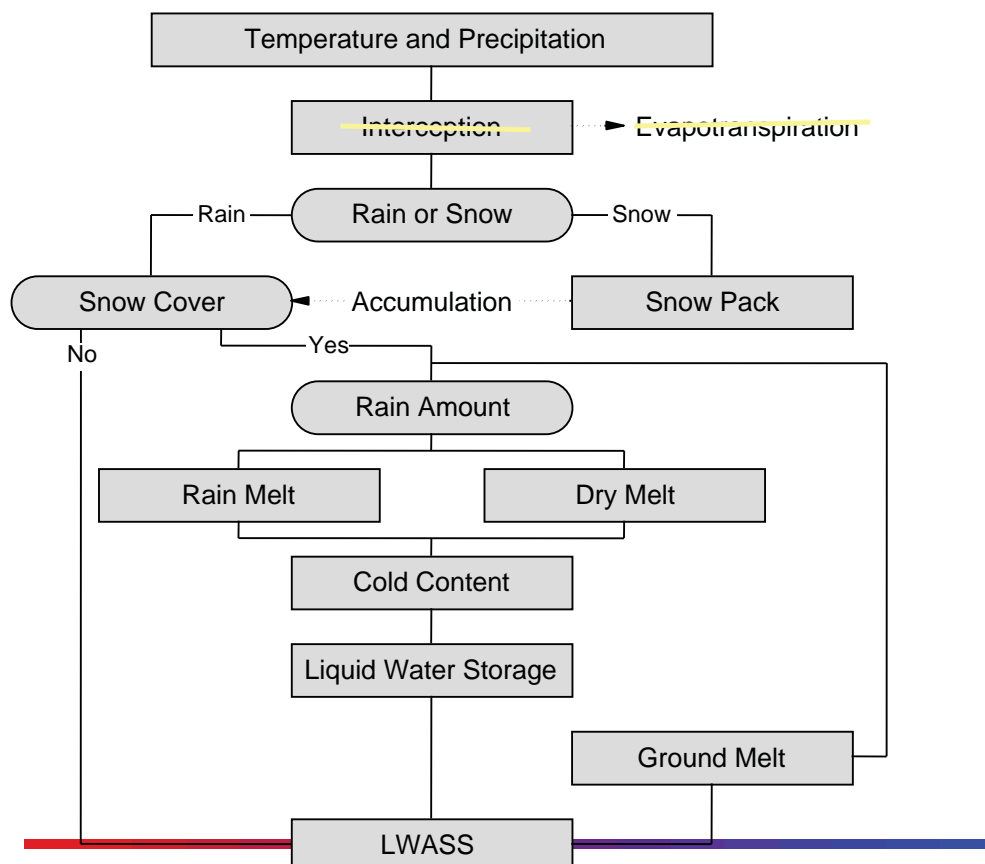
# Ground Melt

- Snow melt that occurs due to heat from the ground beneath the snowpack. Almost always set to zero, especially for relatively shallow, seasonal snow covers (SWE<12 inches)



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# Entering Snow Parameters in HMS

Temp Index

PX Temperature (DEG F)	35.69
Base Temperature (DEG F)	32.6
Wet Meltrate (IN/DEG F-DAY)	0.1912
Rain Rate Limit (IN/DAY)	0.02
ATI-Meltrate Coefficient:	0.98
ATI-Meltrate Function:	snowmelt_atimelt_9697
Meltrate Pattern:	--None--
Cold Limit (IN/DAY)	0.04
ATI-Coldrate Coefficient:	0.90
ATI-Coldrate Function:	snowmelt_aticold_9697
Water Capacity (%)	5.0
Groundmelt Method:	Fixed Value
Groundmelt (IN/DAY)	0



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# Entering Paired Data

Paired Data Table Graph

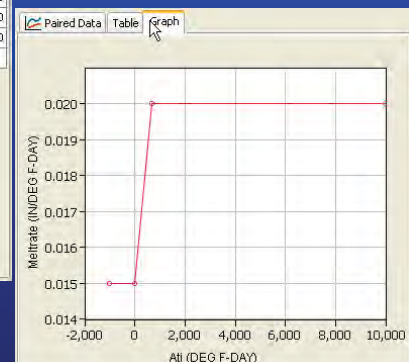
Name: snowmelt\_atimelt\_9697

Description:

Data Source: Manual Entry

Units: DEG F-DAY : IN/DEG F-DAY

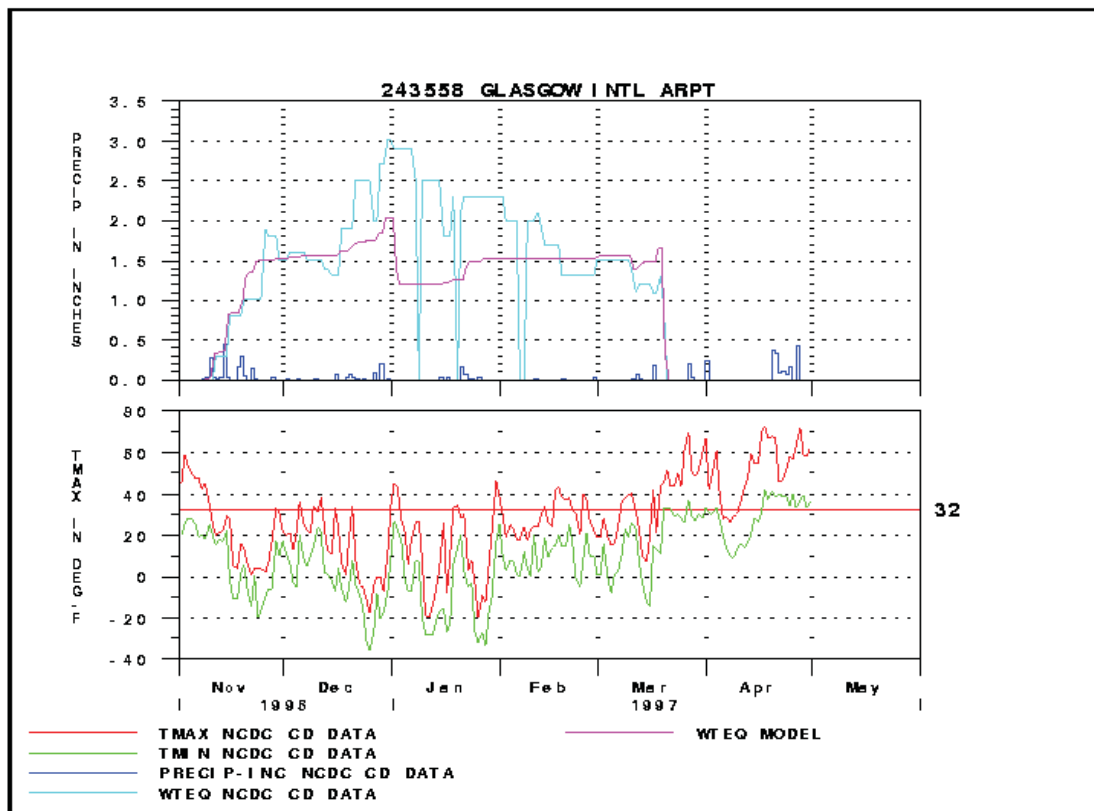
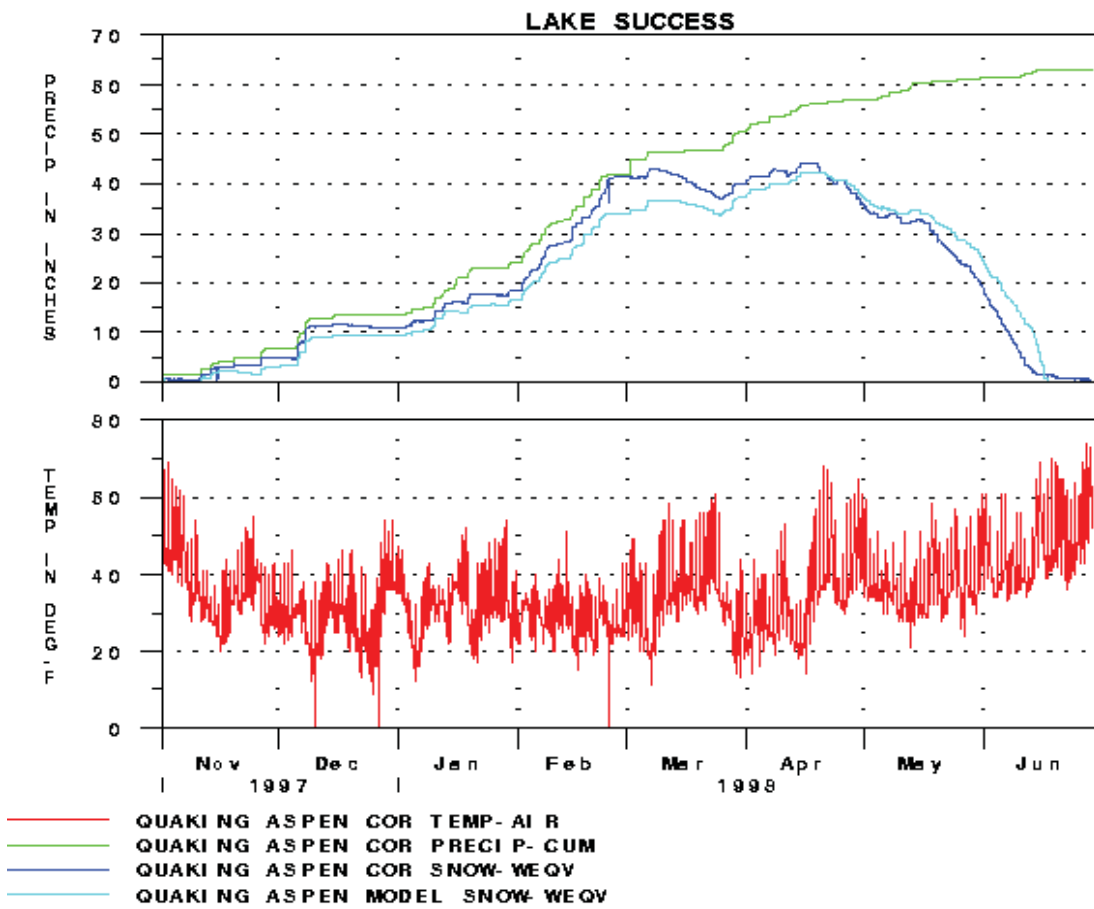
Ati (DEG F-DAY)	Meltrate (IN/DEG F-DAY)
-1000.00	0.0150000
0.00	0.0150000
673.20	0.0200000
9999.00	0.0200000



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# Summary -Terms

- Temperature Index Snow Model
- Complete Energy Budget Snow Model
- Snow Water Equivalent (SWE)
- Cold Content
- Snow Liquid Water Storage
- PX
- Base Temperature
- Rain Rate Limit
- Melt factor
- ATIMR
- Rain melt
- LWASS
- ATICC
- Maximum Liquid Water Capacity
- Ground melt



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## **Appendix B**

## **GeoHMS Processing**

With the wide availability of Geographic Information System (GIS) data layers for the study area, a good proportion of the model set up and parameterization was achieved using ArcHydro, and HEC-GeoHMS, which are publicly available ArcGIS extensions. The primary GIS data sources for the study consisted of a high resolution (3m) Digital Elevation Model (DEM) and stream layer produced by the Niagara Peninsula Conservation Authority; Southern Ontario Land Resource Information System (SOLRIS) land cover layer (Version 2) produced by the Ontario Ministry of Natural Resources (2006) and soils layer obtained from the Ontario Ministry of Agriculture and Food (Niagara Region 1989, Haldimand County 1984, Hamilton-Wentworth 1965). These layers were further processed in a GIS to produce the required HEC HMS model inputs.

### **Primary GIS layers processing**

The digital soil layer from OMAF aggregates county-based soils maps onto a seamless and standardized product. Included in the product are three tables (soil component, soil name file, soil layer file) that can be relationally linked to the spatial data. The soils component attribute table maintains 17 variables for each soil component, which includes a multifaceted variety of soil and soil-related data. Based on their ability to drain precipitation inputs, soils can be categorized into 4 Hydrologic Soil Groups (HSG) ranging from A to D (Appendix C). HSG A soils have the highest infiltration rates, while HSG D soils have the lowest infiltration rates. Each map unit polygon in the component table can be comprised of up to 2 different (HSG) that contribute to the total area of the map unit. The area occupied by each HSG is represented as a percent of the map unit. However, there is no specific information on the location of the individual HSB within any polygon. For example, a BRADY soil series map unit can be comprised of 70 % HSG (A) and 30 % HSG (B). To avail ourselves with the most detailed information for the study, all HEC-HMS model parameter calculations dependent upon HSG were percent-weighted based on HSG.

SOLRIS is a primary data layer that provides a comprehensive landscape level inventory of land use for the study area. The SOLRIS land use classes for the study area are provided in Appendix II. The version in use in this study groups agricultural crop such as corn, grains, wheat, alfalfa, and soybeans into broad agricultural land use classes such as monoculture and annual crops. The level of detail was insufficient for the application of crop specific coefficients required as input for the evapotranspiration calculations in the Priestly-Taylor method used in HEC-HMS. Statistic Canada field crop data (percent by type) at the census consolidated subdivision level was used as a weight to calculate composite crop coefficient values for the SOLRIS agricultural land use classes.

## Derivative GIS layer processing

Setting up the model required the user to select methods to simulate infiltration losses, transform excess precipitation into runoff, represent baseflow contribution to subbasin outflow, and simulate flow in open channels. Each method requires one or more parameters that describe the state of each catchment and allow the model to simulate the hydrologic processes. Table 1 shows the simulation methods used and the required GIS derived parameters. With the exception of the crop coefficient, channel bottom width, and channel side slope all vector products were converted to a 15m raster-based product.

Table 1. HEC-HMS simulation methods used

Category	Method	Required GIS Parameter
Loss	Deficit and Constant	Soil Water Holding Capacity Soil maximum infiltration rate Crop Coefficient (Priestly Taylor)
Transform	SCS Unit Hydrograph	CN, Initial Abstraction, Lag Time
Baseflow	Linear Reservoir	N/A
Routing	Muskingum-Cunge	Channel Side Slope, Channel Bottom Width

### Constant infiltration rates

The deficit and constant loss method assumes that the soil has a set maximum infiltration rate approximated by the saturated soil hydraulic conductivity. Using the information in Appendix C, average maximum infiltration rates were assigned to each polygon in the soil layer based on their HSG.

### Soil water holding capacity

In the simulation, the soil is also assumed to have a fixed water holding capacity, typically affected by the active rooting depth of vegetation and HSG. The soil water holding capacity layer was built by intersecting the SOLRIS land cover and the OMAF soils layers and by assigning soil water holding capacity values from Appendix C to each unique combination of land cover class and soil HSG. OMAF polygons mapped as urban were not included in the procedure; imperviousness is addressed later.

### Crop Coefficient

In the deficit constant method, water is removed from the soil to simulate evapotranspiration. In the model, evapotranspiration was calculated through the Priestly-Taylor. This method requires the use of crop coefficient  $K_c$ , which indicates the ratio of crop potential and grass reference evapotranspiration. Land use layers were created for each day of the year and daily crop coefficients from Appendix D were assigned to the land use classes.

## **CN grid**

CN values are used in the calculation of CN lag time for the SCS Unit Hydrograph transform method. The factors influencing CN values are land cover type, HSG, and Antecedent Soil Moisture Condition (AMC). AMC is an estimate of soil water content prior to the beginning of the simulation period. AMC I reflect soils that are dry but with water content not below the wilting point. AMC II reflects soils having average soil water content, and AMC III reflects soils that have experienced rainfall in the five days previous to the beginning of the simulation period. CN values in the study area were assumed to reflect average soil water content. The CN layer was built by intersecting the SOLRIS land cover and OMAF soil layer and by assigning CN values from Appendix II to each unique combination of land use class and HSG. CN values were not assigned for built-up impervious, built-up pervious, transportation or open water SOLRIS land cover types.

## **Impervious**

HEC-HMS considers an impervious surface as an area in a watershed for which all contributing precipitation runs off, with no infiltration, evaporation, or other volume losses. This surface was built by assigning percentages of 100, 100 and 50 to the transportation, built-up impervious and built-up pervious polygons respectively. All other polygons were assigned a value of 0.

## **Channel width and side slope**

In the model the traditional Muskingum-Cunge routing method was used assuming trapezoidal channel geometry. The method requires the input of channel bottom width, channel side slope, and channel manning roughness coefficient. Channel width for each of the routing reaches was estimated by digitizing cross sections across the channel assuming that the extent (i.e. width) of the water surface on the digital air photos roughly approximates the width of the channel bed. Channel side slope was approximated by digitizing two points at the end of each digitized channel width cross sections using a 2m resolution DEM as a guide. Appropriate channel Manning roughness coefficients from Appendix C were assigned to channel routing reaches following visual channel stream bed condition assessment from 10 and 20 cm resolution digital air photos.

The GIS approach to building a HEC HMS model is generally done in two phases: the terrain processing phase and the model parameterization phase. These are described below.

## **Terrain Processing**

The terrain processing phase requires a terrain model that is hydrologically correct. The terrain is created by integrating a fully connected dendritic stream network into a DEM. This process can be summarized as follows: 1) rasterization of the vector stream network to the same resolution as the DEM, 2) reclassifying the rasterized stream

network by assigning an arbitrary elevation (i.e. 50) value to the cells of the stream network. 3) Subtracting the reclassified grid from the DEM. This has the effect of decreasing the elevation of all DEM cells underlying the stream network by the aforementioned elevation value (50 m). 4) Filling the DEM sinks, thus ensuring that no water is trapped in DEM depressions and that all DEM cells drain to the outlet.

The next steps are the creation of two terrain derivatives from the filled DEM and a series of processing steps to delineate the watershed subbasins. These are performed using the ArcHydro Tools and are briefly outlined below:

- 1) Flow direction grid: Shows the orientation of the DEM cell's to its neighbour steepest down slope.
- 2) Flow accumulation grid: Indicates the number of upstream cells draining to each DEM cell.
- 3) Stream definition: The flow accumulation grid was then used to produce a synthetic stream network by applying a suitable area threshold value. The area threshold value indicates the minimum upstream area required to initiate a synthetic stream network. A 500 ha threshold value was selected so that average catchment size in the study was between 5 and 10 km<sup>2</sup> and lag time for most of these catchments greater than 2 hrs.
- 4) Stream segmentation: The synthetic stream network is divided at the synthetic stream network confluences. All cells belonging to each stream segment are assigned a unique value.
- 5) Catchment Grid Delineation: This step generates a grid representation of a subbasin for each stream segment. All cells belonging to a subbasin are assigned a unique number.
- 6) Catchment Polygon Processing: This step converts the grid representation of the subbasin to a vector representation
- 7) Drainage line processing: This step converts the grid representation of the segmented synthetic stream network into a vector representation.
- 8) Adjoint catchment processing: This step aggregates the upstream subbasins at every stream confluence. This step has no hydrological significance and is done to increase the performance of the point delineation process.

## **Hydrologic Model Creation**

Once the terrain processing is completed, the data required to support model creation and model parameterization can be extracted for the study area using the HEC-GeoHMS tools. The main steps are HEC-HMS model set up, Watershed subbasin



boundary refinement and model parameterization. These steps are briefly outlined below.

### **HEC-HMS model set up**

An HEC-HMS project is created by specifying the outlet point of the study area. During the project generation, the following datasets are created.

- 1) Filled DEM: Hydrologically corrected DEM.
- 2) Raw DEM: Original DEM.
- 3) Flow direction grid
- 4) Flow accumulation grid
- 5) Stream grid: Synthetic stream network in grid representation.
- 6) Stream link grid : Synthetic stream network segmented at confluences.
- 7) Catchment grid: Subbasin extents in grid representation.
- 8) Catchment polygon: Represent the extracted subbasin extent in vector format.
- 9) Rivers: Represents the synthetic stream network in vector format.
- 10) Project point: Represent the watershed outlet.

### **Watershed subbasin boundary refinement**

Once a HMS project has been set up, the watershed subbasin boundaries can be revised. This was done mainly by combining and by subdividing subbasins. Subbasins larger than the 10 km<sup>2</sup> threshold were subdivided at hydrologic control points such as road crossings where changes in flow regime were most likely to occur. Subbasins smaller than the 5 km<sup>2</sup> threshold were merged to adjacent subbasins.

### **HEC-HMS model parameterization**

Once a satisfactory watershed subbasin layout was defined, the next step was the parameterization of the model. Model parameterization is done in two phases. These are the extraction of the watershed physical parameters and the extraction of the model hydrologic parameters. These steps are outlined below.

## **Extraction of the watershed physical parameters**

The HEC-GeoHMS toolbar can compute several topographic related characteristics of streams and subbasins. These include the following:

- 1) River length: Computes the length of each river features.
- 2) River slope: Extracts the upstream and downstream elevation of each stream segment and calculates the river slope.
- 3) Basin slope: Computes the average slope for each subbasin based on an input slope grid that was generated from a 2m DEM.
- 4) Longest flow path: Computes the longest flow path for each subbasin.
- 5) Basin centroid: Calculates the centroid as the center of the longest flow path within the subbasin.
- 6) Centroidal flow path: Calculates the flow path from the projected point of the subbasin centroid on the longest flow path to the subbasin outlet, along the longest flow path route.

## **Extraction of the watershed hydrologic parameters**

In addition to extracting watershed physical parameters the HEC-GeoHMS tool can also extract a number of hydrologic inputs for the HEC-HMS model. The steps involved are briefly outlined below:

- 1) Selection of HMS processes: In this step, the loss, transform, and baseflow type methods for the subbasins and the routing method for the rivers as outlined in Table 1 were selected.
- 2) Subbasin curve number: Computation of the average subbasin curve number.
- 3) Muskingum-Cunge parameters: Allows the selection of the Muskingum-Cunge channel shape. A trapezoidal channel shape was selected for this study.
- 4) CN lag: Computation of the CN lag for each subbasin.

A number of required hydrologic parameters could not be extracted using the HEC-GeoHMS tools. These model parameters were manually calculated or extracted through the development and application of Python scripts. These parameters are listed below:

- 5) Basin imperviousness: Computation of the average subbasin imperviousness (%).

- 6) Basin maximum infiltration rate: Computation of the average subbasin maximum infiltration rate (mm/hr).
- 7) Basin water holding capacity: Computation of the average subbasin maximum water holding capacity (mm).
- 8) Basin Initial Abstraction: The initial abstraction defines the amount of precipitation that must fall before runoff is observed. This was calculated using the following formula:

$$I = 0.2 * ((25400 / [CN]) - 254)$$

Where:

I = initial abstraction (mm)

CN = Subbasin curve number

- 9) Initial Deficit: Initial Deficit represents the empty storage depth (mm) at the beginning of the simulation period. This quantity was set at half of the basin water holding capacity implying average soil moisture content in the soil at the beginning of the simulation period.
- 10) Routing channel bottom width: Computation of the average channel bottom width (m) for each routing reach.
- 11) Routing channel side slope: Computation of the average channel side slope (dimensionless).
- 12) Crop coefficient: Composite crop coefficient values were calculated for each day of the simulation period by calculating an area-weighted value for each catchment.

## **Export Model to HMS**

Before exporting the developed hydrological modelling inputs to an HEC-HMS model input file, the HEC-GeoHMS tools were used to check the GIS layers for stream and basin connectivity, generate HMS schematic, legend, and a background map file.

## **Appendix C**

Average soil infiltration rates based on Hydrologic Soil Group (Haan et al., 1982)

Hydrologic Soil Group	Description	Average Infiltration Rate (mm/hr)
A	Soils having high infiltration rates even when thoroughly wet. These soils consist mainly of deep, well-drained to excessively drained sands or gravels. These soils have a high rate of water transmission and therefore a low runoff potential.	9.51
B	Soils having moderate infiltration rates when thoroughly wet, consisting mainly of moderately deep to deep, moderately well to well-drained soils with moderately fine to moderately coarse textures. These soils have a moderate rate of water transmission.	5.72
C	Soils having slow infiltration rates when thoroughly wet, consisting mainly of either soils with a layer that impedes the downward movement of water or soils with moderately fine or fine textures and slow infiltration rates. These soils have a slow rate of water transmission.	2.54
D	Soils having very slow infiltration rates when thoroughly wet. These are mainly comprised of either clayey soil with high swelling capacity or potential, soils with a high permanent water table, soils with a clay layer at or near the surface, and/or shallow soils over nearly impervious materials. These soils have a very slow rate of water transmission and therefore a high runoff potential.	0.64

Runoff CN number for SOLRIS land use classes and HSG groups.

	<b>Hydrologic Soil Group</b>			
<b>Land use class</b>	<b>A</b>	<b>B</b>	<b>C</b>	<b>D</b>
Annual Crop	67	78	85	89
Bog	72	72	72	72
Coniferous Forest	48	58	70	77
Deciduous Forest	48	58	70	77
Extraction	98	98	98	98
Forest	48	58	70	77
Hedge Rows	48	58	70	77
Idle Land	50	61	74	80
Marsh	85	85	85	85
Mixed Agriculture	64	74	80	85
Mixed Crop	67	78	85	89
Mixed Forest	48	58	70	77
Monoculture	40	62	76	81
Open Shoreline	72	72	72	72
Open Water	100	100	100	100
Orchards	40	62	76	81
Perennial Crop	59	74	83	86
Plantations	38	60	74	80
Rural Land Use	56	70	80	82
Shoreline	72	72	72	72
Swamp	72	72	72	72

Soil water holding capacity (mm) for SOLRIS land use classes and HSG groups.

	<b>Hydrologic Soil Group</b>			
<b>Land use class</b>	<b>A</b>	<b>B</b>	<b>C</b>	<b>D</b>
Annual Crop	75	150	200	150
Bog	250	300	400	350
Built Up Impervious	0	0	0	0
Built Up Pervious	50	75	113	75
Coniferous Forest	250	300	400	350
Deciduous Forest	250	300	400	350
Extraction	0	0	0	0
Forest	250	300	400	350
Hedge Rows	250	300	400	350
Idle Land	100	150	250	200
Marsh	0	0	0	0
Mixed Agriculture	75	150	200	150
Mixed Crop	75	150	200	150
Mixed Forest	250	300	400	350
Monoculture	75	150	200	150
Open Shoreline	0	0	0	0
Open Water	0	0	0	0
Orchards	250	300	400	350
Perennial Crop	100	150	250	200
Plantations	100	150	250	100
Rural Land Use	100	150	250	200
Shoreline	0	0	0	0
Swamp	250	300	400	350
Transportation	0	0	0	0

Channel manning n coefficients under various channel stream bed conditions.

<b>Channel Stream bed condition</b>	<b>Minimum</b>	<b>Average</b>	<b>Maximum</b>
a. Clean, straight, full, no rifts or deep pools	0.025	0.030	0.033
b. Same as above, but more stones and weeds	0.030	0.035	0.040
c. Clean, winding, some pools and shoals	0.033	0.040	0.045
d. Same as above, but some stones and weeds.	0.035	0.045	0.050
e. Same as above, lower stages, more ineffective slopes and sections	0.040	0.048	0.050
f. Same as “d” but more stones	0.045	0.050	0.060
g. Sluggish reaches, weedy, deep pools	0.050	0.070	0.080
h. Very weedy reaches, deep pools, or floodways with heavy stands of timber and brush.	0.070	0.100	0.150



## **Appendix D**



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## Land-Use Evapotranspiration Coefficient Study

Weather INnovations Incorporated (WIN) was contracted by the Niagara Peninsula Conservation Authority (NPCA) to determine coefficients relating evapotranspiration rates to a reference potential evapotranspiration for the highest density land uses in the Niagara Peninsula.

The land uses of greatest interest, the majority of which were identified to be agricultural, were outlined by the NPCA. The crops of greatest density were used to further define the agricultural land use.

**Table 1: Identified Land Use**

<b>Agricultural Land</b>
- Soybeans
- Hay
- Grain Corn
- Winter Wheat
- Vineyards
- Peach
- Pear/Apple/Cherry/Plum
- Grazing Land
<b>Non-Agricultural Land Use</b>
- Idle Land (more than 10 years out of agricultural production)
- Deciduous Forest
- Swamp
- Built-Up Pervious land (sod/grass)
<b>Smaller Coverage Land Uses</b>
-Open/Shallow Water
- Marsh
- Coniferous Forest
- Tallgrass
- Fen
- Bog

The results of many researchers have been used to develop the evapotranspiration coefficients suggested by the Food and Agricultural Organization of the United Nations (FAO), and are accepted internationally. Due to the complexity of measuring evapotranspiration and the variability in the values year to year, there are very few (if any) results indicating a daily estimate for evapotranspiration values. WIN has determined,

where possible, the growth stages for the various agricultural crops listed, based on suggestions by the FAO and other researchers. The step-wise values have been determined where a linear progression occurs from one coefficient to another.

Very few crop coefficients for Ontario are available, and a majority of the land uses identified by the NPCA could not be found in Ontario documentation. As a result, values from the FAO were substituted, except in the situation of grapes and deciduous fruit trees. These values, although based on FAO findings, were more refined by the OMAFRA Best Management - Irrigation Management guide, and the British Columbia Ministry of Agriculture, Food and Fisheries. These two sources were used in combination, to determine the evapotranspiration coefficient for grapes and deciduous fruit trees.

The FAO suggests modifying the mid-growth stage values depending on the minimum daily relative humidity value, and the wind speed (at 2m) based on the region of interest, and the crop height. They suggest this adjustment to both  $K_{c_{mid}}$  and  $K_{c_{end}}$ . The equation indicated is identified for  $K_{c_{mid}}$ , however the equation for  $K_{c_{end}}$  is the same, just with the table value for  $K_{c_{end}}$  substituted in place of  $K_{c_{mid(tab)}}$ .

$$K_{c_{mid}} = K_{c_{mid(tab)}} + [0.04(u_2 - 2) - 0.004[RH_{min} - 45]](h/3)^{0.3},$$

where  $K_{c_{mid(tab)}}$  is the published FAO crop coefficient,  $u_2$  is the wind speed at 2m (in m/s),  $RH_{min}$  is the minimum daily RH value (%), and  $h$  is the crop height (m).

In order to conduct this calculation, wind speed and RH data from a station in the Niagara Peninsula was used for 2006. It is important to note that these values will change yearly depending on the season. A yearly average of the ‘adjusted’  $K_c$  value will be used for the purpose of this project.

The following changes to  $K_{c_{mid}}$  and  $K_{c_{end}}$  were made to the following crops.

**Table 2:** Adjusted  $K_{c_{mid}}$  and  $K_{c_{end}}$  values

Crop	Crop Height (m)	$K_{c_{mid(tab)}}$	Adjusted $K_{c_{mid}}$	$K_{c_{end(tab)}}$	Adjusted $K_{c_{end}}$
Soybeans	0.7	1.15	1.15	0.50	0.51
Winter Wheat	1.0	1.15	1.17	0.32	0.34
Maize	2.0	1.20	1.25	0.48	0.65
Rye Grass Hay	0.3	1.05	1.05	1.00	1.00
Clover Hay	0.6	0.90	0.90	0.85	0.85
Pasture	0.4	0.95	0.95	0.85	0.85
Wetlands	1.5	1.20	1.27	0.30	0.37

FAO provides estimated duration for each crop coefficient. These were taken into consideration when determining the change from one coefficient to another. However, in some instances, alterations were made to better suit the Ontario growing season. Many of the planting dates set for crops are close to the earliest planting date for the crop. These values should be adjusted, if required, to reflect a variety of situations based on planting dates.

The evapotranspiration coefficients for boreal deciduous and coniferous forests were determined from a study by Komastu (2005). The results, from various research projects around the world, were examined for a comparison of coefficients for the Priestly-Taylor model. In order to determine the values for this study, the average of the findings for both the boreal deciduous and the boreal coniferous forests were used. The values for boreal coniferous forest range from 0.38 to 0.69, with an average of 0.55. Values estimating the winter evapotranspiration coefficients for conifer trees could not be determined in the time allotted for this project. As such, evaporative losses of 0.2 were substituted, the value which is currently used for deciduous trees.

The determination for a coefficient for idle land becomes more complex. As the land has been out of agricultural production for 10 or more years, it is assumed that grasses, weeds and native vegetation are now established. The FAO indicates the use of the following equations to determine the mid-season evaporation rates. Due to the lack of information regarding leaf area index (LAI) values, it was indicated by the FAO that full coverage vegetation would have an LAI value of 3. For the scenario of tallgrass, a similar methodology was used. However, a study by Verma and Berry (1997) indicates that the LAI from a tallgrass prairie was 0.2 from mid-March to early May, and ranged from 0.3 to 1.8 in the later part of May (average of 1.05 will be used for this study). At the peak of the season, the range was 2.5 to 2.8 (average of 2.65 used for this study).

$$K_{cb, h} = 1.0 + 0.1h, \text{ for } h \leq 2m$$

$$K_{cb \text{ full}} = K_{cb, h} + [0.04(u_2 - 2) - (0.004(RH_{\min} - 45))](h/3)^{0.3}$$

$$K_{cb \text{ mid}} = K_{c \min} + (K_{c \text{ full}} - K_{c \min})(1 - \exp[-0.7LAI]),$$

where  $K_{c \min}$  is the minimum Kc value for bare soil (ranging from 0.15-0.2). OMAFRA indicates a bare soil coefficient of 0.2. This value was used.

Due to a lack of information regarding evapotranspiration values for swamps, bogs, marshes and fens individually, the value for wetlands was used in all four situations.

Many models used to calculate potential evapotranspiration (ET<sub>o</sub>) utilize a well-watered turf surface as the reference point. The coefficient for the built-up pervious area (e.g. sod/grass) will be 1.

Open water, especially water at a depth greater than 5m, creates a complex situation. The FAO indicates that deep bodies of water experience fluctuating temperatures, but this may not be true of frozen surfaces. During periods of peak evapotranspirative losses, radiation is being absorbed into the water. Therefore, the evaporative losses are less than ET<sub>o</sub>. During cooler temperature periods, the energy exchange is reversed. This causes the evaporation rates to be higher than those for grass, during the same period.

The following tables are daily evapotranspiration coefficients for the identified land use classes.

**Table 3:** Annual evapotranspiration coefficients for soybeans, winter wheat and grain corn.

	Soybeans	Kc	Winter Wheat	Kc	Grain Corn	Kc
Jan. 1	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 2	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 3	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 4	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 5	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 6	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 7	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 8	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 9	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 10	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 11	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 12	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 13	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 14	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 15	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 16	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 17	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 18	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 19	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 20	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 21	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 22	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 23	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 24	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 25	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 26	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 27	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 28	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 29	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 30	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Jan. 31	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 1	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 2	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 3	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 4	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 5	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 6	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 7	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 8	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 9	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 10	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 11	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 12	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 13	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 14	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 15	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 16	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 17	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 18	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 19	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 20	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 21	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 22	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 23	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Feb. 24	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20

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Apr. 25	Bare Soil	0.20	Kdev	0.82	Kini	0.30
Apr. 26	Bare Soil	0.20	Kdev	0.84	Kini	0.30
Apr. 27	Bare Soil	0.20	Kdev	0.85	Kini	0.30
Apr. 28	Bare Soil	0.20	Kdev	0.87	Kini	0.30
Apr. 29	Bare Soil	0.20	Kdev	0.89	Kini	0.30
Apr. 30	Bare Soil	0.20	Kdev	0.90	Kini	0.30
May 1	Bare Soil	0.20	Kdev	0.92	Kini	0.30
May 2	Bare Soil	0.20	Kdev	0.94	Kini	0.30
May 3	Bare Soil	0.20	Kdev	0.95	Kini	0.30
May 4	Bare Soil	0.20	Kdev	0.97	Kini	0.30
May 5	Bare Soil	0.20	Kdev	0.99	Kini	0.30
May 6	Bare Soil	0.20	Kdev	1.00	Kini	0.30
May 7	Bare Soil	0.20	Kdev	1.02	Kini	0.30
May 8	Bare Soil	0.20	Kdev	1.04	Kini	0.30
May 9	Bare Soil	0.20	Kdev	1.05	Kini	0.30
May 10	Bare Soil	0.20	Kdev	1.07	Kini	0.30
May 11	Bare Soil	0.20	Kdev	1.09	Kini	0.30
May 12	Bare Soil	0.20	Kdev	1.10	Kini	0.30
May 13	Bare Soil	0.20	Kdev	1.12	Kini	0.30
May 14	Bare Soil	0.20	Kdev	1.14	Kini	0.30
May 15	Kini	0.40	Kdev	1.15	Kini	0.30
May 16	Kini	0.40	Kmid	1.17	Kdev	0.32
May 17	Kini	0.40	Kmid	1.17	Kdev	0.35
May 18	Kini	0.40	Kmid	1.17	Kdev	0.37
May 19	Kini	0.40	Kmid	1.17	Kdev	0.39
May 20	Kini	0.40	Kmid	1.17	Kdev	0.42
May 21	Kini	0.40	Kmid	1.17	Kdev	0.44
May 22	Kini	0.40	Kmid	1.17	Kdev	0.46
May 23	Kini	0.40	Kmid	1.17	Kdev	0.49
May 24	Kini	0.40	Kmid	1.17	Kdev	0.51
May 25	Kini	0.40	Kmid	1.17	Kdev	0.53
May 26	Kini	0.40	Kmid	1.17	Kdev	0.55
May 27	Kini	0.40	Kmid	1.17	Kdev	0.58
May 28	Kini	0.40	Kmid	1.17	Kdev	0.60
May 29	Kini	0.40	Kmid	1.17	Kdev	0.62
May 30	Kini	0.40	Kmid	1.17	Kdev	0.65
May 31	Kini	0.40	Kmid	1.17	Kdev	0.67
June 1	Kini	0.40	Kmid	1.17	Kdev	0.69
June 2	Kini	0.40	Kmid	1.17	Kdev	0.72
June 3	Kini	0.40	Kmid	1.17	Kdev	0.74
June 4	Kdev	0.42	Kmid	1.17	Kdev	0.76
June 5	Kdev	0.45	Kmid	1.17	Kdev	0.79
June 6	Kdev	0.47	Kmid	1.17	Kdev	0.81
June 7	Kdev	0.49	Kmid	1.17	Kdev	0.83
June 8	Kdev	0.51	Kmid	1.17	Kdev	0.86
June 9	Kdev	0.54	Kmid	1.17	Kdev	0.88
June 10	Kdev	0.56	Kmid	1.17	Kdev	0.90
June 11	Kdev	0.58	Kmid	1.17	Kdev	0.93
June 12	Kdev	0.60	Kmid	1.17	Kdev	0.95
June 13	Kdev	0.63	Kmid	1.17	Kdev	0.97
June 14	Kdev	0.65	Kmid	1.17	Kdev	1.00
June 15	Kdev	0.67	Kmid	1.17	Kdev	1.02
June 16	Kdev	0.70	Kmid	1.17	Kdev	1.04
June 17	Kdev	0.72	Kmid	1.17	Kdev	1.06
June 18	Kdev	0.74	Kmid	1.17	Kdev	1.09
June 19	Kdev	0.76	Kmid	1.17	Kdev	1.11
June 20	Kdev	0.79	Kmid	1.17	Kdev	1.13
June 21	Kdev	0.81	Kmid	1.17	Kdev	1.16
June 22	Kdev	0.83	Kmid	1.17	Kdev	1.18

June 23	Kdev	0.85	Kmid	1.17	Kdev	1.20
June 24	Kdev	0.88	Kmid	1.17	Kdev	1.23
June 25	Kdev	0.90	Kmid	1.17	Kmid	1.25
June 26	Kdev	0.92	Kmid	1.17	Kmid	1.25
June 27	Kdev	0.95	Kmid	1.17	Kmid	1.25
June 28	Kdev	0.97	Kmid	1.17	Kmid	1.25
June 29	Kdev	0.99	Kmid	1.17	Kmid	1.25
June 30	Kdev	1.01	Kmid	1.17	Kmid	1.25
July 1	Kdev	1.04	Kmid	1.17	Kmid	1.25
July 2	Kdev	1.06	Kmid	1.17	Kmid	1.25
July 3	Kdev	1.08	Kmid	1.17	Kmid	1.25
July 4	Kdev	1.10	Kmid	1.17	Kmid	1.25
July 5	Kdev	1.13	Kmid	1.17	Kmid	1.25
July 6	Kmid	1.15	Kmid	1.17	Kmid	1.25
July 7	Kmid	1.15	Kmid	1.17	Kmid	1.25
July 8	Kmid	1.15	Kmid	1.17	Kmid	1.25
July 9	Kmid	1.15	Kmid	1.17	Kmid	1.25
July 10	Kmid	1.15	Kmid	1.17	Kmid	1.25
July 11	Kmid	1.15	Klate	1.14	Kmid	1.25
July 12	Kmid	1.15	Klate	1.12	Kmid	1.25
July 13	Kmid	1.15	Klate	1.09	Kmid	1.25
July 14	Kmid	1.15	Klate	1.06	Kmid	1.25
July 15	Kmid	1.15	Klate	1.03	Kmid	1.25
July 16	Kmid	1.15	Klate	1.01	Kmid	1.25
July 17	Kmid	1.15	Klate	0.98	Kmid	1.25
July 18	Kmid	1.15	Klate	0.95	Kmid	1.25
July 19	Kmid	1.15	Klate	0.92	Kmid	1.25
July 20	Kmid	1.15	Klate	0.90	Kmid	1.25
July 21	Kmid	1.15	Klate	0.87	Kmid	1.25
July 22	Kmid	1.15	Klate	0.84	Kmid	1.25
July 23	Kmid	1.15	Klate	0.82	Kmid	1.25
July 24	Kmid	1.15	Klate	0.79	Kmid	1.25
July 25	Kmid	1.15	Klate	0.76	Kmid	1.25
July 26	Kmid	1.15	Klate	0.73	Kmid	1.25
July 27	Kmid	1.15	Klate	0.71	Kmid	1.25
July 28	Kmid	1.15	Klate	0.68	Kmid	1.25
July 29	Kmid	1.15	Klate	0.65	Kmid	1.25
July 30	Kmid	1.15	Klate	0.62	Kmid	1.25
July 31	Kmid	1.15	Klate	0.60	Kmid	1.25
Aug. 1	Kmid	1.15	Klate	0.57	Kmid	1.25
Aug. 2	Kmid	1.15	Klate	0.54	Kmid	1.25
Aug. 3	Kmid	1.15	Klate	0.52	Kmid	1.25
Aug. 4	Kmid	1.15	Klate	0.49	Kmid	1.25
Aug. 5	Kmid	1.15	Klate	0.46	Kmid	1.25
Aug. 6	Kmid	1.15	Klate	0.43	Kmid	1.25
Aug. 7	Kmid	1.15	Klate	0.41	Kmid	1.25
Aug. 8	Kmid	1.15	Klate	0.38	Kmid	1.25
Aug. 9	Kmid	1.15	Klate	0.35	Kmid	1.25
Aug. 10	Kmid	1.15	Klate	0.33	Kmid	1.25
Aug. 11	Kmid	1.15	Bare Soil	0.20	Kmid	1.25
Aug. 12	Kmid	1.15	Bare Soil	0.20	Kmid	1.25
Aug. 13	Kmid	1.15	Bare Soil	0.20	Kmid	1.25
Aug. 14	Kmid	1.15	Bare Soil	0.20	Klate	1.23
Aug. 15	Kmid	1.15	Bare Soil	0.20	Klate	1.22
Aug. 16	Kmid	1.15	Bare Soil	0.20	Klate	1.20
Aug. 17	Kmid	1.15	Bare Soil	0.20	Klate	1.19
Aug. 18	Kmid	1.15	Bare Soil	0.20	Klate	1.17
Aug. 19	Kmid	1.15	Bare Soil	0.20	Klate	1.16
Aug. 20	Kmid	1.15	Bare Soil	0.20	Klate	1.14



Aug. 21	Kmid	1.15	Bare Soil	0.20	Klate	1.13
Aug. 22	Kmid	1.15	Bare Soil	0.20	Klate	1.11
Aug. 23	Kmid	1.15	Bare Soil	0.20	Klate	1.10
Aug. 24	Kmid	1.15	Bare Soil	0.20	Klate	1.08
Aug. 25	Kmid	1.15	Bare Soil	0.20	Klate	1.07
Aug. 26	Kmid	1.15	Bare Soil	0.20	Klate	1.05
Aug. 27	Kmid	1.15	Bare Soil	0.20	Klate	1.04
Aug. 28	Kmid	1.15	Bare Soil	0.20	Klate	1.02
Aug. 29	Kmid	1.15	Bare Soil	0.20	Klate	1.01
Aug. 30	Kmid	1.15	Bare Soil	0.20	Klate	0.99
Aug. 31	Kmid	1.15	Bare Soil	0.20	Klate	0.98
Sept. 1	Kmid	1.15	Bare Soil	0.20	Klate	0.96
Sept. 2	Kmid	1.15	Bare Soil	0.20	Klate	0.95
Sept. 3	Kmid	1.15	Bare Soil	0.20	Klate	0.93
Sept. 4	Klate	1.13	Bare Soil	0.20	Klate	0.92
Sept. 5	Klate	1.10	Bare Soil	0.20	Klate	0.90
Sept. 6	Klate	1.08	Bare Soil	0.20	Klate	0.89
Sept. 7	Klate	1.05	Bare Soil	0.20	Klate	0.87
Sept. 8	Klate	1.03	Bare Soil	0.20	Klate	0.85
Sept. 9	Klate	1.01	Bare Soil	0.20	Klate	0.84
Sept. 10	Klate	0.98	Bare Soil	0.20	Klate	0.82
Sept. 11	Klate	0.96	Bare Soil	0.20	Klate	0.81
Sept. 12	Klate	0.93	Bare Soil	0.20	Klate	0.79
Sept. 13	Klate	0.91	Bare Soil	0.20	Klate	0.78
Sept. 14	Klate	0.89	Bare Soil	0.20	Klate	0.76
Sept. 15	Klate	0.86	Kini	0.40	Klate	0.75
Sept. 16	Klate	0.84	Kini	0.40	Klate	0.73
Sept. 17	Klate	0.81	Kini	0.40	Klate	0.72
Sept. 18	Klate	0.79	Kini	0.40	Klate	0.70
Sept. 19	Klate	0.76	Kini	0.40	Klate	0.69
Sept. 20	Klate	0.74	Kini	0.40	Klate	0.67
Sept. 21	Klate	0.72	Kini	0.40	Klate	0.66
Sept. 22	Klate	0.69	Kini	0.40	Klate	0.64
Sept. 23	Klate	0.67	Kini	0.40	Klate	0.63
Sept. 24	Klate	0.64	Kini	0.40	Klate	0.61
Sept. 25	Klate	0.62	Kini	0.40	Klate	0.60
Sept. 26	Klate	0.60	Kini	0.40	Klate	0.58
Sept. 27	Klate	0.57	Kini	0.40	Klate	0.57
Sept. 28	Klate	0.55	Kini	0.40	Klate	0.55
Sept. 29	Klate	0.52	Kini	0.40	Klate	0.54
Sept. 30	Klate	0.50	Kini	0.40	Klate	0.52
Oct. 1	Bare Soil	0.20	Kini	0.40	Klate	0.51
Oct. 2	Bare Soil	0.20	Kini	0.40	Klate	0.48
Oct. 3	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 4	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 5	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 6	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 7	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 8	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 9	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 10	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 11	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 12	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 13	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 14	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 15	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 16	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 17	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Oct. 18	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20

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Dec. 17	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 18	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 19	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 20	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 21	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 22	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 23	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 24	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 25	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 26	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 27	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 28	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 29	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 30	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20
Dec. 31	Bare Soil	0.20	Kini	0.40	Bare Soil	0.20

**Table 4:** Annual evapotranspiration coefficients for hay, pasture, and wetlands.

	Hay <sup>+</sup>	Kc	Pasture	Kc	Wetlands	Kc
Jan. 1	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 2	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 3	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 4	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 5	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 6	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 7	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 8	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 9	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 10	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 11	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 12	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 13	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 14	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 15	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 16	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 17	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 18	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 19	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 20	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 21	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 22	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 23	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 24	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 25	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 26	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 27	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 28	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 29	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 30	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Jan. 31	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 1	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 2	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 3	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 4	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 5	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 6	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 7	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Feb. 8	Bare Soil	0.20	Dormant	0.40	Dormant	0.37

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Apr. 9	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 10	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 11	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 12	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 13	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 14	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 15	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 16	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 17	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 18	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 19	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 20	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 21	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 22	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 23	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 24	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 25	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 26	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 27	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 28	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 29	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Apr. 30	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
May 1	Kini	0.95	Kini	0.40	Dormant	0.37
May 2	Kini	0.95	Kini	0.40	Dormant	0.37
May 3	Kini	0.95	Kini	0.40	Dormant	0.37
May 4	Kini	0.95	Kini	0.40	Dormant	0.37
May 5	Kini	0.95	Kini	0.40	Dormant	0.37
May 6	Kini	0.95	Kini	0.40	Dormant	0.37
May 7	Kini	0.95	Kini	0.40	Dormant	0.37
May 8	Kini	0.95	Kini	0.40	Dormant	0.37
May 9	Kini	0.95	Kini	0.40	Dormant	0.37
May 10	Kini	0.95	Kini	0.40	Dormant	0.37
May 11	Kini	0.95	Kdev	0.43	Dormant	0.37
May 12	Kini	0.95	Kdev	0.45	Dormant	0.37
May 13	Kini	0.95	Kdev	0.48	Dormant	0.37
May 14	Kini	0.95	Kdev	0.50	Dormant	0.37
May 15	Kini	0.95	Kdev	0.53	Kini	0.30
May 16	Kini	0.95	Kdev	0.56	Kini	0.30
May 17	Kini	0.95	Kdev	0.58	Kini	0.30
May 18	Kini	0.95	Kdev	0.61	Kini	0.30
May 19	Kini	0.95	Kdev	0.64	Kini	0.30
May 20	Kini	0.95	Kdev	0.66	Kini	0.30
May 21	Kdev	0.95	Kdev	0.69	Kini	0.30
May 22	Kdev	0.96	Kdev	0.71	Kini	0.30
May 23	Kdev	0.96	Kdev	0.74	Kini	0.30
May 24	Kdev	0.97	Kdev	0.77	Kini	0.30
May 25	Kdev	0.97	Kdev	0.79	Kdev	0.33
May 26	Kdev	0.98	Kdev	0.82	Kdev	0.36
May 27	Kdev	0.98	Kdev	0.85	Kdev	0.39
May 28	Kdev	0.98	Kdev	0.87	Kdev	0.43
May 29	Kdev	0.99	Kdev	0.90	Kdev	0.46
May 30	Kdev	0.99	Kdev	0.92	Kdev	0.49
May 31	Kdev	1.00	Kmid	0.95	Kdev	0.52
June 1	Kdev	1.00	Kmid	0.95	Kdev	0.55
June 2	Kdev	1.00	Kmid	0.95	Kdev	0.58
June 3	Kdev	1.01	Kmid	0.95	Kdev	0.61
June 4	Kdev	1.01	Kmid	0.95	Kdev	0.64
June 5	Kdev	1.02	Kmid	0.95	Kdev	0.68
June 6	Kdev	1.02	Kmid	0.95	Kdev	0.71

June 7	Kdev	1.03	Kmid	0.95	Kdev	0.74
June 8	Kdev	1.03	Kmid	0.95	Kdev	0.77
June 9	Kdev	1.03	Kmid	0.95	Kdev	0.80
June 10	Kdev	1.04	Kmid	0.95	Kdev	0.83
June 11	Kdev	1.04	Kmid	0.95	Kdev	0.86
June 12	Kdev	1.05	Kmid	0.95	Kdev	0.89
June 13	Kdev	1.05	Kmid	0.95	Kdev	0.93
June 14	Kdev	1.06	Kmid	0.95	Kdev	0.96
June 15	Kmid	1.06	Kmid	0.95	Kdev	0.99
June 16	Kmid	1.06	Kmid	0.95	Kdev	1.02
June 17	Kmid	1.06	Kmid	0.95	Kdev	1.05
June 18	Kmid	1.06	Kmid	0.95	Kdev	1.08
June 19	Kmid	1.06	Kmid	0.95	Kdev	1.11
June 20	Kmid	1.06	Kmid	0.95	Kdev	1.14
June 21	Kmid	1.06	Kmid	0.95	Kdev	1.18
June 22	Kmid	1.06	Kmid	0.95	Kdev	1.21
June 23	Kmid	1.06	Kmid	0.95	Kdev	1.24
June 24	Kmid	1.06	Kmid	0.95	Kdev	1.27
June 25	Kmid	1.06	Kmid	0.95	Kmid	1.27
June 26	Kmid	1.06	Kmid	0.95	Kmid	1.27
June 27	Kmid	1.06	Kmid	0.95	Kmid	1.27
June 28	Kmid	1.06	Kmid	0.95	Kmid	1.27
June 29	Kmid	1.06	Kmid	0.95	Kmid	1.27
June 30	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 1	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 2	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 3	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 4	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 5	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 6	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 7	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 8	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 9	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 10	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 11	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 12	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 13	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 14	Kmid	1.06	Kmid	0.95	Kmid	1.27
July 15	Klate	1.05	Kmid	0.95	Kmid	1.27
July 16	Klate	1.04	Kmid	0.95	Kmid	1.27
July 17	Klate	1.02	Kmid	0.95	Kmid	1.27
July 18	Klate	1.01	Kmid	0.95	Kmid	1.27
July 19	Klate	1.00	Kmid	0.95	Kmid	1.27
July 20	Kini	0.95	Kmid	0.95	Kmid	1.27
July 21	Kini	0.95	Kmid	0.95	Kmid	1.27
July 22	Kini	0.95	Kmid	0.95	Kmid	1.27
July 23	Kini	0.95	Kmid	0.95	Kmid	1.27
July 24	Kini	0.95	Kmid	0.95	Kmid	1.27
July 25	Kini	0.95	Kmid	0.95	Kmid	1.27
July 26	Kdev	0.96	Kmid	0.95	Kmid	1.27
July 27	Kdev	0.96	Kmid	0.95	Kmid	1.27
July 28	Kdev	0.97	Kmid	0.95	Kmid	1.27
July 29	Kdev	0.98	Kmid	0.95	Kmid	1.27
July 30	Kdev	0.98	Kmid	0.95	Kmid	1.27
July 31	Kdev	0.99	Kmid	0.95	Kmid	1.27
Aug. 1	Kdev	1.00	Kmid	0.95	Kmid	1.27
Aug. 2	Kdev	1.01	Kmid	0.95	Kmid	1.27
Aug. 3	Kdev	1.01	Kmid	0.95	Kmid	1.27
Aug. 4	Kdev	1.02	Kmid	0.95	Kmid	1.27

Aug. 5	Kdev	1.03	Kmid	0.95	Kmid	1.27
Aug. 6	Kdev	1.03	Kmid	0.95	Kmid	1.27
Aug. 7	Kdev	1.04	Kmid	0.95	Kmid	1.27
Aug. 8	Kdev	1.05	Kmid	0.95	Kmid	1.27
Aug. 9	Kdev	1.05	Kmid	0.95	Kmid	1.27
Aug. 10	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 11	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 12	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 13	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 14	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 15	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 16	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 17	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 18	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 19	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 20	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 21	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 22	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 23	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 24	Kmid	1.06	Kmid	0.95	Kmid	1.27
Aug. 25	Klate	1.07	Kmid	0.95	Kmid	1.27
Aug. 26	Klate	1.08	Kmid	0.95	Kmid	1.27
Aug. 27	Klate	1.10	Kmid	0.95	Kmid	1.27
Aug. 28	Klate	1.11	Kmid	0.95	Kmid	1.27
Aug. 29	Klate	1.00	Kmid	0.95	Kmid	1.27
Aug. 30	Kini	0.95	Kmid	0.95	Kmid	1.27
Aug. 31	Kini	0.95	Kmid	0.95	Kmid	1.27
Sept. 1	Kini	0.95	Kmid	0.95	Kmid	1.27
Sept. 2	Kini	0.95	Kmid	0.95	Kmid	1.27
Sept. 3	Kini	0.95	Kmid	0.95	Kmid	1.27
Sept. 4	Kini	0.95	Kmid	0.95	Kmid	1.27
Sept. 5	Kini	0.95	Kmid	0.95	Kmid	1.27
Sept. 6	Kdev	0.96	Kmid	0.95	Kmid	1.27
Sept. 7	Kdev	0.96	Kmid	0.95	Kmid	1.27
Sept. 8	Kdev	0.97	Kmid	0.95	Kmid	1.27
Sept. 9	Kdev	0.98	Kmid	0.95	Kmid	1.27
Sept. 10	Kdev	0.98	Kmid	0.95	Kmid	1.27
Sept. 11	Kdev	0.99	Kmid	0.95	Kmid	1.27
Sept. 12	Kdev	1.00	Kmid	0.95	Kmid	1.27
Sept. 13	Kdev	1.01	Kmid	0.95	Kmid	1.27
Sept. 14	Kdev	1.01	Kmid	0.95	Kmid	1.27
Sept. 15	Kdev	1.02	Kmid	0.95	Kmid	1.27
Sept. 16	Kdev	1.03	Kmid	0.95	Kmid	1.27
Sept. 17	Kdev	1.03	Kmid	0.95	Kmid	1.27
Sept. 18	Kdev	1.04	Kmid	0.95	Kmid	1.27
Sept. 19	Kdev	1.05	Kmid	0.95	Kmid	1.27
Sept. 20	Kdev	1.05	Kmid	0.95	Klate	1.23
Sept. 21	Kmid	1.06	Kmid	0.95	Klate	1.20
Sept. 22	Kmid	1.06	Kmid	0.95	Klate	1.16
Sept. 23	Kmid	1.06	Kmid	0.95	Klate	1.12
Sept. 24	Kmid	1.06	Kmid	0.95	Klate	1.08
Sept. 25	Kmid	1.06	Kmid	0.95	Klate	1.05
Sept. 26	Kmid	1.06	Kmid	0.95	Klate	1.01
Sept. 27	Kmid	1.06	Kmid	0.95	Klate	0.97
Sept. 28	Kmid	1.06	Kmid	0.95	Klate	0.93
Sept. 29	Kmid	1.06	Kmid	0.95	Klate	0.90
Sept. 30	Kmid	1.06	Kmid	0.95	Klate	0.86
Oct. 1	Kmid	1.06	Kmid	0.95	Klate	0.82
Oct. 2	Kmid	1.06	Klate	0.94	Klate	0.79

Oct. 3	Kmid	1.06	Klate	0.94	Klate	0.75
Oct. 4	Kmid	1.06	Klate	0.93	Klate	0.71
Oct. 5	Kmid	1.06	Klate	0.92	Klate	0.67
Oct. 6	Klate	1.05	Klate	0.92	Klate	0.64
Oct. 7	Klate	1.05	Klate	0.91	Klate	0.60
Oct. 8	Klate	1.04	Klate	0.90	Klate	0.56
Oct. 9	Klate	1.04	Klate	0.90	Klate	0.52
Oct. 10	Klate	1.03	Klate	0.89	Klate	0.49
Oct. 11	Klate	1.02	Klate	0.88	Klate	0.45
Oct. 12	Klate	1.02	Klate	0.88	Klate	0.41
Oct. 13	Klate	1.01	Klate	0.87	Klate	0.37
Oct. 14	Klate	1.01	Klate	0.86	Klate	0.34
Oct. 15	Klate	1.00	Klate	0.86	Klate	0.30
Oct. 16	Bare Soil	0.20	Klate	0.85	Dormant	0.37
Oct. 17	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 18	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 19	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 20	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 21	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 22	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 23	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 24	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 25	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 26	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 27	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 28	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 29	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 30	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Oct. 31	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 1	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 2	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 3	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 4	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 5	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 6	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 7	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 8	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 9	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 10	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 11	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 12	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 13	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 14	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 15	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 16	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 17	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 18	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 19	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 20	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 21	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 22	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 23	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 24	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 25	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 26	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 27	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 28	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 29	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Nov. 30	Bare Soil	0.20	Dormant	0.40	Dormant	0.37



Dec. 1	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 2	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 3	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 4	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 5	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 6	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 7	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 8	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 9	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 10	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 11	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 12	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 13	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 14	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 15	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 16	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 17	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 18	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 19	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 20	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 21	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 22	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 23	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 24	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 25	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 26	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 27	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 28	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 29	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 30	Bare Soil	0.20	Dormant	0.40	Dormant	0.37
Dec. 31	Bare Soil	0.20	Dormant	0.40	Dormant	0.37

\*3 cuttings were assumed for Hay

**Table 5:** Annual evapotranspiration coefficients for peaches, apples, cherries, pears, and grapes.

	Peaches*	Apples, Cherries, Pears <sup>‡</sup>	Grapes <sup>a</sup>
Jan. 1	0.2	0.2	0.2
Jan. 2	0.2	0.2	0.2
Jan. 3	0.2	0.2	0.2
Jan. 4	0.2	0.2	0.2
Jan. 5	0.2	0.2	0.2
Jan. 6	0.2	0.2	0.2
Jan. 7	0.2	0.2	0.2
Jan. 8	0.2	0.2	0.2
Jan. 9	0.2	0.2	0.2
Jan. 10	0.2	0.2	0.2
Jan. 11	0.2	0.2	0.2
Jan. 12	0.2	0.2	0.2
Jan. 13	0.2	0.2	0.2
Jan. 14	0.2	0.2	0.2
Jan. 15	0.2	0.2	0.2
Jan. 16	0.2	0.2	0.2
Jan. 17	0.2	0.2	0.2
Jan. 18	0.2	0.2	0.2
Jan. 19	0.2	0.2	0.2
Jan. 20	0.2	0.2	0.2

Jan. 21	0.2	0.2	0.2
Jan. 22	0.2	0.2	0.2
Jan. 23	0.2	0.2	0.2
Jan. 24	0.2	0.2	0.2
Jan. 25	0.2	0.2	0.2
Jan. 26	0.2	0.2	0.2
Jan. 27	0.2	0.2	0.2
Jan. 28	0.2	0.2	0.2
Jan. 29	0.2	0.2	0.2
Jan. 30	0.2	0.2	0.2
Jan. 31	0.2	0.2	0.2
Feb. 1	0.2	0.2	0.2
Feb. 2	0.2	0.2	0.2
Feb. 3	0.2	0.2	0.2
Feb. 4	0.2	0.2	0.2
Feb. 5	0.2	0.2	0.2
Feb. 6	0.2	0.2	0.2
Feb. 7	0.2	0.2	0.2
Feb. 8	0.2	0.2	0.2
Feb. 9	0.2	0.2	0.2
Feb. 10	0.2	0.2	0.2
Feb. 11	0.2	0.2	0.2
Feb. 12	0.2	0.2	0.2
Feb. 13	0.2	0.2	0.2
Feb. 14	0.2	0.2	0.2
Feb. 15	0.2	0.2	0.2
Feb. 16	0.2	0.2	0.2
Feb. 17	0.2	0.2	0.2
Feb. 18	0.2	0.2	0.2
Feb. 19	0.2	0.2	0.2
Feb. 20	0.2	0.2	0.2
Feb. 21	0.2	0.2	0.2
Feb. 22	0.2	0.2	0.2
Feb. 23	0.2	0.2	0.2
Feb. 24	0.2	0.2	0.2
Feb. 25	0.2	0.2	0.2
Feb. 26	0.2	0.2	0.2
Feb. 27	0.2	0.2	0.2
Feb. 28	0.2	0.2	0.2
Mar. 1	0.2	0.2	0.2
Mar. 2	0.2	0.2	0.2
Mar. 3	0.2	0.2	0.2
Mar. 4	0.2	0.2	0.2
Mar. 5	0.2	0.2	0.2
Mar. 6	0.2	0.2	0.2
Mar. 7	0.2	0.2	0.2
Mar. 8	0.2	0.2	0.2
Mar. 9	0.2	0.2	0.2
Mar. 10	0.2	0.2	0.2
Mar. 11	0.2	0.2	0.2
Mar. 12	0.2	0.2	0.2
Mar. 13	0.2	0.2	0.2
Mar. 14	0.2	0.2	0.2
Mar. 15	0.2	0.2	0.2
Mar. 16	0.2	0.2	0.2
Mar. 17	0.2	0.2	0.2
Mar. 18	0.2	0.2	0.2
Mar. 19	0.2	0.2	0.2
Mar. 20	0.2	0.2	0.2

Mar. 21	0.2	0.2	0.2
Mar. 22	0.2	0.2	0.2
Mar. 23	0.2	0.2	0.2
Mar. 24	0.2	0.2	0.2
Mar. 25	0.2	0.2	0.2
Mar. 26	0.2	0.2	0.2
Mar. 27	0.2	0.2	0.2
Mar. 28	0.2	0.2	0.2
Mar. 29	0.2	0.2	0.2
Mar. 30	0.2	0.2	0.2
Mar. 31	0.2	0.2	0.2
Apr. 1	0.2	0.2	0.2
Apr. 2	0.2	0.2	0.2
Apr. 3	0.2	0.2	0.2
Apr. 4	0.2	0.2	0.2
Apr. 5	0.2	0.2	0.2
Apr. 6	0.2	0.2	0.2
Apr. 7	0.2	0.2	0.2
Apr. 8	0.2	0.2	0.2
Apr. 9	0.2	0.2	0.2
Apr. 10	0.2	0.2	0.2
Apr. 11	0.2	0.2	0.2
Apr. 12	0.2	0.2	0.2
Apr. 13	0.2	0.2	0.2
Apr. 14	0.2	0.2	0.2
Apr. 15	0.2	0.2	0.2
Apr. 16	0.2	0.2	0.2
Apr. 17	0.2	0.2	0.2
Apr. 18	0.2	0.2	0.2
Apr. 19	0.2	0.2	0.2
Apr. 20	0.2	0.2	0.2
Apr. 21	0.2	0.2	0.2
Apr. 22	0.2	0.2	0.2
Apr. 23	0.2	0.2	0.2
Apr. 24	0.2	0.2	0.2
Apr. 25	0.2	0.2	0.2
Apr. 26	0.2	0.2	0.2
Apr. 27	0.2	0.2	0.2
Apr. 28	0.2	0.2	0.2
Apr. 29	0.2	0.2	0.2
Apr. 30	0.2	0.2	0.2
May 1	0.3	0.3	0.5
May 2	0.3	0.3	0.5
May 3	0.3	0.3	0.5
May 4	0.3	0.3	0.5
May 5	0.3	0.3	0.5
May 6	0.3	0.3	0.5
May 7	0.3	0.3	0.5
May 8	0.3	0.3	0.5
May 9	0.3	0.3	0.5
May 10	0.3	0.3	0.5
May 11	0.3	0.3	0.5
May 12	0.3	0.3	0.5
May 13	0.3	0.3	0.5
May 14	0.3	0.3	0.5
May 15	0.3	0.3	0.5
May 16	0.3	0.3	0.5
May 17	0.3	0.3	0.5
May 18	0.3	0.3	0.5

May 19	0.3	0.3	0.5
May 20	0.3	0.3	0.5
May 21	0.3	0.3	0.5
May 22	0.3	0.3	0.5
May 23	0.3	0.3	0.5
May 24	0.3	0.3	0.5
May 25	0.3	0.3	0.5
May 26	0.3	0.3	0.5
May 27	0.3	0.3	0.5
May 28	0.3	0.3	0.5
May 29	0.3	0.3	0.5
May 30	0.3	0.3	0.5
May 31	0.3	0.3	0.5
June 1	0.4	0.4	0.68
June 2	0.4	0.4	0.68
June 3	0.4	0.4	0.68
June 4	0.4	0.4	0.68
June 5	0.4	0.4	0.68
June 6	0.4	0.4	0.68
June 7	0.4	0.4	0.68
June 8	0.4	0.4	0.68
June 9	0.4	0.4	0.68
June 10	0.4	0.4	0.68
June 11	0.4	0.4	0.68
June 12	0.4	0.4	0.68
June 13	0.4	0.4	0.68
June 14	0.4	0.4	0.68
June 15	0.4	0.4	0.68
June 16	0.6	0.6	0.68
June 17	0.6	0.6	0.68
June 18	0.6	0.6	0.68
June 19	0.6	0.6	0.68
June 20	0.6	0.6	0.68
June 21	0.6	0.6	0.68
June 22	0.6	0.6	0.68
June 23	0.6	0.6	0.68
June 24	0.6	0.6	0.68
June 25	0.6	0.6	0.68
June 26	0.6	0.6	0.68
June 27	0.6	0.6	0.68
June 28	0.6	0.6	0.68
June 29	0.6	0.6	0.68
June 30	0.6	0.6	0.68
July 1	1.0	1.0	0.77
July 2	1.0	1.0	0.77
July 3	1.0	1.0	0.77
July 4	1.0	1.0	0.77
July 5	1.0	1.0	0.77
July 6	1.0	1.0	0.77
July 7	1.0	1.0	0.77
July 8	1.0	1.0	0.77
July 9	1.0	1.0	0.77
July 10	1.0	1.0	0.77
July 11	1.0	1.0	0.77
July 12	1.0	1.0	0.77
July 13	1.0	1.0	0.77
July 14	1.0	1.0	0.77
July 15	1.0	1.0	0.77
July 16	1.0	1.0	0.77

July 17	1.0	1.0	0.77
July 18	1.0	1.0	0.77
July 19	1.0	1.0	0.77
July 20	1.0	1.0	0.77
July 21	1.0	1.0	0.77
July 22	1.0	1.0	0.77
July 23	1.0	1.0	0.77
July 24	1.0	1.0	0.77
July 25	1.0	1.0	0.77
July 26	1.0	1.0	0.77
July 27	1.0	1.0	0.77
July 28	1.0	1.0	0.77
July 29	1.0	1.0	0.77
July 30	1.0	1.0	0.77
July 31	1.0	1.0	0.77
Aug. 1	1.0	1.0	0.80
Aug. 2	1.0	1.0	0.80
Aug. 3	1.0	1.0	0.80
Aug. 4	1.0	1.0	0.80
Aug. 5	1.0	1.0	0.80
Aug. 6	1.0	1.0	0.80
Aug. 7	1.0	1.0	0.80
Aug. 8	1.0	1.0	0.80
Aug. 9	1.0	1.0	0.80
Aug. 10	1.0	1.0	0.80
Aug. 11	1.0	1.0	0.80
Aug. 12	1.0	1.0	0.80
Aug. 13	1.0	1.0	0.80
Aug. 14	1.0	1.0	0.80
Aug. 15	1.0	1.0	0.80
Aug. 16	1.0	1.0	0.80
Aug. 17	1.0	1.0	0.80
Aug. 18	1.0	1.0	0.80
Aug. 19	1.0	1.0	0.80
Aug. 20	1.0	1.0	0.80
Aug. 21	1.0	1.0	0.80
Aug. 22	1.0	1.0	0.80
Aug. 23	1.0	1.0	0.80
Aug. 24	1.0	1.0	0.80
Aug. 25	1.0	1.0	0.80
Aug. 26	1.0	1.0	0.80
Aug. 27	1.0	1.0	0.80
Aug. 28	1.0	1.0	0.80
Aug. 29	1.0	1.0	0.80
Aug. 30	1.0	1.0	0.80
Aug. 31	1.0	1.0	0.80
Sept. 1	0.95	0.95	0.75
Sept. 2	0.95	0.95	0.75
Sept. 3	0.95	0.95	0.75
Sept. 4	0.95	0.95	0.75
Sept. 5	0.95	0.95	0.75
Sept. 6	0.95	0.95	0.75
Sept. 7	0.95	0.95	0.75
Sept. 8	0.95	0.95	0.75
Sept. 9	0.95	0.95	0.75
Sept. 10	0.95	0.95	0.75
Sept. 11	0.95	0.95	0.75
Sept. 12	0.95	0.95	0.75
Sept. 13	0.95	0.95	0.75

Sept. 14	0.95	0.95	0.75
Sept. 15	0.95	0.95	0.75
Sept. 16	0.95	0.95	0.75
Sept. 17	0.95	0.95	0.75
Sept. 18	0.95	0.95	0.75
Sept. 19	0.95	0.95	0.75
Sept. 20	0.95	0.95	0.75
Sept. 21	0.95	0.95	0.75
Sept. 22	0.95	0.95	0.75
Sept. 23	0.95	0.95	0.75
Sept. 24	0.95	0.95	0.75
Sept. 25	0.95	0.95	0.75
Sept. 26	0.95	0.95	0.75
Sept. 27	0.95	0.95	0.75
Sept. 28	0.95	0.95	0.75
Sept. 29	0.95	0.95	0.75
Sept. 30	0.95	0.95	0.75
Oct. 1	0.83	0.80	0.63
Oct. 2	0.83	0.80	0.63
Oct. 3	0.83	0.80	0.63
Oct. 4	0.83	0.80	0.63
Oct. 5	0.83	0.80	0.63
Oct. 6	0.83	0.80	0.63
Oct. 7	0.83	0.80	0.63
Oct. 8	0.83	0.80	0.63
Oct. 9	0.83	0.80	0.63
Oct. 10	0.83	0.80	0.63
Oct. 11	0.83	0.80	0.63
Oct. 12	0.83	0.80	0.63
Oct. 13	0.83	0.80	0.63
Oct. 14	0.83	0.80	0.63
Oct. 15	0.83	0.80	0.63
Oct. 16	0.83	0.80	0.63
Oct. 17	0.83	0.80	0.63
Oct. 18	0.83	0.80	0.63
Oct. 19	0.83	0.80	0.63
Oct. 20	0.83	0.80	0.63
Oct. 21	0.83	0.80	0.63
Oct. 22	0.83	0.80	0.63
Oct. 23	0.83	0.80	0.63
Oct. 24	0.83	0.80	0.63
Oct. 25	0.83	0.80	0.63
Oct. 26	0.83	0.80	0.63
Oct. 27	0.83	0.80	0.63
Oct. 28	0.83	0.80	0.63
Oct. 29	0.83	0.80	0.63
Oct. 30	0.83	0.80	0.63
Oct. 31	0.83	0.80	0.63
Nov. 1	0.2	0.2	0.2
Nov. 2	0.2	0.2	0.2
Nov. 3	0.2	0.2	0.2
Nov. 4	0.2	0.2	0.2
Nov. 5	0.2	0.2	0.2
Nov. 6	0.2	0.2	0.2
Nov. 7	0.2	0.2	0.2
Nov. 8	0.2	0.2	0.2
Nov. 9	0.2	0.2	0.2
Nov. 10	0.2	0.2	0.2
Nov. 11	0.2	0.2	0.2

Nov. 12	0.2	0.2	0.2
Nov. 13	0.2	0.2	0.2
Nov. 14	0.2	0.2	0.2
Nov. 15	0.2	0.2	0.2
Nov. 16	0.2	0.2	0.2
Nov. 17	0.2	0.2	0.2
Nov. 18	0.2	0.2	0.2
Nov. 19	0.2	0.2	0.2
Nov. 20	0.2	0.2	0.2
Nov. 21	0.2	0.2	0.2
Nov. 22	0.2	0.2	0.2
Nov. 23	0.2	0.2	0.2
Nov. 24	0.2	0.2	0.2
Nov. 25	0.2	0.2	0.2
Nov. 26	0.2	0.2	0.2
Nov. 27	0.2	0.2	0.2
Nov. 28	0.2	0.2	0.2
Nov. 29	0.2	0.2	0.2
Nov. 30	0.2	0.2	0.2
Dec. 1	0.2	0.2	0.2
Dec. 2	0.2	0.2	0.2
Dec. 3	0.2	0.2	0.2
Dec. 4	0.2	0.2	0.2
Dec. 5	0.2	0.2	0.2
Dec. 6	0.2	0.2	0.2
Dec. 7	0.2	0.2	0.2
Dec. 8	0.2	0.2	0.2
Dec. 9	0.2	0.2	0.2
Dec. 10	0.2	0.2	0.2
Dec. 11	0.2	0.2	0.2
Dec. 12	0.2	0.2	0.2
Dec. 13	0.2	0.2	0.2
Dec. 14	0.2	0.2	0.2
Dec. 15	0.2	0.2	0.2
Dec. 16	0.2	0.2	0.2
Dec. 17	0.2	0.2	0.2
Dec. 18	0.2	0.2	0.2
Dec. 19	0.2	0.2	0.2
Dec. 20	0.2	0.2	0.2
Dec. 21	0.2	0.2	0.2
Dec. 22	0.2	0.2	0.2
Dec. 23	0.2	0.2	0.2
Dec. 24	0.2	0.2	0.2
Dec. 25	0.2	0.2	0.2
Dec. 26	0.2	0.2	0.2
Dec. 27	0.2	0.2	0.2
Dec. 28	0.2	0.2	0.2
Dec. 29	0.2	0.2	0.2
Dec. 30	0.2	0.2	0.2
Dec. 31	0.2	0.2	0.2

\* Values for Peaches were determined using the OMAFRA Best Management Practices - Irrigation Management guide for mature fruit trees with permanent sod and herbicide strip. Values for January to April, and November to December were based on the April coefficient, which is equal to that of bare soil. The values from the British Columbia Ministry of Agriculture, Food and Fisheries were used for October. The values used are an average of the coefficients indicated for the 3 regions in British Columbia.

‡The values for all months, excluding October, were determined from OMAFRA Best Management Practices - Irrigation Management guide for mature fruit trees with

permanent sod and herbicide strip. The values for October were determined based on the average values indicated by the 3 regions in British Columbia by the Ministry of Agriculture, Food and Fisheries.

<sup>a</sup>Ontario does not have published values for grapes. The BC Ministry of Agriculture, Food and Fisheries was used to determine these values from May to the end of October. The values are an average of the 3 identified regions in BC. The coefficient for bare soil was used for the off season months.

**Table 6:** Annual evapotranspiration coefficients for deciduous and coniferous forests, and idle land.

	Deciduous Forest	Kc <sup>b</sup>	Coniferous Forest	Kc <sup>b</sup>	Idle Land	Kc <sup>b</sup>
Jan. 1	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 2	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 3	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 4	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 5	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 6	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 7	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 8	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 9	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 10	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 11	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 12	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 13	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 14	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 15	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 16	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 17	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 18	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 19	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 20	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 21	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 22	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 23	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 24	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 25	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 26	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 27	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 28	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 29	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 30	Dormant	0.2	Off Season	0.2	Dormant	0.2
Jan. 31	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 1	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 2	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 3	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 4	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 5	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 6	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 7	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 8	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 9	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 10	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 11	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 12	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 13	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 14	Dormant	0.2	Off Season	0.2	Dormant	0.2



Feb. 15	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 16	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 17	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 18	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 19	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 20	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 21	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 22	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 23	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 24	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 25	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 26	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 27	Dormant	0.2	Off Season	0.2	Dormant	0.2
Feb. 28	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 1	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 2	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 3	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 4	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 5	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 6	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 7	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 8	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 9	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 10	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 11	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 12	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 13	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 14	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 15	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 16	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 17	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 18	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 19	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 20	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 21	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 22	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 23	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 24	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 25	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 26	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 27	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 28	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 29	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 30	Dormant	0.2	Off Season	0.2	Dormant	0.2
Mar. 31	Dormant	0.2	Off Season	0.2	Dormant	0.2
Apr. 1	Kmid	1.09	Kmid	0.55	Kmid	1.00
Apr. 2	Kmid	1.09	Kmid	0.55	Kmid	1.01
Apr. 3	Kmid	1.09	Kmid	0.55	Kmid	1.07
Apr. 4	Kmid	1.09	Kmid	0.55	Kmid	1.06
Apr. 5	Kmid	1.09	Kmid	0.55	Kmid	0.96
Apr. 6	Kmid	1.09	Kmid	0.55	Kmid	0.98
Apr. 7	Kmid	1.09	Kmid	0.55	Kmid	0.93
Apr. 8	Kmid	1.09	Kmid	0.55	Kmid	1.01
Apr. 9	Kmid	1.09	Kmid	0.55	Kmid	1.00
Apr. 10	Kmid	1.09	Kmid	0.55	Kmid	1.03
Apr. 11	Kmid	1.09	Kmid	0.55	Kmid	1.00
Apr. 12	Kmid	1.09	Kmid	0.55	Kmid	1.03
Apr. 13	Kmid	1.09	Kmid	0.55	Kmid	1.02
Apr. 14	Kmid	1.09	Kmid	0.55	Kmid	0.86

Apr. 15	Kmid	1.09	Kmid	0.55	Kmid	0.99
Apr. 16	Kmid	1.09	Kmid	0.55	Kmid	0.99
Apr. 17	Kmid	1.09	Kmid	0.55	Kmid	0.99
Apr. 18	Kmid	1.09	Kmid	0.55	Kmid	1.00
Apr. 19	Kmid	1.09	Kmid	0.55	Kmid	1.02
Apr. 20	Kmid	1.09	Kmid	0.55	Kmid	1.00
Apr. 21	Kmid	1.09	Kmid	0.55	Kmid	0.90
Apr. 22	Kmid	1.09	Kmid	0.55	Kmid	0.85
Apr. 23	Kmid	1.09	Kmid	0.55	Kmid	0.94
Apr. 24	Kmid	1.09	Kmid	0.55	Kmid	1.04
Apr. 25	Kmid	1.09	Kmid	0.55	Kmid	1.06
Apr. 26	Kmid	1.09	Kmid	0.55	Kmid	1.03
Apr. 27	Kmid	1.09	Kmid	0.55	Kmid	1.03
Apr. 28	Kmid	1.09	Kmid	0.55	Kmid	1.07
Apr. 29	Kmid	1.09	Kmid	0.55	Kmid	1.08
Apr. 30	Kmid	1.09	Kmid	0.55	Kmid	1.07
May 1	Kmid	1.09	Kmid	0.55	Kmid	1.04
May 2	Kmid	1.09	Kmid	0.55	Kmid	1.04
May 3	Kmid	1.09	Kmid	0.55	Kmid	1.00
May 4	Kmid	1.09	Kmid	0.55	Kmid	0.98
May 5	Kmid	1.09	Kmid	0.55	Kmid	1.03
May 6	Kmid	1.09	Kmid	0.55	Kmid	1.00
May 7	Kmid	1.09	Kmid	0.55	Kmid	1.04
May 8	Kmid	1.09	Kmid	0.55	Kmid	1.03
May 9	Kmid	1.09	Kmid	0.55	Kmid	1.04
May 10	Kmid	1.09	Kmid	0.55	Kmid	1.02
May 11	Kmid	1.09	Kmid	0.55	Kmid	0.97
May 12	Kmid	1.09	Kmid	0.55	Kmid	0.96
May 13	Kmid	1.09	Kmid	0.55	Kmid	0.97
May 14	Kmid	1.09	Kmid	0.55	Kmid	0.94
May 15	Kmid	1.09	Kmid	0.55	Kmid	0.91
May 16	Kmid	1.09	Kmid	0.55	Kmid	0.96
May 17	Kmid	1.09	Kmid	0.55	Kmid	0.94
May 18	Kmid	1.09	Kmid	0.55	Kmid	0.97
May 19	Kmid	1.09	Kmid	0.55	Kmid	1.05
May 20	Kmid	1.09	Kmid	0.55	Kmid	1.04
May 21	Kmid	1.09	Kmid	0.55	Kmid	1.02
May 22	Kmid	1.09	Kmid	0.55	Kmid	1.03
May 23	Kmid	1.09	Kmid	0.55	Kmid	1.03
May 24	Kmid	1.09	Kmid	0.55	Kmid	0.99
May 25	Kmid	1.09	Kmid	0.55	Kmid	0.88
May 26	Kmid	1.09	Kmid	0.55	Kmid	0.98
May 27	Kmid	1.09	Kmid	0.55	Kmid	0.97
May 28	Kmid	1.09	Kmid	0.55	Kmid	1.01
May 29	Kmid	1.09	Kmid	0.55	Kmid	0.99
May 30	Kmid	1.09	Kmid	0.55	Kmid	0.97
May 31	Kmid	1.09	Kmid	0.55	Kmid	0.88
June 1	Kmid	1.09	Kmid	0.55	Kmid	0.91
June 2	Kmid	1.09	Kmid	0.55	Kmid	0.87
June 3	Kmid	1.09	Kmid	0.55	Kmid	1.01
June 4	Kmid	1.09	Kmid	0.55	Kmid	1.01
June 5	Kmid	1.09	Kmid	0.55	Kmid	1.00
June 6	Kmid	1.09	Kmid	0.55	Kmid	1.00
June 7	Kmid	1.09	Kmid	0.55	Kmid	0.99
June 8	Kmid	1.09	Kmid	0.55	Kmid	0.98
June 9	Kmid	1.09	Kmid	0.55	Kmid	1.08
June 10	Kmid	1.09	Kmid	0.55	Kmid	1.00
June 11	Kmid	1.09	Kmid	0.55	Kmid	0.97
June 12	Kmid	1.09	Kmid	0.55	Kmid	0.98

June 13	Kmid	1.09	Kmid	0.55	Kmid	0.98
June 14	Kmid	1.09	Kmid	0.55	Kmid	1.03
June 15	Kmid	1.09	Kmid	0.55	Kmid	1.04
June 16	Kmid	1.09	Kmid	0.55	Kmid	1.06
June 17	Kmid	1.09	Kmid	0.55	Kmid	1.03
June 18	Kmid	1.09	Kmid	0.55	Kmid	0.97
June 19	Kmid	1.09	Kmid	0.55	Kmid	1.00
June 20	Kmid	1.09	Kmid	0.55	Kmid	0.97
June 21	Kmid	1.09	Kmid	0.55	Kmid	1.02
June 22	Kmid	1.09	Kmid	0.55	Kmid	0.95
June 23	Kmid	1.09	Kmid	0.55	Kmid	0.93
June 24	Kmid	1.09	Kmid	0.55	Kmid	0.97
June 25	Kmid	1.09	Kmid	0.55	Kmid	0.93
June 26	Kmid	1.09	Kmid	0.55	Kmid	0.89
June 27	Kmid	1.09	Kmid	0.55	Kmid	0.95
June 28	Kmid	1.09	Kmid	0.55	Kmid	0.96
June 29	Kmid	1.09	Kmid	0.55	Kmid	0.98
June 30	Kmid	1.09	Kmid	0.55	Kmid	1.04
July 1	Kmid	1.09	Kmid	0.55	Kmid	1.03
July 2	Kmid	1.09	Kmid	0.55	Kmid	0.97
July 3	Kmid	1.09	Kmid	0.55	Kmid	1.00
July 4	Kmid	1.09	Kmid	0.55	Kmid	1.00
July 5	Kmid	1.09	Kmid	0.55	Kmid	0.98
July 6	Kmid	1.09	Kmid	0.55	Kmid	1.01
July 7	Kmid	1.09	Kmid	0.55	Kmid	1.02
July 8	Kmid	1.09	Kmid	0.55	Kmid	1.03
July 9	Kmid	1.09	Kmid	0.55	Kmid	0.94
July 10	Kmid	1.09	Kmid	0.55	Kmid	0.92
July 11	Kmid	1.09	Kmid	0.55	Kmid	0.88
July 12	Kmid	1.09	Kmid	0.55	Kmid	0.96
July 13	Kmid	1.09	Kmid	0.55	Kmid	0.94
July 14	Kmid	1.09	Kmid	0.55	Kmid	0.97
July 15	Kmid	1.09	Kmid	0.55	Kmid	0.99
July 16	Kmid	1.09	Kmid	0.55	Kmid	1.01
July 17	Kmid	1.09	Kmid	0.55	Kmid	1.01
July 18	Kmid	1.09	Kmid	0.55	Kmid	0.92
July 19	Kmid	1.09	Kmid	0.55	Kmid	0.95
July 20	Kmid	1.09	Kmid	0.55	Kmid	0.92
July 21	Kmid	1.09	Kmid	0.55	Kmid	0.90
July 22	Kmid	1.09	Kmid	0.55	Kmid	0.94
July 23	Kmid	1.09	Kmid	0.55	Kmid	1.01
July 24	Kmid	1.09	Kmid	0.55	Kmid	0.99
July 25	Kmid	1.09	Kmid	0.55	Kmid	1.01
July 26	Kmid	1.09	Kmid	0.55	Kmid	0.96
July 27	Kmid	1.09	Kmid	0.55	Kmid	0.94
July 28	Kmid	1.09	Kmid	0.55	Kmid	0.95
July 29	Kmid	1.09	Kmid	0.55	Kmid	0.91
July 30	Kmid	1.09	Kmid	0.55	Kmid	0.95
July 31	Kmid	1.09	Kmid	0.55	Kmid	0.99
Aug. 1	Kmid	1.09	Kmid	0.55	Kmid	1.01
Aug. 2	Kmid	1.09	Kmid	0.55	Kmid	0.90
Aug. 3	Kmid	1.09	Kmid	0.55	Kmid	1.00
Aug. 4	Kmid	1.09	Kmid	0.55	Kmid	0.95
Aug. 5	Kmid	1.09	Kmid	0.55	Kmid	0.98
Aug. 6	Kmid	1.09	Kmid	0.55	Kmid	1.00
Aug. 7	Kmid	1.09	Kmid	0.55	Kmid	1.02
Aug. 8	Kmid	1.09	Kmid	0.55	Kmid	1.01
Aug. 9	Kmid	1.09	Kmid	0.55	Kmid	0.98
Aug. 10	Kmid	1.09	Kmid	0.55	Kmid	1.01

Aug. 11	Kmid	1.09	Kmid	0.55	Kmid	1.01
Aug. 12	Kmid	1.09	Kmid	0.55	Kmid	1.01
Aug. 13	Kmid	1.09	Kmid	0.55	Kmid	0.97
Aug. 14	Kmid	1.09	Kmid	0.55	Kmid	1.02
Aug. 15	Kmid	1.09	Kmid	0.55	Kmid	0.98
Aug. 16	Kmid	1.09	Kmid	0.55	Kmid	1.00
Aug. 17	Kmid	1.09	Kmid	0.55	Kmid	0.96
Aug. 18	Kmid	1.09	Kmid	0.55	Kmid	0.88
Aug. 19	Kmid	1.09	Kmid	0.55	Kmid	0.93
Aug. 20	Kmid	1.09	Kmid	0.55	Kmid	1.00
Aug. 21	Kmid	1.09	Kmid	0.55	Kmid	0.99
Aug. 22	Kmid	1.09	Kmid	0.55	Kmid	1.01
Aug. 23	Kmid	1.09	Kmid	0.55	Kmid	0.95
Aug. 24	Kmid	1.09	Kmid	0.55	Kmid	0.91
Aug. 25	Kmid	1.09	Kmid	0.55	Kmid	0.92
Aug. 26	Kmid	1.09	Kmid	0.55	Kmid	0.92
Aug. 27	Kmid	1.09	Kmid	0.55	Kmid	0.91
Aug. 28	Kmid	1.09	Kmid	0.55	Kmid	0.93
Aug. 29	Kmid	1.09	Kmid	0.55	Kmid	0.98
Aug. 30	Kmid	1.09	Kmid	0.55	Kmid	0.97
Aug. 31	Kmid	1.09	Kmid	0.55	Kmid	0.97
Sept. 1	Kmid	1.09	Kmid	0.55	Kmid	0.93
Sept. 2	Kmid	1.09	Kmid	0.55	Kmid	0.90
Sept. 3	Kmid	1.09	Kmid	0.55	Kmid	0.92
Sept. 4	Kmid	1.09	Kmid	0.55	Kmid	0.88
Sept. 5	Kmid	1.09	Kmid	0.55	Kmid	0.92
Sept. 6	Kmid	1.09	Kmid	0.55	Kmid	0.95
Sept. 7	Kmid	1.09	Kmid	0.55	Kmid	0.99
Sept. 8	Kmid	1.09	Kmid	0.55	Kmid	0.91
Sept. 9	Kmid	1.09	Kmid	0.55	Kmid	0.98
Sept. 10	Kmid	1.09	Kmid	0.55	Kmid	0.98
Sept. 11	Kmid	1.09	Kmid	0.55	Kmid	0.90
Sept. 12	Kmid	1.09	Kmid	0.55	Kmid	0.91
Sept. 13	Kmid	1.09	Kmid	0.55	Kmid	0.88
Sept. 14	Kmid	1.09	Kmid	0.55	Kmid	0.88
Sept. 15	Kmid	1.09	Kmid	0.55	Kmid	0.87
Sept. 16	Kmid	1.09	Kmid	0.55	Kmid	0.94
Sept. 17	Kmid	1.09	Kmid	0.55	Kmid	1.00
Sept. 18	Kmid	1.09	Kmid	0.55	Kmid	0.98
Sept. 19	Kmid	1.09	Kmid	0.55	Kmid	1.00
Sept. 20	Kmid	1.09	Kmid	0.55	Kmid	0.97
Sept. 21	Kmid	1.09	Kmid	0.55	Kmid	0.92
Sept. 22	Kmid	1.09	Kmid	0.55	Kmid	0.93
Sept. 23	Kmid	1.09	Kmid	0.55	Kmid	1.01
Sept. 24	Kmid	1.09	Kmid	0.55	Kmid	1.00
Sept. 25	Kmid	1.09	Kmid	0.55	Kmid	0.96
Sept. 26	Kmid	1.09	Kmid	0.55	Kmid	1.01
Sept. 27	Kmid	1.09	Kmid	0.55	Kmid	0.92
Sept. 28	Kmid	1.09	Kmid	0.55	Kmid	0.94
Sept. 29	Kmid	1.09	Kmid	0.55	Kmid	0.95
Sept. 30	Kmid	1.09	Kmid	0.55	Kmid	0.95
Oct. 1	Kmid	1.09	Kmid	0.55	Kmid	0.95
Oct. 2	Kmid	1.09	Kmid	0.55	Kmid	0.91
Oct. 3	Kmid	1.09	Kmid	0.55	Kmid	0.98
Oct. 4	Kmid	1.09	Kmid	0.55	Kmid	0.98
Oct. 5	Kmid	1.09	Kmid	0.55	Kmid	0.93
Oct. 6	Kmid	1.09	Kmid	0.55	Kmid	0.95
Oct. 7	Kmid	1.09	Kmid	0.55	Kmid	0.92
Oct. 8	Kmid	1.09	Kmid	0.55	Kmid	0.95

Oct. 9	Kmid	1.09	Kmid	0.55	Kmid	0.94
Oct. 10	Kmid	1.09	Kmid	0.55	Kmid	0.93
Oct. 11	Kmid	1.09	Kmid	0.55	Kmid	0.98
Oct. 12	Kmid	1.09	Kmid	0.55	Kmid	1.01
Oct. 13	Kmid	1.09	Kmid	0.55	Kmid	1.00
Oct. 14	Kmid	1.09	Kmid	0.55	Kmid	1.01
Oct. 15	Kmid	1.09	Kmid	0.55	Kmid	0.99
Oct. 16	Kmid	1.09	Kmid	0.55	Kmid	0.98
Oct. 17	Kmid	1.09	Kmid	0.55	Kmid	0.94
Oct. 18	Kmid	1.09	Kmid	0.55	Kmid	0.90
Oct. 19	Kmid	1.09	Kmid	0.55	Kmid	0.93
Oct. 20	Kmid	1.09	Kmid	0.55	Kmid	0.91
Oct. 21	Kmid	1.09	Kmid	0.55	Kmid	0.92
Oct. 22	Kmid	1.09	Kmid	0.55	Kmid	1.00
Oct. 23	Kmid	1.09	Kmid	0.55	Kmid	0.97
Oct. 24	Kmid	1.09	Kmid	0.55	Kmid	0.99
Oct. 25	Kmid	1.09	Kmid	0.55	Kmid	0.95
Oct. 26	Kmid	1.09	Kmid	0.55	Kmid	0.92
Oct. 27	Kmid	1.09	Kmid	0.55	Kmid	0.95
Oct. 28	Kmid	1.09	Kmid	0.55	Kmid	1.07
Oct. 29	Kmid	1.09	Kmid	0.55	Kmid	1.01
Oct. 30	Kmid	1.09	Kmid	0.55	Kmid	1.03
Oct. 31	Kmid	1.09	Kmid	0.55	Kmid	0.99
Nov. 1	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 2	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 3	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 4	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 5	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 6	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 7	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 8	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 9	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 10	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 11	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 12	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 13	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 14	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 15	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 16	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 17	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 18	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 19	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 20	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 21	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 22	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 23	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 24	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 25	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 26	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 27	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 28	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 29	Dormant	0.2	Off Season	0.2	Dormant	0.2
Nov. 30	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 1	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 2	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 3	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 4	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 5	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 6	Dormant	0.2	Off Season	0.2	Dormant	0.2

Dec. 7	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 8	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 9	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 10	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 11	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 12	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 13	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 14	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 15	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 16	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 17	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 18	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 19	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 20	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 21	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 22	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 23	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 24	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 25	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 26	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 27	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 28	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 29	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 30	Dormant	0.2	Off Season	0.2	Dormant	0.2
Dec. 31	Dormant	0.2	Off Season	0.2	Dormant	0.2

<sup>b</sup>Based on the information available regarding the evapotranspiration of deciduous and coniferous forest, and idle land, step-wise changes in values were used. A growing season of April 1 to Oct. 31 was estimated in order to include bud development as well as killing frost at the end of the season. Should these values not be appropriate for the region of concern or vary seasonally, they should be adjusted accordingly.

**Table 7:** Annual evapotranspiration coefficients for tallgrass, built-up pervious areas, and open water.

	Tallgrass <sup>c</sup>	Kc <sup>d</sup>	Built-Up Pervious Area	Kc <sup>d</sup>	Open Water (Shallow)	Open Water (>5m depth) <sup>e</sup>
Jan. 1	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 2	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 3	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 4	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 5	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 6	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 7	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 8	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 9	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 10	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 11	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 12	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 13	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 14	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 15	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 16	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 17	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 18	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 19	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 20	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 21	Dormant	0.20	Reference	1.0	1.05	1.25
Jan. 22	Dormant	0.20	Reference	1.0	1.05	1.25

[illegible]

Mar. 23	Kini	0.36	Reference	1.0	1.05	1.25
Mar. 24	Kini	0.36	Reference	1.0	1.05	1.25
Mar. 25	Kini	0.36	Reference	1.0	1.05	1.25
Mar. 26	Kini	0.37	Reference	1.0	1.05	1.25
Mar. 27	Kini	0.38	Reference	1.0	1.05	1.25
Mar. 28	Kini	0.37	Reference	1.0	1.05	1.25
Mar. 29	Kini	0.37	Reference	1.0	1.05	1.25
Mar. 30	Kini	0.37	Reference	1.0	1.05	1.25
Mar. 31	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 1	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 2	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 3	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 4	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 5	Kini	0.36	Reference	1.0	1.05	1.25
Apr. 6	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 7	Kini	0.36	Reference	1.0	1.05	1.25
Apr. 8	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 9	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 10	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 11	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 12	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 13	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 14	Kini	0.35	Reference	1.0	1.05	1.25
Apr. 15	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 16	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 17	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 18	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 19	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 20	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 21	Kini	0.37	Reference	1.0	1.05	1.25
Apr. 22	Kini	0.36	Reference	1.0	1.05	1.25
Apr. 23	Kini	0.35	Reference	1.0	1.05	1.25
Apr. 24	Kini	0.36	Reference	1.0	1.05	1.25
Apr. 25	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 26	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 27	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 28	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 29	Kini	0.38	Reference	1.0	1.05	1.25
Apr. 30	Kini	0.38	Reference	1.0	1.05	1.25
May 1	Kini	0.38	Reference	1.0	1.05	0.65
May 2	Kini	0.38	Reference	1.0	1.05	0.65
May 3	Kini	0.38	Reference	1.0	1.05	0.65
May 4	Kini	0.37	Reference	1.0	1.05	0.65
May 5	Kini	0.37	Reference	1.0	1.05	0.65
May 6	Kini	0.38	Reference	1.0	1.05	0.65
May 7	Kini	0.37	Reference	1.0	1.05	0.65
May 8	Kini	0.38	Reference	1.0	1.05	0.65
May 9	Kini	0.38	Reference	1.0	1.05	0.65
May 10	Kini	0.38	Reference	1.0	1.05	0.65
May 11	Kini	0.37	Reference	1.0	1.05	0.65
May 12	Kini	0.37	Reference	1.0	1.05	0.65
May 13	Kini	0.36	Reference	1.0	1.05	0.65
May 14	Kini	0.37	Reference	1.0	1.05	0.65
May 15	Kdev	0.85	Reference	1.0	1.05	0.65
May 16	Kdev	0.83	Reference	1.0	1.05	0.65
May 17	Kdev	0.86	Reference	1.0	1.05	0.65
May 18	Kdev	0.85	Reference	1.0	1.05	0.65
May 19	Kdev	0.86	Reference	1.0	1.05	0.65
May 20	Kdev	0.91	Reference	1.0	1.05	0.65



May 21	Kdev	0.91	Reference	1.0	1.05	0.65
May 22	Kdev	0.90	Reference	1.0	1.05	0.65
May 23	Kdev	0.90	Reference	1.0	1.05	0.65
May 24	Kdev	0.90	Reference	1.0	1.05	0.65
May 25	Kdev	0.87	Reference	1.0	1.05	0.65
May 26	Kdev	0.81	Reference	1.0	1.05	0.65
May 27	Kdev	0.87	Reference	1.0	1.05	0.65
May 28	Kdev	0.87	Reference	1.0	1.05	0.65
May 29	Kdev	0.89	Reference	1.0	1.05	0.65
May 30	Kdev	0.88	Reference	1.0	1.05	0.65
May 31	Kdev	0.87	Reference	1.0	1.05	0.65
June 1	Kmid	1.19	Reference	1.0	1.05	0.65
June 2	Kmid	1.22	Reference	1.0	1.05	0.65
June 3	Kmid	1.18	Reference	1.0	1.05	0.65
June 4	Kmid	1.32	Reference	1.0	1.05	0.65
June 5	Kmid	1.32	Reference	1.0	1.05	0.65
June 6	Kmid	1.31	Reference	1.0	1.05	0.65
June 7	Kmid	1.31	Reference	1.0	1.05	0.65
June 8	Kmid	1.30	Reference	1.0	1.05	0.65
June 9	Kmid	1.29	Reference	1.0	1.05	0.65
June 10	Kmid	1.38	Reference	1.0	1.05	0.65
June 11	Kmid	1.31	Reference	1.0	1.05	0.65
June 12	Kmid	1.28	Reference	1.0	1.05	0.65
June 13	Kmid	1.29	Reference	1.0	1.05	0.65
June 14	Kmid	1.29	Reference	1.0	1.05	0.65
June 15	Kmid	1.33	Reference	1.0	1.05	0.65
June 16	Kmid	1.34	Reference	1.0	1.05	0.65
June 17	Kmid	1.36	Reference	1.0	1.05	0.65
June 18	Kmid	1.33	Reference	1.0	1.05	0.65
June 19	Kmid	1.28	Reference	1.0	1.05	0.65
June 20	Kmid	1.31	Reference	1.0	1.05	0.65
June 21	Kmid	1.28	Reference	1.0	1.05	0.65
June 22	Kmid	1.33	Reference	1.0	1.05	0.65
June 23	Kmid	1.26	Reference	1.0	1.05	0.65
June 24	Kmid	1.23	Reference	1.0	1.05	0.65
June 25	Kmid	1.27	Reference	1.0	1.05	0.65
June 26	Kmid	1.24	Reference	1.0	1.05	0.65
June 27	Kmid	1.20	Reference	1.0	1.05	0.65
June 28	Kmid	1.26	Reference	1.0	1.05	0.65
June 29	Kmid	1.26	Reference	1.0	1.05	0.65
June 30	Kmid	1.29	Reference	1.0	1.05	0.65
July 1	Kmid	1.34	Reference	1.0	1.05	0.65
July 2	Kmid	1.33	Reference	1.0	1.05	0.65
July 3	Kmid	1.28	Reference	1.0	1.05	0.65
July 4	Kmid	1.30	Reference	1.0	1.05	0.65
July 5	Kmid	1.31	Reference	1.0	1.05	0.65
July 6	Kmid	1.29	Reference	1.0	1.05	0.65
July 7	Kmid	1.32	Reference	1.0	1.05	0.65
July 8	Kmid	1.32	Reference	1.0	1.05	0.65
July 9	Kmid	1.34	Reference	1.0	1.05	0.65
July 10	Kmid	1.25	Reference	1.0	1.05	0.65
July 11	Kmid	1.23	Reference	1.0	1.05	0.65
July 12	Kmid	1.19	Reference	1.0	1.05	0.65
July 13	Kmid	1.26	Reference	1.0	1.05	0.65
July 14	Kmid	1.25	Reference	1.0	1.05	0.65
July 15	Kmid	1.27	Reference	1.0	1.05	0.65
July 16	Kmid	1.29	Reference	1.0	1.05	0.65
July 17	Kmid	1.32	Reference	1.0	1.05	0.65
July 18	Kmid	1.32	Reference	1.0	1.05	0.65

July 19	Kmid	1.23	Reference	1.0	1.05	0.65
July 20	Kmid	1.26	Reference	1.0	1.05	0.65
July 21	Kmid	1.23	Reference	1.0	1.05	0.65
July 22	Kmid	1.21	Reference	1.0	1.05	0.65
July 23	Kmid	1.25	Reference	1.0	1.05	0.65
July 24	Kmid	1.32	Reference	1.0	1.05	0.65
July 25	Kmid	1.29	Reference	1.0	1.05	0.65
July 26	Kmid	1.32	Reference	1.0	1.05	0.65
July 27	Kmid	1.27	Reference	1.0	1.05	0.65
July 28	Kmid	1.25	Reference	1.0	1.05	0.65
July 29	Kmid	1.26	Reference	1.0	1.05	0.65
July 30	Kmid	1.22	Reference	1.0	1.05	0.65
July 31	Kmid	1.26	Reference	1.0	1.05	0.65
Aug. 1	Kmid	1.29	Reference	1.0	1.05	0.65
Aug. 2	Kmid	1.32	Reference	1.0	1.05	0.65
Aug. 3	Kmid	1.21	Reference	1.0	1.05	0.65
Aug. 4	Kmid	1.30	Reference	1.0	1.05	0.65
Aug. 5	Kmid	1.26	Reference	1.0	1.05	0.65
Aug. 6	Kmid	1.29	Reference	1.0	1.05	0.65
Aug. 7	Kmid	1.31	Reference	1.0	1.05	0.65
Aug. 8	Kmid	1.33	Reference	1.0	1.05	0.65
Aug. 9	Kmid	1.32	Reference	1.0	1.05	0.65
Aug. 10	Kmid	1.29	Reference	1.0	1.05	0.65
Aug. 11	Kmid	1.31	Reference	1.0	1.05	0.65
Aug. 12	Kmid	1.32	Reference	1.0	1.05	0.65
Aug. 13	Kmid	1.31	Reference	1.0	1.05	0.65
Aug. 14	Kmid	1.27	Reference	1.0	1.05	0.65
Aug. 15	Kmid	1.32	Reference	1.0	1.05	0.65
Aug. 16	Kmid	1.29	Reference	1.0	1.05	0.65
Aug. 17	Kmid	1.30	Reference	1.0	1.05	0.65
Aug. 18	Kmid	1.26	Reference	1.0	1.05	0.65
Aug. 19	Kmid	1.19	Reference	1.0	1.05	0.65
Aug. 20	Kmid	1.24	Reference	1.0	1.05	0.65
Aug. 21	Kmid	1.30	Reference	1.0	1.05	0.65
Aug. 22	Kmid	1.30	Reference	1.0	1.05	0.65
Aug. 23	Kmid	1.31	Reference	1.0	1.05	0.65
Aug. 24	Kmid	1.26	Reference	1.0	1.05	0.65
Aug. 25	Kmid	1.22	Reference	1.0	1.05	0.65
Aug. 26	Kmid	1.23	Reference	1.0	1.05	0.65
Aug. 27	Kmid	1.23	Reference	1.0	1.05	0.65
Aug. 28	Kmid	1.22	Reference	1.0	1.05	0.65
Aug. 29	Kmid	1.24	Reference	1.0	1.05	0.65
Aug. 30	Kmid	1.28	Reference	1.0	1.05	0.65
Aug. 31	Kmid	1.27	Reference	1.0	1.05	0.65
Sept. 1	Klate	0.86	Reference	1.0	1.05	0.65
Sept. 2	Klate	0.84	Reference	1.0	1.05	0.65
Sept. 3	Klate	0.82	Reference	1.0	1.05	0.65
Sept. 4	Klate	0.84	Reference	1.0	1.05	0.65
Sept. 5	Klate	0.81	Reference	1.0	1.05	0.65
Sept. 6	Klate	0.84	Reference	1.0	1.05	0.65
Sept. 7	Klate	0.85	Reference	1.0	1.05	0.65
Sept. 8	Klate	0.87	Reference	1.0	1.05	0.65
Sept. 9	Klate	0.83	Reference	1.0	1.05	0.65
Sept. 10	Klate	0.87	Reference	1.0	1.05	0.65
Sept. 11	Klate	0.87	Reference	1.0	1.05	0.65
Sept. 12	Klate	0.82	Reference	1.0	1.05	0.65
Sept. 13	Klate	0.83	Reference	1.0	1.05	0.65
Sept. 14	Klate	0.81	Reference	1.0	1.05	0.65
Sept. 15	Klate	0.81	Reference	1.0	1.05	0.65

Sept. 16	Klate	0.80	Reference	1.0	1.05	0.65
Sept. 17	Klate	0.85	Reference	1.0	1.05	0.65
Sept. 18	Klate	0.88	Reference	1.0	1.05	0.65
Sept. 19	Klate	0.87	Reference	1.0	1.05	0.65
Sept. 20	Klate	0.88	Reference	1.0	1.05	0.65
Sept. 21	Klate	0.87	Reference	1.0	1.05	0.65
Sept. 22	Klate	0.84	Reference	1.0	1.05	0.65
Sept. 23	Klate	0.84	Reference	1.0	1.05	0.65
Sept. 24	Klate	0.89	Reference	1.0	1.05	0.65
Sept. 25	Klate	0.88	Reference	1.0	1.05	0.65
Sept. 26	Klate	0.86	Reference	1.0	1.05	0.65
Sept. 27	Klate	0.89	Reference	1.0	1.05	0.65
Sept. 28	Klate	0.83	Reference	1.0	1.05	0.65
Sept. 29	Klate	0.85	Reference	1.0	1.05	0.65
Sept. 30	Klate	0.85	Reference	1.0	1.05	0.65
Oct. 1	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 2	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 3	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 4	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 5	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 6	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 7	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 8	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 9	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 10	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 11	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 12	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 13	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 14	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 15	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 16	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 17	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 18	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 19	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 20	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 21	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 22	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 23	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 24	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 25	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 26	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 27	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 28	Klate	0.36	Reference	1.0	1.05	1.25
Oct. 29	Klate	0.38	Reference	1.0	1.05	1.25
Oct. 30	Klate	0.37	Reference	1.0	1.05	1.25
Oct. 31	Klate	0.38	Reference	1.0	1.05	1.25
Nov. 1	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 2	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 3	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 4	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 5	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 6	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 7	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 8	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 9	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 10	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 11	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 12	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 13	Dormant	0.20	Reference	1.0	1.05	1.25

Nov. 14	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 15	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 16	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 17	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 18	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 19	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 20	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 21	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 22	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 23	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 24	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 25	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 26	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 27	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 28	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 29	Dormant	0.20	Reference	1.0	1.05	1.25
Nov. 30	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 1	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 2	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 3	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 4	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 5	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 6	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 7	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 8	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 9	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 10	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 11	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 12	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 13	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 14	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 15	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 16	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 17	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 18	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 19	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 20	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 21	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 22	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 23	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 24	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 25	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 26	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 27	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 28	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 29	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 30	Dormant	0.20	Reference	1.0	1.05	1.25
Dec. 31	Dormant	0.20	Reference	1.0	1.05	1.25

<sup>c</sup>A LAI of 0 was used from January 1 to March 14, 0.2 from March 15 to May 15, 1.05 from May 15 to May 31, and 2.65 for June, July and August. An LAI of 1.05 was used from Sept. 1 to Sept. 30, and 0.2 from Oct. 1 to Oct. 31; 0 was used for the remainder of the year. The crop height used was 1.5m.

<sup>d</sup>Based on the information available regarding the evapotranspiration of deciduous and coniferous forest, and idle land, step-wise changes in values were used.

<sup>e</sup>The dates chosen for the change in values for open water >5m are an estimated time as to when the average daily temperature exceeds water temperature (May 1) and when the

average daily temperature is below water temperature (Oct. 1). These values should be adjusted should the conditions vary by the season or by region.

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Disclaimer: The evapotranspiration coefficients indicated in this report are estimated from published literature produced in various climatic regions. These regions do not necessarily reflect the situations found in the Niagara Peninsula. The reported values are based on “well-watered” soil conditions and dry plant canopies. Information on dormant season (winter) ET is very limited. Any errors in the published literature may be reflected in the values presented in this report. The evapotranspiration coefficients reported are the best estimates available, but they should be used with the full recognition of these limitations.